

**GEOPHYSICAL INVERSION IN AN INTEGRATED EXPLORATION
PROGRAM: EXAMPLES FROM THE SAN NICOLÁS DEPOSIT**

By

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Abstract

The recent ability to produce three-dimensional physical property models of the subsurface from surface geophysical data, coupled with an increasing need to explore for minerals in concealed terranes, results in geophysical inversions providing more significant information to the exploration team. This thesis examines the role that geophysical inversion can play in an integrated mineral exploration program, and the impact it can have on the results. As an example, geophysical data from the San Nicolás copper-zinc massive sulphide deposit in Mexico are inverted. Such deposits are distinguished by high density, magnetic susceptibility, conductivity, and chargeability values.

Within the framework of an integrated exploration program (which is communicated through the use of a flowchart) many different, and hard-to-interpret, geophysical data sets are given geologic context. This is achieved by interpreting physical property models that have been generated by inversion modeling. The aim of generating these models, and interpreting geological information from them, is to: 1) assist mineral exploration in the area around the deposit, and roughly define areas of mineralization that can be used as a starting point for more detailed modeling ; 2) define the size and location of the San Nicolás deposit; 3) further improve the models and delineate ore more accurately by combining the detailed modeling with core physical property measurements; and 4) determine how best to use the information that has been acquired through modeling to find additional sulphide ore-bodies, including those that may be deeper than the existing deposit.

Density and magnetic susceptibility distribution models, inverted from regional gravity and magnetic data respectively, define large-scale structures that reflect the tectonic

setting of the region. Several distinct anomalies that exhibit high density and magnetic susceptibility values are identified. Since massive sulphides are often dense and magnetic, a correlation method is employed that determines volumes that have high density and magnetic susceptibility. The correlation procedure isolates five anomalies. Two of these are easily dismissed as having poor exploration potential, and two of the remaining anomalies, one of which is the San Nicolás deposit, are associated with mineralization.

At a more detailed scale, the deposit is well defined by gravity, magnetic, CSAMT, and IP methods individually. However, a drill hole that is targeted on the intersection of these favorable physical property distributions would have intersected the heart of the deposit. This demonstrates the advantages of using these methods in concert. Other sources of information, such as core physical property measurements and geologic constraints, are also used to improve modeling results. The inclusion of data from a single drill-hole is shown to significantly enhance detailed physical property distributions, and produces models that correlate better with known mineralization.

Finally, forward modeling of the physical property models is used to demonstrate how deep ore can be detected by implementing different survey designs that increase signal-to-noise of the data. This newly acquired information can then be used in the next step of the exploration program as the search for mineralization continues in the surrounding area.

In addition to the main focus of the thesis, it was found that polarizable material is needed in order to fit time-domain, airborne electromagnetic data collected over the deposit. This demonstrates the potential for detecting chargeable bodies from the air, and has significant implications for future mineral exploration.

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Dedication

I would like to dedicate this thesis to my family who support me in exploring the riches of the world.

Chapter 1

Introduction

Geophysical techniques can be used to detect and delineate mineral deposits. This has become more evident in recent years for the following two reasons:

1. The recently developed ability to invert geophysical data to recover three-dimensional physical property models.
2. The increasing need to explore in concealed terranes.

As a result, geophysics is providing a larger amount of information than in the past, which can impact an exploration program. At this time, it is important to re-evaluate how geophysical inversion can be employed in an integrated exploration program that includes geological, geophysical, geochemical, and remote sensing components. The possible role of geophysics in an integrated exploration program is outlined by a flowchart in Figure 1.1. The aim of this thesis is to demonstrate the use of geophysical inversion techniques at various stages of the exploration program. As an example, the results of inverting geophysical data from the San Nicolás copper and zinc deposit, in Mexico, are presented.

An exploration methodology is presented that outlines the steps that might be taken in an exploration program. Within this larger framework, the role of geophysics is defined. The geophysical role is an integrated part of the program with geologic and physical property information being essential for the optimal use of geophysics. This geophysical exploration methodology is followed throughout the thesis from defining an exploration model to delineating the deposit.

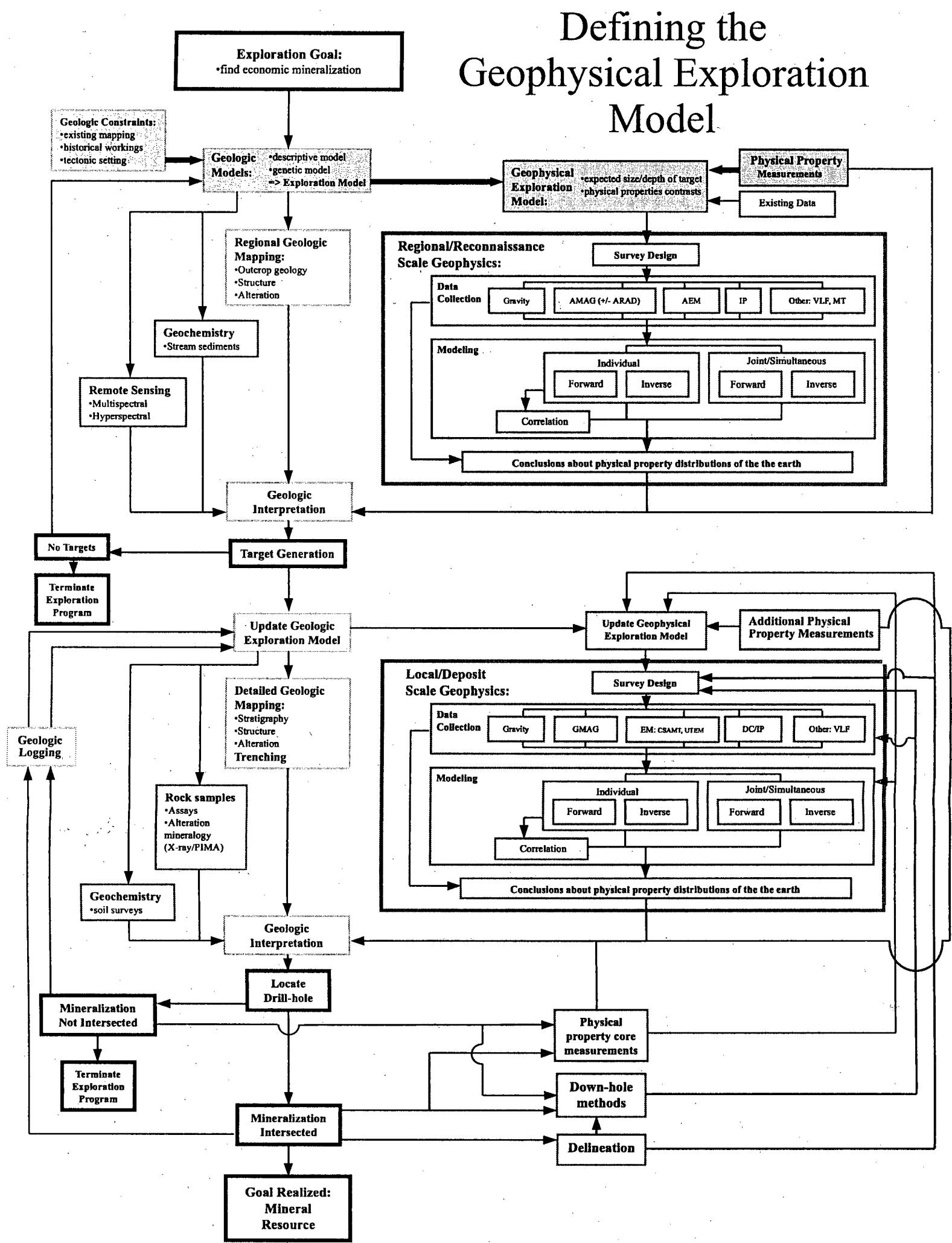


Figure 1.1: Geophysical Exploration Methodology Flowchart: Defining the geophysical exploration model.

In this thesis, sophisticated inversion techniques are used to generate digital, physical property models of the subsurface from geophysical data. These techniques are implemented at the regional scale to identify additional areas that might host mineralization, and at the deposit scale in order to define the deposit and delineate mineralization. Physical property measurements made on drill-core have been integrated in the inversion modeling at the delineation stage. The models are more easily interpreted than the data, and can be related to geology using physical property measurements of local rocks. Furthermore, different inversion models are combined to establish a more comprehensive understanding of the subsurface. Interpretations of the inverted models are proposed based on the geologic information presented in the thesis.

In addition, the San Nicolás deposit is almost entirely bounded to the east by a southwest-dipping fault, which could have been a feeder structure. Mineralization continues to follow the fault at depth in an unconstrained part of the deposit called the "keel". Modeling has been performed to demonstrate the sensitivity of different survey designs to this deep high-grade ore-zone.

The remainder of this chapter introduces: (1) the exploration methodology that is followed throughout this thesis, (2) the geologic background of San Nicolás, which is used to produce a geologic model, (3) the physical properties of rock units at San Nicolás, (4) the geophysical exploration model, (5) the data that have been collected, and (6) the tools that are used to model the geophysical data.

1.1 Exploration methodology

Mineral exploration programs commonly employ a staged approach when assessing large amounts of land or data. The aim of this approach is to systematically discern the most

prospective targets from a large amount of land without allowing any economic mineralization to be overlooked. These targets can then be evaluated with a local exploration program and drill-testing.

The flowchart in Figure 1.1 starts with an exploration goal and proceeds with defining a geologic exploration model on which a regional scale exploration program is based. The geologic exploration model provides exploration criteria, which are an estimate of the essential controls on mineralization (favorable stratigraphy, association with faults, etc.). These criteria are based on descriptive models (including mineralogy) and possibly genetic models that provide information about how the deposit might have formed. Also the model can be defined from knowledge about existing deposits in the area, and their tectonic setting.

Geophysical exploration methodology requires a geophysical exploration model. This builds on the geologic model and, with the inclusion of physical property information, defines geophysical exploration criteria. The criteria provide the expected physical properties of the mineralization itself, or physical properties of critical geologic features associated with mineralization.

At the regional scale the aim is to identify smaller, more focused, areas of interest. Different geophysical techniques can be used with survey designs that are tailored to the geophysical exploration model. Modeling of the geophysical data can produce inferences about the physical property distribution of the earth that, along with physical property information, will aid in regional geologic interpretations. The regional program may result in generation of targets that warrant more detailed investigation.

As shown in the flowchart, target generation is followed by the local- or deposit-scale exploration program. The local geophysical program will require different survey techniques and designs to those used at the regional scale, and this will be based on an updated geophysical exploration model. At this stage, geophysical inversion can play a

role in defining more detailed aspects of the subsurface, such as the depth and size of a body of favorable physical properties, or the geometry of deposit-critical geologic features. Along with physical property information of the local rocks, geophysical inversion models contribute to local-scale geologic interpretation. This geologic interpretation can result in the spotting of drill-holes.

The iterative nature of an integrated exploration program can continue during drilling. Information from drill-holes or drill-cores can be used to further refine geologic and geophysical exploration models. This will result in better inversion modeling of the physical property distributions. Subsequently, these models can better guide drilling, which could result in defining mineralization with fewer drill-holes. Furthermore, surveys can be designed to delineate specific areas of mineralization that may have been intersected, but not completely defined by drilling.

This exploration model, while only outlining two stages of the exploration process, allows for several iterations and further focusing of a program through the use of *feedback* paths.

Within the flowchart, emphasis is placed on how geophysics and inversion is integrated into such a program, but this is not meant to reflect a reduced importance of the other methods. It is also acknowledged that the exploration model may be over-simplified; however, it does demonstrate how geophysics has been used at San Nicolás.

1.2 San Nicolás

San Nicolás, owned by Teck Corporation and Western Copper Holdings Ltd., is an unmined, volcanic-hosted, massive sulphide deposit containing ore-grade copper and zinc with associated gold and silver. It is located in Zacatecas State, Mexico (Figure 1.2), a region that, until the discovery of San Nicolás and the nearby El Salvador deposit,

was unrecognized as having potential to host such deposits [Johnson et al., 2000]¹. The reserve estimates are 72 million tonnes grading 1.4% copper and 2.27% zinc, making it the largest massive sulphide deposit yet discovered in Mexico, and a world class deposit.

1.2.1 Geology

Discovery of economic ore is of primary importance. Although we hope to detect massive sulphide mineralization using geophysics directly, the geologic environment must be considered in order to effectively use geophysical information. An understanding of the regional geology can focus exploration efforts in more favorable structural and stratigraphic settings, thus expediting the whole exploration process. The association of mineralization with geologic features that could be identified using geophysics, such as: favorable horizons, plutons, faults, shallow intrusions, rhyolite domes, and alteration zoning, aid in producing informed exploration decisions. Discrimination of geologic features that could be mistaken for sulphide ore is also an important part of the exploration process. Furthermore, knowledge of the local geology, and the associated physical property values, allow for expected geophysical contrasts to be determined, and appropriate geophysical methods to be applied. The geophysical data that have been collected is not only used for direct detection of massive sulphide ore, but for providing information that can help form a geologic model of the area.

A background to the geologic setting and deposit genesis is given below, while more detailed regional- and local-scale geologic information is given at the beginning of Chapters 3 and 4 respectively.

¹The area has been known mostly for the historically worked, epithermal silver and gold deposits.



Figure 1.2: Map of Mexico (upper panel) and Zacatecas State (lower panel) showing the location of San Nicolás.

Geologic setting

The paleotectonic regime in which the San Nicolás deposit formed is a transition from a back-arc basin to an arc, and possibly a fore-arc setting², as documented by Danielson [Danielson, 2000]. The deposit is hosted in marine volcanic and sedimentary rocks, locally known as the Chilitos Formation, of Upper Jurassic to Lower Cretaceous age, [Johnson et al., 2000]. Mineralization is predominately found along a *time-line* that occurs between the formation of rhyolitic lava domes and the deposition of mafic extrusives above. These rocks include: mafic, intermediate, and rhyolite flows and shallow intrusions, felsic tuffs, volcanoclastic sediments, mudstone, chert, and limestone.

This package of rocks was subsequently accreted to North America; possibly in the mid-Cretaceous. Northeast-directed thrusting of the Late Cretaceous to Early Tertiary resulted in the juxtaposition of the Chilitos formation against carbonates and siliciclastic sedimentary rocks. Throughout the region, the rock assemblages are intruded by Late Cretaceous to Early Tertiary granitic plutons, and are unconformably overlain by Tertiary aged felsic volcanic flows and tuffs, [Johnson et al., 2000]. North-, northwest-, and west-striking faults cut the assemblage; the age and displacement of these faults are unknown, although they could be synvolcanic extensional or postvolcanic strike-slip faults, [Johnson et al., 2000]. A section of rocks, which includes the Chilitos Formation, has also been uplifted to form a horst, within which the San Nicolás deposit is found. Tertiary volcanic rocks have intruded along the northwest-striking faults, and along north-striking faults at the west of the horst.

The deposit, itself, is bounded to the southwest by the northeast flank of a rhyolite

²A back-arc to fore-arc setting refers to relative locations within an island arc system. Island arcs are a chain of islands, such as the Aleutian Islands, that form at the convergent margin of two oceanic plates. They are often associated with deep oceanic trenches, and volcanic activity. A fore-arc setting would exhibit compressional tectonic forces, while a back-arc setting exhibits tensional tectonic forces.

dome, and by a southwest dipping fault to the northeast. The deposit is copper-rich towards the bottom with a high-grade polymetallic cap at the top. Also, the sulphide is granular and fragmented in the lower zone and laminated in the upper zone, typical of VMS deposits. The two zones could have different physical properties that geophysics could help differentiate. Silicification and barite are present along the fault that bounds the deposit to the northeast. Extensive hydrothermal alteration of the host rocks is also observed; this could affect physical properties of the rocks.

Deposit Genesis

Massive sulphide deposits are often formed in the proximity to plate margins [Lydon, 1988]. Mineral-rich hydrothermal fluids, either originating from a magmatic source or circulating in the sediments and being driven by the heat from a magmatic source, react with sea water and deposit sulphide minerals. These deposit types are usually quite small and the duration of mineral deposition short-lived. Several deposits can be found in the same area often in association with submarine volcanic flows and sediments [Lydon, 1988]. San Nicolás, although probably formed by a similar process, is larger than usual and was possibly formed in a tectonic environment not usually associated with massive sulphide deposits. This makes San Nicolás an unusual massive sulphide deposit, and provides a new and exciting type of exploration target for this region. Consequently, the regional potential for discovering other deposits is high because exploration programs in this area have not previously focused on this type of deposit.

Geologic Exploration Model

The main features of the deposit that could be critical to similar deposits being found comprises the geologic exploration model. At San Nicolás this could include the following:

1. Favorable host stratigraphy between the rhyolitic and mafic volcanics that corresponds to a *time-line*
2. Association of the deposit with a quartz-rhyolite dome complex.
3. The shallow depth of the Chilitos formation due to the presence of a horst.
4. Faults that may have provided conduits for hydrothermal fluids.

1.3 Physical Properties

Geophysics can help define the physical property distribution of the subsurface. This is useful when objects can not easily be investigated firsthand, and have physical properties that distinguish them from their surroundings. Estimates of physical property measurements are also needed to relate inversion results to geology. Massive sulphides are commonly known to have contrasting values in density, conductivity, chargeability, and possibly magnetic susceptibility. These physical properties stem from the minerals which constitute the ore. In the case of San Nicolás, the ore is primarily made up of pyrite, chalcopyrite, and sphalerite, with some magnetite (not a sulphide mineral, although frequently associated with sulphide assemblages), and possibly pyrrhotite. The total sulphide content is greater than fifty percent by volume within the main ore body.

From the mineralogy described above, the deposit is expected to have the following physical attributes:

1. High density (a minimum of 3.8 g/cc [Ward, 1966])
2. High magnetic susceptibility (due to the presence of magnetite and pyrrhotite)
3. High conductivity (this variable property often depends on the connectivity of the

Rock Type	Density (g/cm ³)	Magnetic Susceptibility (S.I. $\times 10^{-3}$)	Resistivity (Ohm-m)	Chargeability (mV/V)
Tertiary Breccia	2.3	0-5	20	10-30
Mafic Volcanics	2.7	2	80	30-80
Sulphide	3.5	10	20-30	200+
Quartz rhyolite	2.4	0-1	100	10-20
Graphitic Mudstone	2.4	0	100+	30-70

Table 1.1: Estimated physical properties for the five major rock types encountered at San Nicolás. The massive sulphide has contrasting values in all physical properties apart from resistivity. A low resistivity value of 20 Ohm-m for sulphides, similar to the Tertiary breccia, makes detection of the deposit more difficult when using some geophysical methods.

conducting paths³)

4. High chargeability due to the presence of metallic minerals

[Telford, Geldart, and Sheriff, 1990c]

Apart from the sulphide ore, the geological units that are of significant interest are: the Tertiary overburden, which is conductive⁴ and contains some susceptible units; a mafic volcanic unit that is magnetically susceptible; and dense mafic intrusions. Table 1.1 provides physical property information for the major rock types found at San Nicolás⁵, while Appendix B contains a complete review of physical property values used to constrain the models and interpret the inversion results.

³many sulphides, excluding sphalerite, are semiconductors with conductivities often above 1000 S/m in their pure form [Sherman, 1997]

⁴Conductivity (σ [Siemens per meter]) is the inverse of resistivity (ρ [Ohm-meters]).

⁵The density of the sulphide has been reduced due to the presence of semi-massive sulphide material.

1.4 Geophysical Exploration Model

Based on the physical property values and the geologic exploration model we can define a geophysical exploration model (Figure 1.1). For San Nicolás this can be defined as follows:

1. Bodies with contrasting density, chargeability, conductivity and magnetic susceptibility will be good candidates for mineralization.
2. In combination, the physical property contrasts can be used to differentiate between mineralization and shallow intrusives that may be just dense and magnetic.
3. The density contrast between the host volcanic rocks and the overburden will help define the horst, which is important in bringing the favorable horizon near to the surface.

1.5 Data

Many geophysical data have been collected over San Nicolás, and the surrounding area. Data were collected at San Nicolás during the last five years, and provided to U.B.C. by Teck Corporation. The data used in this study include: airborne electromagnetic and magnetic data, gravity, ground magnetic, CSAMT (controlled source audio-frequency magnetotelluric), IP (induced polarization) data. The locations of these data sets are shown in Figure 1.3. Plots of the data are presented in Appendix D along with a summary of survey specifications for each data set (Table D.1). In order to limit the number of plots in the main text I present only: (1) physical property inversion models, (2) geologic plan maps and sections, and (3) interpretative figures. Geological information, as interpreted from drill holes, is also used to demonstrate the success of inversions performed, and imply the benefits of inversion methods in exploration. Unless stated otherwise, all coordinates

are given in meters and are defined on a local San Nicolás grid system. The conversion between the local grid system and the UTM Zone 13, NAD 27 system that was employed, is given in Appendix A.

1.6 Modeling

Geophysical modeling involves relating, or mapping, data (a numerical representation of a physical signal that has been measured or observed) to a numerical representation of a physical property distribution within the earth, known as a model. The mapping procedure can occur in both directions.

1.6.1 Forward modeling

Forward modeling is a procedure in which a unique set of synthetic geophysical data are calculated based on a known physical property distribution. As well as being an inherent part of the inversion process, forward modeling plays an important role in designing geophysical surveys, and testing geologic scenarios.

1.6.2 Inverse modeling

Inversion is the process of constructing physical property models of the earth based on information contained in geophysical data. Unlike forward modeling, the inversion problem is often ill-posed. An infinite number of models could have produced the data. We are able to choose one model by only considering those with simple structural characteristics. Inversion models are generally much easier to interpret than the original data and provide a superior understanding of the subsurface. A description of the inversion procedure, as it applies to this thesis, is given in Chapter 2 and the inversion parameters used to generate the physical property distributions presented are given in Table E.1.

1.7 Outline of the thesis

The following provides a synopsis of each chapter. Having already defined the geophysical exploration model, the following chapters (apart from Chapter two, which introduces geophysical inversion) proceed through the geophysical exploration steps while emphasizing the role of inversion. The work is presented within the context of San Nicolás.

1.7.1 Inversion methodology

Chapter two provides an overview of the inversion process. This is needed in order for the inversion models to be given their correct context regarding the information they supply to the geologist.

1.7.2 Regional-scale modeling

The regional scale for this thesis is defined by the coverage of gravity and airborne magnetic and electromagnetic data. This large-scale region is shown by the upper plot of geology with data boundaries in Figure 1.3. Not all of the airborne magnetic and electromagnetic, and reconnaissance IP data sets cover such a large area as the regional-scale coverage, so an intermediate-scale is introduced and shown in the center plot in Figure 1.3. Inversions presented in Chapter three are performed at either the regional or intermediate scales. In the context of this thesis, only inversion models of geophysical data are used to answer relevant regional geologic questions and focus exploration efforts.

1.7.3 Deposit-scale modeling

This is the scale at which anomalies are explored and deposits are found, or grossly delineated. More detailed surveys are carried out on the best areas and targets that were identified at the previous stage of the exploration process. It may be, that work carried

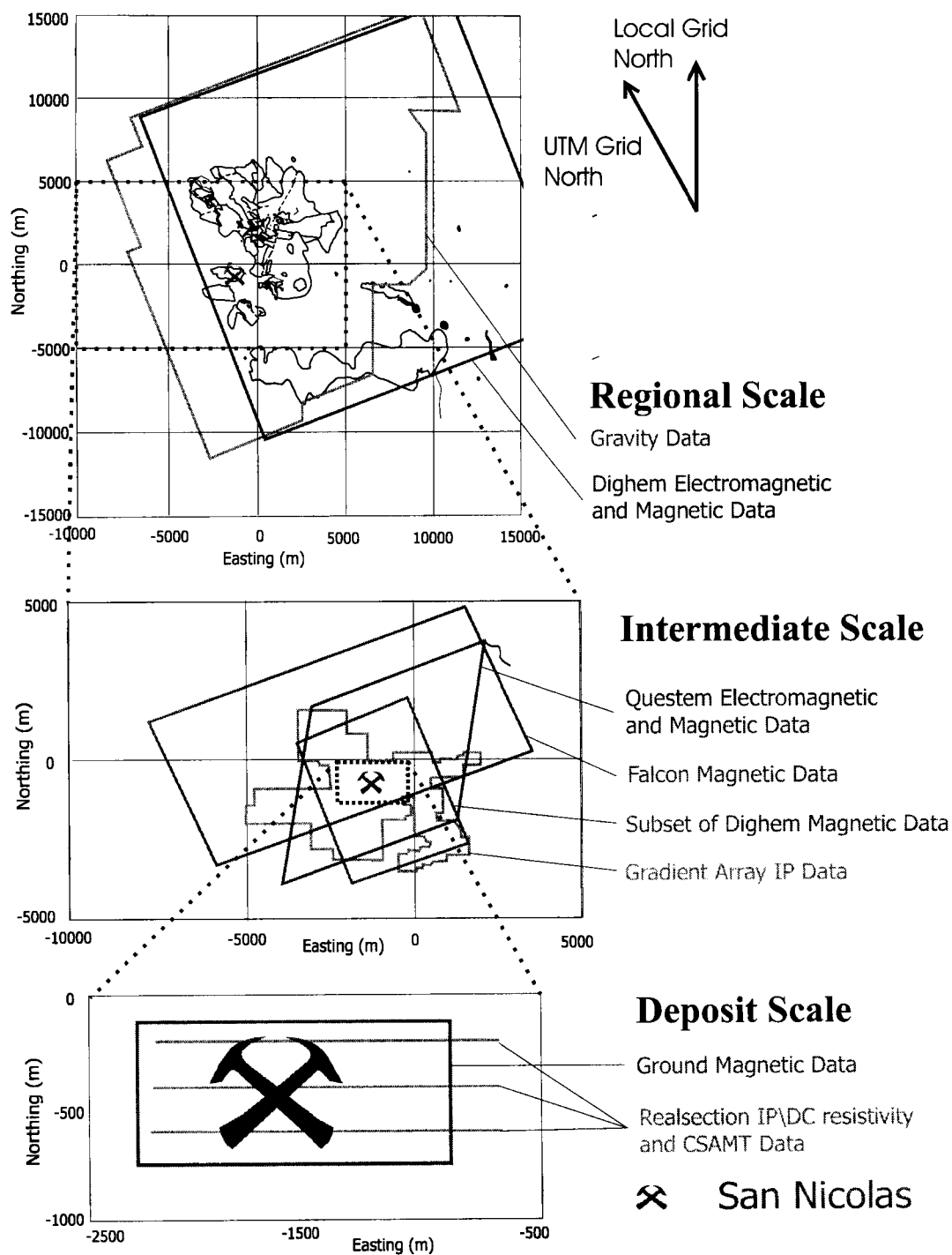


Figure 1.3: Data coverage over the San Nicolás deposit and the surrounding area. Top: the regional-scale area with data coverage and outcrop geology (legend shown). Center: Intermediate-scale data coverage shown on top of outcrop geology. Bottom: deposit-scale data coverage. The size of the rock-hammer symbol approximately corresponds to the extents of the surface projection of the San Nicolás deposit.

out on a more regional scale is sufficient to identify a discrete target, but more detailed information about the deposit is needed to determine whether it is worth drilling with one or five drill-holes, or at all. Alternatively, drilling may already have commenced and the geologist may require information about the geometry of the deposit before drilling the next hole. These scenarios provide the motivation behind the deposit-scale inversions of gravity, magnetic, electromagnetic, and induced polarization data that are presented in Chapter four. This work is performed at the scale depicted by the lower plot in Figure 1.3.

1.7.4 Detailed modeling using drill-hole information

Chapter five addresses the use of a priori information in refining the physical property models constructed at the local-scale. Density and magnetic susceptibility measurements provide information about the earth that is independent of the data. Including this information in the inversion process forces only models that are consistent with known physical property distributions to be constructed. Such a priori information could be used to help spot a second drill-hole based on the results of the first, at the early stages of drilling, or could help model what is below a drilled out portion of a deposit at the end of the first phase of drilling. Furthermore, inversions with a priori information could be used to incrementally update physical property models with new information as each drill-hole is completed and logged. The evolving physical property models could help guide a drill program that would potentially result in less drilling needed to delineate the mineralization or conclusively test the anomaly.

1.7.5 Detection of deep ore: survey design

Forward modeling is an important tool that generates synthetic geophysical data based on physical property distribution models. Chapter six employs forward modeling to examine

the success (based on signal-to noise estimates) of different survey designs in detecting the keel and any extensions of mineralization that may exist. Forward modeling can also be used to test the geophysical response of different geological scenarios against field data.

1.8 Summary

The flowchart of Figure 1.1 provides the framework around which this thesis is based. The flowchart is presented at the beginning of each results chapter in order to set the scene and provide an exploration context for the results.

Equipped with knowledge of the geologic environment, physical properties, and inversion methods, geophysical inversions are performed on data from San Nicolás. The inversion results demonstrate the assistance that geophysical inversion can provide to geologists regarding a variety of issues that range from regional exploration to detection and delineation of deep ore. Understanding the geology, and providing physical property models with a geologic context, are essential at all stages of the exploration process and are important to this study. Geologic information is also used to determine the success of using geophysics alone to define the deposit.

Chapter 2

Geophysical Inversion

2.1 Introduction

This chapter aims to construct a framework of understanding about the inversion process so that important issues such as *non-uniqueness* and *data misfit* are appreciated, and the inversion models are given their correct context regarding the information they supply to the geologist. Geophysical inversion produces physical property models from geophysical data, whereas forward modeling produces data from a physical property model of the earth. Several factors influence the inversion procedure and, therefore, the outcome of the model. These factors include: (1) estimating data errors, (2) using different reference models, (3) choosing smoother models or smaller model types, and (4) choosing a different balance between honoring the data and introducing features in the model. This chapter will briefly address the aforementioned issues, along with providing a synthetic example, in order to demonstrate the construction of the physical property models presented in this thesis.

2.2 Inversion Basics

2.2.1 Data

What are the data? This may seem like a basic question, but it is one that needs to be explicitly answered in order for an inversion to be successful. Data include not only the actual datum values, such as the strength of a magnetic field or the voltage between

two electrodes, but, an error associated with each datum (where possible), locations of transmitters and receivers, orientations of instruments, data normalizations or changes in units, and detailed knowledge of any processing that is applied to the data. These are all crucial elements that define what the datum is and how it is connected to the spatial distribution of physical property values within the earth.

Other information that can be thought of as a form of data could be physical property measurements made directly on the earth or any geologic knowledge that is pertinent to the region of the earth that produced the geophysical data. In the context of this thesis I deem such information to be *a priori*, that is information known prior to inverting the geophysical data. We can use these data to constrain and guide inversions.

2.2.2 Discretization and forward modeling

The earth is a continuous medium having physical property values everywhere. We can describe the physical properties of the earth as a continuous function of position, $m(x, y, z)$. When modeling, we usually simplify the earth through discretization, that is, dividing the earth into many cells whose physical property values are constant and whose size is smaller than the resolving power of the experiment or survey. For the inversions presented in this thesis, the cells are cuboidal and aligned with an orthogonal coordinate system. A 3D grid system or *mesh* defines the discretization of the model (Figure 2.1). The earth is now represented by a column vector $\mathbf{m} = (m_1, m_2, \dots, m_M)^T$ of length M (the number of cells), where each element of the vector corresponds to a the physical property value of a different cell in the model. Forward modeling then consists of solving a system of equations to predict the responses at each observation location based on the physical property distribution of the discretized earth.

For a linear system of equations, the data are calculated by performing a matrix-vector multiplication. The matrix notation of the discretized forward problem is stated

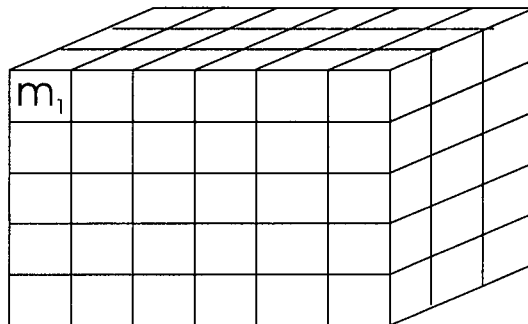


Figure 2.1: The earth is discretized and represented by an orthogonal 3D mesh.

as:

$$\mathbf{d} = \mathbf{G}\mathbf{m}, \quad (2.1)$$

where \mathbf{G} is the forward operating matrix (of size $N \times M$), \mathbf{m} is a column vector comprised of M model parameters, and \mathbf{d} is a column vector of length N , containing the observations, or data. For the inversion problems addressed in this thesis, the number of data N is much less than the number of model parameters M .

2.2.3 Non-uniqueness and the Model Objective Function

In the inverse problem we have the data (\mathbf{d}) and a forward operator (\mathbf{G}), and we want to find the model (\mathbf{m}). However, as with any data set, we are limited to dealing with a finite number of inaccurate observations. This means there are many more model parameters than data, which makes the system underdetermined. Also, we don't want to produce a model that will exactly reproduce the observed data, as this will also reproduce noise. Therefore the data don't constrain all of the model parameters uniquely, and there are an infinite number of models that, when forward modeled, will reproduce, or *fit*, the data to a specified degree.

In order to choose one model from the infinite many that fit the data, we select the type of model we desire using a model objective function (sometimes called a model norm). The inverse problem is posed as an optimization problem by minimizing the model objective function subject to the requirement that the recovered model produces data that adequately resembles the original observations. When a model objective function is suitably designed the recovered model has simple structural characteristics that should emulate geology. Flexibility should be built into the model objective function so most simple geologic situations can be accommodated. The model objective function for a continuous 3D model $m(x, y, z)$ can be written as:

$$\begin{aligned} \phi_m = & \alpha_s \int_V w_s (m(x, y, z) - m(x, y, z)_{ref})^2 dv + \alpha_x \int_V w_x \left(\frac{\partial m(x, y, z)}{\partial x} \right)^2 dv \\ & + \alpha_y \int_V w_y \left(\frac{\partial m(x, y, z)}{\partial y} \right)^2 dv + \alpha_z \int_V w_z \left(\frac{\partial m(x, y, z)}{\partial z} \right)^2 dv, \end{aligned} \quad (2.2)$$

with the equivalent matrix notation written as:

$$\Phi_m = \|\mathbf{W}_m(\mathbf{m} - \mathbf{m}_{ref})\|^2. \quad (2.3)$$

The first term in equation 2.2 measures the closeness of the model to a reference model ($m(x, y, z)_{ref}$), while the following terms measure the smoothness of the model in the x , y , and z directions. Coefficients α_s , α_x , α_y , and α_z are used to weight the different parts of the model objective function relative to each other. This enables the model objective function to recover models that are close to the reference model, or preferentially smooth in one direction, by only changing one or two parameters. Spatially dependent weighting functions w_s , w_x , w_y , and w_z are used to weight the importance of one model parameter versus another. The reference and weighting functions allow for *a priori* information to be incorporated into the inversion that may come from other geological or geophysical sources. Bounds can be implemented to limit the range of values a cell can exhibit and

dip information can also be included [Li and Oldenburg, 2000b]. Equation 2.3 is the matrix notation of the model objective function, where \mathbf{W}_m is a weighting matrix of size $M \times M$ that includes all the weighting and derivative functions of the model objective function, and $\|\cdot\|$ is the l_2 norm.

2.2.4 Signal-to-noise and data misfit

Signal-to-noise plays an important part in inversion because, no matter how good the field work and instrumentation, the data always contain noise. It follows that we do not wish to produce a model that can exactly reproduce the data as it would also reproduce the noise (known as *over-fitting* the data). Unfortunately noise levels are often unknown or only approximately known, but some estimate as to either absolute or relative noise levels is usually required. In order to measure how well the model response, or predicted data, *fit* the observations, we design a data misfit function as stated below.

$$\Phi_d = \sum_{i=1}^N \left(\frac{\mathbf{d}_i^{obs} - \mathbf{d}_i^{pre}}{\epsilon_i} \right)^2 \quad (2.4)$$

$$= \|\mathbf{W}_d(\mathbf{d} - \mathbf{d}^{pre})\|^2, \quad (2.5)$$

where N is the number of data, \mathbf{d}_i^{obs} is the i^{th} observed datum, \mathbf{d}_i^{pre} is the i^{th} predicted datum generated by forward modeling the model, and ϵ_i represents the standard deviation of the i^{th} datum, and is included in \mathbf{W}_d when the expression is written using matrix notation (Equation 2.5).

We assume the noise is Gaussian in nature, has zero mean, and is non-correlative. This being said, if noise levels are well estimated, the data misfit should be approximately equal to the number of observations ($\Phi_d \approx N$). When data have been over fit ($\Phi_d \ll N$), the misfit is too small, and models often display erroneous structural complexity. This complexity is needed to fit noise in the signal and has no reflection on the structure of

the earth. If the data misfit is too large ($\Phi_d \gg N$), the model often lacks structure signifying that not all of the information that the data contain about the earth has been extracted. In practice, several inversions are generally needed in order to produce a model that has some structure, as defined by signal in the data, but not too much structure or complexity, as originating from noise in the data. Estimating the correct noise levels often comes down to the experience of the user. Obviously, the better the signal-to-noise ratio and the noise estimates of the data, the better the inversion model.

2.2.5 Optimization

The inversion procedure is stated as the following optimization problem: find a model that minimizes the model objective function subject to fitting the data to a specified tolerance. The objective function is comprised of the data misfit and the model objective function (Equation 2.6), with β representing a trade-off, or regularization, parameter. Minimization of the objective function will generate the model.

$$\Phi(m) = \Phi_d + \beta\Phi_m \quad (2.6)$$

Choosing the regularization parameter is key to producing a model that has the maximum amount of information about the earth while containing the minimum amount of erroneous structure that relates to noise in the data. When noise levels are known, the regularization parameter can be determined using the discrepancy principle [Hansen, 1998]. Methods that can be applied when noise levels are unknown include: (1) generalized cross-validation (GCV), which is based on the principal that a good model would be able to predict a missing datum¹ [Haber, 1997], (2) the L-curve criteria that searches

¹For the generalized cross validation method to be most effective, a sense of the relative size of data errors is necessary, although the actual magnitude of the error is not needed.

for the corner of the curve when data misfit and model objective function values are plotted for different values of the trade-off parameter², both of which are described by Hansen [Hansen, 1998], and (3) manually performing inversions with different noise estimates. In practice it is often useful to apply GCV or L-curve methods to get an idea of the approximate magnitude of the errors, and then use the discrepancy principle to slightly increase the misfit.

2.2.6 Appraisal

When a model is generated via the inversion procedure it is often desirable to determine how good the model is, that is, how well does it represent the earth? As geophysics is usually a minimally invasive technique we don't tend to dig up the earth, but revert to the data and model to answer this question. What is really being asked is: what can be concluded about the earth, given the data? As it is already known that the constructed model is not the only one that will fit the data, a range of models that fit the data could be generated. Construction of many models using inversion could be quite time intensive, although a rudimentary exploration of this *model space* is a good idea. This can be accomplished by changing the model objective function. The user soon gets an idea of features in the model that don't change (which are therefore probably required in order for the model to fit the data) and features that do change (which are not constrained by the data). It is important to be aware that the model may not define the earth well in all places, and over-interpretation of features that are not constrained by the data is

²On a log-log scale, a plot of the data misfit (ϕ_d) verses model norm (ϕ_m) for different values of the trade-off parameter (β) produces a curve which resembles the letter *L*. In the upper part of the curve, a decreasing value of β will produce large decreases in data misfit for small increases in model norm. Conversely, along the lower part of the curve, as β decreases further a small decrease in data misfit will correspond to a large increase in model complexity, as measured by the model norm. A model that corresponds to the point of maximum curvature will have a reasonable amount of complexity but should not over fit the data.

possible. As an alternative to defining how good the model is, hypothesis testing can be performed to confirm, or eliminate, geological scenarios of interest.

Other, more sophisticated methods of appraisal, which are not used in this thesis, might include analysis of averaging functions [Backus and Gilbert, 1968] and point spread functions of the model resolution matrix [Menke, 1984].

2.3 A synthetic example

In order to demonstrate the inversion process, and the level of understanding that is needed in order to gain the most from an inversion result, a synthetic example has been devised. A density-contrast model is used to forward model synthetic, vertical-component, gravity data. The data are contaminated with noise and then inverted using *GRAV3D* (Appendix E.2) to recover a density-contrast model.

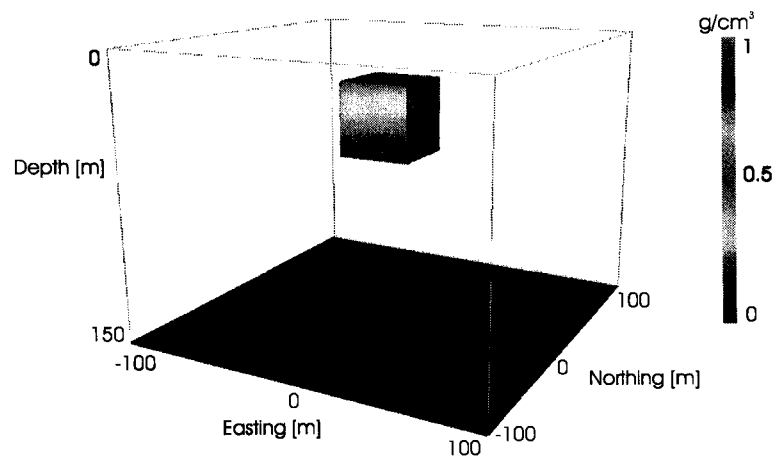


Figure 2.2: Perspective view of synthetic density-contrast model. The model has been volume rendered so that cells with density contrasts less than 1g/cm^3 have been removed.

2.3.1 The data

To start, the data are generated by making measurements over a three-dimensional density contrast distribution (Figure 2.2), much like we would in the field. This is performed using a forward modeling procedure that is described in Appendix C.1. Figure 2.2 shows a volume-rendered, perspective view of the synthetic density contrast model. The model consists of 7200 cells of size 10m x 10m x 5m. The feature we wish to model is a 40m x 40m x 40m cube located 20m below the origin with a density-contrast of 1 g/cm³. The cube is hosted in a background of 0 g/cm³ density-contrast.

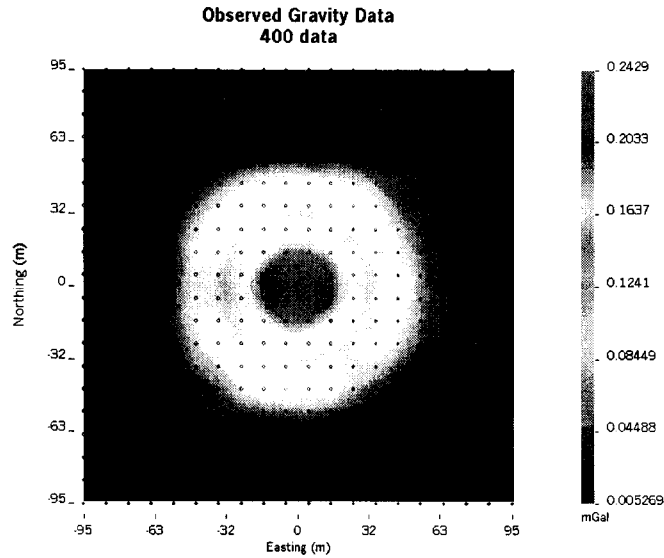


Figure 2.3: Surficial synthetic gravity observations with 3% + 0.0005 mGal random noise added.

Four hundred synthetic gravity measurements were made at the surface on a regular 10m grid. The data were contaminated with 3% + 0.0005 mGal random noise (Figure 2.3) and standard deviations were assigned equal to 3% of the magnitude of each datum value. The minimum standard deviation was limited to 0.0005 mGal.

2.3.2 Presentation and interpretation of inversion models.

Mineral deposits and the geologic structures can exhibit very three-dimensional geometries. The geology of San Nicolás and the surrounding area is no exception, and much confidence is gained when we are able to model the earth in three-dimensions. Models often contain complicated 3D structures of varying physical property values that are difficult to present in a document. In this thesis 3D models are presented using a combination of perspective view images, plan views at different elevations, and cross-sections. Visualization tools such as: coloration based on physical property values, shading, volume rendering, and view angle are factors that can change what is perceived when viewing a model. As a result, an attempt is made to objectively convey the salient features of each model with one or two images. For comparison between models, north-facing cross-sections at -400m north, and west-facing cross-sections at -1700m east are frequently presented. Furthermore, in order to address non-uniqueness issues, information may be distilled from the analysis of several models generated using different objective functions. Care is needed when making inferences about geology from an inversion model, and when the models are discussed in the text the reader may assume the features have been appraised using the methods outlined in Section 2.2.6.

2.3.3 Recovered models

The inversion problem is solved by minimizing Equation 2.6. The solution depends on the definition of the model objective function (Φ_m), the data misfit (Φ_d), and the regularization parameter (β) that controls the trade-off between misfit and structure. The following sections show how changing the desired misfit or alpha coefficients in the model objective function effects the recovered model. Examples of changing the reference model and assigning bounds on some of the model cells are also shown.

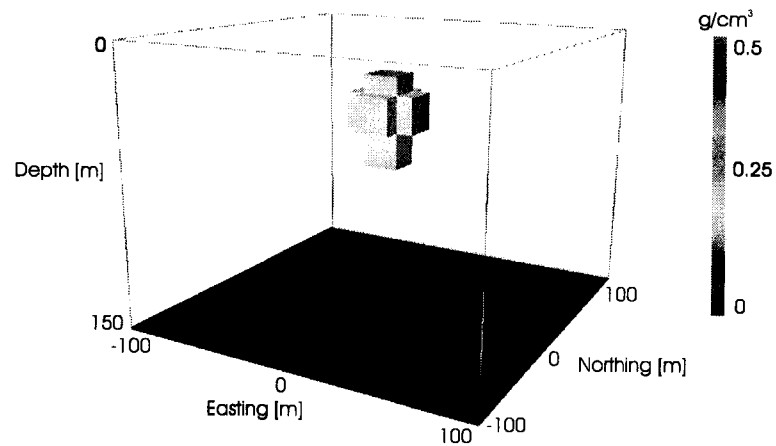


Figure 2.4: Perspective view of recovered density-contrast model. The model is viewed from the southeast (local coordinate system), and is volume-rendered with a cut off at 0.3 g/cm^3 .

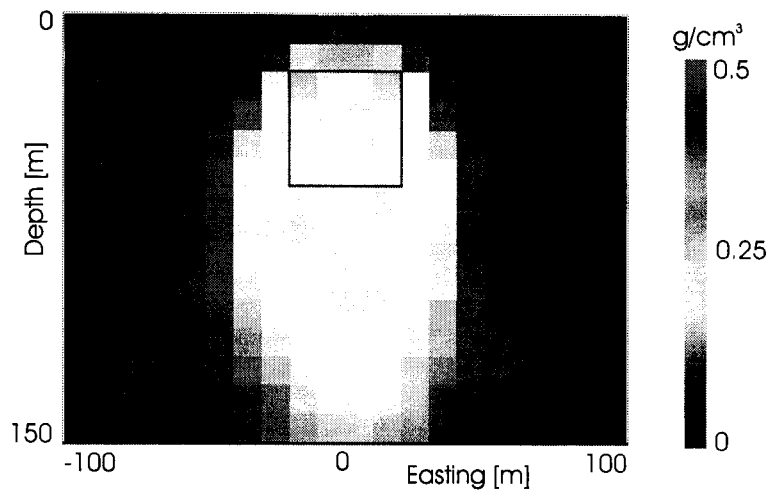


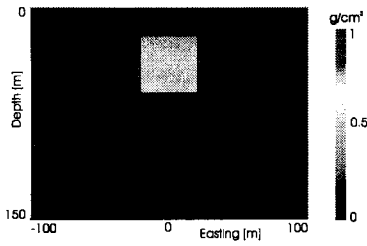
Figure 2.5: North-facing cross-section through recovered density-contrast model with location of true model shown.

A generic model

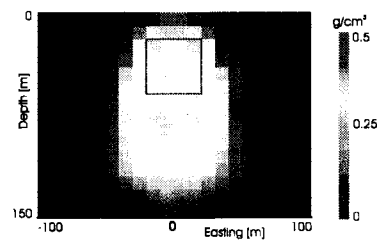
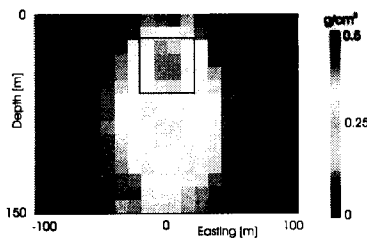
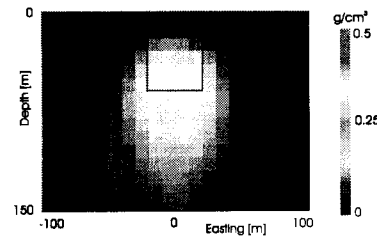
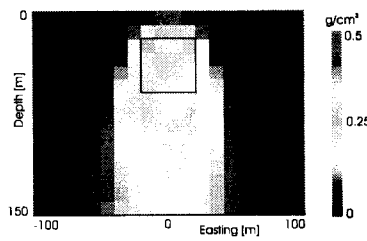
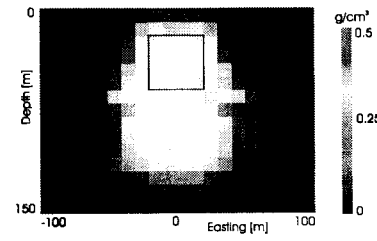
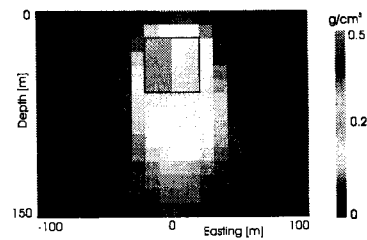
For this first example the errors are assigned appropriately and so we have an estimate of the desired misfit. A model is recovered that fits the data to a value equal to the number of data ($\Phi_d = N$). Figure 2.4 shows a perspective view of the model recovered. The model has been volume-rendered using a cut-off value of 0.3 g/cm^3 . The view angle is the same as that used in Figure 2.3, but the coloration has been changed. A north-facing cross-section through the model is shown in Figure 2.5 with the outline of the true model also displayed. As can be seen from these two images, the recovered model has a concentration of anomalous density coincident with the location of the true model. The recovered model seems smeared out relative to the true model, especially at depth, producing an anomaly of larger volume, and, as a result, lower density contrast values. This is because the model objective function was designed to generate a model that is close to a reference density-contrast model of 0 g/cm^3 , and have smooth features. The density values of the recovered model are smaller than the true model; however, the total mass within the model is the same.

Effects of over- or under-fitting the data

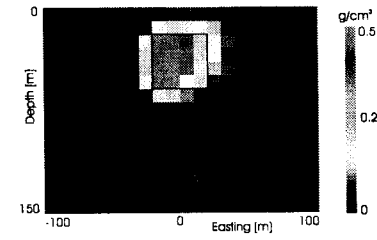
Figure 2.6 shows north-facing cross-sections through the true model and seven recovered density contrast models from inversions of the synthetic gravity data. Each model is overlain with an outline of the true model, and all of the models are plotted on the same colour scale (0 to 0.5 g/cc). For comparison, Figures 2.6(a) and 2.6(b) show cross-sections of the true and recovered model with $\Phi_d = N$ respectively. Figure 2.6(c) shows a model chosen with the data misfit one tenth of the number of data. Because the correct noise levels were assigned to the data, the model over-fits the data. In comparison to Figure 2.6(b), where the data misfit was appropriate, the density-contrast values are slightly



(a) True density-contrast model

(b) Model *fits* the data: $\Phi_d = N$ (c) Model *over fits* the data: $\Phi_d = 0.1N$ (d) Model *under fits* the data: $\Phi_d = 5N$ (e) Model smoother in z-direction: $\alpha_z > \alpha_x, \alpha_y$ (f) Model smoother in horizontal direction: $\alpha_x, \alpha_y > \alpha_z$ 

(g) Reference model is half of true model



(h) Half of model bounded

Figure 2.6: North-facing cross-sections through true model and seven recovered models for seven different inversions. All of the inversion models are shown on the same color-scale with dark blue representing 0 g/cm^3 and below, and magenta representing 0.5 g/cm^3 and above. The original model, an outline of which is shown on all the recovered models, is shown at a different scale.

closer to the value of the true model. However, erroneous artifacts, represented by more complex structures, are produced in the recovered model as the noise in the data manifests itself. Conversely, the density contrasts are lower when the data misfit is too large, as shown by the model in Figure 2.6(d). This model was produced with a misfit of five times the number of data. The under-fitting arises because the model objective function has more control over the model produced when the data misfit is larger, resulting in a smaller and smoother model with more simple structure. If noise levels were truly unknown then both of these models would be an equally viable a representation of the earth. In such cases determining how well the data have been fit becomes more difficult. GCV and L-curve methods can be useful in such cases as the models are chosen based on methods other than data misfit.

Effects of changing the model objective function

Figures 2.6(e) and 2.6(f) show the results of altering the model objective function to produce models smoother in the vertical direction and smoother in the horizontal directions, respectively. This is done simple by changing the relative values of the α coefficients in equation 2.2. As can be seen, the model shown in figure 2.6(e) is elongated in the vertical direction, which extends the anomaly at depth, while the model shown in figure 2.6(f) is truncated at depth but has broader horizontal extents. Both models still have a concentration of mass in the vicinity of the true model. The models also both fit the data to the same degree ($\Phi_d = N$) so, in the absence of any external biases, both are an equally reasonable representation of the earth. This is a demonstration of non-uniqueness. Together, the models produced represent a very brief exploration of model space.

Including a priori information

As mentioned in section 2.2.1, other forms of information (such as that from a drill hole) can be incorporated into the inversion. If we assume that half of the cube is known to exist prior to inverting the data, we can use this information directly in the inversion. In the first case a reference model is constructed that is 1 g/cm^3 where the cube is known to exist, and 0 g/cm^3 everywhere else. The result of using the reference model with the same parameters (data misfit and α coefficients) as those used in the original inversion (figure 2.5) is given in figure 2.6(g). The recovered model is slightly improved compared to the one generated without the reference model in that the anomaly doesn't have values extending to such depths. This result could be changed further by altering the relative weighting of each cell to the reference model using the weighting matrix in equation 2.3. This would allow regions of the model that are known prior to inverting the data to be weighted more strongly with respect to the reference model than other cells. An alternative method is to restrict the range of values a cell can assume. In figure 2.6(h) a model was constructed using this method. For the part of the cube known to exist, the model was restricted to ranges between 0.9 and 1.1 g/cm^3 , for the rest of the model space the bounds were set from 0 to 1 g/cm^3 . Most of the mass is concentrated in the correct place. It seems evident that this technique has produced a very good model that is the most representative of the true model. This demonstrates how constraining the model in one area can improve the model in the other areas.

2.4 Conclusions

Inverting geophysical data to construct physical property models can allow for greater understanding of the subsurface. However, in order to get the most from the inversion process the following issues must be appreciated: (1) the data have to be fully understood

and, where possible, errors should be assigned; (2) fitting the data to an appropriate degree can be difficult when errors are unknown; (3) users should be aware of non-uniqueness and the role that the model objective function plays; and (4) significant understanding of the main model features can be obtained through rudimentary exploration of model space.

The issues addressed above, and in the synthetic example, underlie the importance of accurate data and the need for well-designed surveys that constrain models leaving as little room for ambiguity, and non-uniqueness, as possible.

Chapter 3

Regional Scale Modeling

3.1 Introduction

This chapter demonstrates the use of regional geophysical inversion in an exploration program. The goals of the regional geophysical program have been determined by the geologic and geophysical exploration models of Chapter one. This chapter presents the results of inverting regional geophysical data provided to U.B.C. by Teck Corporation, and includes a preliminary geologic interpretation based on physical property values and geologic information. The steps that are followed in this chapter are within the context of the exploration model, and are shown by shaded boxes in Figure 3.1.

Extensive Tertiary and Quaternary cover makes exposure of the bedrock limited. As well as obscuring the sulphide-hosting Chilitos Formation, this has resulted in a limited understanding of the regional geology and hampers exploration for VMS deposits. Geophysical methods have been applied in order to shed light on the regional geology and search for massive sulphides underlying the overburden.

Regional-scale inversion of gravity, magnetic, electromagnetic, and IP data is performed in order to generate areas for detailed follow-up exploration. The generation of target areas is carried out by identifying physical property distributions that fit the expected response from massive sulphide deposits and by assigning regional geologic context to physical property features observed in the models. Both of these approaches can help focus exploration efforts.

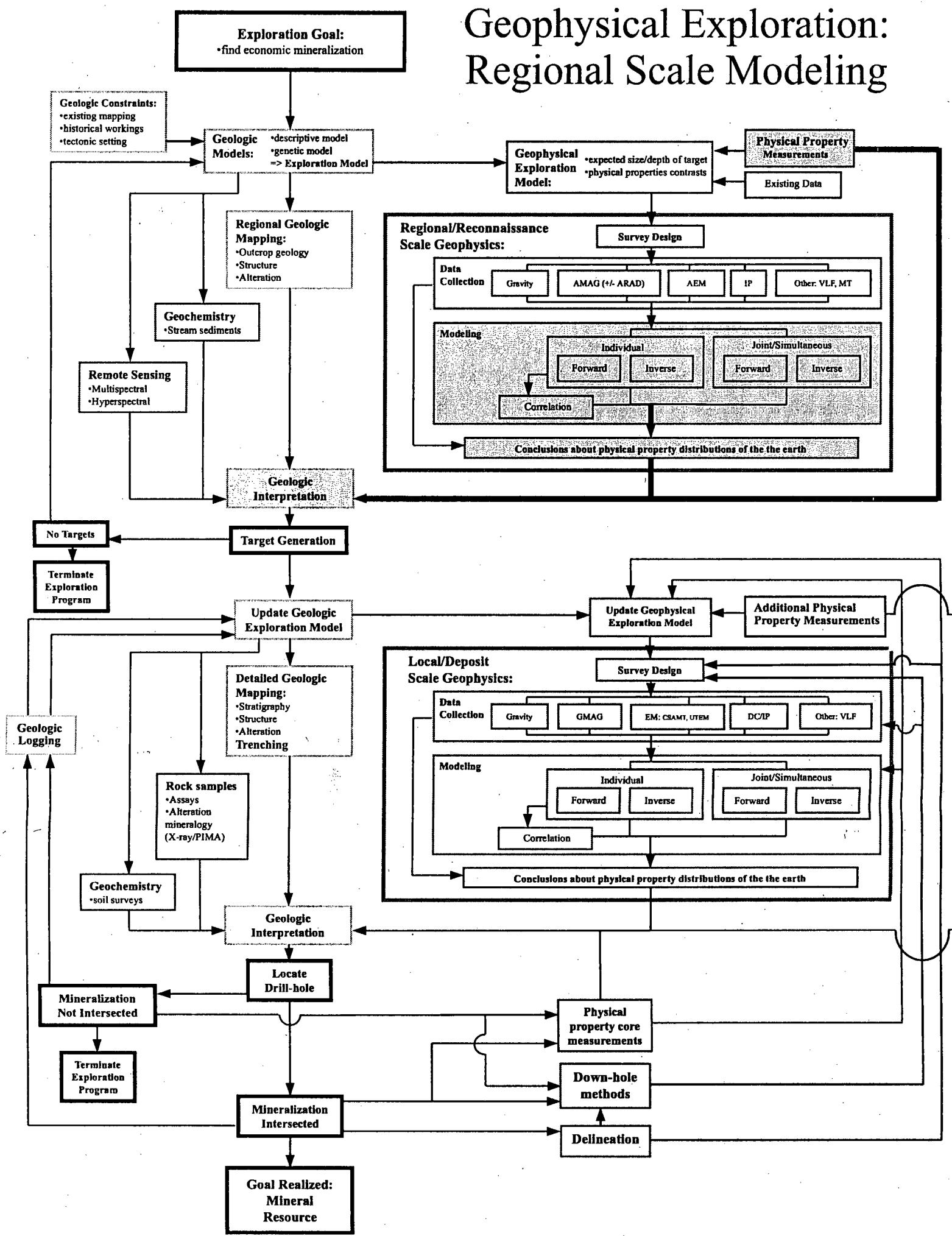


Figure 3.1: Geophysical Exploration Methodology Flowchart: Regional Scale Modeling.

In the first instance, correlation methods are applied to three-dimensional models to determine regions of favorable combinations of physical property values. This involves a rudimentary spatial correlation of cells in inversion models that exhibit physical property values comparable to the physical properties of massive sulphide, as determined in section 1.3.

In the second instance, geologic features are interpreted from models, and the relation of features to mineralization is determined by using conceptualized geologic models. Interpretation of the models is performed while keeping in mind the notions of non-uniqueness and fitting the data.

Ideally, the two methods will complement each other by generating fewer, but better, exploration targets than when one method is used in exclusion. This approach will hopefully result in a more efficient use of exploration resources.

3.2 Regional geology

A regional map showing mapped surficial geology is presented in Figure 3.2. Apart from the geological setting that was described in Chapter 1, and the direct detection of massive sulphides, features that are of regional interest in this chapter include: (1) the varying thickness and physical properties of the overburden, (2) the location of shallow mafic intrusions, (3) the extent of a horst block, and (4) widespread fault systems and how they might relate to mineralization. Some of these features will undoubtedly be related to each other because faults define the boundaries of the horst, and overburden thickens away from the horst. The scenarios mentioned above represent a subset of the many possible geologic attributes of the area that maybe related to mineralization. However, the geologic conceptual model considered in this thesis assumes that the Chilitos formation is where we expect to discover deposits like San Nicolás, and faults play an

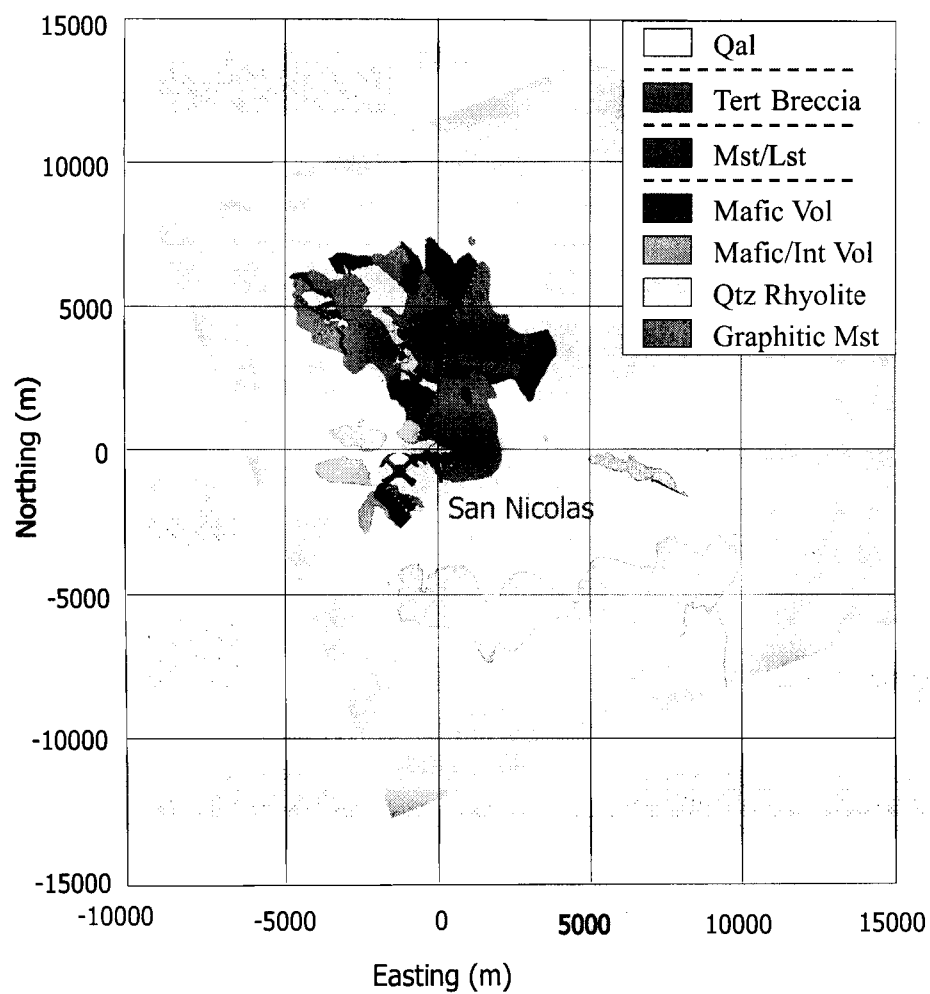


Figure 3.2: Plan-view of regional geologic outcrop at San Nicolás. Modified from Johnson et al., 2000. Most of the region is overlain by Quaternary alluvium. Inset shows a key to the stratigraphic units.

important role in delivering mineral-laden hydrothermal fluids to areas where deposition occurs. Proximity to heat sources could also be significant to the creation of deposits.

(1) Overburden:

The overburden varies in thickness throughout the region, and can change even on a local scale. An increased thickness in overburden could produce a gravity variation that could mask the response of high density sulphides, while false anomalies could be caused by a thinning overburden. Away from the deposit the overburden also contains large volumes of Tertiary volcanic tuff that has high magnetic susceptibility¹. Identifying the overburden, even with varying properties, plays an important part of geophysical regional exploration by allowing false anomalies, and features that could potentially mask a deposit, to be determined. In addition, the possible propagation through the overburden of reactivated, deeper faults can give clues about the underlying structure.

(2) Mafic Intrusions:

A small, intrusive, mafic body outcrops to the northeast of the deposit (shown on Figure 3.2). The mafic material has a high density and, due to the high iron content of mafic rocks, could also have high magnetic susceptibility values. These intrusives have the potential for producing physical distributions similar to massive sulphide concentrations when using gravity and magnetic methods. Other methods are needed in order to discriminate between false anomalies and possible mineralization.

Large, magnetically susceptible, intrusive bodies could also be present at depth. It is

¹Another consideration to the physical characterization of the overburden involves the interaction of water. Seasonal changes in the level of the water table may change the conductivity, and IP responses have been observed coinciding with the boundary between areas of contrasting hydraulic permeability within the overburden. Both of these phenomena, while not being entirely explained or understood in this thesis, may have the potential for producing anomalous responses that could be mistaken for responses from mineralized bodies.

expected that the deposit is proximal to a former heat source that would have driven the circulation of hydrothermal fluids. The relationship between a heat source and the deposit maybe very complicated and could depend, among other things, on the availability of fluid conduits. Nevertheless, it is hoped the identification of large intrusive bodies will make the regional picture more complete and allow geologists to make their own hypotheses about the importance of such a feature.

(3) Horst:

The oldest exposed rocks in the region are Upper Triassic phyllite, slate, quartzite and marble of the Zacatecas Formation ([Danielson, 2000]). The host rocks of the Chilitos formation lay unconformably on these metamorphic rocks. These rocks outcrop to the north and south of the deposit on a large horst-block of upthrust geologic section. The formation of the horst is assumed to be more recent than the deposition of the volcanic rocks of the Chilitos Formation, and more recent than the genesis of the deposit. The proximity of favorable host rocks to the surface throughout the horst makes defining the limits of the horst intrinsic to the exploration process. The gravity data show high values where these rocks outcrop and it is expected that these data will be instrumental in defining the horst.

(4) Faults:

Finally, the data imply the presence of several similarly orientated fault sets; faults that may have played a role in the circulation of hydrothermal fluids. If a fault can be shown to be related to the formation of the deposit, then identification of similar faults becomes an important part of exploring the surrounding area for more deposits. Regional inversions can help define these faults and determine their association with the deposit. It is hoped that airborne EM data will give insight into fault orientations nearer the surface, while

gravity and magnetic methods could give information about faults at depth ². While the existence of faults will define the extents of the horst, faults internal to the horst are also of interest.

3.3 Introduction to the inversion results.

The following sections describe the results of inverting the regional geophysical data sets. Theory and methodology behind each data type is contained in Appendix C, and descriptions about the inversion programs used are in (Appendix E). The main features of each inversion model are described and, where appropriate, geologic information is inferred in an attempt to address the issues mentioned above. Information pertaining to each data set is provided in Table D.1, while Table E.1 shows the inversion parameters used for each inversion. Every attempt is made to provide a clear image of the feature or features being discussed in the text although it is realized that this may not be achieved when only a few images of each model are provided. Interpretations are made from the model displayed and also from the numerous other models that were generated with different, yet just as *reasonable*, model objective functions and data misfit values.

3.4 Density-contrast models

Quantec Geofísica de Mexico and Geociencias Consultores collected gravity data at San Nicolás in 1998. Traditional corrections, including instrument drift, tidal, and Bouguer, were applied to the data by Quantec; however, no terrain corrections were included

²The inversions of the gravity and magnetic data are designed, through implementation of the model objective function, to produce small or smooth models. Ideally a blocky model is sought when the aim is the definition of faults. However, physical property transitions in the models may still be indicative of fault contacts and their orientation.

because of the generally low topographic relief in the majority of the survey area (Figure D.1). Neglecting the terrain correction is probably valid except in the vicinity of a hill located at 7000m East and -1000 North on the local coordinate system. After an initial inversion it was realized that data in and around the hill were erroneous. The decision was made to discard them. The resultant data, further altered by removing their mean value, are shown in Figure D.2.

Surface gravity data contain relatively strong information about lateral variations in density but they don't have any inherent depth resolution. Therefore a depth distribution comes from incorporating an additional depth-weighting into the inversion. The weighting procedure has been developed through the use of synthetic forward modeling and inversion [Li and Oldenburg, 1998b]. *GRAV3D*, a program developed by the UBC-GIF, was used for inverting the gravity data. A brief overview of this program is given in Appendix E.2. For the inversion, each datum was supplied with a location and assigned a standard deviation of 0.01 mGal. Topographic relief was included in the inversion procedure. Table E.1 provides more details about the inversion parameters, as it does for all of the inversion results presented in this thesis.

The minimum data misfit that was achieved, without producing too much complexity in the model, was about twenty five times the number of data. This number seems larger than expected even if the data errors were assigned inappropriately³. The reason for this may reside with the discretization of the model. In order to model such a large volume, within a reasonable amount of time, a cell size of 250m by 250m by 100m was used for the majority of the model. The data were collected on intervals ranging from 100m to 50m. As a result, shallow density variations that may have been detected by the data, would

³Errors of 0.01 mGal magnitude correspond to the normal operating sensitivity of most gravimeters. Applying gravity corrections (Instrument, Bouguer, etc.) to the data does not usually introduce significant errors that would degrade the accuracy of the data.

not be modeled by a cell of uniform density-contrast and the model will not be able to fit the data. However, it is expected that large, regional-scale features are acceptably modeled.

3.4.1 Regional gravity inversion results

A plan-view of the regional density-contrast model is shown in Figure 3.3. All the cells that have a density-contrast of 0.1 g/cm^3 or lower have been removed. Outcrop geology is overlaid and a boundary shows the extent of data coverage. A cut-off of 0.1 g/cm^3 is somewhat arbitrary because the transition from areas of low density-contrast to areas of high density-contrast is smooth. However, this does show the distribution of the denser materials, and the western edge of the dominant north-south feature moves only slightly when the cut-off value is varied. Outcropping rocks that correspond to the Chilitos and Zacatecas Formations roughly coincide with the surface projection of the northern half of a large dense north-south body, which approximately extends from -5,000m to 2,000m east, and -10,000m to 10,000m north. The continuation of this high density feature to the south, together with knowledge of the regional geology, suggests the rocks persist under the Tertiary and Quarternary overburden to about -10,000m north. This large feature probably represents uplifted geologic section and is interpreted as a horst. A north-south trending string of high density anomalies seems to mark the western edge of the horst. For the southern half of the horst, the east and west edges show a north-northeast orientation. This orientation is consistent with a mapped fault located between -750m east, 0m north, and 1,500m east, 7,000m north, which is also shown in Figure 3.2).

Four northwest-southeast linear anomalies are also present east of the main horst feature. Westerly projections of these linear features coincide with breaks and inflections along the east and west edges of the horst, suggesting a relationship between the features. The relationship described above can be explained with northwest-southeast faults that

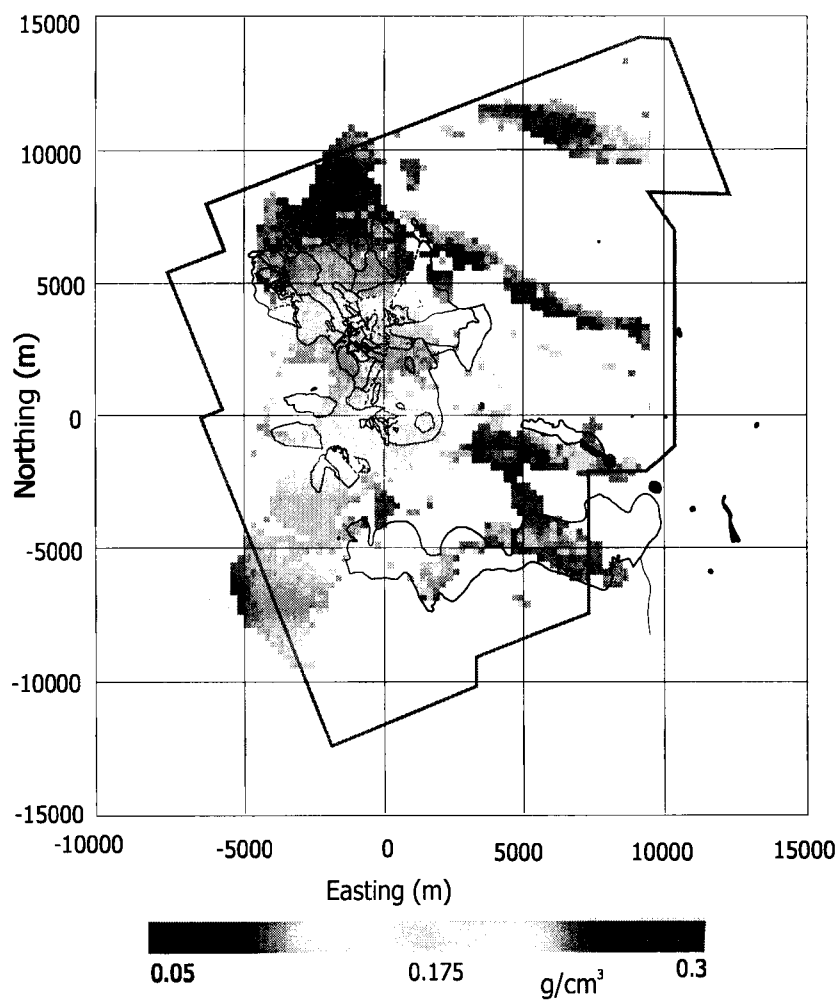


Figure 3.3: Plan view of the regional density contrast model with cells below a value of 0.1 g/cm^3 invisible. Outcrop geology is overlaid, and the boundary of the gravity survey is represented by a purple line. All coordinates are in the local grid system.

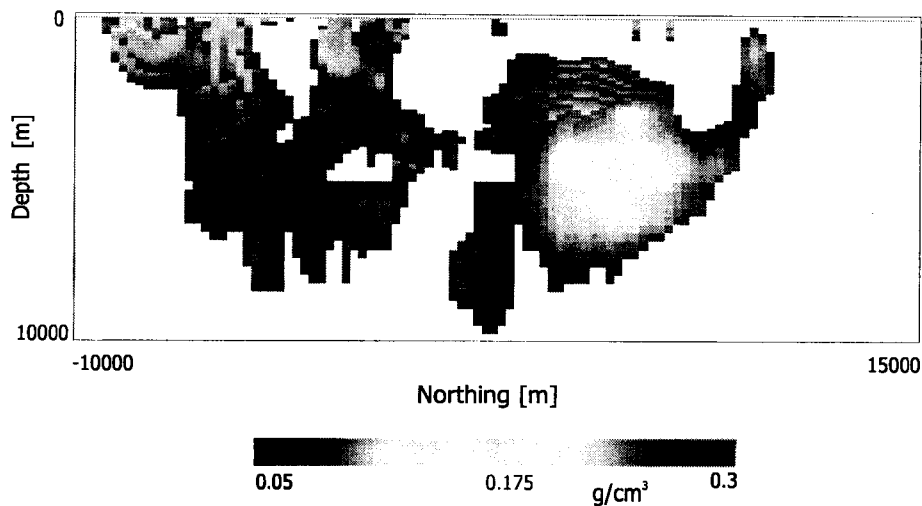


Figure 3.4: West-facing cross section of the regional density contrast model with cells below a value of 0.1 g/cc invisible. This cross-section is located at -1600m east and shows an interpreted horst.

laterally off-set sections of the horst. The linear features themselves, could be faults containing higher density materials.

A west-facing cross-section located at -1600m east shows a side view of the horst (Figure 3.4). The image shows an east projection of material between -1,600m and -3,000m east that has a density-contrast of greater than 0.1 g/cm³. The image shows that the horst, as defined by the density cut-off, is not continuous. In fact three sections of the horst could be interpreted, with the two most southerly sections being joined at depth and the northern section being larger and deeper. This agrees with the idea of northwest-southeast faults being present, as the faults could have been responsible for changing the elevation of the rocks as well as offsetting them laterally.

Figure 3.5 shows the inferred locations of faults in the area based on evaluation of this model. The straight trace of the northwest-southeast faults, and the cross-cutting of the east and west flanks of the horst, imply an age relationship between faults in the

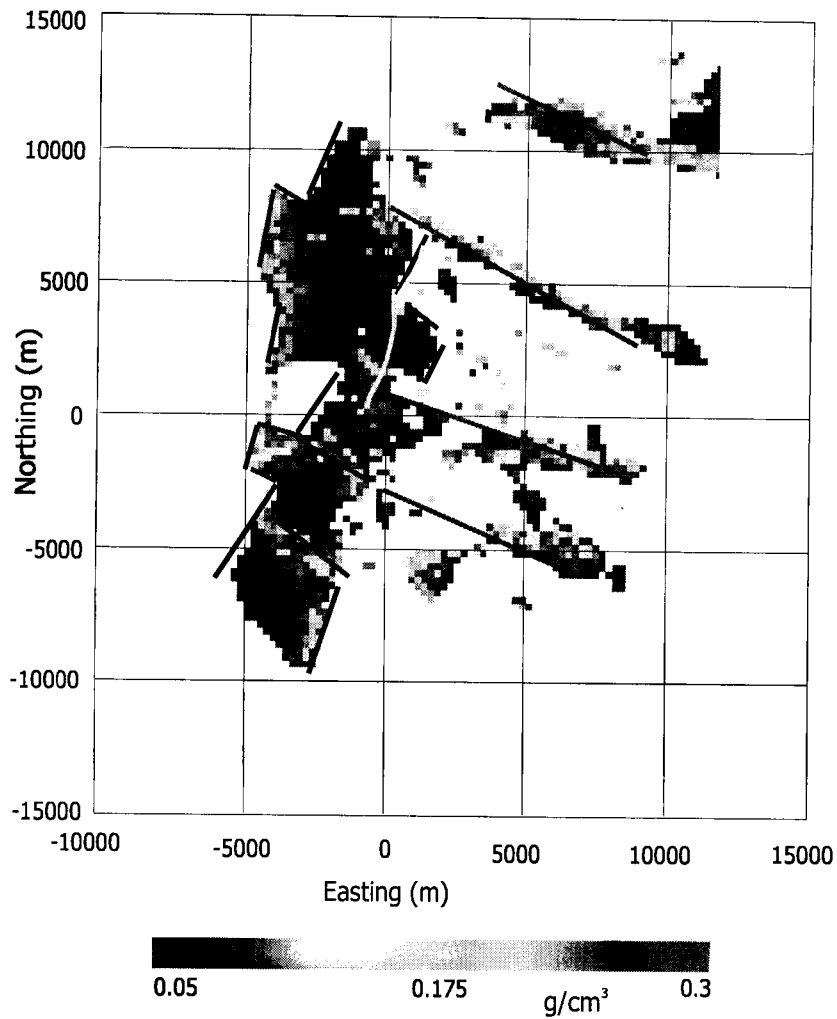


Figure 3.5: Surface projection of the regional density-contrast model with interpreted faults overlaid. Northwest-southeast trending faults are shown in red, while north-northeast trending faults are shown in black. The mapped fault is shown in yellow. Material with a density-contrast of less than 0.1 g/cm^3 has been removed.

area, with the north-northeast fault being older than the northwest-southeast faults.

Figure 3.6 shows a horizontal slice through the regional density contrast model at an elevation of 1700m (400m below the surface). The main features are the linear trends of high density that correspond with the east and west edges of the horst and northwest-southeast trending linear features. It is interesting that high densities would be associated with faults. An explanation for this could be the presence of higher density material within faults. This could be representative of igneous rocks that were intruded along the fault.

An anomaly located at -1,600m east and -400m north represents the deposit. This is a small anomaly and selecting this feature over the others would be difficult. Based on the interpreted faults that are shown in Figure 3.5, the deposit lies within a coherent block of bedrock. This suggests that any faults associated with the genesis of the deposit are not detected and are probably older than the formation of the horst. Searching for density anomalies that are not coincident with younger, interpreted faults seems like a reasonable way to search for massive sulphides. However, other interpretations might place faults directly through the deposit.

A north-facing cross-section through the regional density-contrast model is shown in Figure 3.7. The negative density-contrast located near the surface at 6,000m east is in the vicinity of the hill where data were removed prior to inverting, and is thought to be an artifact created by not terrain correcting the data around the hill. The deposit is shown as a discrete density high near the surface. No significant depth extent to the deposit has been modeled. Another isolated density high is shown at about 4,000m east, this is one of the northwest-southeast linear features. A lateral change in density contrast at -6,000m east represents a fault, which is probably representative of the near vertical western flank of the horst block. The eastern flank of the horst block maybe dipping to the east.

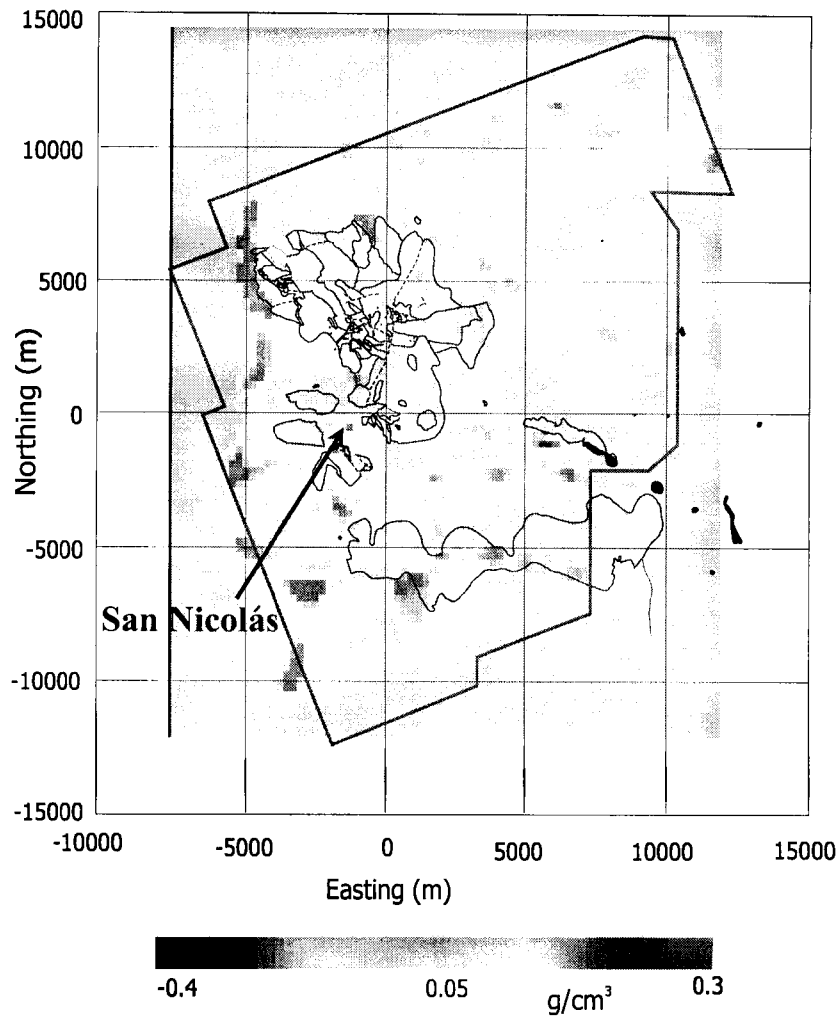


Figure 3.6: Horizontal slice through the regional density-contrast model at an elevation of 1700m (400m below the surface at the location of San Nicolás). An outline of geology outcrops is overlaid, and the extent of data coverage is shown by a purple line.

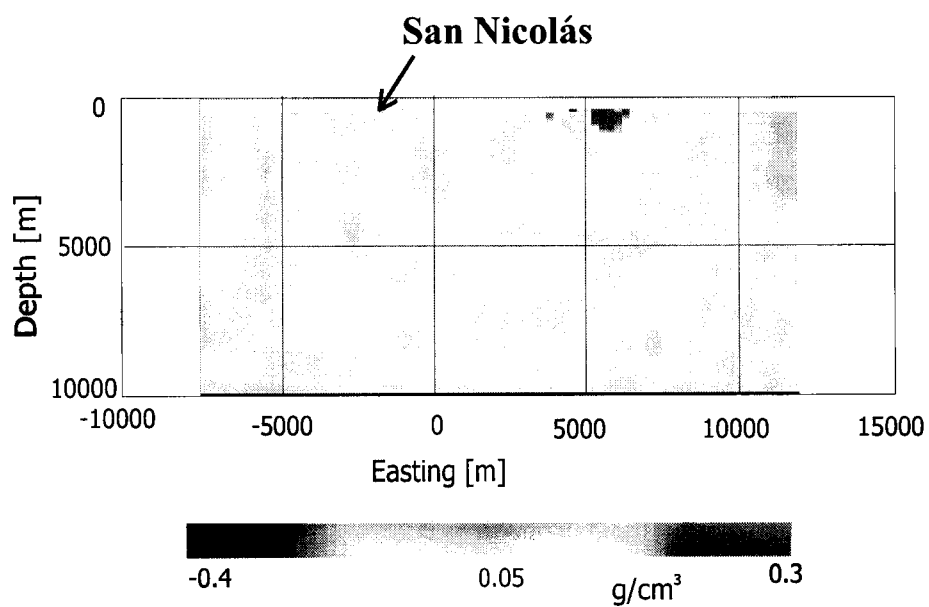


Figure 3.7: North facing cross-section through a regional density contrast model at a northing of -400m (coincident with the deposit).

At depths greater than 4 kilometers, a large amount of higher density material is present in the model, and it is coincident with the bounds of the horst block as defined from the shallower parts of the model. The deep large high density features, needed to explain the high gravity anomalies shown in the data, could be uplifted limestone, mudstone, or metamorphic rocks of the Zacatecas Formation. All of these rocks could have higher densities than the overlying volcanic and sedimentary units. In addition, the deep-rooted, large bodies at depth could represent large intrusive bodies rich in mafic material and high in density.

3.4.2 Summary of regional gravity results

The density contrast model is consistent with the previously interpreted geology and has contributed to regional geological understanding with the following interpretations: (1) the delineation of a horst block, (2) the interpretation of faults below the overburden that cross-cut the horst (3) a possible discrimination of discrete high density anomalies similar in nature to the deposit, and (4) the identification of deep-rooted bodies that could represent large amounts of sedimentary or intrusive igneous rocks, or both.

3.5 Magnetic susceptibility models

This section addresses the the inversion of regional magnetic data.

Total-field magnetic data is inverted with the aim of producing a three-dimensional magnetic susceptibility distribution of the earth. Background information on the magnetic method is given in Appendix C.2. Airborne methods are employed at San Nicolás to provide generalized geologic information on large areas in a short period of time as well as identify anomalous regions of high susceptibility that might indicate mineralization. The data are displayed in Appendix D. Three data sets are inverted and subsequently

combined to form a single susceptibility distribution model of the subsurface.

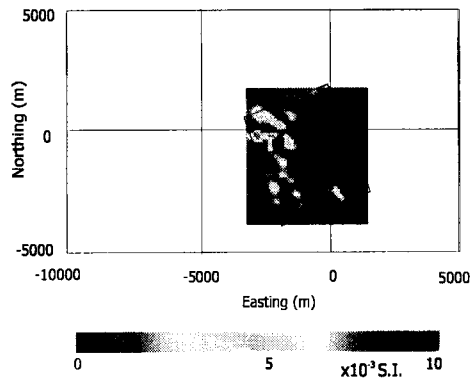
Surface magnetic data also have no inherent depth resolution and so, as with the gravity inversion, depth weighting is needed.

3.5.1 Data

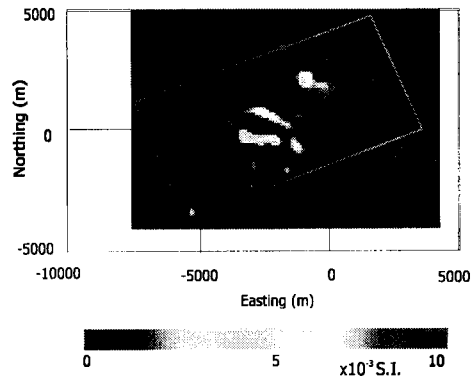
The three airborne magnetic data sets include:

1. Digheem collected airborne magnetic data while also collecting frequency-domain electromagnetic and radiometric data. The Digheem data has the most extensive coverage of the three data sets. Northwest-southwest trending features that can be interpreted from the regional density-contrast models are also apparent in Figure D.3. This *texture* betrays the regional structural setting. A subset of observations from the Digheem data set was chosen for inversion (Figure D.4).
2. The airborne magnetic data, shown in Figure D.5 were acquired by BHP at the same time as airborne gravity-gradient data were being recorded with their Falcon system.
3. Fugro collected airborne magnetic data (Figure D.6) while collecting time-domain electromagnetic data with their Questem system.

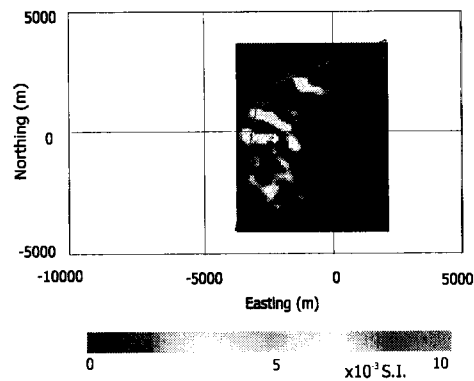
The coverage and survey specifications for each data set are slightly different although similar features are observed. The deposit is represented in the data by a subtle magnetic high at about -1600m east and -400m north, and the response from highly magnetic Tertiary tuffs are seen to the north west of the deposit. All of the data sets had regional fields removed prior to inversion and standard deviations of 10nT were assigned to each datum.



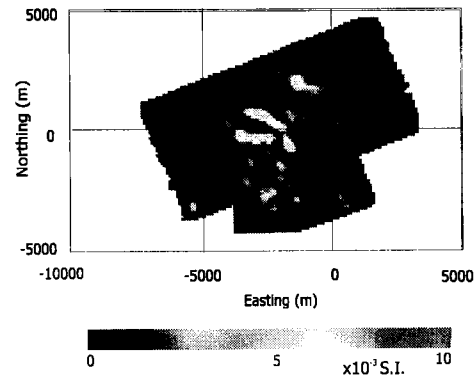
(a) Dighem susceptibility model.



(b) BHP susceptibility model.



(c) Fugro susceptibility model.



(d) Composite susceptibility model.

Figure 3.8: Comparison of individual magnetic susceptibility inversion models with composite model of mean values. The models are shown in plan-view looking at a depth of 300m (within the depth of the deposit).

3.5.2 Composite susceptibility model

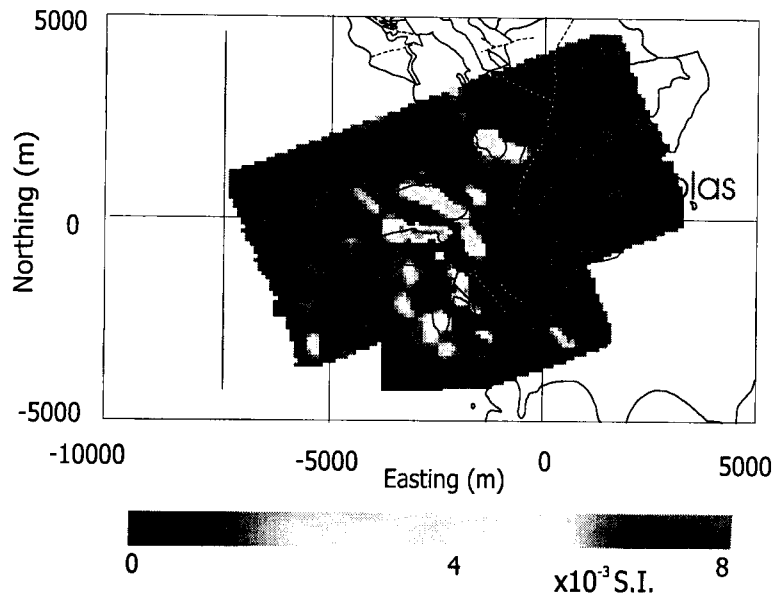
Initially an attempt was made to invert all three data sets simultaneously. The results from this inversion were not successful; this is possibly due to the different regional fields that were removed prior to inverting. Consequently each data set was inverted separately. A single susceptibility model was produced from the three individual inversion models by placing them all on the same mesh and calculating mean value of each cell. Volumes of the model that do not lie directly below the data coverage have been removed. Figure 3.5.1 shows a comparison between the three individual models and the resulting model.

3.5.3 Regional susceptibility model analysis

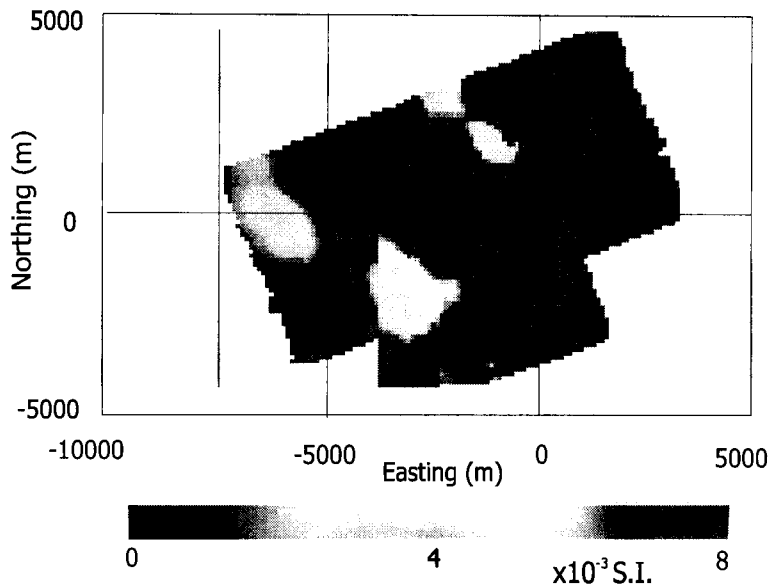
Figure 3.9(a) shows a plan view of the regional magnetic susceptibility model at a depth of 300m⁴. At this depth the deposit is modeled as a discrete anomaly located at about -1700m east and -500m south (labeled). However, features with the highest susceptibility include two elongated regions to the west and northwest of the deposit (between -4000m and -2000m east, and -500 and 750m north), and a discrete anomaly located at -1500m east and 2000m north. The two elongate features correspond with out-cropping Tertiary breccia. The other anomaly to the north (located at -1500m east and 2000m north) corresponds with out-cropping mafic volcanics. Except for a small anomaly located at -2000m east and -2000m north, other mafic volcanic outcrops have low susceptibilities. Several other low magnitude anomalies of unknown origin are also observed at this depth.

At a depth of 1200m, as shown in Figure 3.9(b), four large features are observed. An anomaly at -1500m east and 2000m north is probably an extension of the anomaly seen at 300m depth, and could represent an intrusive body. A large anomaly is observed at this depth, located between -4000 and -2000m, and -1000m and -3000m. It is interesting

⁴The deposit is approximately located between depths of 200m and 500m.



(a) Composite regional susceptibility model:
300m below the surface.



(b) Composite regional susceptibility model:
1200m below the surface.

Figure 3.9: Plan-view of horizontal slices through the composite magnetic susceptibility model. The upper panel shows a slice at 300m depth, while the lower panel shows a slice through the model at 1200m depth.

to note that a shallow ring of small, moderate (2×10^{-3} S.I.) susceptibility concentrations coincides with the edge of this body (Figure 3.9(a)). This volume of anomalous susceptibility could be a deep rooted intrusion, and the shallower anomalies may have originated from this deep source. Alternatively, this susceptible body might represent magnetic units below the Chilitos Formation.

The anomaly located between -4000 and -2000m, and -1000m and -3000m also has a distinct northern edge that is orientated in a northwest-southeast direction. This orientation corresponds to those interpreted from the regional density-contrast model. This suggests that it was in place prior to a faulting event or its emplacement was controlled by an existing fault. The western edge of this susceptible body is also distinct but is likely an artifact of combining the three separate susceptibility models.

Anomalous susceptibility distributions have also been modeled at -6000m east, 0m north, and -2500m east, 3000m north. Both of these are constructed to fit incomplete anomalies at the edge of the data coverage. They could also represent deep intrusive bodies or magnetic sequences within the Zacatecas Formation.

Some of the features mentioned above are observed in the north- and west-facing cross-sections of Figure 3.10. The north-facing cross-section is located at -400m south and corresponds to the location of the deposit. The deposit is clearly represented by a discrete susceptibility anomaly (labeled). The anomaly near -4000m east is the southern boundary of the Tertiary tuffaceous unit. The large anomaly in the far west of the cross-section, between depths of 100m and 2500m, corresponds to the most westerly anomaly seen in figure 3.9(b). This anomaly does not extend to great depths although it might extend beyond the modeled volume.

The west-facing cross section shows the deposit which is connected to a smaller magnetic feature to the north. The high susceptibility concentration at -1500m east and 2000m north is clearly seen on this cross-section, along with the large deep body to the

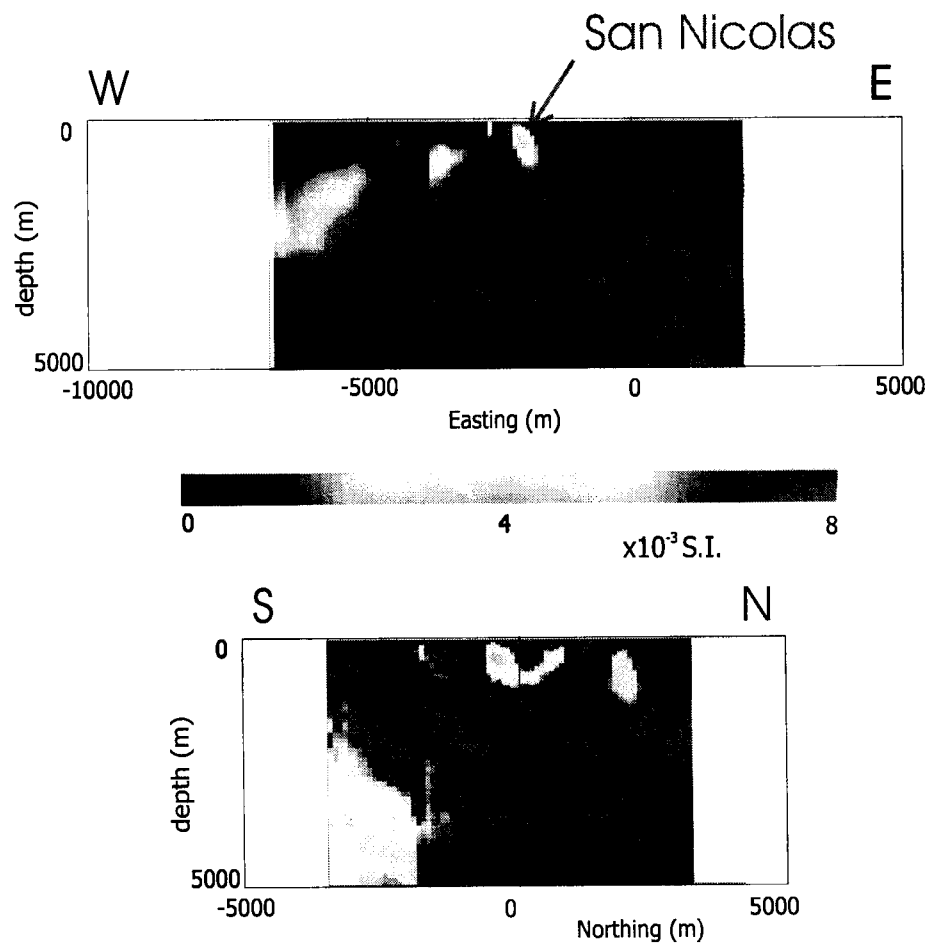


Figure 3.10: North- and west-facing cross-sections (upper and lower sections respectively) through the magnetic susceptibility model generated by combining inversion results from Dighem, Fugro, and BHP airborne magnetic data. The north-facing cross-section is positioned at -400m south, while the west-facing cross-section is positioned at -1600m east.

south. This large, most southerly anomaly is the only anomaly that persists below a depth of 2500m and continues to the bottom of the mesh at a depth greater than 5000m.

3.5.4 Geologic interpretation

Mafic rocks make up part of the volcanic section within which the deposit was formed. The anomalies observed in the mafic volcanics could represent concentrations of magnetite or pyrrhotite associated with massive sulphides, as seen at San Nicolás. Alternatively, the mafic rocks may naturally exhibit a higher magnetite content⁵.

The Tertiary volcanic breccia, while predominately being comprised of subordinate volcanic debris, contain some ash-flow tuffs that are magnetically susceptible⁶. These tuffs are responsible for the elongate anomalies seen in figure 3.9(a).

The extension of the susceptibility values to the north of the deposit may be indicative of an extension of mineralization and warrants further investigation.

The large susceptibility features seen at depth could indicate large intrusive bodies. This is especially true when the relationship between some of the deeper and shallower anomalies are considered. It seems unlikely that thick sections of susceptible, unaltered volcanic rocks reside at depth. Evidence for the alteration or metamorphism of rocks on a regional scale comes from the rocks of the Zacatecas Formation, which underly most of the region.

Some linear features can be seen in the model, such as an approximate north-northwest line of susceptibility anomalies, which includes the deposit, (seen in figure 3.9(a)), and well defined boundaries to some features that share the same orientation. These linear

⁵Volcanic rocks in the San Nicolás region generally exhibit subdued magnetic susceptibility values, possibly due to the effects of wide spread alteration. Less altered volcanic rocks could possess higher magnetic susceptibility values.

⁶Drill hole Sal-133 intersected Tertiary breccia to a depth of more than 400m [Jansen, 2000]. Susceptibility measurement made on the core exceed 2×10^{-3} S.I. between a depth of 200m and 300m (Appendix B.3.1).

features might represent faults although collaborative information would be necessary to confirm this.

3.6 Regional conductivity models

Airborne electromagnetic data are used to determine the regional conductivity structure of the subsurface and are usually acquired contemporaneously with airborne magnetic data. The development of airborne electromagnetic systems has been hand-in-hand with the exploration of massive sulphides as the targets are usually very conductive and can be hosted in quite resistive rocks. Dighem's frequency-domain system and Fugro's Questem time-domain system were employed at San Nicolás and modeling of both these data sets was performed.

3.6.1 Frequency-domain data

An overview of the frequency-domain, airborne electromagnetic method is given in Appendix C.3.1. The overview describes the nature of the data and Appendix E.5 gives a description of *EM1DFM*, the algorithm that is used to invert frequency-domain EM data to recover a one-dimensional conductivity distribution at each station. The data cover a very large area (Appendix D) and it was hoped that a regional conductivity model of the earth could be produced to complement the 3D density contrast and magnetic susceptibility information that was obtained by inverting the gravity and magnetic data. Unfortunately the inversions have shown that the data are not sensitive to conductive features below the overburden. This lack of sensitivity was determined by performing several inversions with different reference models and observing the depth at which the model reverts to the value of the reference model. Both GCV and L-curve methods of choosing the appropriate misfit were used and generate similar outcomes. This suggests

that inappropriate error assignments do not contribute to the shallow penetration. The depth at which the data were deemed insensitive to the model was about 150m.

The reason for the observed sensitivity of this system is probably because of the high overburden conductivity, the frequencies employed, and the geometry of the system. Insight into the observed depth of penetration can be explained by considering the *skin depth* for a plane wave in a homogenous medium [Ward and Hohmann, 1987]⁷. For a 30 Ohm-m medium at the lowest frequency of the system (466 Hz), the skin depth is approximately 127m. This implies that the signal has dramatically deteriorated by the time it passes through the conductive over burden and reaches the volcanic sequence. At higher frequencies the signal will be even more deteriorated by passing through the overburden.

Aside from the lack of penetration, the survey is useful in that it covers a large area and provides detailed information about the thickness of the overburden. Additionally features can be interpreted in the overburden that might be the manifestation of deeper structures. After the 1D inversion models were produced an interpolation routine was used to construct a 3D model⁸. The 3D models are presented below in the same manner that the 3D density and magnetic susceptibility models are presented, although it should be kept in mind that they have been produced by stitching together many 1D conductivity models.

⁷The skin depth (δ) is the depth at which a sinusoidal electric field (from a planar source in a whole-space) diminishes to about 37% of it's original strength, and is approximately stated as $\delta = 503 \sqrt{\frac{\rho}{f}}$ [m], where ρ is the resistivity, and f is the frequency [Ward and Hohmann, 1987]. In the Digheem system we are not dealing with plane waves but a dipolar source. Also the geometry between source and receiver influences depth of penetration. However, this approximation can give an idea of the strength of different frequency signals in different mediums.

⁸The interpolation routine uses triangulation to assign values to cells at each depth. The routine was applied to the Log_{10} of the conductivity values in order to reduce homogeneity artifacts.

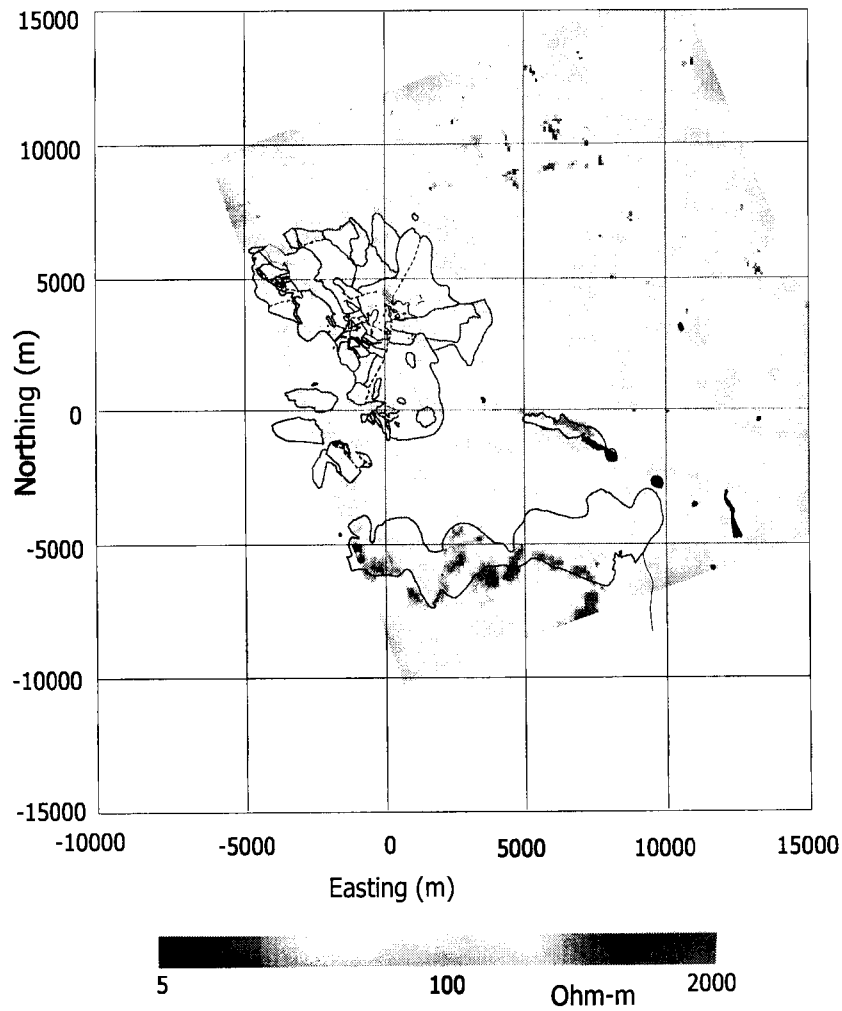


Figure 3.11: Horizontal slice 60m below the surface of the 3D resistivity model generated by interpolating between lines of 1D models inverted from Dighem frequency-domain airborne electromagnetic data.

A horizontal slice through the stitched conductivity⁹ model is shown in Figure 3.11. The slice is at a depth of 60m below the surface and it clearly maps out the outcropping geology with high resistivity values, and the overburden with low resistivity values. Outcropping volcanic units of the Tertiary overburden exhibit low resistivity values. This is especially significant in the south, where a tuffaceous unit outcrops. No anomaly is detected in the vicinity of the San Nicolás deposit.

Although this model doesn't seem that useful in the direct search for massive sulphides, the method can be used as a reconnaissance mapping tool for determining the extent of cover and inferring regional structures. Linear features seen in the model have the same orientation as structures observed in the regional density contrast model. Specifically, three discrete resistive regions are defined that correspond well with the limits of the horst (as defined by the density contrast model). Northwest-southeast trending linear resistivity contrasts are seen that may also relate to the structural nature of the region.

3.6.2 Time-domain data

Airborne time-domain electromagnetic (TEM) surveys are used at the regional or reconnaissance stage of exploration programs. They often enjoy more success than frequency-domain systems at resolving the conductivity structure of the earth below conductive overburdens, and can be used to detect conductors as well as delineate structures. This section details the inversion of such a data set collected by Fugro in 1998. A cursory investigation is also made into the possible cause of negative data that is observed at late time-channels.

Time-domain electromagnetic data were collected using Fugro's Questem system. The nature of the data and how they are collected is discussed in Appendix C.3.2, while

⁹The conductivity (σ) models are presented as resistivity (where $\rho = 1/\sigma$) for consistency with models presented later in the thesis.

the specifications particular to this survey are given in Table D.1. The x-, y-, and z-components of the rate of change of the magnetic field ($\frac{dB}{dt}$) are measured as a function of time, but only the vertical component is considered here. One stacked time-decay series ($\frac{dB_z(t)}{dt}$ measured in fifteen time-windows) is collected for each station. Each time-decay series was inverted using *TEM1D* to recover a layered-earth conductivity model. The resulting 1D models were combined, using an interpolation routine, to produce a 3D conductivity structure of the earth. An overview of *TEM1D* is given in Appendix E.6.

A horizontal slice through the “stitched” conductivity model at a depth of 200m is presented in Figure 3.12. The 3D conductivity model shows the distribution of conductive overburden and maps out the outcrops quite effectively, similar to the frequency-domain results presented in Section 3.6.1. However, not much structure is modeled below the overburden, and the absence of a conductive body in the vicinity of the deposit is conspicuous.

It was expected that data collected with this system would be sensitive to conductivity structure below the overburden¹⁰, and would probably result in a more deeply resolved inversion model than that recovered from the inversion of the Digheem frequency-domain data. The reason for this inconsistency might lie in how standard deviations were assigned to the data to account for the “noise”.

¹⁰The depth of the peak amplitude of a decaying and propagating, impulsive, electromagnetic plane-wave, in a whole-space, at a given time, can be stated as: $z = \left(\frac{2t}{\mu\sigma}\right)^{1/2}$ [m] [Ward and Hohmann, 1987]. This is sometimes referred to as the diffusion depth. Although this depth would be different for a layered earth, and a magnetic dipole source, as used in the Questem survey, this gives a general idea of the penetration of time-domain electromagnetic systems. For an overburden with a conductivity of 20 Ohm-m, the diffusion depth is 280m for late times, which is well below the top of the deposit.

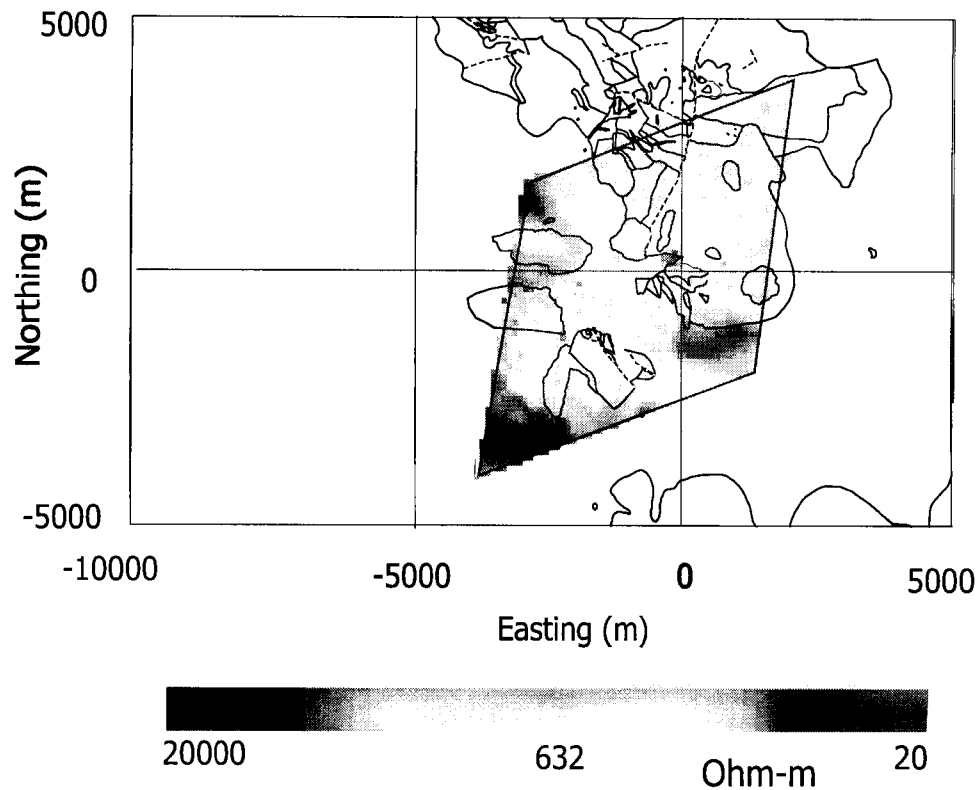


Figure 3.12: Plan view of a slice through the 3-D resistivity model generated by interpolating between lines of 1D models from the inversion of Questem time-domain airborne electromagnetic data. The slice (at a depth of 200m below the surface) shows, in red, the thicker areas of overburden that continue to this depth. The underlying, more resistive, volcanic rocks are represented by cooler colors.

Negative data

A significant amount of negative data in late time-channels, initially thought to be noise, was present throughout the survey. Large errors were assigned to these data. Assigning these errors had the effect of making the inverted models insensitive to the negative data, but also to structures at depth. As a result, the resulting model only detects the shallow feature, namely the overburden.

Sixty percent of the stations contained negative data, usually at late time-windows. However, sixteen percent of the stations contained negative data from time-windows that correspond to earlier times (time-channels 8 to 11); channels that usually contain little noise. Furthermore, these negative values persisted locally from station to station and appear to be above the noise levels, particularly over the deposit. A line of Questem, z-component data is presented in Figure 3.13, along with a total magnetic field profile, and the results of inverting the data using *TEM1D*. A significant negative z-component response is observed over the deposit in time-channels 10 to 15 (annotated with a number 1). Smaller depressions that are possibly above the noise level, are also observed at late times (labeled with 2, 3, and 3). Line-by-line inspection of the whole data set showed that no other late-time response was as compelling as that seen over the deposit.

An example of time-decay data that exhibit negative values for one station over the deposit is provided in Table 3.1 and plotted in Figure 3.14. Although this time-decay series is from over the deposit, it's characteristics are typical of other time-decay series in that it exhibits a faster decay at late time and then changes sign. The negative data are part of a coherent decay pattern observed in the time-decay series.

It is strongly suspected that the negative data are a real part of the signal and are not noise. Such negative data are commonly referred to as *IP effects*. Negative data occur when the rate of change of the magnetic field, as measure by a receiver coil, opposes

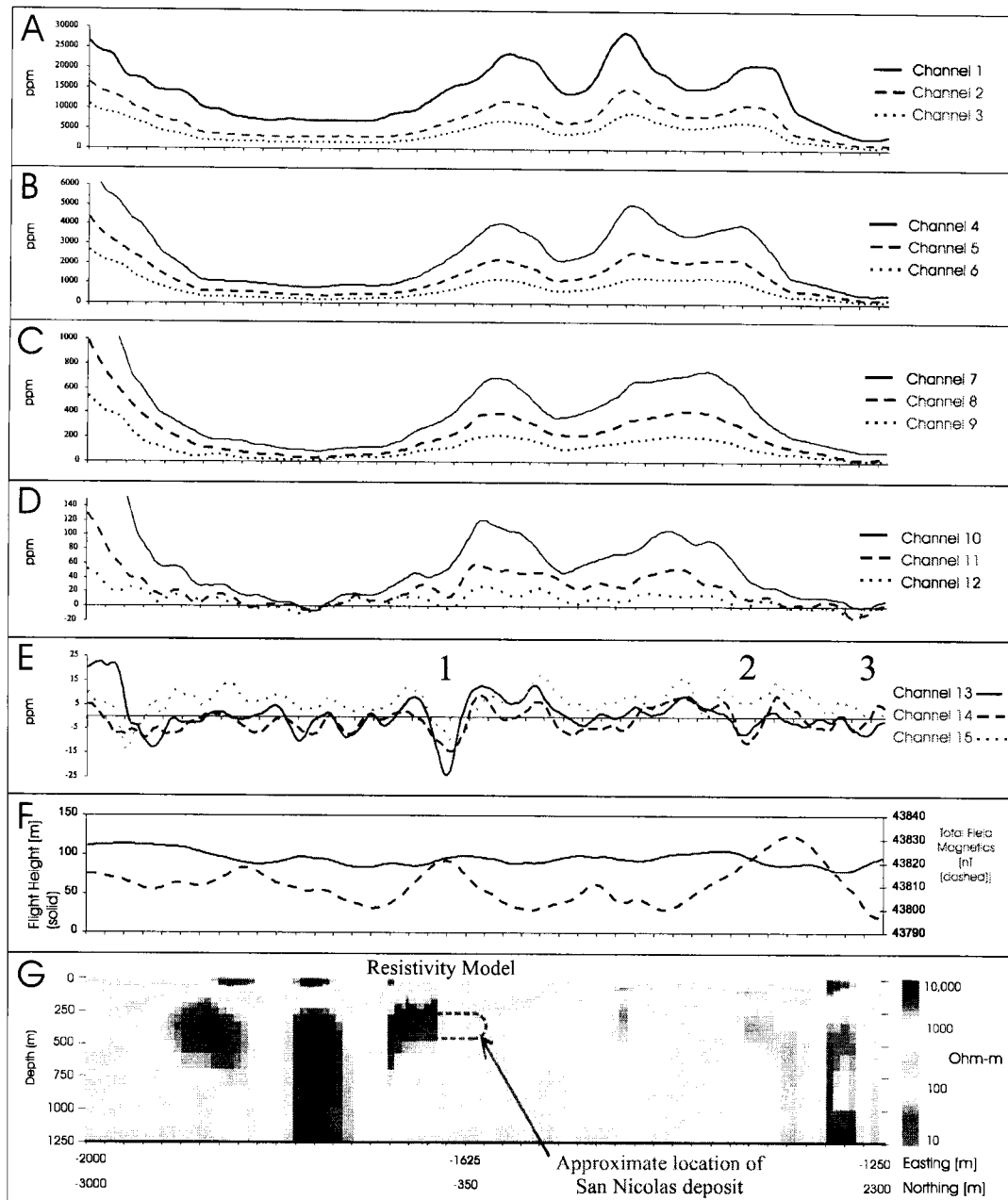


Figure 3.13: Vertical component, time-channel data, flight height, total field magnetics, and resistivity inversion model for line 20110 of the Questem data. The data are presented as parts-per-million of the peak z-component of the waveform, with channels 1 and 15 corresponding to the earliest and latest times respectively. The inversion model is produced by “stitching” together 1D models. The location of the deposit corresponds to a coherent pattern of negative data values in time-channels 12 through 15 (1), and a magnetic high. Other regions of negative data that could be above the noise level are also observed (2, 3 and 4).

Time Channel	Window center time (msec)	z-component (μV)	z-component (ppm)
1	0.375	0.080425279	15834.6
2	0.531	0.038734626	7626.3
3	0.688	0.022706049	4470.5
4	0.922	0.012765264	2513.3
5	1.234	0.006915682	1361.6
6	1.625	0.003740238	736.4
7	2.094	0.002087504	411.0
8	2.641	0.001168190	230.0
9	3.344	0.000640473	126.1
10	4.281	0.000316427	62.3
11	5.531	0.000087868	17.3
12	7.094	-0.000008127	-1.6
13	9.047	-0.000083297	-16.4
14	11.469	-0.000070091	-13.8
15	14.359	-0.000042156	-8.3

Table 3.1: Questem time-decay ($\frac{dB_z}{dt}$) data for one station located over the deposit. The center time of each time channel is given (time after turn-off of primary field) along with the z-component data in micro-volts and in ppm, as normalized by the peak z-component of the primary field.

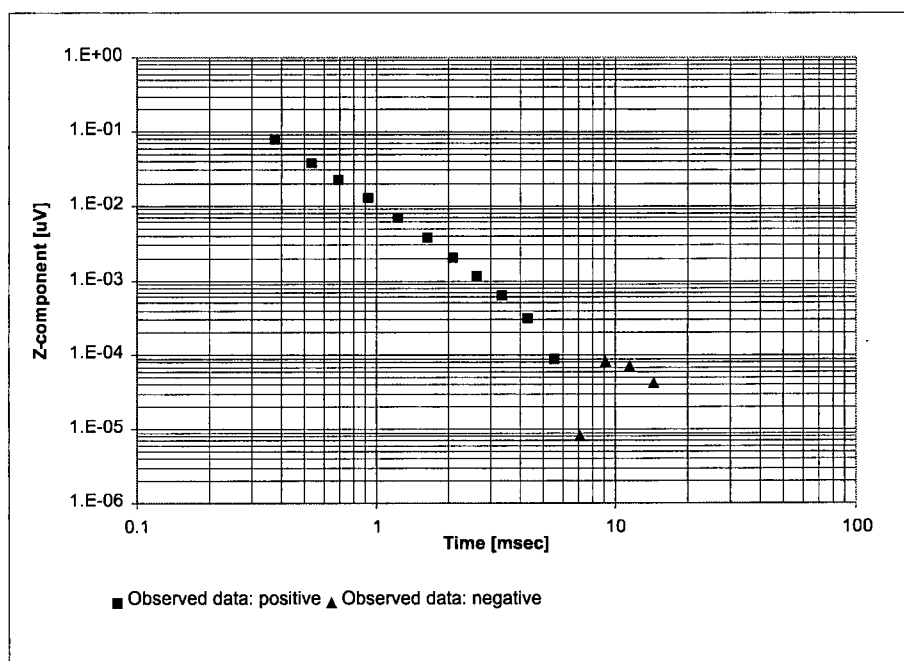


Figure 3.14: Logarithmic plot of Questem vertical-component time-decay data given in Table 3.1. Positive data are shown by squares, and negative data (the last four) are shown with triangles. The data are from over the deposit and show a typical decay pattern for stations that contains negative data.

that of the primary field. This could be caused by geometric effects distorting the field, or by the presence of polarizable material that reverses the field. In addition, for 1D conductivity structures with realistic ranges of magnetic susceptibility, negative data can only be produced when polarizable is present material [Weidelt, 1982]. Flis, Newman, and Hohmann [Flis et al., 1989] have also shown IP effects in TEM data to be caused by polarizable, or chargeable, material at the surface. As the massive sulphides at San Nicolás are expected to be chargeable, further investigation is performed on the negative data, through the use of 1D forward modeling.

3.6.3 Forward modeling the Questem data

Forward modeling is performed in an attempt to reproduce negative data that are observed at late times in the Questem data set. Two investigations are undertaken; one to determine if the geometry of the earth can produce the negative data, and one to determine the role of polarizable material in producing these data.

Forward modeling a 3D conductivity structure

The affect of geometry is investigated by forward modeling synthetic Questem data over a three-dimensional conductivity structure. The conductivity model is generated by digitizing geologic cross-sections and assigning conductivity values to different rock types based on physical property measurements (Appendix B.4). The conductivity sections are interpolated to produce a three-dimensional model (Figure B.5 shows a north-facing cross-section through the conductivity model at -400m south). *EH3DTD* (described in Appendix E.9) was used to collect synthetic $\frac{dB_z}{dt}$ data at twenty four stations, spaced 100m apart along one east-west line, over the conductivity structure. The same survey parameters (transmitter-receiver geometry, wave-form, time-channels) were used to collect the synthetic data as those employed by the Questem system.

All of the synthetic data were positive. Moreover, the synthetic time-decay data over the deposit exhibit a slower decay at late times, which is in contrast to that observed in the field data. The decay-rate is also slower than that observed over a half-space. This result suggests that the negative data were not due to the geometry of the subsurface.

1D forward modeling of polarizable material

The role polarizable material might play in producing negative data is investigated by forward modeling one-dimensional, layered-earth structures that contain polarizable layers¹¹.

A 1D forward modeling program was available that uses the Cole-Cole [Pelton et al., 1978] model to represent complex conductivities¹². Models were produced in an attempt to match time-decay data observed at one station over the deposit. The time-decay chosen for forward modeling analysis was recorded with the receiver located at -1600N and -500N and is representative of many of the decay curves over the deposit. This data is presented in Table 3.1 Figure 3.14. The observed data show a reversal in polarity at the twelfth time-channel (about 7 msec).

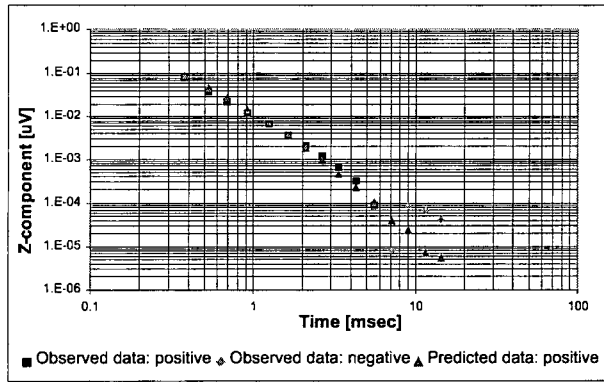
Figure 3.15 shows observed and predicted time-decay data and the corresponding real

¹¹In this case, polarization refers to a relaxation mechanism at the pore-scale that occurs when alternating, or time-dependent currents are present. Polarizable material has frequency-dependent, or complex, conductivity.

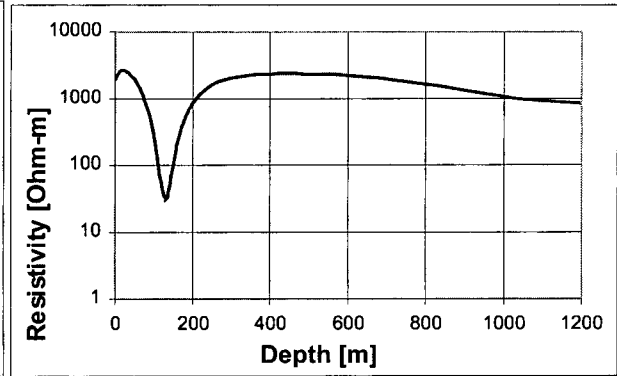
¹²Complex conductivity is introduced in the model by using the empirical Cole-Cole model. This is a simple relaxation model that allows conductivity to change with frequency based on a few parameters. An equation for the Cole-Cole model is given below:

$$Z(\omega) = R_0 \left[1 - m \left(1 - \frac{1}{1 + (i\omega\tau)^c} \right) \right]$$

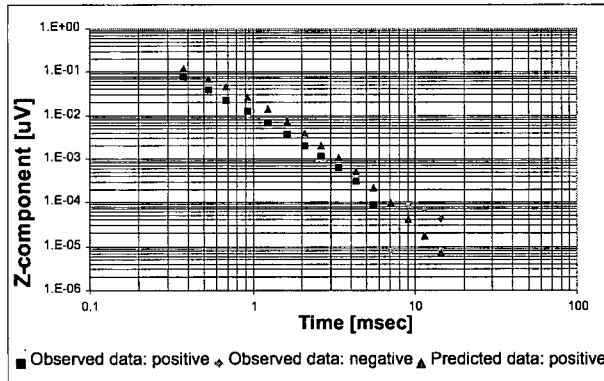
where $Z(\omega)$ is the frequency-dependent impedance, ω is the frequency, R_0 is the zero-frequency resistivity, c controls the rate at which the relaxation occurs, τ is a time constant that determines when the relaxation occurs, and m is the chargeability, and relates the zero-frequency resistivity to the resistivity at infinite frequency.



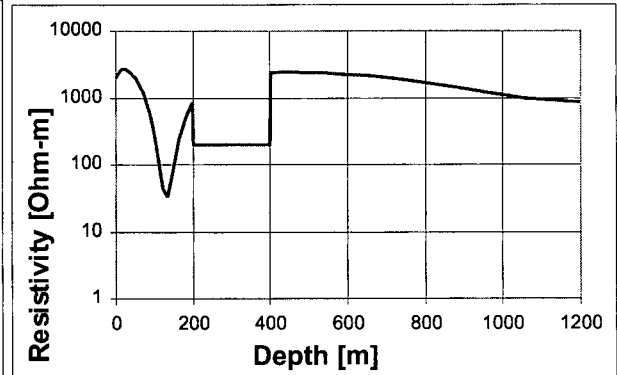
(a) Observed and predicted time-decay data. The predicted data are from the *TEM1D* inversion.



(b) *TEM1D* inverted conductivity model.



(c) Observed and predicted time-decay data for conductivity model with additional conductor at depth.



(d) Conductivity model with additional conductor at depth.

Figure 3.15: Observed and predicted time-decay data, and corresponding 1D conductivity models for Questem station located over the deposit. Frequency-independent conductivity layers do not reproduce the observed negative data.

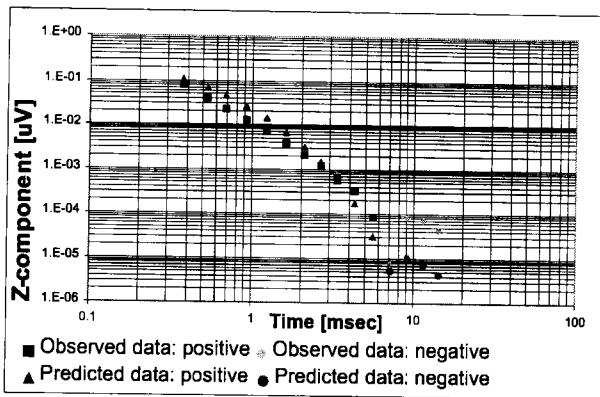
(frequency independent) one-dimensional conductivity models. Figure 3.15(a) shows the observed field data and predicted data from the inversion using *TEM1D*. The predicted data fit the early time data very well, negative signals are not produced at late times. As a result the overburden is well represented in the model, but most depth information is not realized.

Based on the expected conductivity structure of the deposit from physical property measurements, a conductive layer of 200 Ohm-m was introduced into the model between 200m and 400m depth. This is shown in Figure 3.15(d), and the data resulting from this model is shown in Figure 3.15(c). The introduction of a conductor into the model increases the magnitude of the data, and a worse fit to the observations is produced.

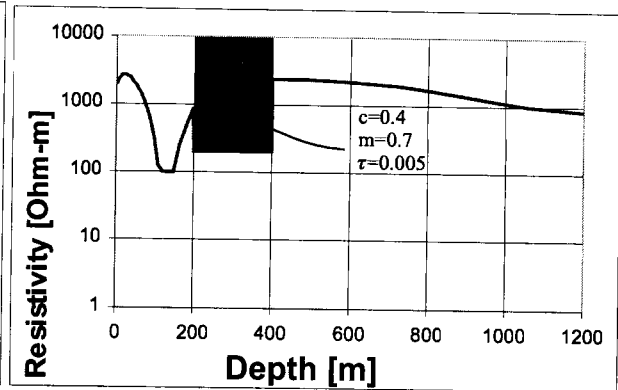
When the conductive layer at depth is assigned a complex (frequency-dependent) conductivity, the data values are reduced, negative data are produced, and the character of the observed time-decay is emulated. The results from two complex conductivity models are shown in Figure 3.16.

Model "A" (Figure 3.16(b)) consists of the one-dimensional real conductivity distribution that was recovered from the inversion with a layer of complex conductivity between depths of 200m and 400m. The conductivity of the overburden was also decreased to 100 Ohm-m in order to more closely fit the early-time data. This model produces data (Figure 3.16(a)) that emulate the observed time-decay data. The predicted late-time data are negative and have the same decay pattern as the observed data. The early-time data are also a *good* fit.

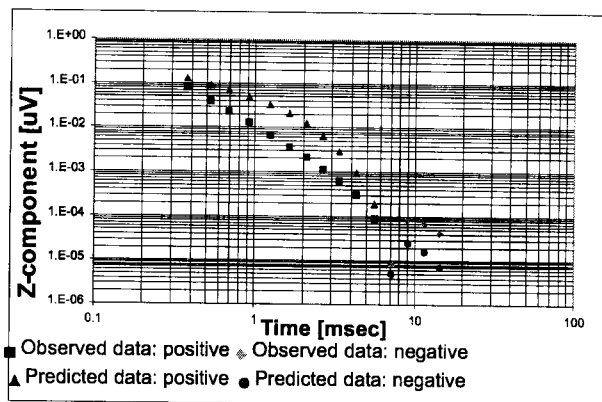
Model "B" (Figure 3.16(d)) consists of the one-dimensional real conductivity distribution that was recovered from the inversion with a layer of complex conductivity between depths of 350m and 480m. This model produces predicted data (Figure 3.16(c)) that also emulate the observed time-decay data. The predicted late-time data are negative and have the same decay pattern as the observed data. The late time data also have a better



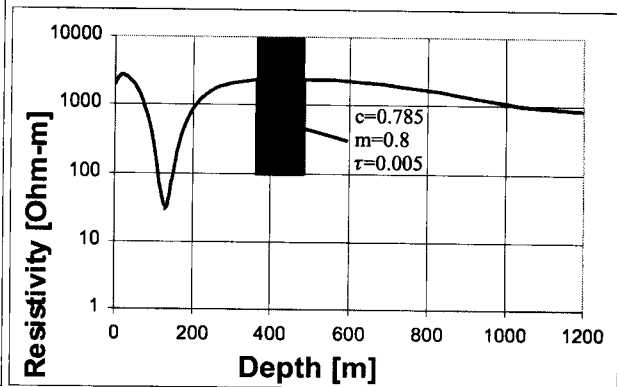
(a) Observed field data and predicted time-decay data for complex conductivity model A.



(b) Complex conductivity model A. A layer of complex conductivity (Cole-Cole parameters are shown) is present between depths of 200m and 400m.



(c) Observed field data and predicted time-decay data for complex conductivity model B.



(d) Complex conductivity model B. A layer of complex conductivity (Cole-Cole parameters are shown) is present between depths of 350m and 480m.

Figure 3.16: Observed and predicted data, and 1D conductivity models for Questem station located over the deposit. A layer with complex (frequency-dependent) conductivity (based on the Cole-Cole dispersion model), positioned at the approximate depth of the deposit, can produce data that has the same character as the observations. The conductivity of the overburden was also decreased in an attempt to more closely fit the early-time data.

fit to the observations than those in Figure 3.16(a). However, the early-time data do not fit the observed data so well.

The Cole-Cole parameters were manipulated to produce data that resemble the observations. Although the data have not been fit exactly, the results seem acceptable for only having introduced one conductive layer. The Cole-Cole parameters used were $R_0 = 200$ Ohm-m, $c = 0.4$, $\tau = 5$ msec, and $m = 0.7$, for model "A", and $R_0 = 100$ Ohm-m, $c = 0.785$, $\tau = 5$ msec, and $m = 0.8$, for model "B". The parameters correspond quite well with numbers determined by Pelton [Pelton et al., 1978] for massive sulphides, although values of c less than 0.55 are expected. It is not unreasonable to expect that, by slightly changing the conductivities and Cole-Cole parameters at different depths, a model can be found that fits the data within the expected noise levels and has chargeable material at depth.

Discussion on inverting complex conductivity data

Since a forward modeling is available that produces time-decay data from complex conductivities, it might seem reasonable to use this to invert the data. However, there are issues that should be considered before this is undertaken. First the appropriateness of the Cole-Cole model should be demonstrated. Pelton [Pelton et al., 1978] has shown that many sulphide minerals cannot be described by one set of Cole-Cole parameters alone. The use of two or more Cole-Cole models is necessary to fit laboratory observations.

Secondly, inverting for more unknowns could introduce *more* non-uniqueness into the problem. That is, each layer of complex conductivity will have its own frequency dependence, and will be characterized by four parameters, σ , m , c , and τ . In contrast, each frequency-independent layer needs only one parameter (σ) for full characterization. This means that the number of unknowns will increase by a factor of four if only one Cole-Cole model is employed.

Alternatively, so as not to be restricted by any parameterized frequency-dependent models, real and imaginary parts of the complex conductivity could be inverted for at each frequency. Routh (1999) investigated this for multiple grounded sources in order to remove coupling effects from IP data. However, data from one source may not contain enough information to constrain both depth and frequency-dependent information.

Negative data are in the noisiest part of the time-decay (late times). These data may not be sufficient to resolve between different combinations of the Cole-Cole parameters. In such cases, additional information would be needed to constrain or eliminate some of the parameters from the inversion. Better constraints could come from the data collected with TEM systems with increased sensitivity and better sampling of the decay series. Alternatively, rock samples in the area may reveal only small variations in some of the Cole-Cole parameters, thus allowing them to be constrained or removed from the inversion altogether.

The discussion so far has only dealt with one-dimensional complex conductivity structures. It is possible that no combination of complex conductivity layers will fit the data. The combined effect of three-dimensional geometries and complex conductivities is unknown, but may be necessary in order predict some decay-series.

3.6.4 Summary of regional conductivity methods.

Both frequency-domain data and early-time, time-domain data are able to model the overburden quite well, and regional structures that manifest themselves in the overburden structure can be identified. While it is not expected that the frequency-domain data can provide much more information than that which has already been modeled, the time-domain data has significant potential for providing more information about the subsurface. Forward modeling has shown that the decay response produced over the deposit could be explained with a buried, polarizable conductor. While other models

might also be able to explain the observed data, and cannot be ruled out, the concept of detecting massive sulphides with this data becomes conceivable. A benefit of accounting for polarizable material would be in modeling non-polarizable, conductive bodies more accurately because this would effectively remove IP effects that come from geologic noise. Until an inversion routine that models complex conductivity is available, it is likely the majority of the data can only be inverted to define the overburden, and complex conductivity forward modeling analysis could be applied to select parts of the data.

3.7 Chargeability models

At San Nicolás, gradient-array and “Real-section” array configuration IP data were acquired by *Quantec*. An overview of how data are generated using the IP method is given in Appendix C.6. The gradient-array data is used for regional modeling as it covers a relatively large area (Figure D.10). The large chargeability anomaly, indicated by a strong positive anomaly at -1600m east, -400m north, is associated with the deposit. There is no doubt that this is valuable information regarding the existence of an ore-body. In fact, this was the data that prompted the discovery of San Nicolás. In the gradient array configuration, the current electrodes are located outside of a rectangular area to be surveyed and the IP potential data are collected using a roving-receiver dipole within the rectangular area. The data are generally used only as a mapping tool because detailed depth information cannot be obtained without multiple transmitter locations. However, the similarity of gradient array IP data to magnetic field data allows for the data to be inverted using *MAG3D* through which depth information is introduced with sensitivity weighting. A description of how gradient IP data can be inverted using *MAG3D* is given in Appendix F, along with an example. Inversion parameters such as noise estimates and cell sizes are provided in Table E.1.

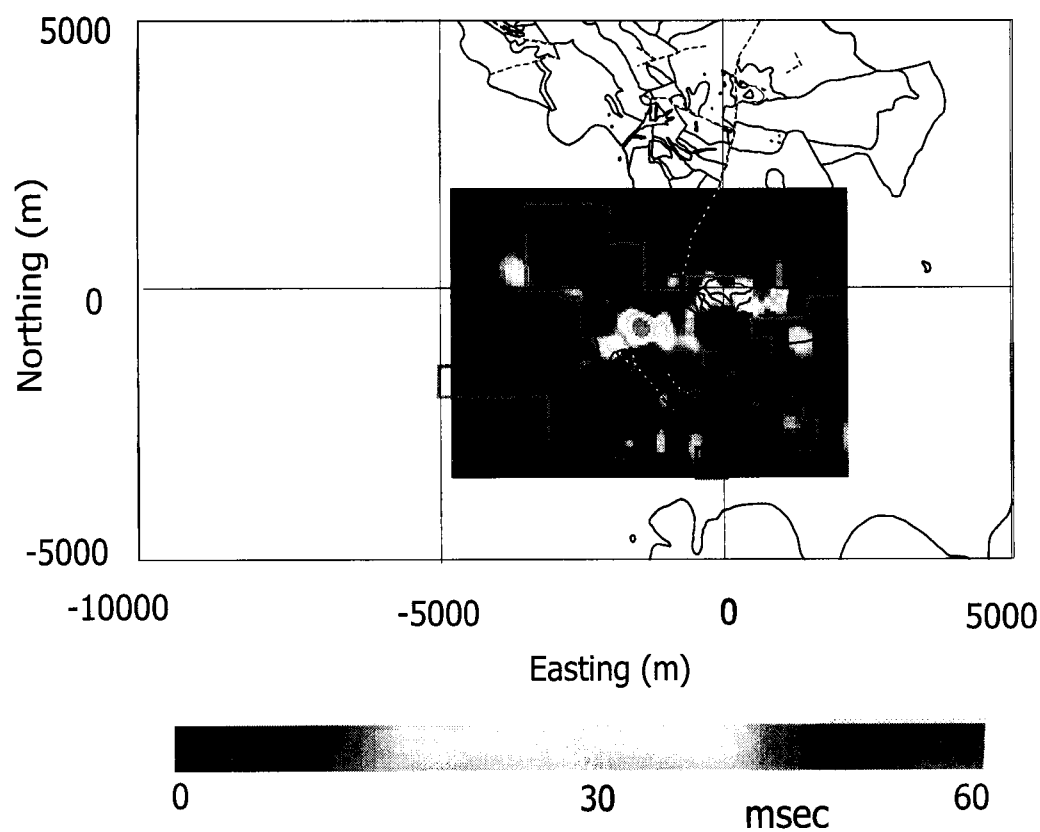


Figure 3.17: Plan-view of a slice at 300m below the surface through the chargeability model generated by inverting gradient-array IP data as magnetic data using *MAG3D*. The deposit is the largest chargeable body detected by the survey. The extents of the survey are shown by the red outline and outcrop geology is shown in black.

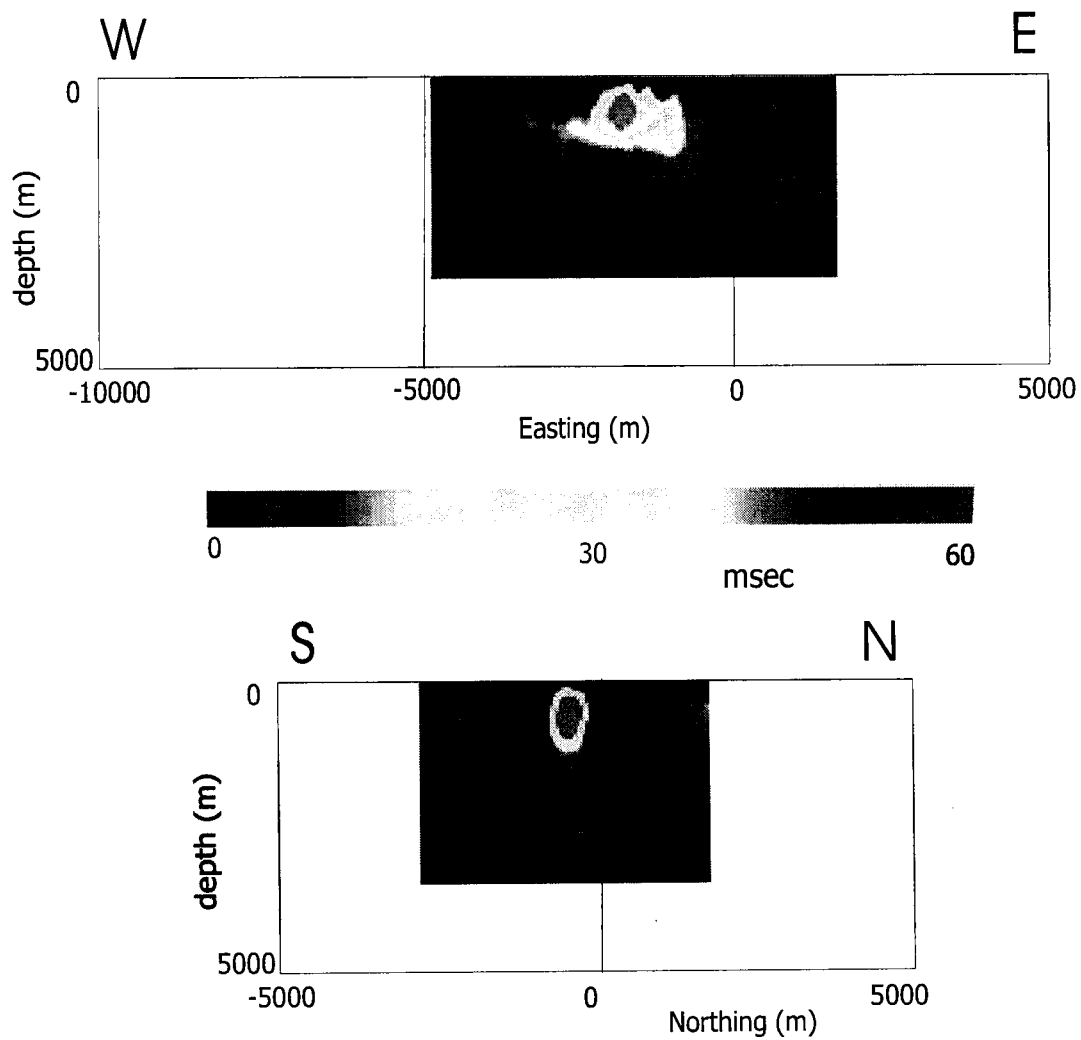


Figure 3.18: North- and west-facing cross-sections (upper and lower sections respectively) through the chargeability model. The chargeable anomaly is well defined yet extends below the elevation of the deposit, as defined by drilling.

Figures 3.17 and 3.18 show the modeled distribution of chargeable material on slices through the deposit in plan and cross-sectional views. As expected from the data, the anomaly corresponding to the deposit is the largest feature seen in the model. In plan view, the anomaly seems to be concentrated slightly south of the deposit and has extensions of lesser chargeable material to the east and the southwest (the eastern extension of chargeability anomaly is also shown in the north-facing cross-section). The deposit does have mineralized extensions to the east and southwest, and although the elevation of this mineralization does not quite correspond with the chargeability model, the model maybe sensitive to these areas. Of course, the distribution of chargeable material may not be one-to-one with ore. Chargeable minerals could have been deposited within the surrounding host-rocks that might result in anomalies peripheral to the deposit.

The modeled concentration of chargeable material also extends deeper than the extent of the deposit, as defined by drilling. This is probably a result of having no depth constraints within the data themselves. However, the overall results suggest that the assumptions made in order to invert the data using *MAG3D*, were appropriate in this case.

A cluster of smaller chargeability anomalies are located to the northeast of San Nicolás centered at about 500m east and -200m north. This is in the vicinity of El Salvador, a smaller lense of massive sulphide hosted within the same volcanic rocks¹³. The only other anomaly of significant magnitude is a small, shallow discrete body located at about -1200m east and -2800m north. This anomaly has a magnitude of about 20 msec and is located on the edge of the survey. This anomalous region of chargeable material coincides with regions of anomalous density and susceptibility and does contain mineralization. An analysis of this anomaly is presented in the next section — physical property correlation.

¹³El Salvador was discovered prior to the discovery of San Nicolás.

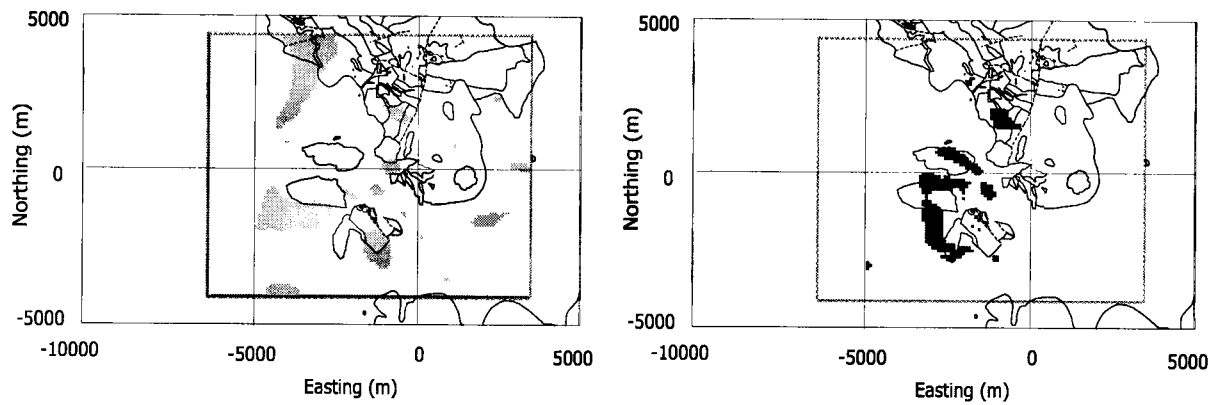
The only anomalies that have been modeled correspond to mineralization. However, away from this area, chargeability anomalies have been associated with some of the Tertiary rocks. In a more widespread survey, other methods such as gravity would be needed to distinguish between mineralization and Tertiary rocks.

3.8 Physical property correlation

The regional physical property models that have been generated contain three-dimensional structures that are, hopefully, easier to interpret and relate to geology than the original data. However, combining multiple sources of information, in order to target areas of favorable physical properties, becomes more involved. As a result, models are evaluated in three-dimensions using two simple spatial correlation methods. For each correlation method the physical property models are first imported on to a common mesh. This allows for direct correlation between cell values of different property types.

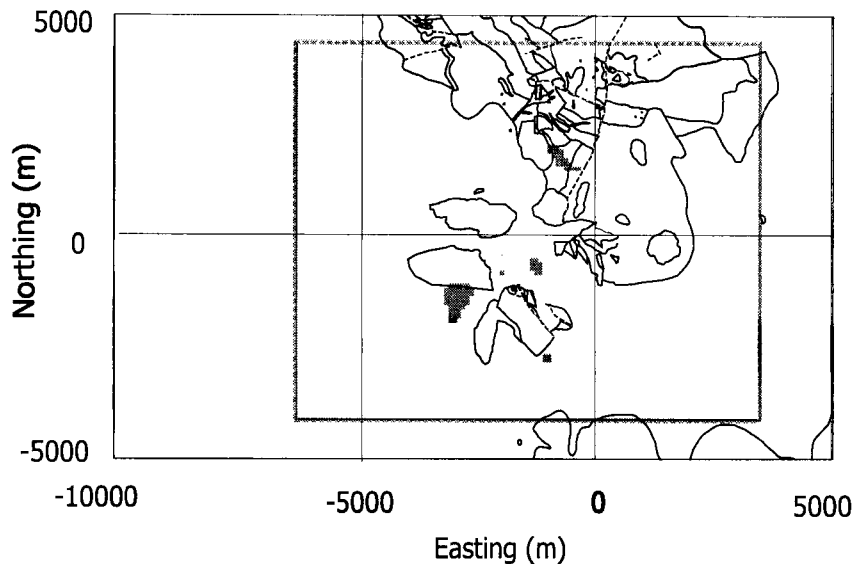
For the first method, cut-off values for the physical properties are determined based on the expected physical property values and the recovered values over San Nicolás. For the density-contrast model a cut-off of 0.13 g/cm^3 is used and a cut-off value of 0.0026 S.I. is used for the regional susceptibility model. The models have cells removed that correspond to values below these cut-off levels. As a result only the cells with values above the cut-off remain. These relate to volumes of high density-contrast and magnetic susceptibility. A model is then generated, on the same mesh, that represents the intersection of regions of high density-contrast and high magnetic susceptibility. The resulting *cut-off* correlation model should be useful in locating more massive sulphides because it represents regions with density and magnetic susceptibility values similar to that of the deposit.

Figure 3.19 shows projection-to-surface plan-views of density contrast, magnetic susceptibility, and correlation models of the upper 1000m. Only the upper 1000m of the



(a) Cells with a density contrast greater than 0.13 g/cm³

(b) Cells with magnetic susceptibility greater than 0.0026 S.I.



(c) *Cut-off* correlation model: regions high in density and magnetic susceptibility.

Figure 3.19: Plan-views of density contrast, magnetic susceptibility, and *cut-off* correlation models. The upper images show the projection-to-surface of cells, within the first 1000m from the surface, that exhibit high density contrast and magnetic susceptibility values. The lower image shows regions that are high in both density and susceptibility. The image is a projection to surface of the volume intersection of the upper models. A blue border shows the extents of the correlation analysis.

earth is considered, because finding ore below that depth is not within the objective of most exploration programs. Cells below a value of 0.13 g/cm^3 in the density contrast model are invisible, as are cells below a value of 0.0026 S.I. in the magnetic susceptibility model. The *cut-off* correlation model (Figure 3.19(c)) shows five separate anomalies, one of which coincides with the deposit at -1600m east, -400m north. The most northern anomaly lies coincident with out cropping mafic volcanics. The small (one model cell), shallow anomaly about 750m to the west-southwest of San Nicolás, upon field inspection, proved to be a small unmapped outcrop of unaltered intrusive mafic material. The large anomaly to the west is the only dense-susceptible body that extends beyond a depth of 1000m. This is the intersection of the large deep intrusive features seen in the regional density-contrast and magnetic susceptibility models.

The remaining anomaly is located at -1300m east, -2700m north within the first 100m from the surface, and is coincident with the chargeability anomaly mentioned in section 3.7. Three drill-holes have been completed each to a depth of 450m around, but not directly into, this anomaly. Of the three drill-holes only one encountered mineralization similar to that found at San Nicolás¹⁴. If the chargeability model was correlated with both the susceptibility and the density models, the resulting *three-way* correlation model would contain only two anomalies; one being the deposit and the other being this small occurrence of mineralization. This is an important result and demonstrates how the inversion models can be used together to find mineralization. However, the known mineralization at El Salvador has not been found with this method.

The method presented above has been successful in reducing the number of targets, however, it is dependent on the cut-off values and doesn't indicate whether one area is more dense or susceptible than another. Consequently, a second correlation method

¹⁴Hole SAL-61 intersected 0.7% copper at 430m depth and 0.2% zinc at 207m.

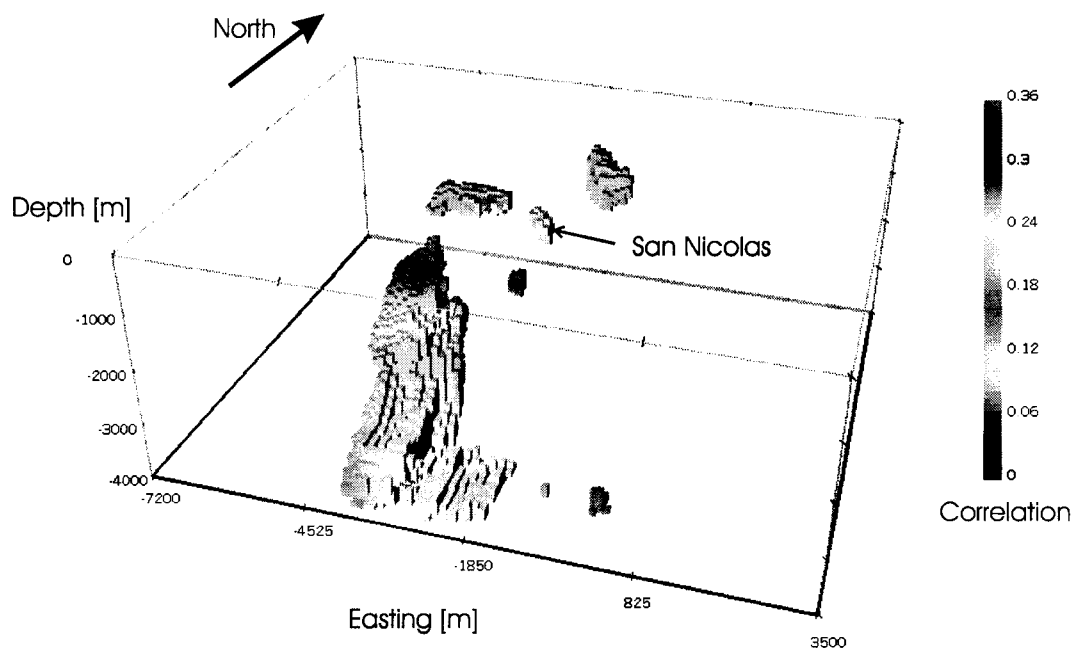
was employed. For the second method, the density-contrast and magnetic susceptibility models are normalized to each of their range of values. This means that high density and high susceptibility regions will have a value of one, while low density and susceptibility values will be represented by zero. The product of these models is then calculated. The resulting model (Figure 3.20) will have high values that corresponds to high density and susceptibility values.

Figure 3.20(a) shows a perspective view of the product of normalized density-contrast and magnetic susceptibility values. Again, the number of anomalies has been significantly reduced with this method. The model has similar features to that shown with the previous correlation method, however one important difference is present. A volume has been produced to the northwest of the deposit that corresponds to very high magnetic susceptibility but only moderate density-contrast values. This is a limitation of this method, however most of the other anomalies produced are from a combination of the two physical properties. The deposit along with a dense mafic intrusion to the north and a deep intrusion to the south have all been modeled. In addition, the small anomaly, located at -1000m E -2700m N, which also corresponds to the chargeability anomaly, is also present.

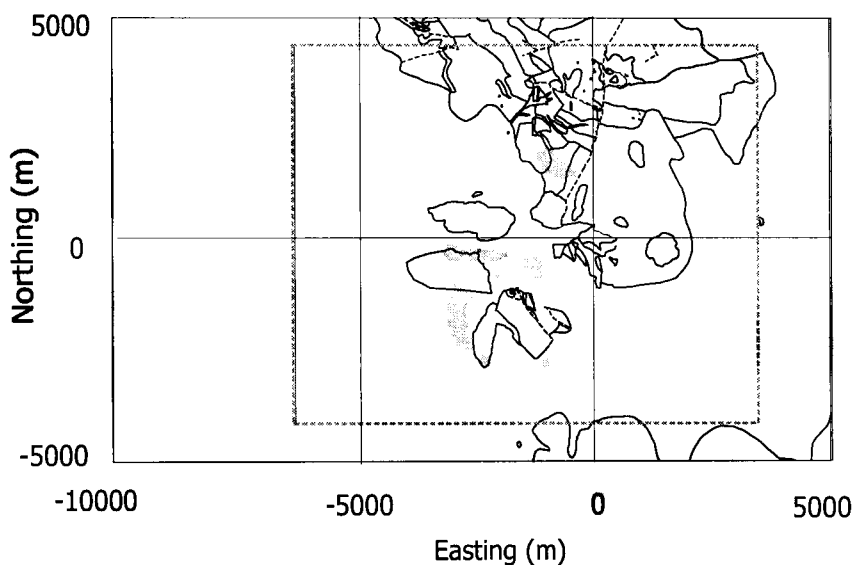
Both of the correlation methods work well at San Nicolás because all of the physical properties considered give high values over the deposit. If low or intermediate physical property values are of interest then further manipulation of the models would be required.

3.9 Conclusions

Inversion of regional geophysical data sets has produced large-scale physical property distributions of the earth. Together with physical property values and an understanding of the geologic setting, these models have been interpreted and given a geologic context.



(a) Northwest-facing perspective view of density-susceptibility correlation values. The model has been volume-rendered at a value of 0.2.



(b) Projection-to-surface of density-susceptibility correlation values within the first 1000m. Cells with a value of less than 0.2 have been removed. A blue border shows the extents of the correlation analysis.

Figure 3.20: Normalized density-susceptibility regional correlation values. The values are the product of normalized regional density-contrast and magnetic susceptibility models.

This interpretation would be much harder, if not impossible, without considering three-dimensional physical property distributions. Geologic features such as the extents of a horst, the detection and definition of a deep rooted intrusion, the identification of faults of two different ages, and the detection of shallow bodies that could be mafic intrusions have all contributed to the regional geologic understanding of the area.

In addition to the regional geologic interpretation, the models have been correlated in order to directly detect mineralization. The correlation of density and susceptibility models results in a reduced number of regions that might warrant further investigation. When combined with results from the inversion of gradient array chargeability data only two anomalous regions would remain, both of which contain mineralization. This method of selecting targets would hopefully be as effective in areas further from the deposit as it is here.

False anomalies can occur from all of the methods used, as a result there isn't one combination of physical properties that will detect mineralization without missing some that is already known. As a result different priorities of targets could be determined. These could be based on the different intersections of dense, magnetic, and susceptible material. If a comprehensive conductivity model could be constructed for the same region the correlation results could change and a better discrimination method might result.

Within the context of our exploration program, regional inversion modeling of geophysical data has been successful in targeting areas for detailed work based on geologic interpretations and direct detection methods. An integrated geophysical approach has been shown to be most efficient at selecting targets. When combined with results from other exploration methods such as geological mapping, and geochemical surveys, the target selection process could be even more refined.

In addition, the knowledge obtained from this work can be used to update the exploration models and thus refine exploration directives.

Chapter 4

Deposit Scale Modeling

4.1 Introduction

The work presented in chapter three was successful in directly detecting mineralization and contributing to the regional geologic picture. From the correlation methods presented possible targets could be identified. Ground surveys are now conducted in order to provide more information about these targets. The flowchart in Figure 4.1 shows how the work presented in this chapter is integrated into the second part of an exploration program, and builds on the knowledge already obtained.

This chapter deals with detailed geophysical modeling of the deposit, using finely discretized meshes, for further physical property characterization and enhanced definition. Results from the regional modeling are used for removal of regional trends from the gravity data prior to inversion, and for constructing reasonable starting and reference models when inverting the other data. Along with the gravity and gradient-array IP data, which was used for the regional modeling, more localized data sets are considered, including ground magnetics, CSAMT, and realsection-array IP.

The goal is to invert the data to see how much information about the deposit can be recovered. Unlike our example exploration program, we are in the position to check the inversion results of this chapter with geologic sections that have been interpreted from drill-hole information. However, the models are constructed through inversion of the geophysical data alone.

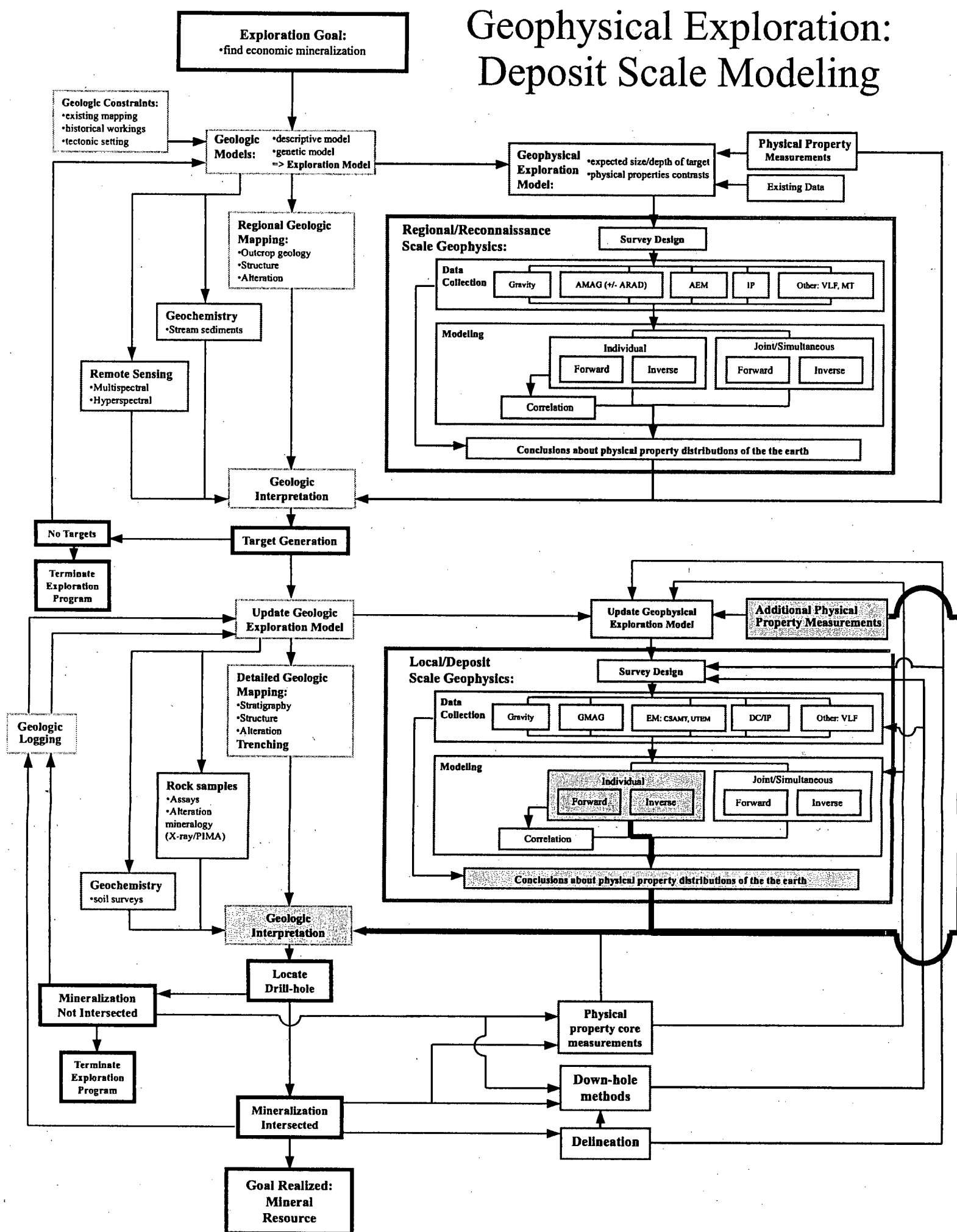


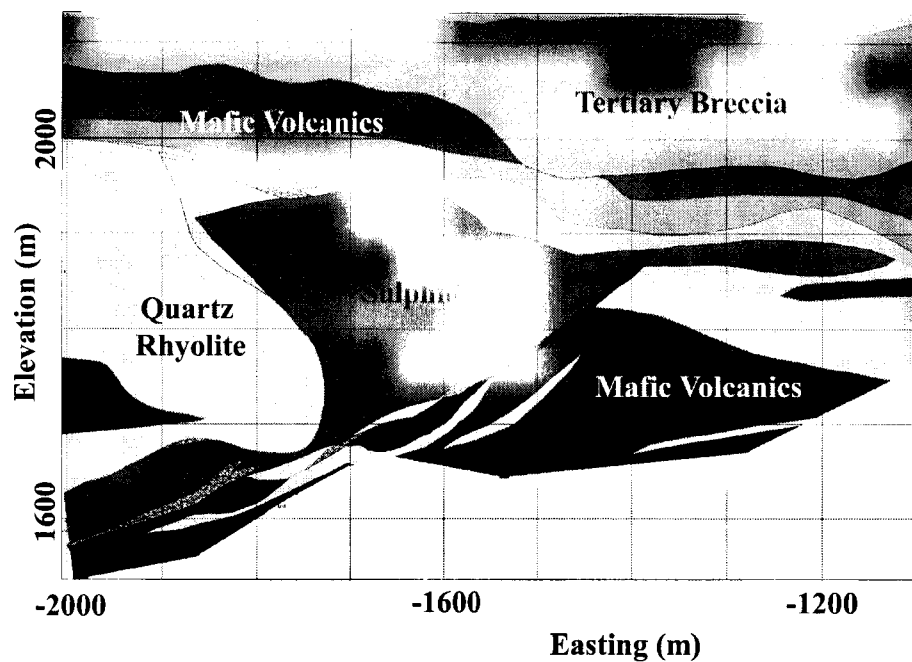
Figure 4.1: Geophysical Exploration Methodology Flowchart: Deposit scale modeling.

4.1.1 Local geology

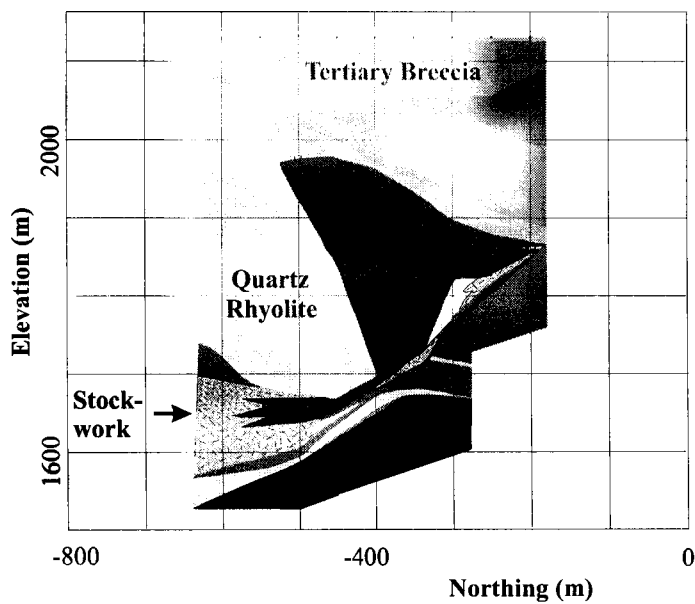
Figure 4.2 shows the local geology in the vicinity of the deposit as interpreted from drilling. The deposit is hosted in a series of interwoven mafic and felsic volcanic rocks of the Upper Jurassic to Lower Cretaceous Chilitos Formation, which lie unconformably over graphitic mudstones of the Zacatecas Formation. The deposit is almost entirely bounded to the east by a southwest-dipping fault, which could have been a feeder structure. Mineralization continues to follow the fault at depth in an unconstrained part of the deposit referred to as the *keel*. The deposit is flanked by a thick package of rhyolites to the west that contains a steeply dipping north-south fault (not shown on this section). The volcanic succession that envelops the deposit is overlain by a Tertiary-aged breccia overburden, which varies in thickness from 50 to 150m. The breccia includes tuffs and clasts derived from the underlying volcanics. Outcrops of the breccia have been mapped to the northwest, but an overlying thin veneer of Quaternary alluvium is present in the vicinity of the deposit. Hydrothermal alteration, which may have caused a change in physical property values, is prevalent throughout the deposit and surrounding geology. Figure 4.3 shows a simplified stratigraphic section as interpreted in the vicinity of the deposit.

Introduction to inversion results

The models constructed by inverting local-scale geophysical data are presented in the following sections. Only the models are presented; the data are displayed in Appendix D. The models are presented in plan, cross-sectional, and perspective views in much the same way as Chapter 3. The geologic sections shown in Figure 4.2 are shown with each inversion result in order to evaluate the success of the inversion. All of the inversion parameters used are provided in Table E.1.



(a) North-facing geologic cross-section located at -400m north.



(b) West-facing geologic cross-section located at -1700m east.

Figure 4.2: Geologic cross-sections through the San Nicolás deposit, modified from sections provided by Teck Corp. The main geologic units are labeled.

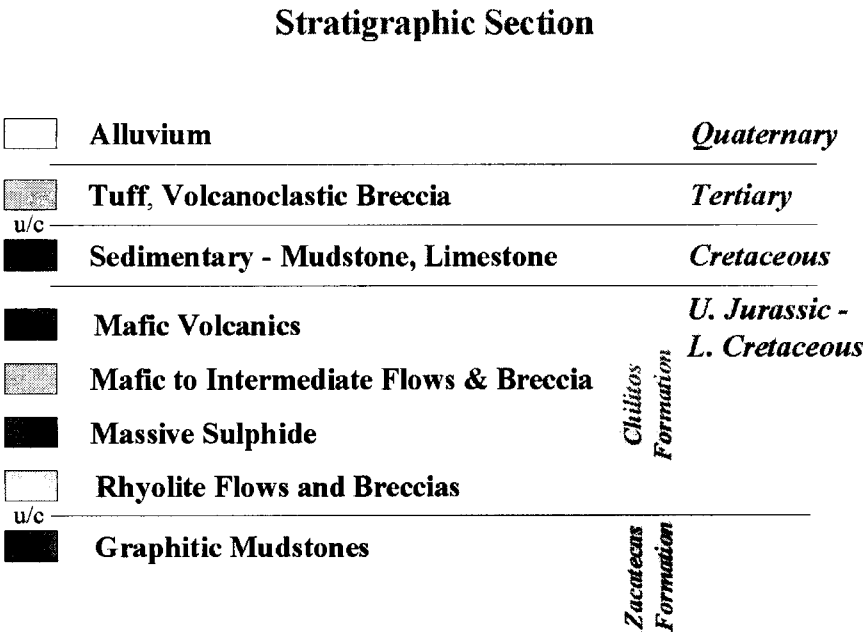


Figure 4.3: Simplified stratigraphic section in the vicinity of the San Nicolás deposit, modified from [Danielson, 2000].

4.2 Density

Inverting the gravity data on a local-scale requires a two-pass procedure. First, a “regional” field is removed from a subset of the original data, and secondly, the residual data are inverted on a finely discretized mesh, centered on the deposit. The removal of regional signal from the local data is undertaken in order to remove signal in the data that doesn’t originate from the localized region of interest, thus allowing the inversion to be concerned with only the deposit and the immediately surrounding volume.

4.2.1 Regional Removal

The regional density contrast model created in Chapter 3 is used to generate a regional gravitational field over a 1.8 km x 2.4 km area centered on the deposit. This is done by calculating the response of the regional density contrast model with a volume removed (Figure 4.4). The removed volume encompasses the deposit and two smaller, dense bodies. The resulting regional signal is subtracted from the original observations and residual data are produced that can be used for the detailed inversion (Figure D.11). Li and Oldenburg [1998a] describe removal of regional magnetic fields using this technique.

4.2.2 Inversion of Residual Gravity Data

Data errors of 0.01 mGal were assigned to the residual gravity data and the 3D inversion was carried out with *GRAV3D* using the parameters listed in Table E.1¹. The observed data are reproduced very well by responses from the detailed density contrast model (Figure 4.5) and the errors that were assigned to the data seem to have been appropriate.

¹The regional density contrast model used to create the regional response is based on a mesh with large cells (100m x 100m x 50m). Consequently, this is a coarse representation of the earth, and the regional gravity field contains small inaccuracies that will propagate through to the residual gravity data. In order to account for these inaccuracies, padding cells are used in the inversion of the residual gravity.

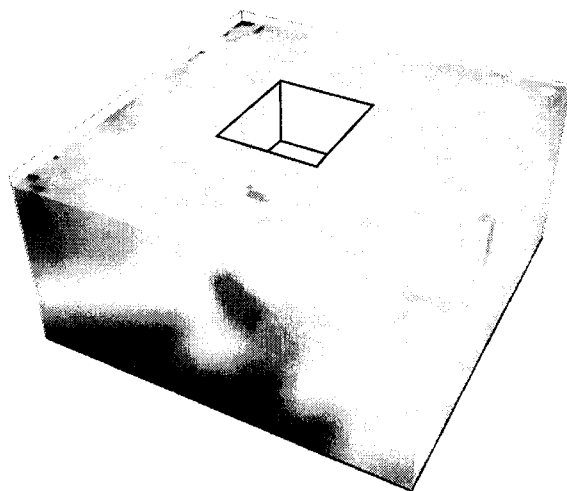


Figure 4.4: Perspective view, facing north-west, of regional density-contrast model used for calculation of the *regional* signal. The volume that has been removed contains the deposit and two smaller high density anomalies, and is the volume on which the more detailed inversion is focused.

The central density contrast anomaly represents the San Nicolás deposit and correlates well with the dense massive sulphides. This is an improvement over the regional model probably because the deposit is defined with small cells, and discretization is not affecting the results. Figure 4.6 shows cross-sections of the density-contrast model overlain with geologic boundaries. The centroid of the anomalous density coincides with the San Nicolás deposit and corresponds well with the dense massive sulphides. The lateral dimensions of the main body of the deposit are reasonably well defined, and the density anomaly doesn't continue to the surface, as it does in the regional inversion. Large density-contrast values are present above the expected elevation of the deposit, and neither the keel nor the thin extension of sulphides to the east have been modeled in this inversion. However, this is a useful result when one considers the lack of depth information contained in the data, and locating a drill-hole based on this model would

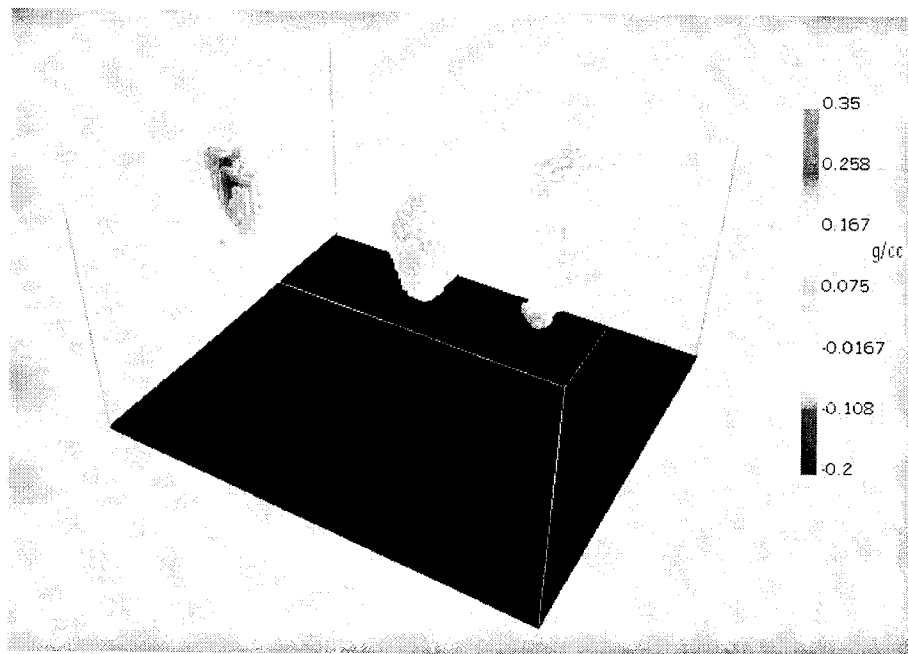


Figure 4.5: South-east perspective view of the detailed, density-contrast model shown as a volume-rendered isosurface plot using a cut-off of 0.18 g/cc. Cells with density contrasts less than this value are invisible.

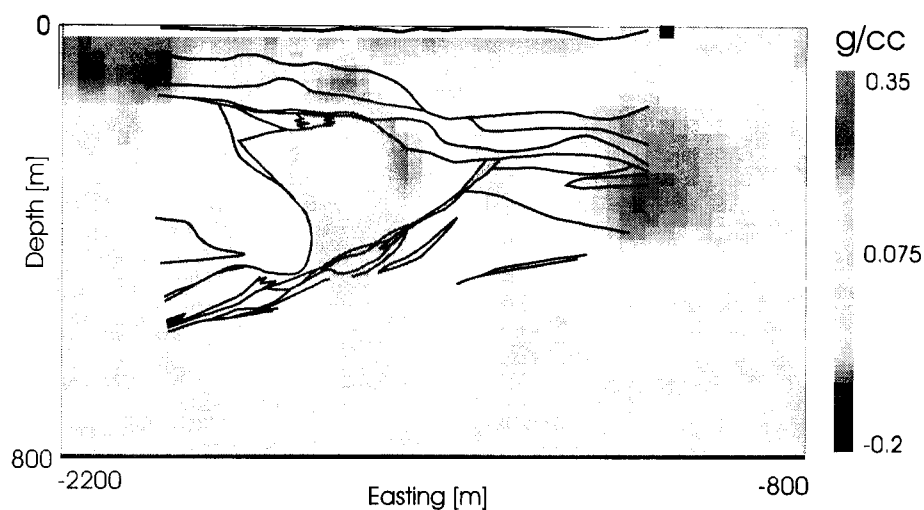
lead to a successful intersection of the ore-zone. Two other density anomalies are also well defined by the model. The north-eastern anomaly was drilled and intersected 400m of intrusive, mafic rock. Field inspection of the western anomaly revealed similar mafic intrusives outcropping at the surface.

4.3 Magnetic Susceptibility

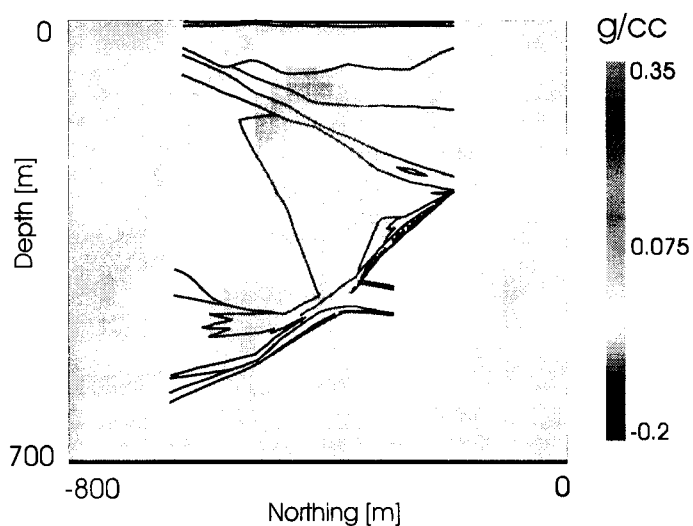
Finely sampled, magnetic field data were available for detailed modeling in the form of ground magnetic data. A base-level of 44,000 nano-Tesla was removed from the 614 diurnally corrected data as a processing step prior to inverting. In addition, several data in the center of the survey that were contaminated with cultural noise, such as a fence or steel-cased drill hole, were discarded. The 3D magnetic inversion was performed using *MAG3D* and the predicted data were in good agreement with the observations (Figure D.12).

A volume-rendered representation of the 3D susceptibility structure is shown in Figure 4.7. The cut-off value is 5×10^{-3} S. I. A distinct body of higher susceptibility is modeled. The majority of the susceptibility high is coincident with the deposit, however, high values continue to the north (as seen in the data and the regional model). A good correlation between magnetic susceptibility values and geology in the vicinity of the deposit can be seen in the north-facing cross section presented in Figure 4.8. The high magnetic susceptibilities associated with the deposit align well with the boundaries of the sulphide body. A concentration of high susceptibility is present in the western half of the deposit.

As with the regional magnetic susceptibility model (Figure 3.10), there appears to be a continuation of moderately susceptible material to the east and below the deposit. Mafic rocks below the ore-bounding fault zone are less altered, and may be the source of



(a) North-facing cross-section through density-contrast model located at -400m north. Geology from Figure 4.2(a) is overlain.



(b) West-facing cross-section through density-contrast model located at -1700m east. Geology from Figure 4.2(b) is overlain.

Figure 4.6: Cross-sections through deposit-scale density-contrast model at the San Nicolás deposit.

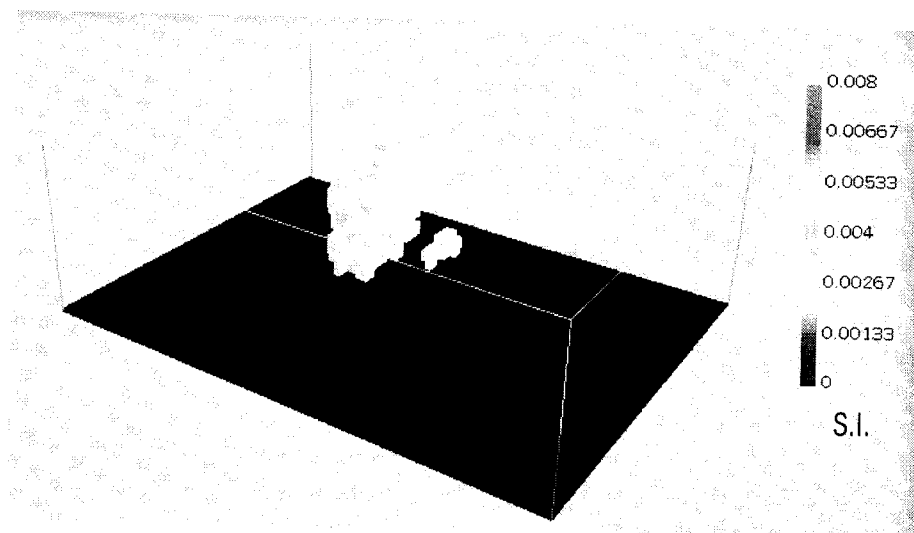
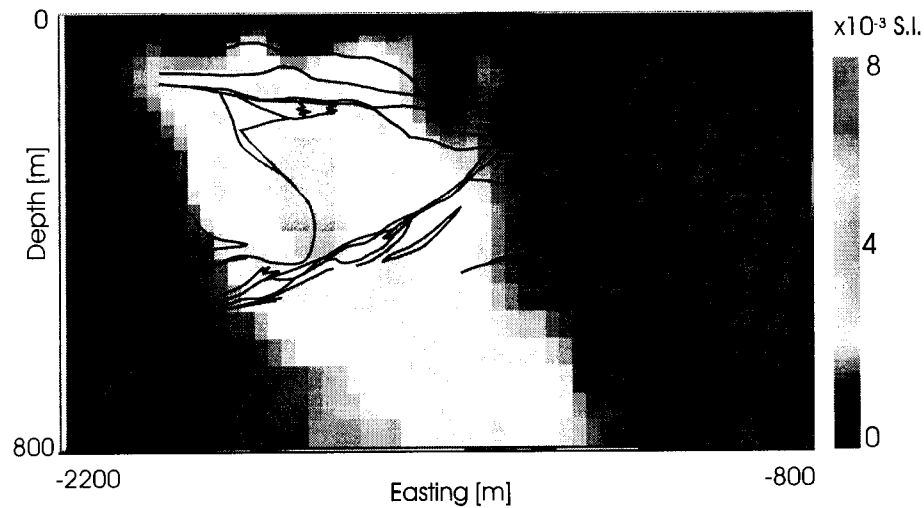


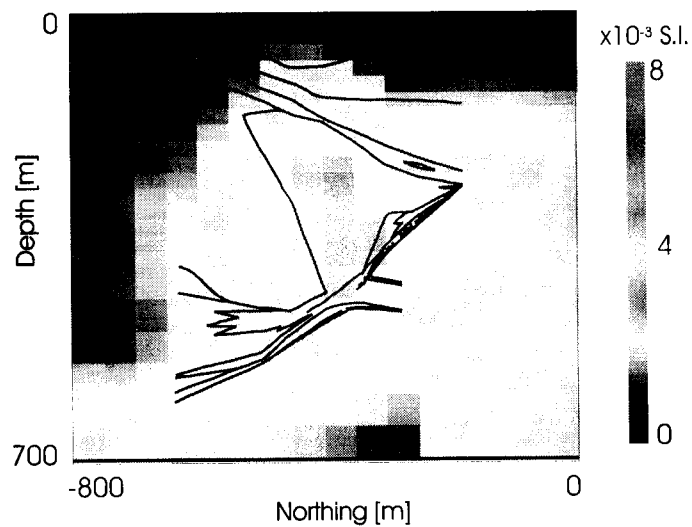
Figure 4.7: Perspective view of the detailed, magnetic susceptibility model.

this susceptible material. There is no indication of the keel or the eastern extension of sulphide in this model. The result is a smoothed version of the true susceptibility values, and it does not contain highly detailed structural information, particularly at depth.

A reasonable magnetic susceptibility model has been recovered from the ground magnetic data, however, it appears that obtaining information about the location of sulphides from the model is not a straight forward exercise. Evidence from physical property measurements (Appendix B.3.1) show that magnetic minerals, while initially expected to be associated with sulphides, are not distributed throughout the deposit. Moreover, they can also be present in host rocks and might have been deposited by hydrothermal events not associated with the deposit itself. This might explain the high susceptibility values that persist to the north, and those that persist to the east and below the deposit, and the lower values that are modeled in the eastern half of the deposit.



(a) North-facing cross-section through susceptibility model located at -400m north. Geology from Figure 4.2(a) is overlain.



(b) West-facing cross-section through susceptibility model located at -1700m east. Geology from Figure 4.2(b) is overlain.

Figure 4.8: Cross-sections through deposit-scale magnetic susceptibility model at the San Nicolás deposit.

4.4 Conductivity Data

Three lines of controlled source audio-frequency magneto-telluric (CSAMT) data were collected over the San Nicolás deposit. The CSAMT method is an electromagnetic technique that uses a grounded dipole current source². At San Nicolás, CSAMT data were collected with the receiver lines orientated perpendicular to perceived geologic strike. It is not obvious from the observed apparent resistivity or phase data (Figures D.13 and D.14) that a conductive ore deposit is present.

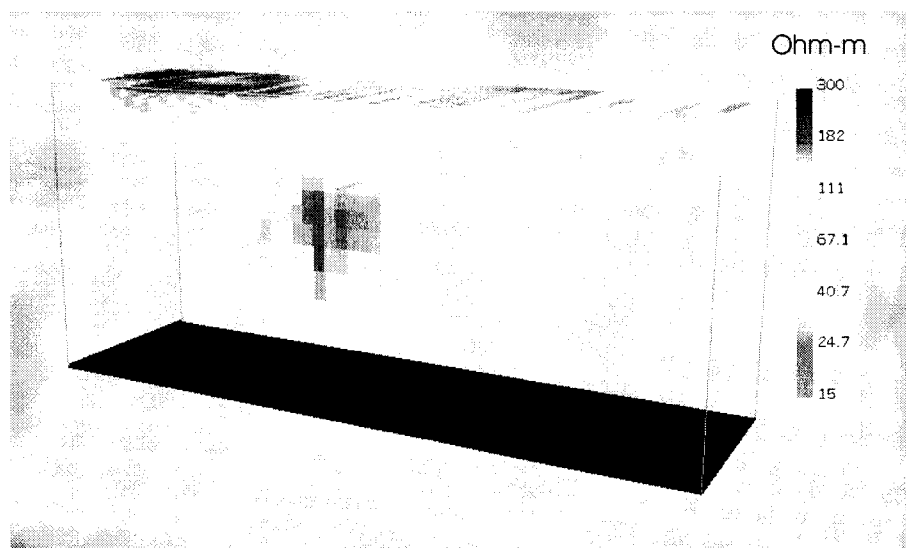


Figure 4.9: Perspective view of resistivity model created by interpolating between lines of “stitched” 1D models. The inversion has successfully discriminated between the overburden and the deposit.

Apparent resistivity and phase data are inverted to recover a one-dimensional, resistivity model beneath each station. The one hundred and eighty 1D models were “stitched” together to form a 2D model along each of the three lines of data. An interpolation routine was used to fill in the volumes between the lines and a three-dimensional resistivity

²An overview of the CSAMT method is given in Appendix C.4.

model was produce. A volume-rendered image of the model, with an isosurface cutoff of 30 Ohm-m, is shown in Figure 4.9. The large conductive feature at depth is the deposit. It is separated from a conductive overburden by an intervening layer of higher resistivity. The cross-section of the resistivity model, shown in Figure 4.10, shows that the deposit boundaries interpreted from the inversion agree reasonably well with those inferred from drilling.

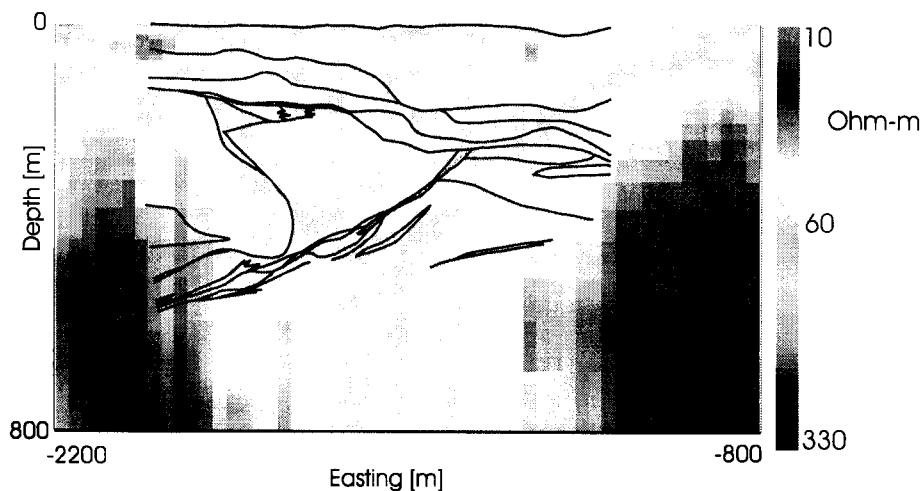


Figure 4.10: Cross section of resistivity model with geology. The resistivity model was produced by “stitching” together 60 1D inversion models.

4.4.1 Testing the 3D resistivity model

In order to test the validity of the one-dimensional inversion, the “stitched” inversion models and a 3D physical property model³ were forward modeled using *EH3D* (see Appendix E.8). Apparent resistivity and phase data were calculated from electric and magnetic field measurements and are plotted in Figure 4.11. The forward-modeled data from

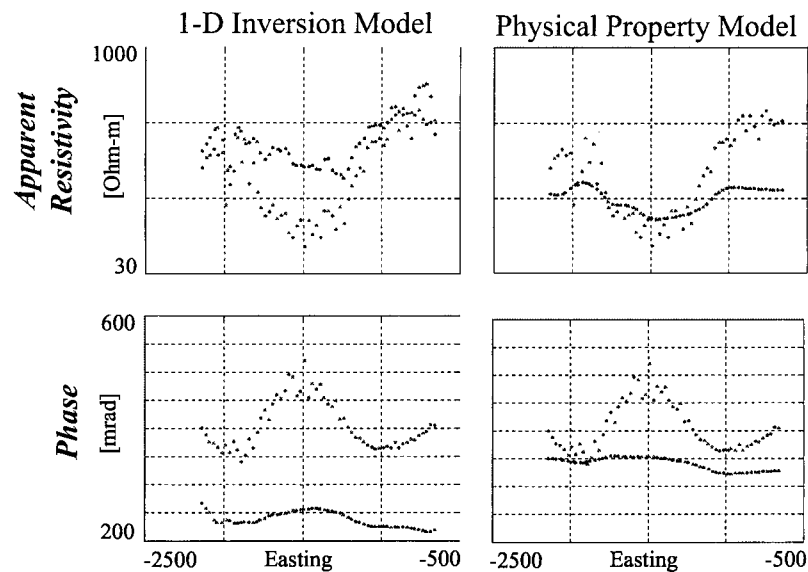
³The physical property model was generated by digitizing the geologic sections, assigning conductivity values based on the physical property measurements, and interpolating between the sections to create a 3D model (Appendix B.5).

the inversion models replicated the original data quite well at higher frequencies (Figure 4.11(b) - left plots). At low frequencies, the 3D conductivity model produced data that came closer to the observations than the “stitched” inversion models (Figure 4.11(a) - right plots). This result is plausible when one considers the geometry of the deposit and the volume of the subsurface that low frequency measurements will sample. At low frequencies the measurements will sample a larger volume, including the 3D deposit, and this information is hard to represent in only one dimension. Conversely, at high frequencies the part of the subsurface that the measurement has sampled will be more one-dimensional. Since the high frequency data are sensitive to shallow features, this suggests that structure in the upper region of the “stitched” inversion models is valid, and the structure of the physical property model is more representative of the deposit at depth. Discrepancies between the observed data and the forward modeled physical property data possibly arise due to inaccurate assignment of conductivity values.

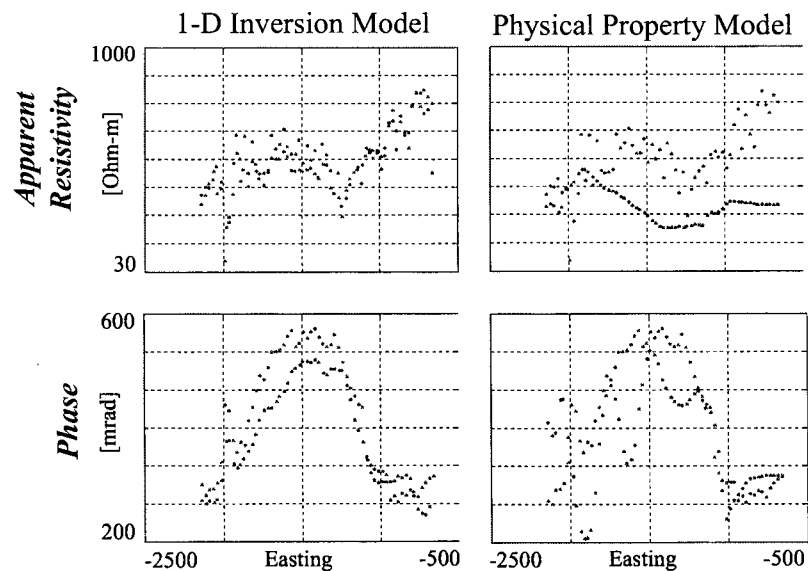
The forward modeling mentioned above suggests that this data set might produce a model with greater resolution at depth if it were to be modeled in three dimensions. With the inversion codes available, the raw electric and magnetic field data, from which the apparent resistivity and phase data were derived, are needed in order to invert this data in 3D. Unfortunately, at this time, these raw data have not been recovered. Of specific interest in a 3D inversion would be detection of the deep high-grade keel. A sensitivity analysis on the detectability of the keel using different CSAMT and CSEM (Controlled Source Electromagnetic) survey designs is presented in Section 6.4.

4.5 Chargeability

In this section “Real-section” array IP data is simultaneously inverted along with a subset of the gradient array data in order to produce a chargeability model that has more detail,



(a) 8Hz data



(b) 256Hz data

Figure 4.11: Comparison of CSAMT field data with forward modeled data from the *stitched* 1-D inversion resistivity model (left plots) and a physical property resistivity model (right plots). The field observations are displayed in red while the modeled data is shown in blue. At 8Hz (upper plots), the physical property model fits the field data well while the 1-D inversion model fits the field data better at a frequency of 256Hz (lower plots).

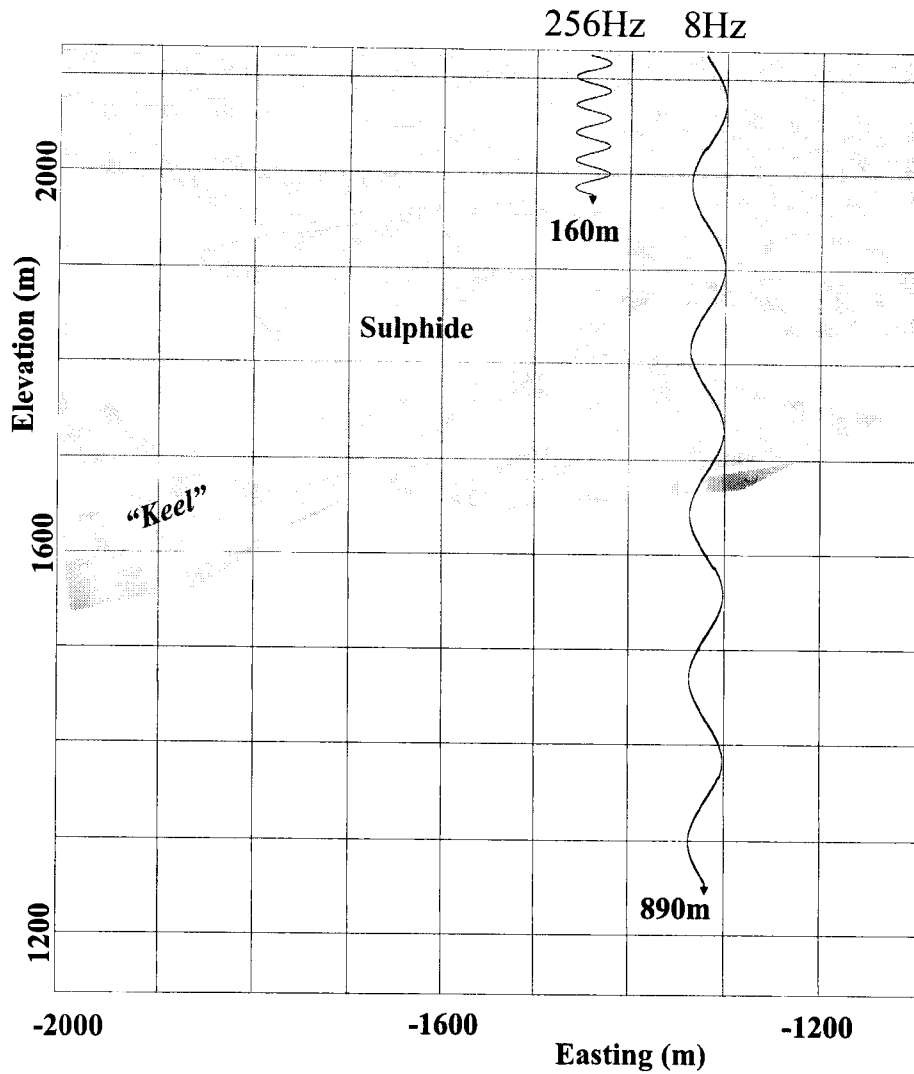


Figure 4.12: North-facing, cross-sectional diagram of the geology at San Nicolás showing the skin-depths associated with propagating plane-waves at frequencies of 256Hz and 8Hz. The skin-depths were calculated for a 25 Ohm-m whole space. The higher frequency signal samples geology that is much more “one-dimensional” than that sampled by the low frequency signal.

and better depth information, than the regional model that was produced using *MAG3D* in Section 3.7.

4.5.1 Chargeability data

In real-section surveys, the voltages from consecutive dipoles are measured and plotted in a pseudosection format. The current transmitter electrodes straddle the potential electrode array and the transmitter electrode spacing is continually reduced to change current flow in the subsurface. Data are recorded only at those potential electrodes lying interior to the transmitters. The configuration is a modified version of that employed in gradient-array surveys but should have better depth information due to the increase in number, and varied spacing of the transmitter electrodes.

The calculated chargeability is plotted beneath the receiver dipole and each row of the resultant pseudosection-type plot corresponds to a different location of the current electrodes. The final plot has an inverted appearance compared to pseudosections obtained with more traditional pole-dipole or dipole-dipole plots. The Realsection plots for San Nicolás are shown in Figure D.15. It is apparent from the data that a chargeable body is present, but these pseudosections provide no tangible information about the depth of the target. That can only be obtained through inversion. The inversion procedure for chargeability is a three-step process. First the resistivity of the model volume needs to be calculated. Next, the sensitivities needed for the IP inversion are calculated using the resistivity model. Lastly, the IP data are inverted using *IPINV3D* to recover a 3-D chargeability model. The details for the IP inversion are provided in Table E.1.

4.5.2 The resistivity model

DC resistivity data are collected at the same time as IP data and, as codes are available to invert these data in 3D, they are usually used to construct the resistivity model needed

for the sensitivity calculation when inverting IP data. At San Nicolás, these data do not produce a resistivity model that corresponds to the expected resistivity distribution. A surficial conductive layer was modeled that resembles the overburden, but no conductive body is recovered in the vicinity of the deposit. One possible explanation could be that the currents emitted from the transmitters do not penetrate into the volcanic rocks and are only concentrated within the more conductive overburden; in other words, taking the path of least electrical resistance. This would seem to make sense, although some currents would have to penetrate to the deposit in order for a chargeability response to be recorded (assuming the deposit is the only chargeable source)⁴.

Only an approximate conductivity structure is needed for the sensitivity calculations, so chargeability inversions were performed using sensitivities calculated from the inversion of DC resistivity data. This produced a chargeability model that corresponded well with the deposit, but seemed to have layered artifacts corresponding to the Real-section transmitter locations.

However, other, more accurate, conductivity models were available; the conductivity model produced from the inversion of C.S.A.M.T. data, and a model constructed by digitizing the geologic sections and assigning conductivities based on physical property measurements. Both of these conductivity models were used to construct sensitivities for the inversion of the IP data. The resulting models did not have the artifacts observed before. The model produced using the CSAMT conductivity model is presented in this section as it contains no apparent artifacts and also is derived only from geophysical data collected at the surface.

⁴The phenomenon of a disconnected conductive and chargeable body, lying below a conductive overburden, producing a negligible DC resistivity yet observable IP response has been observed before and is documented by Brescianini et al. [Brescianini et al., 1992].

4.5.3 Removal of electrode effects

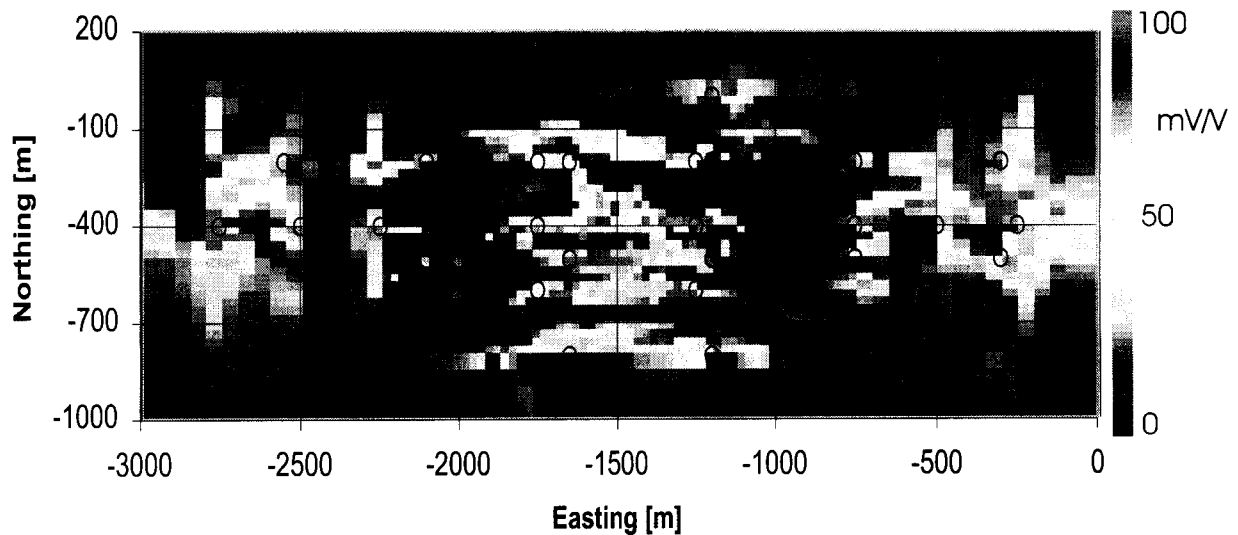
When inverting IP data using *IPINV3D*, especially data that have limited inherent depth information, such as gradient-array data, accumulations of chargeable material are often observed around the transmitter electrodes. This “electrode-effect” often leads to the recovery of bodies with smaller chargeability values than expected. For this reason, an empirical weighting function is applied in the inversion of the San Nicolás data that penalizes against chargeability values near the transmitter electrodes. The weighting function is included in the spatially dependent weighting matrix (w_s) introduced in Equation 2.2. It is designed to strongly weight the model close to a reference model of zero at the transmitters. Appendix G shows how this weighting function has been derived using synthetic models.

The result of implementing this weighting function on the model is shown in Figure 4.13. Most of the surficial chargeability has been reduced. The magnitude of chargeability material at the anomaly that corresponds to the deposit has doubled from about 25 msec to 50 msec.

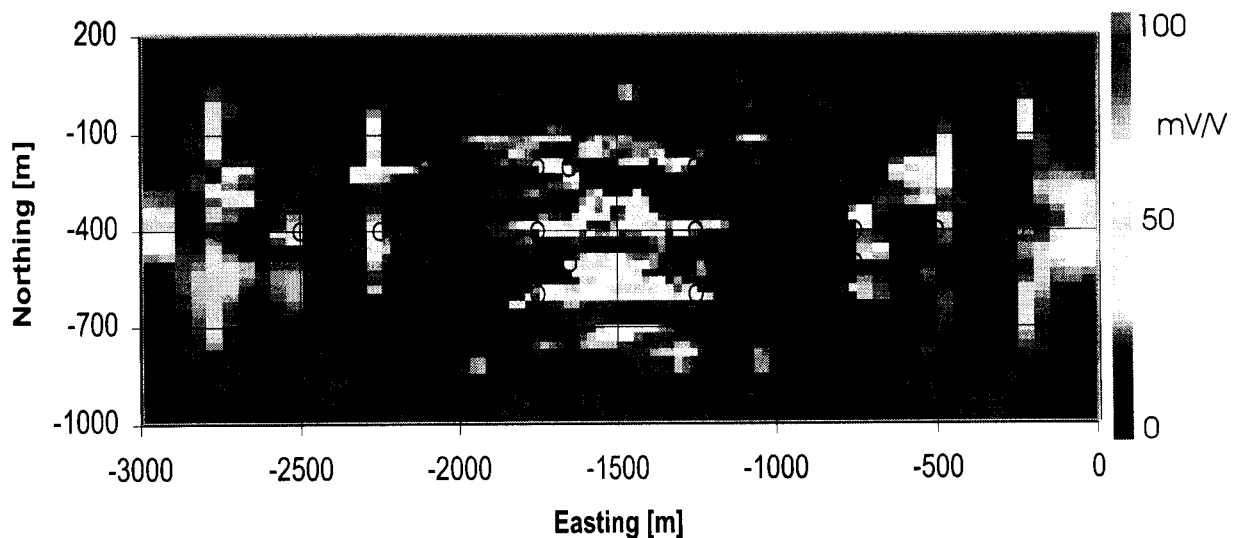
An alternative method to using this weighting function would be to weight the sensitivity matrix used in the inversion. This could make cells near the transmitter locations less sensitive to the data, and cells further away more sensitive to the data.

4.5.4 The chargeability model

Figure 4.14 shows a perspective view of the recovered chargeability model. The spatial distribution of the high chargeability values agrees quite well with the location of the deposit. The high chargeability values are off-set slightly to the south in relation to the deposit (this is also seen in the observed real-section data in Figure D.15). The high chargeability values present at the surface are probably remaining artifacts of the



(a) Plan-view of recovered chargeability model at the surface. The black circles show the location of the transmitters; these locations correspond with most of the high chargeability material that has been recovered at the surface.



(b) Plan-view of recovered chargeability model at the surface. The black circles show the location of the transmitters. Much of the surficial chargeable material that was modeled has been subdued with a weighting function.

Figure 4.13: Plan-views of the recovered chargeability models at the surface. The high chargeability values that correspond with electrode positions have been reduced through the use of a weighting function.

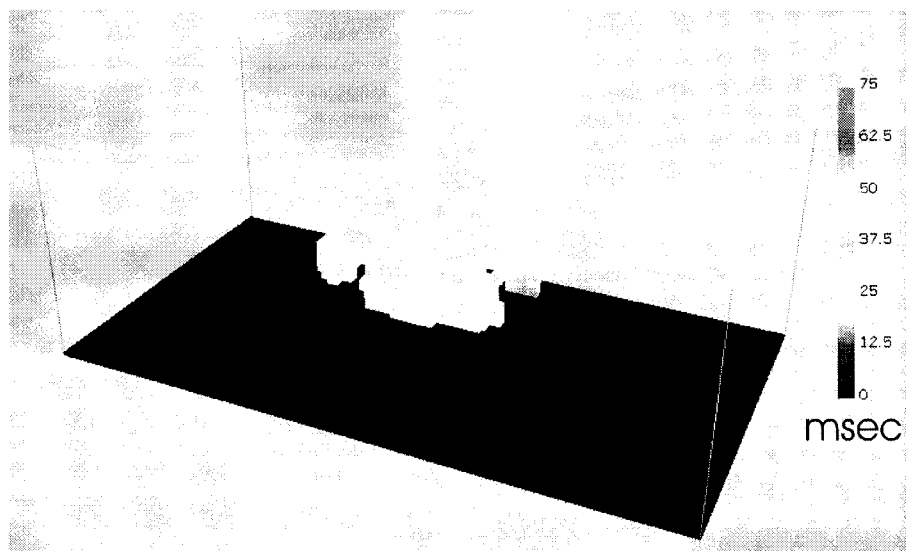
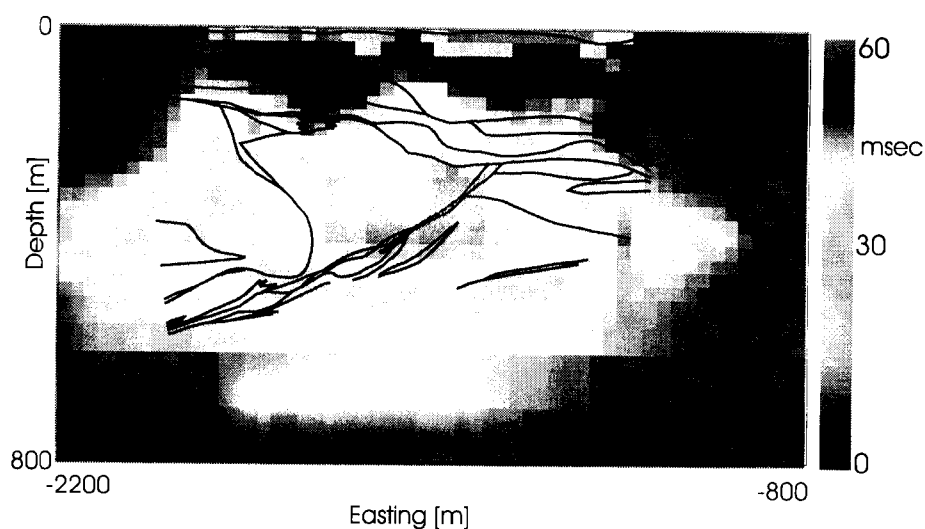


Figure 4.14: Perspective view of the volume-rendered chargeability model with a cut-off value of 40 msec.

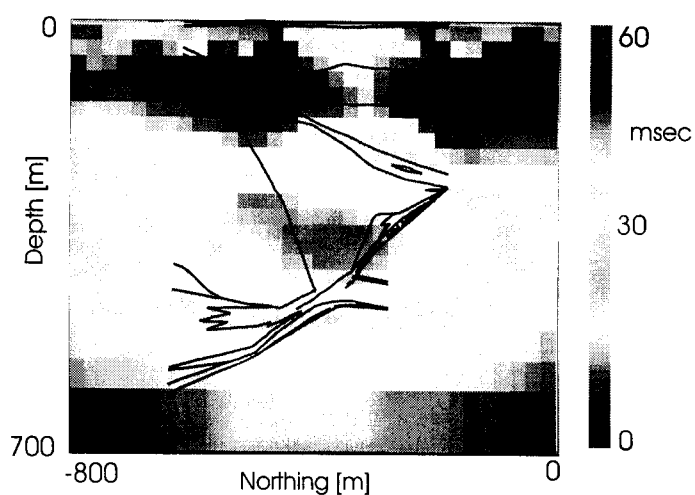
electrode-effect mentioned above, although they could be real. By viewing a north-facing cross-section (Figure 4.15), one can see a concentration of chargeable material centered on the south-west dipping fault. An exploration of model-space was conducted by changing the inversion parameters and the chargeable material around the fault was consistent for all models produced. The fault zone is known to contain semi-massive sulphides and stringer zones, which contain high-grade veins mixed in with host rock. It might be possible that this fault zone could be more chargeable than other parts of the deposit.

4.6 Conclusions

All of the models produced at the deposit-scale build on the understanding obtained at the regional scale. The density model has been refined, through removal of a regional signal and finer discretization, in order to produce an anomaly that better represents the deposit. A much more detailed magnetic susceptibility model has also been produced.



(a) North-facing cross-section through chargeability model located at -400m north. Geology from Figure 4.2(a) is overlain.



(b) West-facing cross-section through chargeability model located at -1700m east. Geology from Figure 4.2(b) is overlain.

Figure 4.15: Cross-sections through deposit-scale chargeability model at the San Nicolás deposit.

However, the new model raises questions about the relationship between the distribution of magnetic minerals and massive sulphides at San Nicolás. The distribution of magnetic material is offset from the distribution of ore in what can be described as an over-printing effect. This over-printing effect suggests that the magnetic content of San Nicolás was produced after the deposition of the sulphides in a later hydrothermal event. Layered-earth inversions of the CSAMT data have produced a conductivity model that differentiates between the deposit and the overburden. When the models are "stitched" together to produce a 3D model, they have been shown to fit the data well at high frequencies. The simultaneous inversion of gradient-array data with "Real-section" IP data has produced a chargeability model that defines the deposit well, and has high values co-located with the main south-east dipping fault zone. These models were all constructed by inverting the geophysical data alone, and information in the form of physical property measurements and geologic interpretations have only been used to demonstrate the success of these models, and the validity of the techniques that produced them.

This work clearly demonstrates how geophysics can be used as part of a more focused exploration effort to spot drill-holes. All of these methods have proved successful in defining the deposit at the correct depths and lateral extents. Furthermore, the models start to produce more detailed information regarding the geometry of the deposit. The gravity inversion seems to be sensitive to the keel and the chargeability model produces increased values around the fault-zone.

The next step in the exploration program would be to use information that might result from drilling these well defined targets. The next chapter addresses this in an attempt to further refine the models and delineate more ore by including physical property information directly in the inversion procedure.

Chapter 5

Detailed Modeling: Incorporating *a priori* information

5.1 Introduction

Inversion of surface geophysical data to produce deposit-scale physical property models has been successful in recovering the general features of the San Nicolás deposit (as shown in Chapter 4). These models could have been used to successfully direct a drilling program that would have resulted in the discovery of San Nicolás. The next stage in the exploration process is to further delineate any mineralization that may have been found. Geophysical inversion can play a role at this delineation scale, as outlined by the shaded elements of the flowchart presented in Figure 5.1. If a target is drill tested, and whether or not mineralization is intersected, each drill-hole can provide additional information that can be used to further refine the inversion models. The inclusion of this data-independent information will help eliminate models inconsistent with known geology that fit the data [Scales and Tenorio, 2001]¹.

The additional information can be in the form of analysis performed on the drill-core itself, down-hole geophysical measurements, or even interpreted geology that can be assigned physical property values (this last option would obviously not provide as much information unless the physical properties were very constrained before hand). These updated models can, in turn, guide further drilling. This iterative process can expedite

¹In a similar manner additional information would be an important addition to inversions performed on in-mine data sets where mine geologists or mining engineers might require detailed information about parts of the earth where drilling has not yet reached. Again, the inclusion of known physical property information would constrain the result and focus the inversion on areas of interest.

a drilling program and more effectively define a resource.

The extensive drilling program that was completed at San Nicolás (Figure 5.2 shows the location of drill-holes that make up over 25,000m of drilling²) presents an opportunity to improve upon the deposit-scale models by including additional information about the subsurface in the inversion process.

The idea of including drill-hole information to constrain an inversion seems entirely acceptable, however, some major issues are raised when this put into practice. The issues include how to incorporate detailed point measurements in a large-scale model, and how to perform inversions that are consistent with the drill-hole information. This chapter presents a first look at these issues. The results demonstrate the potential benefits that may arise from including information from one drill-hole, and those that may arise from including information from many drill-holes. The process of generating these models identifies potential research areas.

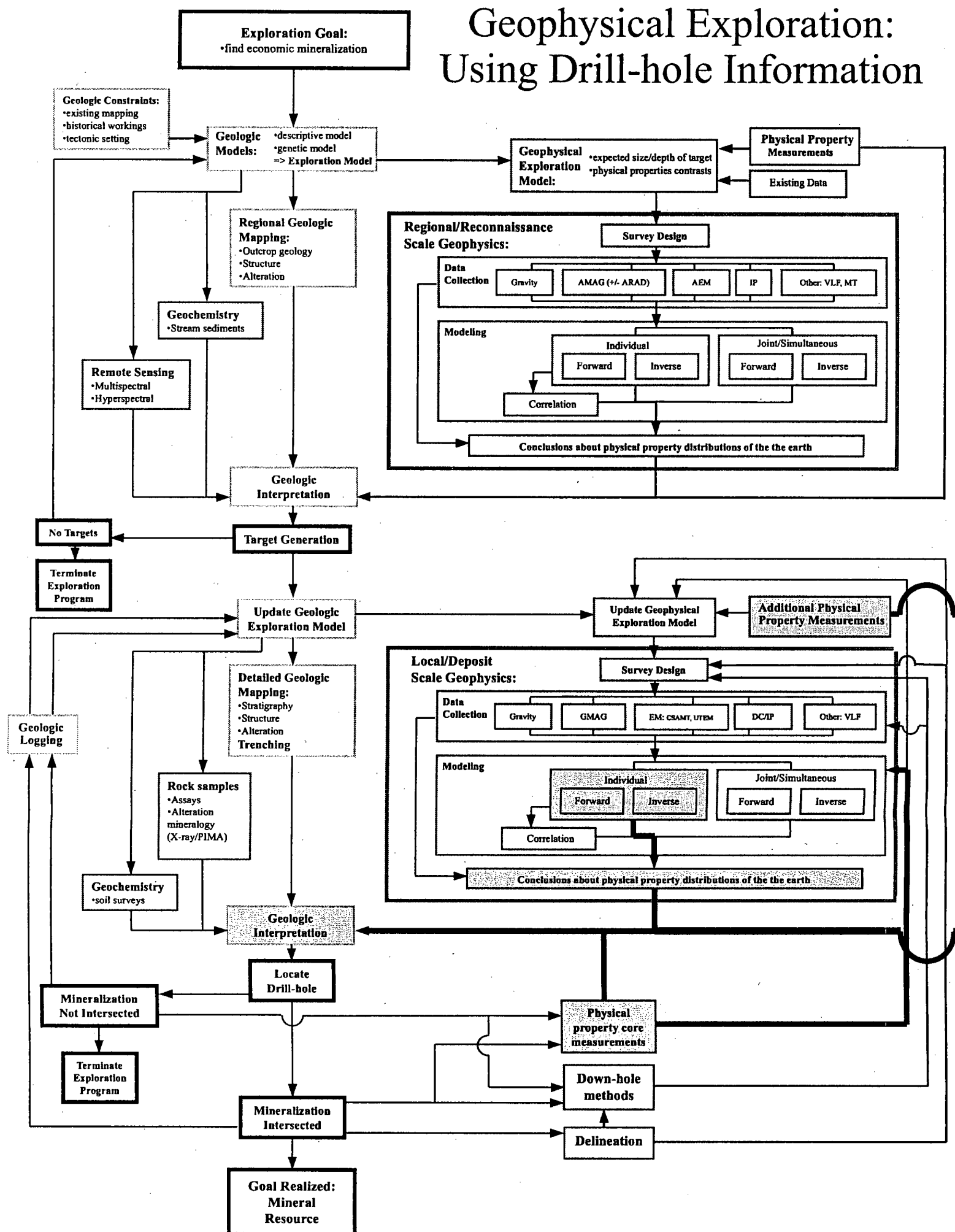
5.1.1 A priori information at San Nicolás

As well as providing geologic information about the subsurface, core samples were measured for density and magnetic susceptibility physical property values. Density measurements were made on sulphides from most drill-holes in order to estimate mine tonnages. The susceptibility measurements were made by myself during a visit to San Nicolás in February of 2001. Over one thousand magnetic susceptibility measurements were made on core samples from thirteen drill-holes. The measurements were made on all rock types in the local geologic section.

These physical property measurements are considered to be a source of explicit *a priori* information, that is, information that can provide direct knowledge about the physical

²The drill-hole notation of SAL represents *Salvador*—the San Nicolás discovery came from a drilling program that was initiated on the El Salvador project.

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property distribution prior to inverting the data. For this study, *a priori* information is incorporated in the inversion of gravity and magnetic data by directly constraining the range of values model cells can assume.

For the inversion of gravity and magnetic data, the *a priori* information is used in two stages. First, information from one drill-hole that intersects the deposit is used to constrain the inversion. This is done to determine if constraining part of the deposit can improve definition to other parts of the deposit. Secondly, all the available drill-hole information is included. This is performed in an attempt to identify any sulphide ore that has not been already drilled. If the physical property distribution of the deposit is well accounted for, it is hoped other parts of the model might be more accurately represented, thus improving the chances of finding more ore. Including *a priori* information in these two scenarios demonstrates the approaches that might be used at different stages of a drilling program.

5.2 Inversion of gravity observations

Density measurements were made on core samples of the sulphide ore, and some of the host rock material, that came from all of the drill-holes that intersected sulphide ore, as well as a few other holes. This is performed for reserve estimates, however, it provides an excellent source of information for constraining inversions. Measurements were made about every 1.5m along the core and a plot of density verses depth, with geology and assay results, for drill-hole SAL-31 is shown in Figure 5.3. The higher density values clearly correspond to the massive sulphides³, which have an average value of about 4.3 g/cm^3 , while the density of the mafic rocks below the deposit are about 2.8 g/cm^3 .

³The copper and zinc assay results show a change in ore-grade with depth, having larger concentrations of zinc at the top of the deposit and more copper present towards the base. The copper assay values seem to display a negative correlation with density; this probably reflects the relative concentrations of pyrite and chalcopyrite.

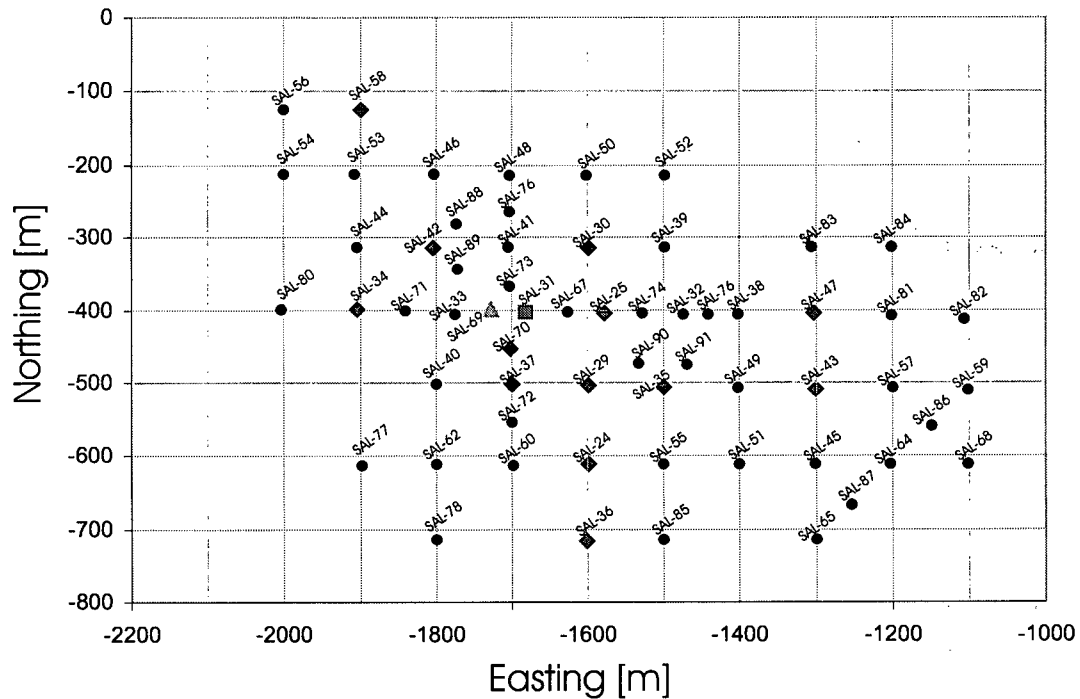


Figure 5.2: Plan-view of drill-hole locations. Density measurements on core from drill-hole SAL-31 are used to set bounds on the model values in Section 5.2.2 (shown by a red square), while measurements made on core from all of the drill-holes shown (apart from SAL-90 and SAL-91) are used to provide *a priori* density contrast information in Section 5.2.3. Susceptibility measurements made on drill-cores from SAL-69 are used to set bounds in Section 5.3.3 (shown by a green triangle), while bounds are assigned in Section 5.3.4 based on magnetic susceptibility values made on cores from thirteen drill-holes that are denoted by a blue diamond.

Discussion of bounds

The density measurements are used to constrain the density inversion model through the use of bounds that limit the range of values a cell can assume. The thought of bounding a 25m x 25m x 25m cell from values measured along drill-core that is five or six centimeters in diameter raises questions about the reliability of the process. There is obviously a discrepancy between the information we have and the information we require in order to constrain the inversion. This discrepancy is both a scale issue, and an issue relating to lateral inhomogeneities.

The information recorded in fine detail by the core measurements cannot be adequately represented by one cell value. In order to alleviate this problem we have two options. Either we can reduce the size of the cells or we can try to scale-up the drill hole information. Wherever possible it would be better to reduce the size of the cells so there is less inconsistency in scale between drill-hole information and the model. A smaller cell will probably contain a smaller range of values as sampled by a drill hole, and will probably be less likely to have large lateral inhomogeneities.

Unfortunately, the size of the model being used at San Nicolás was already large (close to half a million cells) so the decision was made to scale-up the drill-hole information. In doing so we lose valuable information.

5.2.1 Setting bounds on the density-contrast model

To bound a model, we require upper and lower bounds for each model cell. Drill-core density measurements provide very detailed and localized information (values every 1.5m) whereas cells within the density-contrast model exhibit constant values throughout 25m x 25m x 25m volumes. Therefore, processing of these core measurements is necessary in order for this information to be consistent with the density-contrast model, and to

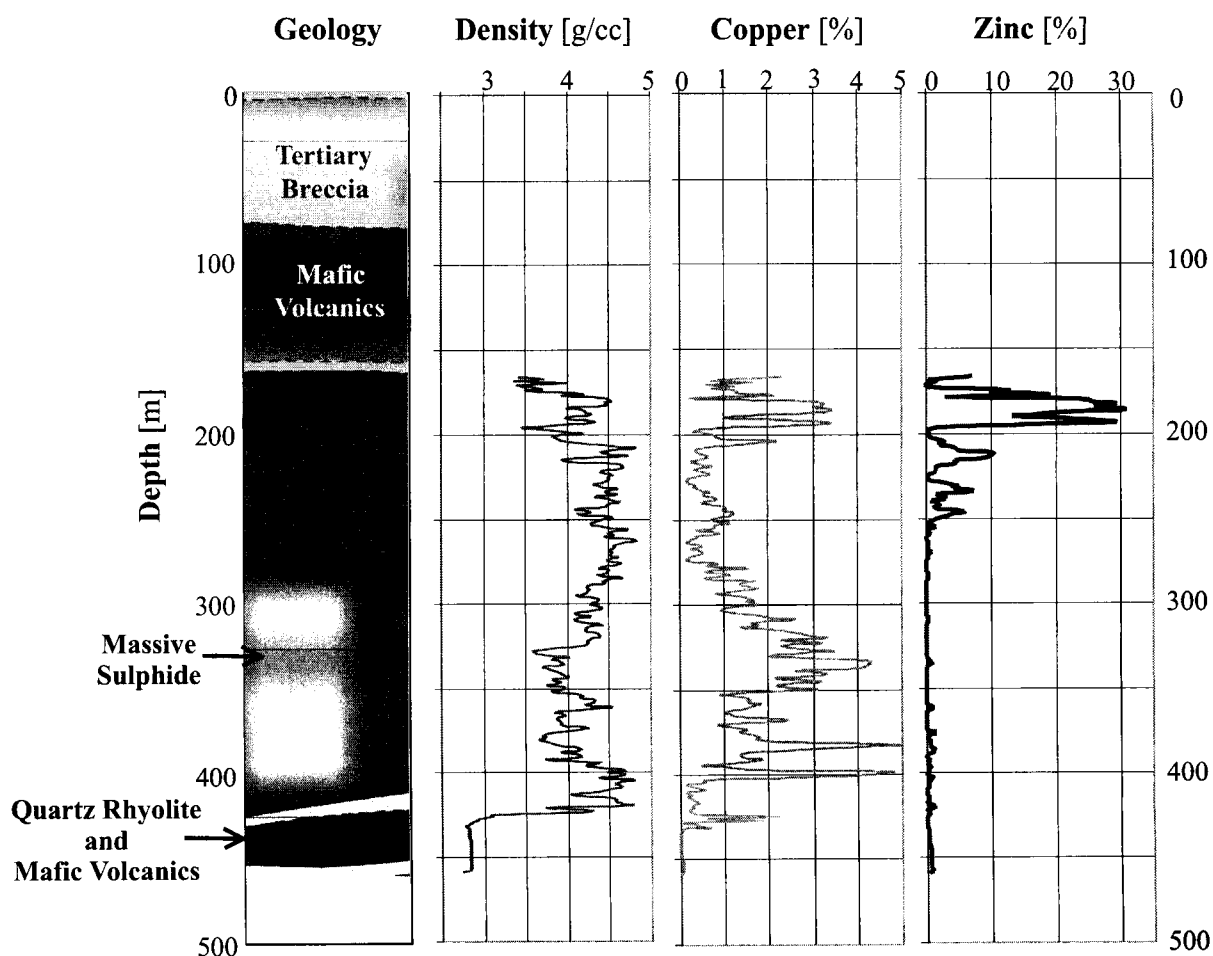


Figure 5.3: Drill-hole SAL-31. Displayed as a function of depth from left to right: 1) Geology, 2) density (g/cm^3), 3) percent copper, and 4) percent zinc. The average density of the sulphide is $4.2 g/cm^3$ and the density of the rocks below the deposit is approximately $2.8 g/cm^3$.

obtain the minimum and maximum density-contrast bounds for each model cell. Furthermore, the core measurements are absolute density values, while the model values are density contrasts relative to an unknown background value. To that end, inversions were conducted with a background value removed from the core density measurements. Considering the volcanic host rocks, it is expected that a background value of between 2.35 g/cm^3 and 2.5 g/cm^3 is reasonable.

One method of setting bounds on a model cell could be to assign the minimum and maximum core density values that are located within the model cell. It doesn't seem realistic, however, to expect the cell, which has a consistent value throughout, to exhibit these extreme values; a cell value closer to the mean core-density might be more probable.

The core-density measurements were sampled regularly throughout the vertical extent of each cell, and from these measurements the mean density value from the core was determined. The upper and lower bounds were calculated by adding and subtracting a small density variation from the mean. This variation could be based on the standard deviation of the measurements. However, the uncertainty in the result would decrease as more samples are used. For independent random variables, the uncertainty associated with the mean would be the standard deviation divided by the square-root of the number of samples. Although, as more samples are used the likelihood of them being independent will probably reduce.

In addition, other factors have to be taken into account. We are only sampling the cell at one position in x and y . The drill-hole would sample about three parts-per-million of the model cell (a five centimeter diameter drill-hole intersecting a $25\text{m} \times 25\text{m} \times 25\text{m}$ model cell) and, although we can consider vertical density variations within the cell at one horizontal location, we know nothing about lateral density variations within the cell. We might expect horizontal inhomogeneities to exist that could increase the range of values that the cell might have to represent. As a result, the difference between the upper and

lower bounds would probably increase. Apart from a knowledge of the geological setting, we have no way of knowing the possible homogeneities that might exist close to a drill hole.

At San Nicolás, the variation from the mean value ranged from 0.05 g/cm^3 to 0.2 g/cm^3 for different inversion runs. In addition, if the core-measurements only populated half of the cell height, the variation from the mean was doubled relative to cells with core-measurements throughout the cell; and cells with no core-measurements were assigned upper and lower density-contrast bounds of 5 g/cm^3 and 2 g/cm^3 respectively. These values might be a bit small although, based on the discussion above, it is hard to accurately determine these bounds at such large scales. This is one area where further investigation is needed.

Implementing the bounds

Once values for the upper and lower bounds have been determined, they are used as inputs in the inversion process. *GRAV3D* uses a logarithmic barrier method, directly incorporated in the objective function, to implement bounds on the model during inversion. An equation for the objective function with the barrier term is given in Equation 5.1. This function is minimized for subsequently decreasing values of the λ while keeping the regularization parameter (β) constant. When λ reaches zero the resulting model honors the *a priori* information provided.

$$\Phi(\lambda) = \Phi_d + \beta\Phi_m - 2\lambda \sum_{j=1}^M [\ln(\rho_j - \rho_j^{\min}) + \ln(\rho_j^{\max} - \rho_j)], \quad (5.1)$$

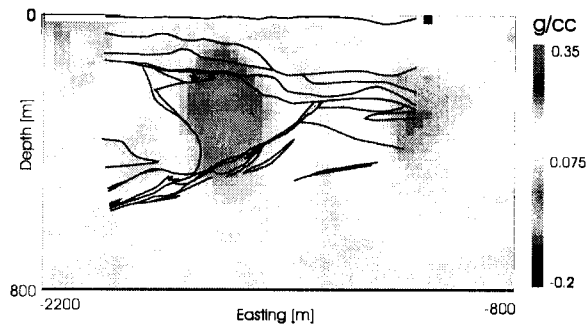
where: M is the number of model cells, ρ_j is the density of the j^{th} cell, and ρ_j^{\max} and ρ_j^{\min} are the upper and lower bounds of the j^{th} cell respectively.

For the subsequent inversions presented, unless otherwise stated, the same inversion parameters were used as those presented in Section 4.2.2 and Table E.1. That is, the same noise levels, reference model, and choice of regularization parameter are used.

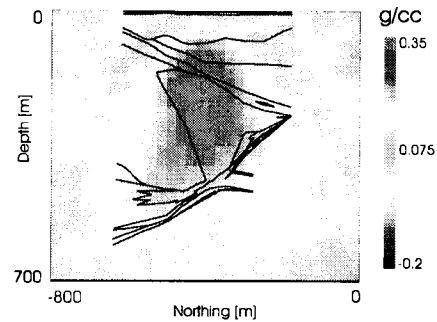
5.2.2 A priori density values from drill-hole SAL-31

Using the method described above, density values from drill-hole SAL-31 were used to constrain the deposit scale density-contrast model. The bounds implemented assume a background density value of 2.5 g/cm^3 , and constrain the model along a vertical column of cells, from 150m to 475m depth. The bounded cells correspond mostly with massive sulphides, and about twenty-five meters of interlayered, quartz rhyolite and mafic volcanics. The inversion result, displayed as east-west, and north-south cross-sections in Figure 5.4, shows an improvement over the results of inverting the gravity data without a priori information. The anomaly that corresponds with the sulphide ore has a higher density-contrast throughout most of the deposit, not just around the bounded cells (the density contrast increases from 0.3 g/cm^3 to 0.4 g/cm^3 75m away from bounded cells). These higher density-contrast values make the anomaly less spread out, and are consistent with expected values and a priori information. The model has also been well constrained along the fault that marks the lower extent of the deposit, and there is much less mass contained in the model below the deposit than previously modeled. Furthermore, the small density-contrast anomaly that was thought to correspond with the keel, in the inversion results presented in Section 4.2.2, is again present.

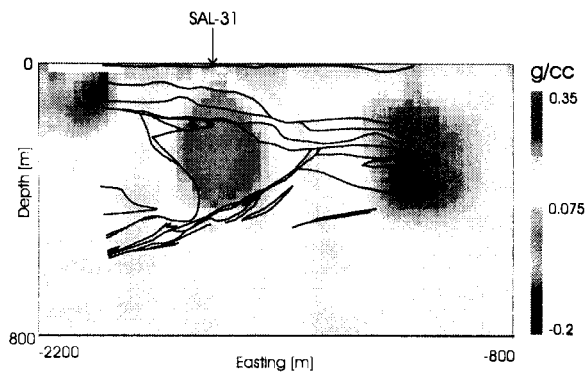
The model is more affected by the introduction of a priori information from drill-hole SAL-31. Changes over the original inversion are observed vertically above and below, and laterally adjacent to the bounded cells. This modeling suggests that high density material is not present below the fault, which is in contrast to the previous gravity inversion result that had a more significant continuation of high density-contrast values at depth.



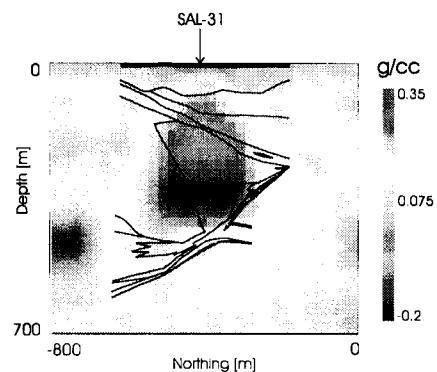
(a) Initial density model: north-facing cross-section with geology overlaid.



(b) Initial density model: west-facing cross-section with geology overlaid.



(c) Model bounded using density values from drill-hole SAL-31: north-facing cross-section with geology overlaid.



(d) Model bounded using density values from drill-hole SAL-31: west-facing cross-section with geology overlaid.

Figure 5.4: North- and west-facing cross-sections through recovered deposit-scale density-contrast models. The upper cross-sections show the results from the original deposit-scale gravity inversion, described in Section 4.2.2. The lower cross-sections show the results of inverting the gravity data with the model bounded using the density measurements made on core from drill-hole SAL-31. Geologic contacts have been overlaid.

Away from drill-hole, the model has not been significantly changed, and the next section addresses the results of incorporating all the available drill-hole information in an attempt to further improve the model, and determine the density structure below, and to the side of where drill-holes have sampled the earth.

Summary of using a single borehole to constrain a gravity inversion

The models presented above show the results of inverting gravity data with additional information obtained from a single drill-core. The results have improved over the previously inverted model and now we investigate the use of multiple drill-hole information.

5.2.3 A priori density values from all drill-holes

Figure 5.5 shows the north-facing interpreted geology section at -400S with the location of drill-holes. These drill-holes, along with many others to the north and south, are used to bound the density-contrast model during the inversion of gravity data, by providing explicit *a priori* density information. This is performed in order to account for the density distribution of the deposit and, hopefully, enable other parts of the model to be more accurately defined, particularly any extension to the keel.

The same method of bounding the model cells as described previously, is employed. Different inversions were performed using background densities of 2.35 g/cm^3 and 2.5 g/cm^3 , and bounds calculated with deviations of between 0.05 g/cm^3 and 0.2 g/cm^3 above and below the mean value for each cell. All the results were very comparable, showing no major differences, and the results of using a background density of 2.35 g/cm^3 with a 0.1 g/cm^3 deviation from the mean are presented in Figure 5.6.

As expected, the deposit is now very well defined by the density-contrast model, and the density structure matches the interpreted geology. Also, the density anomaly that corresponds with the deposit exhibits even greater density-contrast values, which is

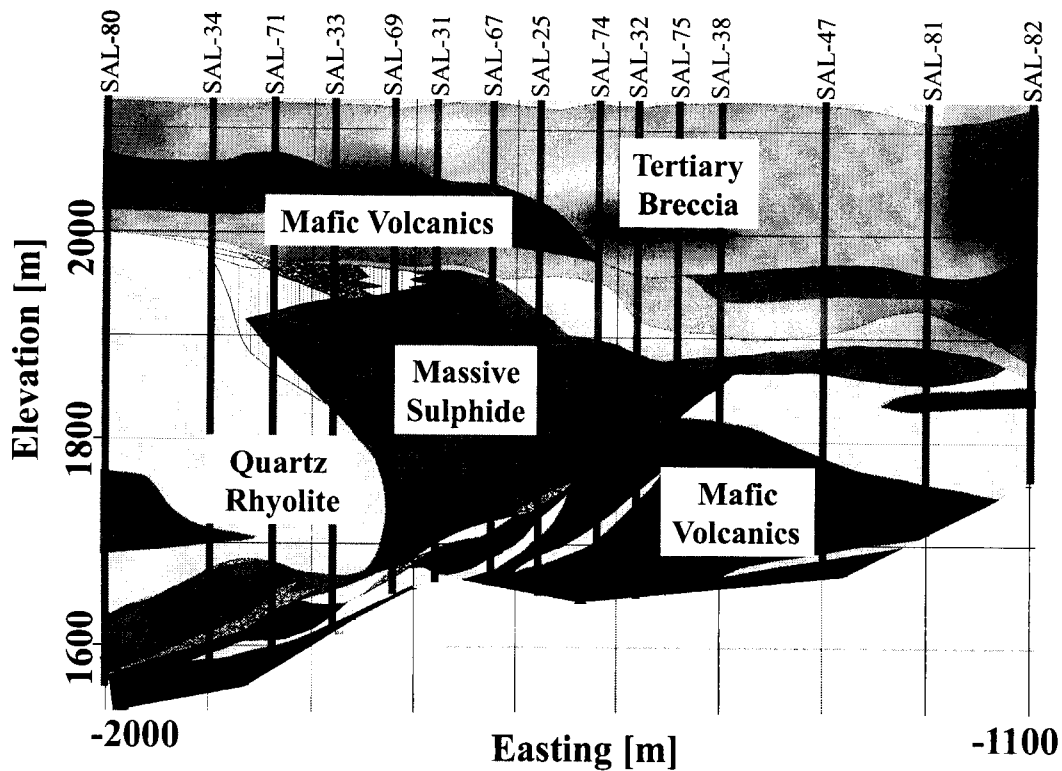


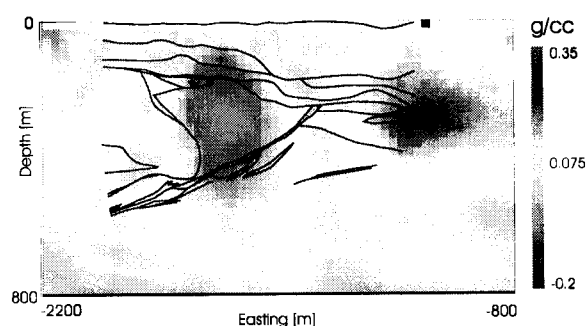
Figure 5.5: North-facing cross-section of interpreted geology with drill-hole locations. Density measurements made on core from the drill-holes were used to bound model cells during the inversion of gravity data. Only the drill-holes along line -400m S are displayed, although many more were used to bound the model.

consistent with the *a priori* information. Below the penetration depth of the drill holes the model is very uniform; giving more credence to the absence of high-density sulphides at depth. Apart from where *a priori* information has been explicitly incorporated, the keel has not been well defined as a distinct density-contrast structure. The smooth, deep, south-west extension to the density anomaly seen in the previous inversions of this data has been replaced by a few cells of high density-contrast. In addition there is no extension of the keel present in this model. This does not mean the keel terminates at the limit of the drilling. The result merely indicates a model can exist that doesn't exhibit an extension to the keel, and meets the following requirements: (1) it fits the data to a reasonable degree, (2) it is consistent with the *a priori* density information provided, and the way in which it has been implemented.

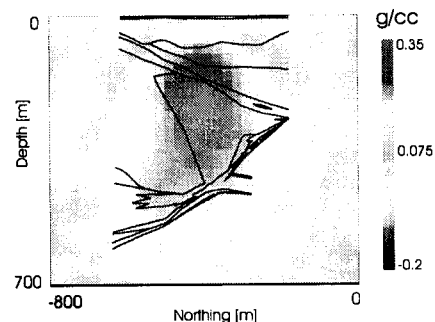
Summary of using multiple borehole to constrain a gravity inversion

Density contrast models have been produced by inverting gravity data while constraining the range of values some of the models can exhibit. This has been useful in defining the deposit, but more importantly accounting for the mass of the deposit will increase the resolution away from the drill-holes, especially below the deposit. In this case, a density anomaly that was observed when drill-hole information was not included virtually disappeared when the drill-hole information was included.

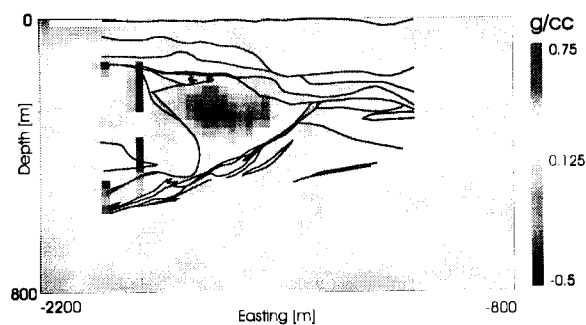
The recovered model does seem to be discontinuous in some areas. This is mainly due to the fact that in cells next to those that are bounded, the bounds are opened right up to accept almost any reasonable density value. Constraining cells next to those that contain drill-hole information is a philosophical one and relates back to the issue of lateral inhomogeneities that are expected in the geology. A discussion is presented at the end of this chapter that deals with some of the issues regarding the discontinuous nature of the recovered models.



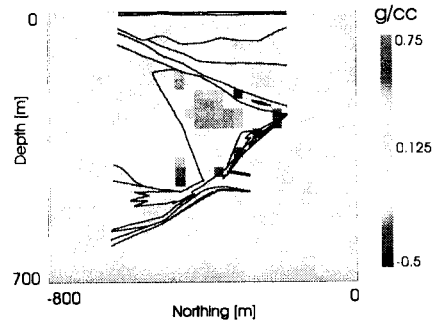
(a) Initial density model: north-facing cross-section with geology overlaid.



(b) Initial density model: west-facing cross-section with geology overlaid.



(c) Density-contrast model bounded by all drill-holes: north-facing cross-section with geology overlaid.



(d) Density contrast model bounded by all drill-holes: west-facing cross-section with geology overlaid.

Figure 5.6: Cross-sections through recovered deposit-scale density-contrast models. The upper cross-sections show the results from the original deposit-scale gravity inversion, described in Section 4.2.2. The lower cross-sections show the results of inverting the gravity data with the model bounded by density measurements made on core from all drill-holes. Interpreted geologic contacts are shown on each cross-section. Note: the lower cross-sections are shown with a range of density-contrast that is different from those previously presented.

5.3 Magnetic Susceptibility

Inversion of ground magnetic data has produced a model that shows part of the deposit is magnetically susceptible, parts of the deposit are void of susceptible material, and some rocks outside of the deposit are also likely to be magnetically susceptible. Over one thousand magnetic susceptibility measurements, made on core samples from thirteen drill-holes, confirm the complexity of the susceptibility distribution (Appendix B.3.1). Figure 5.7 shows magnetic susceptibility values along with geology for core from the four drill-holes: SAL-25, SAL-34, SAL-47, and SAL-70. This drill-hole information, and that from other drill-holes, is used to constrain the inversion of ground magnetic data in an attempt to improve upon the existing models. Firstly, measurements from drill-hole SAL-69 are included, and secondly, all the available measurements from the drill-holes are included.

5.3.1 Core Measurements

A KT-9 magnetic susceptibility meter was used to make measurements on drill-core samples at San Nicolás. The instrument gets 90% of its signal from the first 20mm of the sample, can be calibrated for core measurements, and has a maximum sensitivity of 1×10^{-5} S.I. units. The instrument was also calibrated in free air before each reading.

Several test measurements showed that the susceptibility distribution within the deposit could dramatically change from less than 1×10^{-4} to over 0.1 S.I. within a few centimeters. As time constraints did not permit very fine sampling of the core, I took readings approximately every 3 meters in rocks that were expected to exhibit low susceptibilities, every meter in rocks with high or variable susceptibilities, and several readings were collected either side of faults and geologic contacts. Tests were conducted over different rock types to see how effectively this sampling characterized the finer susceptibility

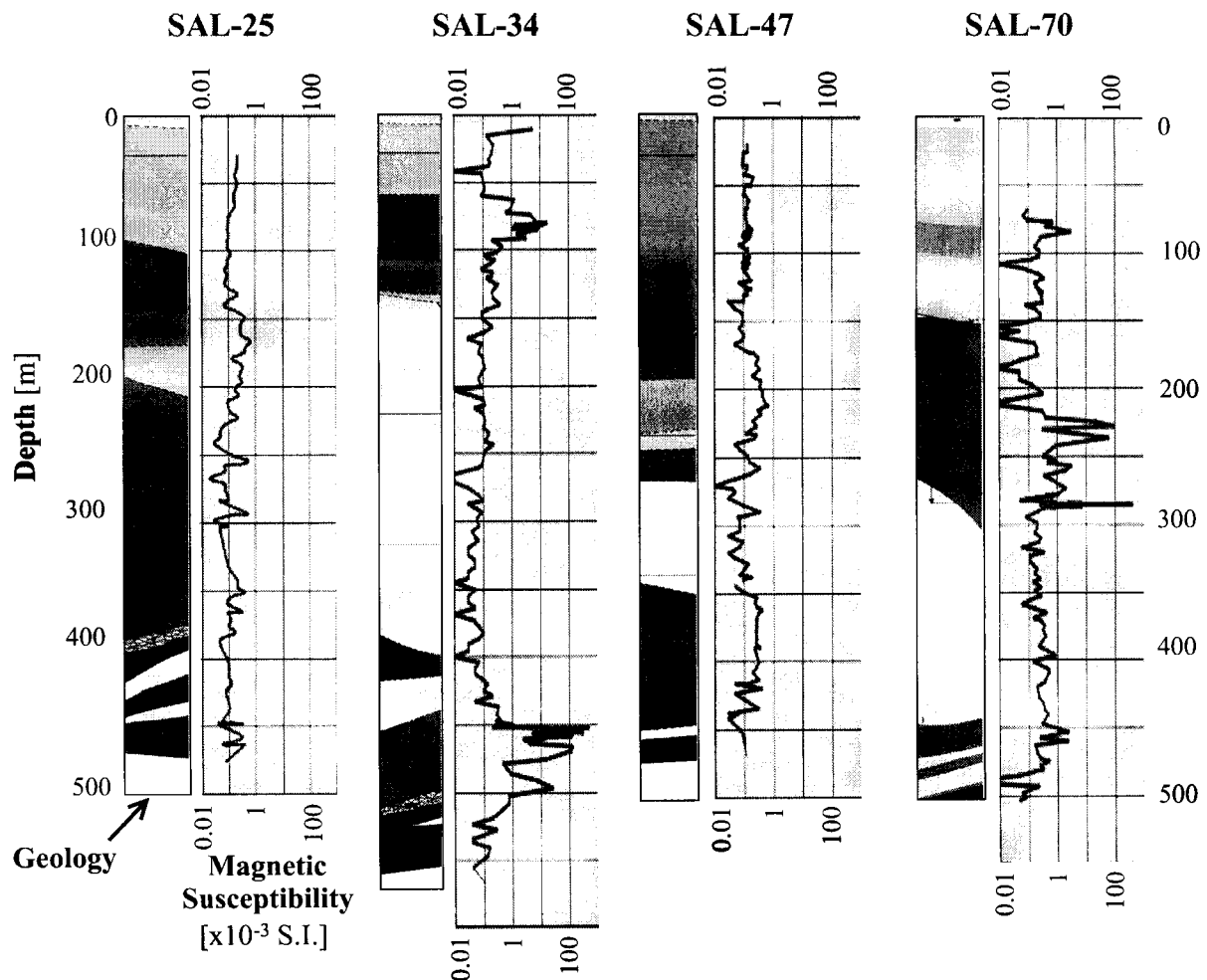


Figure 5.7: Magnetic susceptibility core measurements and geology for drill-holes SAL-25, SAL-34, SAL-47, and SAL-70. Note: the upper-most geologic unit is the Tertiary breccia, although it may appear a similar colour to the sulphide ore, which is shown in red. Reference the geologic section and recognize it as an interpretation.

distribution, the results of which are presented in Appendix B.3.1.

5.3.2 Bounding the magnetic susceptibility model

The core susceptibility measurements are used to set upper and lower bounds on model cells, which are implemented using the logarithmic barrier method. As with the density measurements, many core susceptibility readings were used to make up one value for each model cell around which bounds were centered. Unlike the density measurements, however, no background level need be removed prior to including this data in the inversion. Also the samples were not evenly collected along the core. The uneven sampling had to be taken into account when calculating the composite susceptibility value to use for each cell (Appendix B.3.2).

Upper and lower bounds were centered around the composite susceptibility value. As with the density measurements, the magnitude that the upper and lower bounds vary from the composite value is debatable. Most of the core susceptibility values were around 0.1×10^{-3} S.I. with some values as high as 1×10^{-3} S.I. and a few as high as 100×10^{-3} S.I.. The value that was chosen to vary the bounds by was 5×10^{-4} S.I., however, more information might be contained in the inversion by using a bound that was related to the core measurements themselves.

5.3.3 Magnetic inversion using *a priori* information from SAL-69

Constraining the inversion of ground magnetic data by using bounds on the susceptibility model has produced a model of the San Nicolás deposit that has significant differences with respect to that which was produced without *a priori* information. Figure 5.8 shows a comparison between the magnetic susceptibility models recovered with and without *a priori* information from SAL-69.

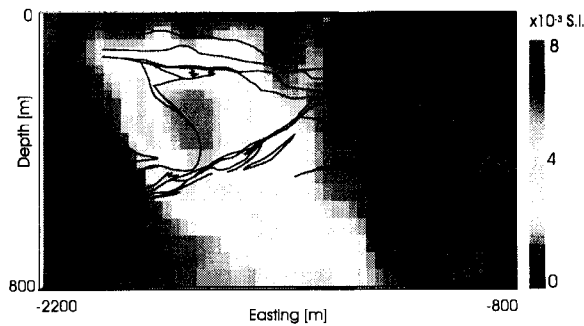
When the model is bounded, the concentration of susceptible material in the west part of the deposit exhibits higher susceptibility values, occupies less volume, and is tightly constrained between the depths of 275m and 350m. Moderate ($2-3 \times 10^{-3}$ S.I.) susceptibility values are still present below the deposit and the north extension of the susceptibility anomaly is also unaffected by including the *a priori* information.

5.3.4 Magnetic inversion using *a priori* information from all drill-holes

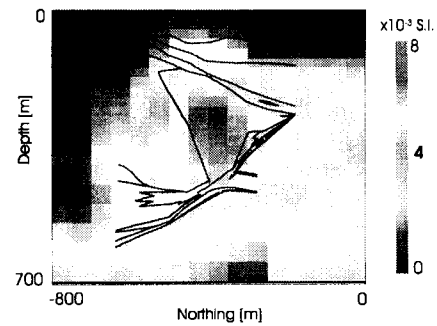
A priori susceptibility information was used from all twelve drill-holes measured in the inversion of ground magnetic data. The results show a susceptibility distribution that is concentrated in the southwest part of the main ore-zone, the keel, and to the northwest of the delineated deposit. Figure 5.9. The additional *a priori* information helps constrain the susceptibility distribution within the main ore-zone to a small volume that is located against the rhyolite dome to the west and the southwest dipping fault to the northeast. Also, high susceptibility values have been introduced in the keel that have not been previously modeled. These high susceptibility values in the keel are introduced through the presence of drill-hole information but persist in cells next to the drill-holes where the model was not constrained. This suggests that drill-hole information is improving the model in cells that are not bounded by drill-hole information.

A continuation of high susceptibility values persist to the north and moderately high values are present below the deposit. Less-altered, mafic volcanic rocks could be responsible for the susceptible concentrations below the deposit.

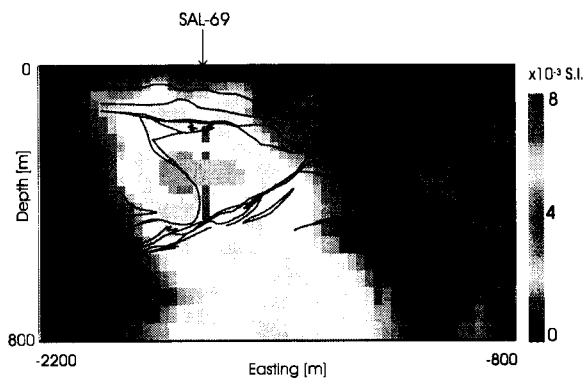
Different inversion models were constructed using different alpha coefficients to smooth the model in different directions. The results show that the effect the bounded cells have on laterally surrounding cells is dependent on the relative amount of horizontal (x-component and y-component) weighting that is introduced. This could be used to construct more coherent models when geologic stratigraphy can be assumed.



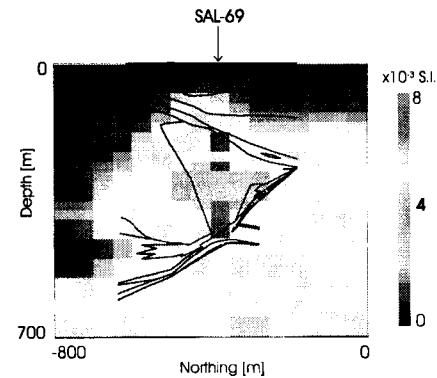
(a) Initial magnetic susceptibility model: north-facing cross-section



(b) Initial magnetic susceptibility model: west-facing cross-section

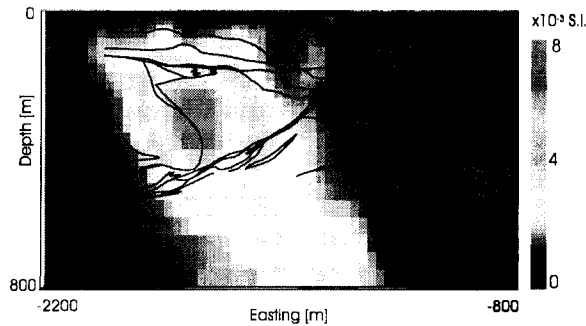


(c) Magnetic susceptibility model bounded by SAL-69: north-facing cross-section

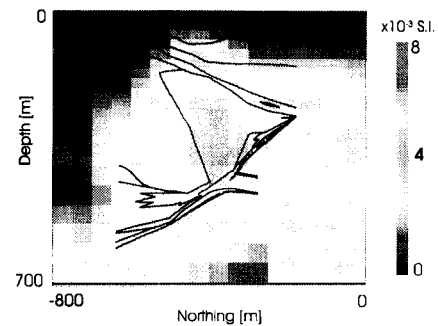


(d) Model bounded by SAL-69: west-facing cross-section

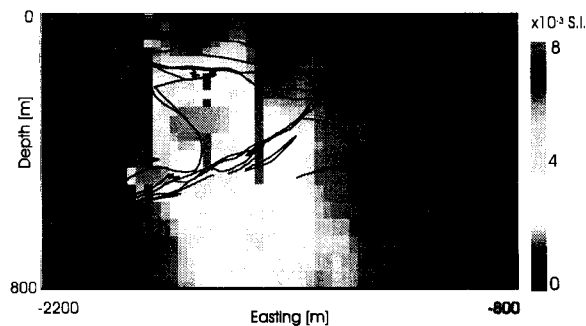
Figure 5.8: North- and west-facing cross-sections through recovered deposit-scale magnetic susceptibility models. The upper cross-sections show the results from the original deposit-scale magnetic susceptibility inversion, described in Section 4.3. The lower cross-sections show the results of inverting the magnetic data with the model bounded by magnetic susceptibility measurements made on core from drill-hole SAL-69. An outline of geology interpreted from drill-holes is overlaid on each section.



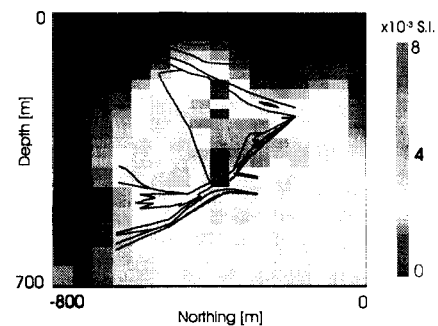
(a) Initial magnetic susceptibility model: north-facing cross-section



(b) Initial magnetic susceptibility model: west-facing cross-section



(c) Magnetic susceptibility model bounded by thirteen drill-holes: north-facing cross-section



(d) Model bounded by thirteen drill-holes: west-facing cross-section

Figure 5.9: North and west-facing cross-sections through recovered deposit-scale magnetic susceptibility models. The upper cross-sections show the results from the original deposit-scale magnetic inversion, described in Section 4.3. The lower cross-sections show the results of inverting the magnetic data with the model bounded by magnetic susceptibility measurements made on core from twelve drill-holes. Geologic contacts as interpreted from drill-holes are overlaid on each section.

5.4 Discussion on the discontinuous nature of the models

When examining Figures 5.4(c), 5.4(d), 5.6(c), 5.6(d), 5.8(c), 5.8(c), 5.9(c), and 5.9(d), it seems apparent that models produced with *a priori* information are somewhat blocky and discontinuous. This is consistent with the *a priori* information that reflects discrete geologic boundaries, but it is not consistent with the smooth requirements of the model objective function. In fact, the blockish nature of the bounded model is counteracting the efforts of the model objective function to produce a smooth model. A more consistent way of producing a model would be to have both the model bounds and the model objective function acting in concert; both seeking a model of the same type. This could be achieved by changing the bounds to be smooth, or by changing the model objective function to allow blocky models. In this case, changing the bounds would be dishonoring the *a priori* information, so the appropriate alternative would be to favor blocky models. This can be accomplished in the following two ways: (1) the W_x , W_y , and W_z weighting matrices could be changed so the model doesn't have to be smooth in the vicinity of sharp discontinuities, as determined from the *a priori* data, and (2) blocky, or piecewise-continuous, models could be sought through the use of the L_1 norm in the model objective function, as discussed in [Farquharson and Oldenburg, 1998].

5.5 Conclusion

The introduction of *a priori* information into an inversion can improve the resulting model. For the potential field examples presented here, the results have improved around where drill-hole information has been included, and below the known parts of the deposit. All of the results produced fit the data to the same degree to that used in Chapter 4.

With the incorporation of information from one drill-hole, expected physical property values were increased over much of the deposit. The density contrast was reduced below

the deposit, while the susceptibility distribution was constrained to a range of depths within the deposit. This demonstrates how *a priori* core measurements could be used in the early stages of drilling to help direct subsequent holes.

Incorporating all *a priori* information has produced models that define the deposit, in the case of the density-contrast model, and define a localized concentration of susceptibility in the western part of the deposit, in the case of the susceptibility model. The *a priori* information is useful in defining the deposit so that parts of the signal represented by the data can be accounted for, and areas that have not been drilled may be more accurately modeled. The gravity inversion has shown a decrease in density values below the deposit that suggests there is no high density sulphide ore below the level of drilling. None of the models presented exhibit an extension to the keel but the data may not be sensitive to such a small, and deep, body (the sensitivity of the data to the keel is addressed in Chapter 6).

The issues involved with incorporating drill-hole information have been presented and only a first pass solution has been formulated. It is understood that the work presented here represents a preliminary attempt to resolve these issues but also demonstrates the potential benefits that can come about if drill-hole information can be included.

Issues that still require investigation include:

1. *A priori* information can originate from finely sampled measurements made on drill-core. This information is processed in order for it to be consistent with the large model cell volumes that are used to constrain the inversion using bounds. This processing is not trivial. Errors are introduced that relate to the sampling size of the cells, variations of physical property within a cell along a drill hole and lateral variations that can only be assumed. In addition, an estimate of this error is needed to determine upper and lower limits of the bounds. It should also be realized that

the finely sample drill-hole information is itself a coarse representation of the actual susceptibility distribution within the earth (Appendix B.3.1).

2. Removal of regional gravity signals and assignment of appropriate background density values are issues that need to be addressed when using density information in gravity inversions.
3. The recovered models that have been bounded have a blocky appearance. This could be due to the inappropriate application of a smooth component model objective function when a block model is known to exist. Methods of recovering models that are consistent with known geology has a greater scope than just incorporating bounds in a model. We have to address how the model is constructed.

The difficulties in implementing drill-hole information accurately in inversions suggests that actual down-hole measurements might be better suited to producing inversion models. In such cases all the data will be consistent with each other.

In summary, this chapter has addressed the methods that might be employed to incorporate drill-hole information in an inversion. This study reveals important aspects that require further investigation. However, removing ambiguities from an inversion model by constraining a model could have major benefits. Superior models can be produced through the use of just one drill-hole. This can lead to more accurate spotting of drill-holes in an iterative exploration program.

Chapter 6

Forward Modeling and Survey Design: Detecting the keel

6.1 Introduction

Chapter five began looking at the delineation of mineralization. That work is continued here. In the context of an exploration program (Figure 6.1) we address the issue of delineation by trying to conduct surveys that are adequately sensitive to the target. In the case of San Nicolás, detecting a deep ore zone involves modification of survey designs.

The main objective of delineation is to define the major features of the deposit. At San Nicolás a major feature is the keel. The keel makes up about one sixth of the volume of the deposit, is located below and to the west of the main ore-zone, and was found through drilling.

Despite the success of inverting surficial geophysical data at San Nicolás to define the deposit, the keel was not readily identified (although possibly by the gravity inversion). The primary reason for this is that the signal (anomalous field) from the keel is generally small. We quantify these signals through forward modeling.

As will be shown, some of the data collected at San Nicolás were probably not adequate to detect the keel. This work leads naturally to the question: "What data are required to find the keel?" As a result, the focus of this chapter is in two parts.

Firstly, the signal from the keel is quantitatively evaluated. This is performed through forward modeling of synthetic physical property models with and without the keel. The difference between the two responses is compared with assumed noise levels in order to

determine if the response from the keel is observable.

The second focus of this chapter is to obtain data that are substantially influenced by the keel. For gravity and magnetic data synthetic measurements are made down-hole. The simulations provide insight about the increase in signal strength as readings are made closer to the keel. For electromagnetic and DC/IP data (active source methods), there is much more flexibility for designing an informative survey. This is because both sources and receivers can be placed close to the keel in order to increase signal strength. The methodology for designing optimum surveys is a separate research project; one that is not addressed in this thesis. However, computing the response of the keel for simple survey design changes is performed. This may provide a foundation on which a more comprehensive analysis of the survey design problem can be addressed.

Finding these small ore-zones¹ is more difficult in the absence of a large amount of known sulphide, as it becomes harder to instigate follow-up, detailed exploration. As a result, a method of detecting these from the surface becomes important.

Information about the keel is obtained from geology sections that were interpreted from drill-holes. This information, along with linear interpolation methods, is used to construct physical property models, or *rock-models* (digitized on 25m x 25m x 25m cells), on which forward modeling can be conducted. The main ore-zone, containing massive sulphide, semi-massive sulphide, and sulphide stringer zones, is modeled as having a volume of about seventeen million cubic meters, and is located between a depth of 175m and 425m. The keel, on the other hand, is modeled as having a volume of about three million cubic meters (one sixth the size of the main zone), and is located between depths of 425m and 550m. This helps put in perspective the different signals that might come from the keel and the main-ore zone, and suggests that different surveys maybe

¹Smaller amounts of ore, that are not economically feasible to mine on their own, could be economically viable because of the mining infrastructure that might be in place due to San Nicolás.

Geophysical Exploration: Survey design for the delineation of deep ore

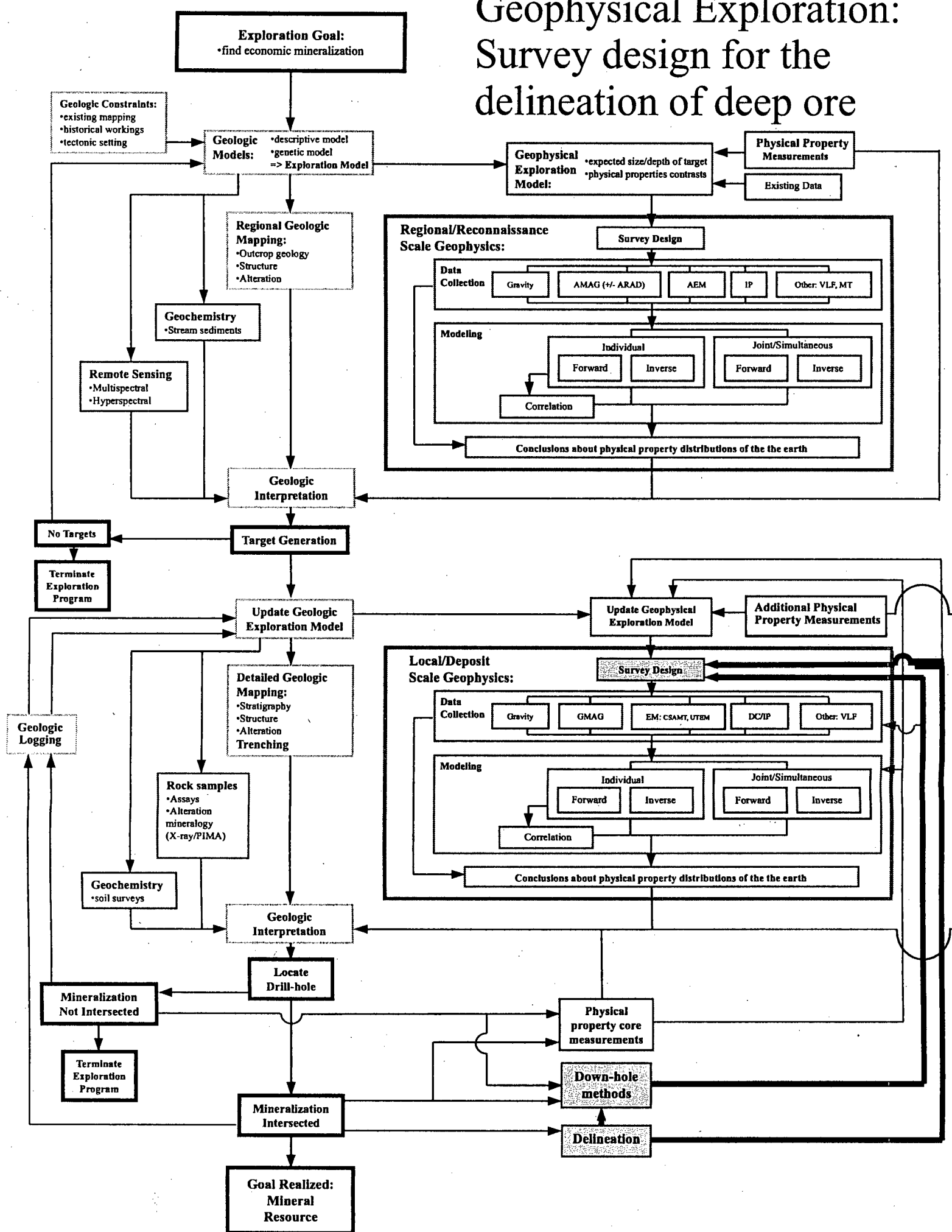


Figure 6.1: Geophysical Exploration Methodology: Survey design

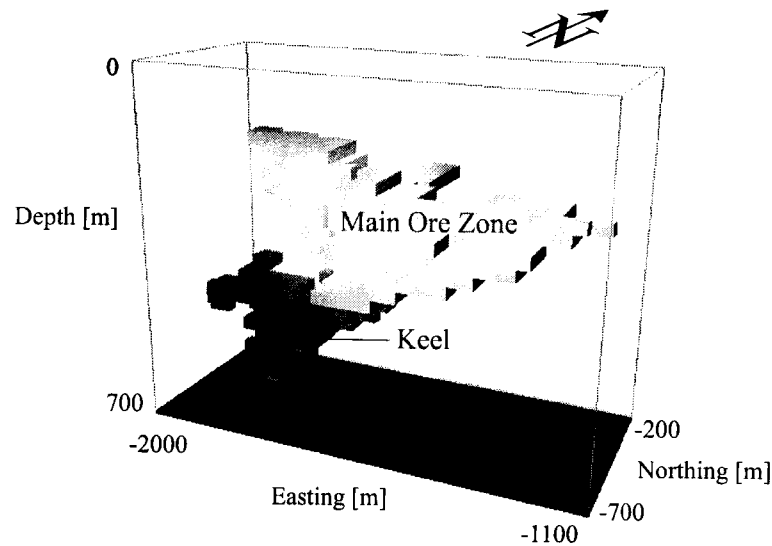


Figure 6.2: Size and position of the keel relative to the deposit.

needed in order to detect the keel.

6.2 Gravity Observations

Inversion of gravity observations in Section 4.2 have produced a model that, along with approximately defining the main ore zone, had a distribution of higher density-contrast values in the vicinity of the keel. This suggests that the data contain information about the keel above the noise level. In this section the magnitude of the gravitational response of the keel is investigated by performing synthetic gravity surveys on digitized density rock-models. The results are compared to the estimated instrument sensitivity and noise levels in order to see if gravity readings could be used to constrain an inversion model of the keel. A synthetic survey is conducted that is similar to that conducted in the field, and several down-hole gravity surveys are also performed.

As we are not able to change the mass or location of the keel, the only way to

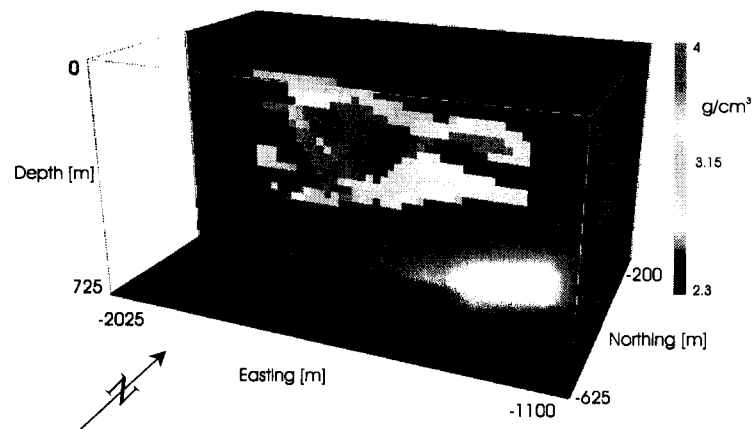
increase the signal from the keel is to decrease the distance from the receiver to the keel. To that end, synthetic down-hole gravity measurements are made in order to determine how near to the keel a drill-hole need pass for a discernable signal to be observed.

6.2.1 Simulation of gravity field observations

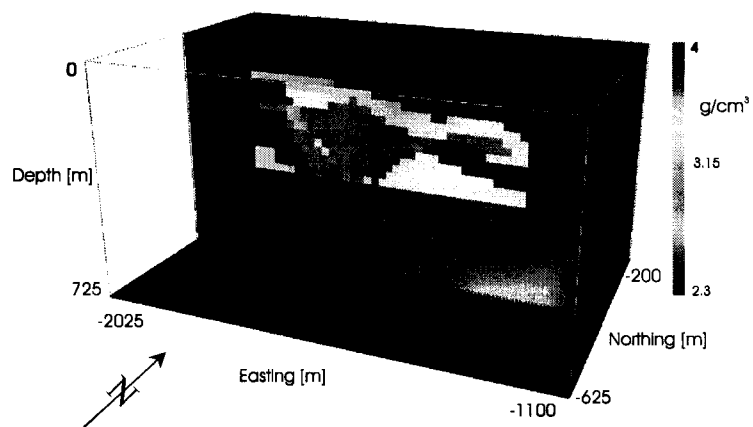
Two sets of synthetic gravity data are made on different density rock-models. One model contained density distributions of the main ore-zone and the keel, while the other model had density distributions of only the main ore-zone. By comparing gravity observations made on each model one can determine the magnitude of the response of the keel. This response is compared with estimated noise levels of between 0.02 and 0.05 mGal. In order to confidently distinguish the response of the keel in the data, signal-to-noise levels of greater than 3 are assumed to be needed, therefore, a gravitational response from the keel of greater than 0.06 mGal is required.

The 3D density rock-model was created by digitizing geologic cross-sections provided by Teck, assigning density values to the rock units, and interpolating between the 2D sections. The density value assigned to each geologic unit is determined from the measurements made on the core, but small density variations within each unit are not taken into account (a constant value of 4 g/cm³ is assigned to the massive sulphide ore). The density rock-model, which covered only a limited volume, was then embedded within a background of 2.5 g/cm³ material. Two density rock-models were constructed (Figure 6.3); one with the keel, as digitized from the geologic sections, and one with all density values below a depth of 425m replaced by with values of 2.5 g/cm³. Replacing the density values below 425m effectively removes the keel from model and, although some minor density variations are also changed during this process, it is assumed that the keel represents most of the anomalous mass at these depths.

Five hundred and eighty-five synthetic gravity observations were made on the surface



(a) Perspective view of density rock-model showing an east-west section at -400m north.



(b) Perspective view of density rock-model with the keel removed. The model is cut at -400m north to reveal an east-west cross-section.

Figure 6.3: The two density rock-models that were used to generate synthetic gravity measurements.

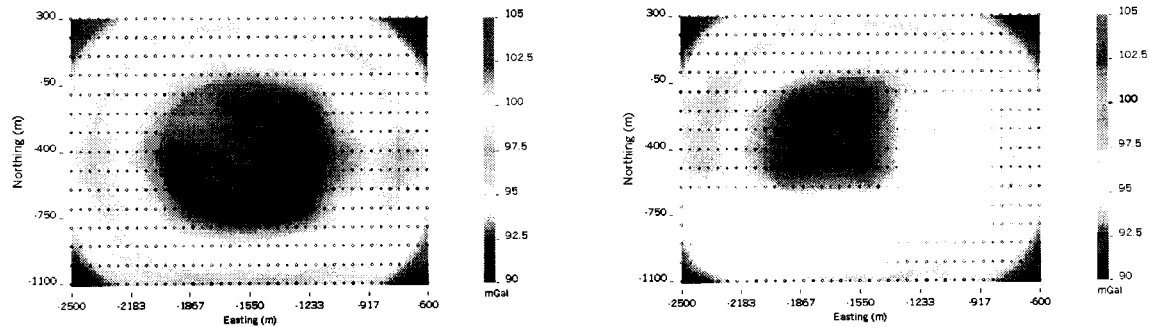
of each density rock-model. The observations, shown in Figure 6.4, have a station spacing of 50m and a line spacing of 100m, which is similar to some of the more closely spaced readings taken at San Nicolás. The difference between the observations made on the model with the keel and those made on the model with the keel removed is shown in Figure 6.4(c). The difference ranges in magnitude from 0.004 to 0.12 mGal and, in comparison to estimated noise levels (between 0.01 and 0.05 mGal), the response from the keel gives a discernable response. Assuming the densities were appropriately assigned to the rock-model, this has the important implication that information specific to the keel could be extracted from surface gravity data, and is consistent with the inversion model, which exhibits an extension to the deposit where drilling indicates the presence of the keel.

6.2.2 Synthetic, downhole, gravity observations

Accurate determination of the keel's density distribution would likely come from inverting gravity observations that contain information about the keel significantly above noise levels. If we assume that noise levels are constant, the only other option would be to increase the gravitational response of the keel. Because we are dealing with a potential field, this can only be done by making observations nearer to the source; we do this by performing down-hole measurements. The use of drill-holes in geophysical surveys is a logical progression of survey design at San Nicolás because of the extensive drilling that has occurred, and although down-hole gravity measurements are not often carried out in the minerals industry², they have the potential to produce data with low noise levels because of the quiet environment in which they are made.

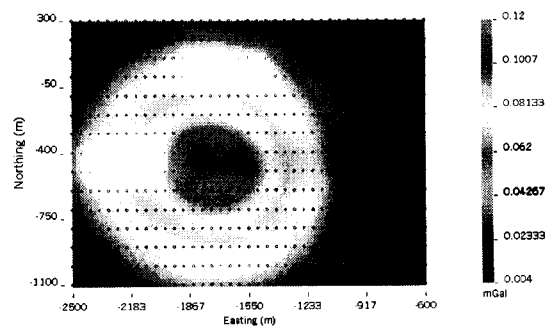
Six synthetic drill holes were used to collect down-hole gravity measurements, at 10m

²Down-hole gravity measurements are not often performed due to the expense of the instrumentation and the possibility of losing it down a hole.



(a) Synthetic gravity data produced from density rock-model.

(b) Synthetic gravity data produced from density rock-model without keel.



(c) Difference between synthetic gravity data with keel and synthetic gravity data without keel.

Figure 6.4: Synthetic gravity data produced from a density rock-model of San Nicolás. Data are produced from models with and without the keel. The difference is also plotted.

intervals, to a depth of 650m on the rock-model. Figure 6.5 shows the locations of these drill-holes plotted on a west-facing cross-section through the density rock-model.

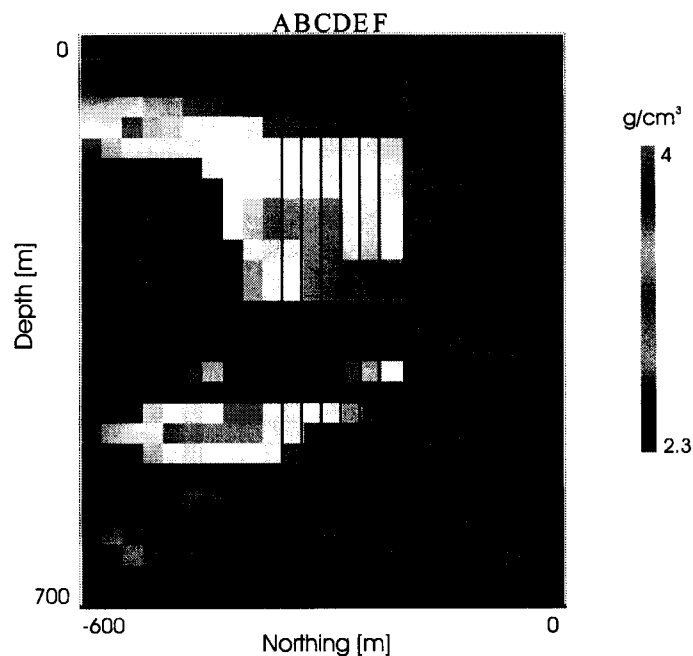


Figure 6.5: West-facing cross-section of density rock-model. The main ore-zone is represented by the higher density values in the top half of the model while the keel is represented by the anomaly in the lower half of the model. The locations of six drill-holes, within which synthetic gravity measurements are made, are also shown. Drill-hole A is 25m north of the keel and each subsequent drill-hole is positioned 25m further north than the last one.

The difference between the gravitational response of the model with the keel and the model with the keel removed is shown for measurements made down the six drill-holes in Figure 6.6. Drill-hole A is located on the north edge of the keel and, at a depth of between 450m and 500m, which corresponds to the depth of the keel, the vertical gravitational response changes from 0.7mGal to about -0.5 mGal. This response is very much above the expected noise levels, and would definitely indicate an anomalous mass. As measurements are made down drill-holes further to the north, the response from the

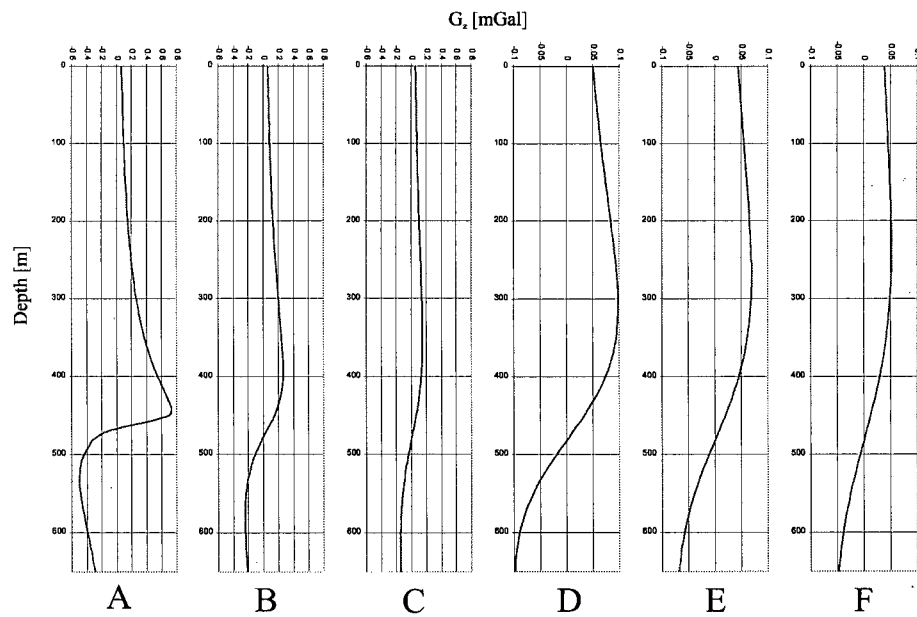


Figure 6.6: Down-hole plots of synthetic gravity data made in six drill-holes. The data are the difference between the response of a density rock-model with the keel, and a model with the keel removed. Note the change in horizontal scale for plots C, E, and F.

keel is seen to smooth out and decrease in value to ± 0.05 mGal in drill F, 250m north of the keel. The magnitude of response from the keel in drill-hole F is close to the expected noise levels. It is likely that gravity data collected in drill-holes further than 150m from the keel would have difficulty in detecting the keel from the noise.

6.2.3 Summary of gravity forward modeling.

A simulation of the observations made in the field show that surface gravity data could contain sufficient information for the keel to be determined. This result adds credence to the inversion results obtained in Section 4.2. With assumed noise levels, down-hole measurements should be made within about 150m of the location of the keel.

6.3 Magnetic Observations

Inversion of the ground magnetic data was not successful in modeling the keel, even though magnetic susceptibility measurements made on drill-core show the keel to have high susceptible values. Forward modeling of magnetic data is performed to determine if the field ground magnetic data that were inverted in Section 4.3 could detect the keel. In addition, down-hole synthetic measurements are made in an attempt to increase signal levels of the magnetic response of the keel, and determine how near such down-hole measurements need be in order to detect the keel.

Forward modeling is conducted on a magnetic susceptibility rock-model. As outlined in Appendix B, the susceptibility rock-model is constructed using interpreted geologic sections and assigned magnetic susceptibility values based on core measurements³. A second susceptibility model is constructed which is the same as the one just described,

³Even though core susceptibility measurements show that the distribution susceptible material isn't very predictable within the deposit, it does seem to be consistently high within the keel, especially along the fault.

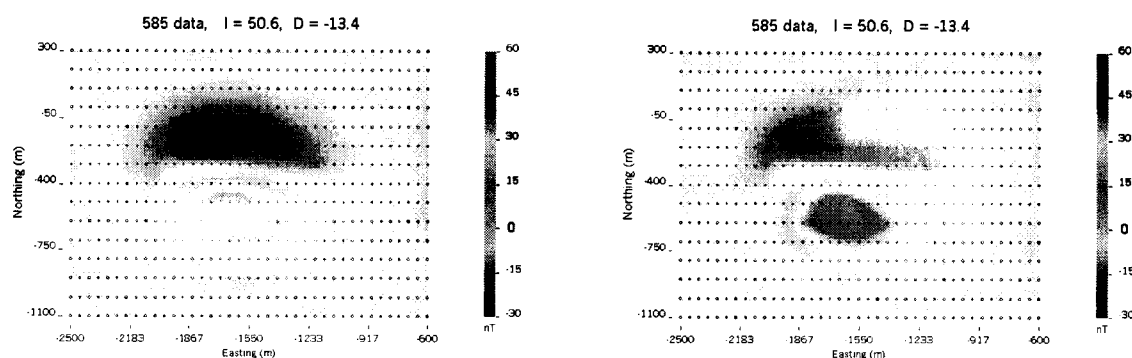
except all susceptibility values below 425m are replaced with a value of zero. The difference between the data collected on each of these models will be used to determine the magnitude of the response of the keel. For the field ground magnetic data a noise level of 2nT was used during the inversions. This resulted in a model that corresponded with the deposit yet was devoid of overly complex structures. Based on this information, a signal level of about 5nT is expected from sources in order for them to be adequately discerned from the noise.

6.3.1 Simulating the ground magnetic field data

Five hundred and eighty-five surface measurements of the total magnetic field are calculated over both susceptibility rock-models. The inducing field strength and orientation was approximately that in which the original field data were collected. The synthetic data are shown in Figure 6.7 along with a plot of the difference between the two. The response from the keel (as depicted by the plot of the difference between the two data sets) ranges from about -1 to 3 nT in magnitude. This signal level is above a noise level of 2 nT, but only exhibits a signal-to-noise ratio of 1.5; smaller than that required for confident detection.

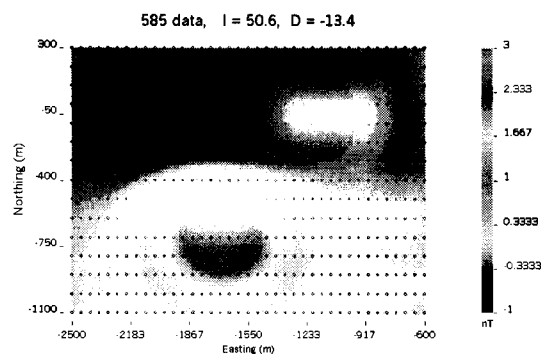
6.3.2 Synthetic downhole magnetic observations

So how can magnetic field observations be used to detect and model the keel? The observations need to contain information about the keel above the expected noise levels. As the source of the magnetic field cannot be moved nearer the receiver, the receiver is moved nearer the source; below the surface. Synthetic magnetic field observations are made below the surface in four drill-holes that extend to a depth of 650m. The drill-holes, shown on a cross-section of the susceptibility rock-model in Figure 6.8, are spaced 50m apart and are located at -1800m west, and from -350m north (against the northern edge



(a) Synthetic magnetic data produced from susceptibility rock-model.

(b) Synthetic magnetic data produced from susceptibility rock-model without keel.



(c) Difference between synthetic magnetic data with keel and synthetic magnetic data without keel.

Figure 6.7: Synthetic magnetic data produced from a susceptibility rock-model of San Nicolás. Data are produced from models with and without the keel. The difference is also plotted.

of the keel) to -200m north. Sixty-six measurements are down the hole to a depth of 650m. The measurements are made every 10m.

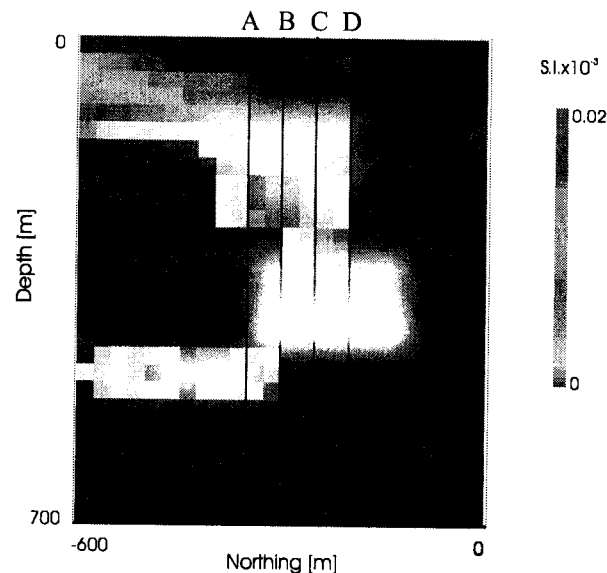
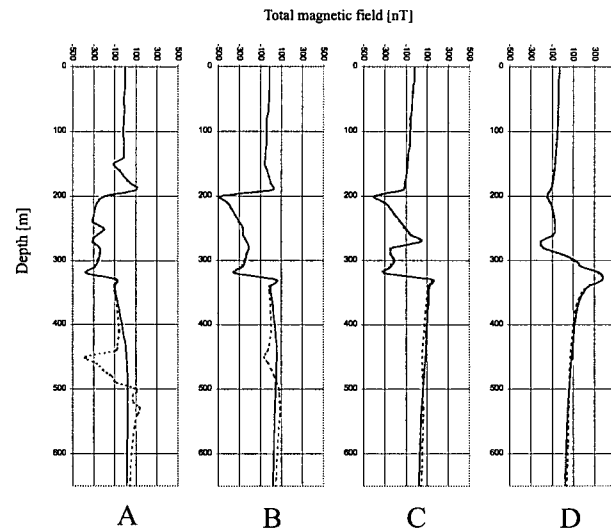
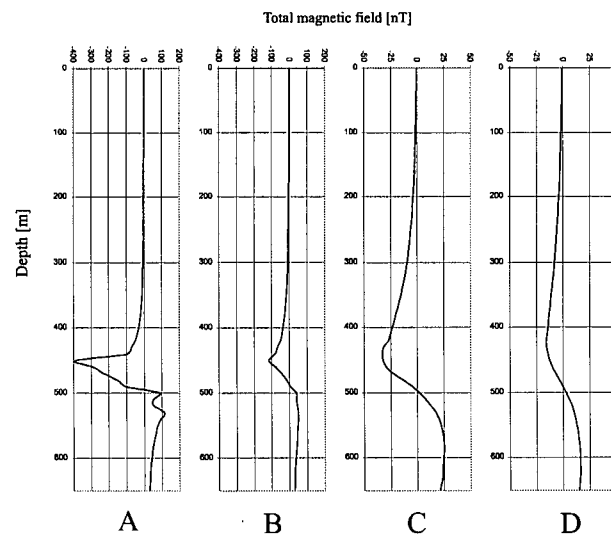


Figure 6.8: West-facing cross-section of the magnetic susceptibility rock-model with the keel shown to the lower left. Four drill-holes are also shown, within which synthetic total magnetic field measurements are performed.

The total field magnetic response for observations made down-hole on both susceptibility models are shown in Figure 6.9(a). The response from the model with the keel is shown as a dashed line, while the response from the model without the keel is shown as a solid line. The difference between the two responses are plotted in Figure 6.9(b). The signal that emanates from the keel is easily observed in data from drill-hole A at a depth of between 450 and 500m. The observations made down drill-holes at more northern locations show the signal from the keel decreasing to a magnitude of about ± 15 nT at a distance of 150m from the northern edge of the keel. According to the noise and signal-to-noise levels prescribed above, this response should be sufficient to detect the keel.



(a) Down-hole, synthetic total magnetic field data. For each hole the response from the model containing the keel is shown by a dashed line, and the response from the model without the keel is shown by a solid line.



(b) Difference between down-hole data collected on a model with the keel and those collected on a model without the keel. This is equal to the response from the keel alone. Note the change of scale for plots from drill-holes C and D.

Figure 6.9: Synthetic total-field, down-hole, magnetic data produced from a magnetic susceptibility rock-model of San Nicolás. Data are produced from models with and without the keel. The difference is also plotted. Drill-hole A is just north of the keel and drill-holes B, C, and D are each 50m further north than the previous one so that drill-hole D is 150m north of the keel.

6.3.3 Summary of magnetic forward modeling

Forward modeling of total field magnetic data over a susceptibility rock-model of San Nicolás has produced results that suggest noise levels are too high for surface data to detect the keel. This result is consistent with the inversion results that did not model the keel. Also, drill-hole data collected at a distance of 200m from the keel will still contain information about the keel that is sufficiently above assumed noise levels.

6.4 CSAMT/CSEM observations

One-dimensional inverse modeling of CSAMT data collected over San Nicolás has, when the models are 'stitched' together, successfully produced a conductivity structure of the subsurface that defines the ore-body and the Tertiary overburden. A conductivity distribution that models the keel, however, was not recovered. Through the forward modeling presented in Section 4.4.1 it has already been shown that 3D modeling is probably necessary in order to recover all the information contained in the CSAMT data, especially information about structure at depth encoded in lower frequency data (0.5 Hz to 32 Hz). But, even if this data were to be modeled in three dimensions, would it contain information about the keel that is above expected noise levels?

In order to address this question, and examine ways of increasing signal-to-noise levels of the response from the keel, forward modeling of CSAMT data with different transmitter-receiver layouts is performed over a conductivity rock-model of San Nicolás. Although the field CSAMT data were collected at fifteen frequencies ranging from 0.5 Hz to 8192Hz, forward modeling is only presented for a frequency of 16Hz. This frequency is chosen because the plane-wave skin-depth for this frequency, within a 50 Ohm-m half-space (a reasonable approximation for the conductive environment at San Nicolás), is about 900m. Therefore, signals at this frequency would probably contain sufficient

information about the keel. Of course, as the geometry of the survey changes, the plane-wave approximation will become less valid, but an investigation into the fields generated at 16Hz may give some insight into the problem of detecting the keel using the CSAMT method.

When inverting the field data, standard deviations of between 2 and 5 percent were assigned to the apparent resistivity data, and between 1 and 2 degrees were assigned to the phase. For this sensitivity study of the keel, apparent resistivity and phases are not calculated, rather the data remain in terms of magnitudes of electric and magnetic field components. For the purpose of drawing conclusions here, it is assumed that noise levels of greater than 5 percent are necessary for a coherent signal from the keel to be useful in detecting the keel.

6.4.1 Simulated CSAMT observations

In a similar manner as before, rock-models were made for conductivity structures with and without the keel. The rock-model was embedded in a half-space of 150 Ohm-m material and a perspective view of the conductivity rock-model (with keel), cut at -400m south to reveal the deposit, is shown in Figure 6.10. The location of the transmitter (grounded at each end) as used in the field CSAMT survey is also shown. Receiver locations were centered over the deposit along lines -600m, -400m, and -200m north. Using *EH3D* (Appendix E.8), electric and magnetic fields were modeled for the conductivity models due to a current source of 1 Amp, oscillating at 16Hz. As with the field survey, real and imaginary parts of the x-component electric field (E_x) and y-component magnetic field (H_y) were measured (an orthogonal coordinate system is orientated with the local coordinate system). These measurements were made at 264 stations located at the surface on a regular, 50m x 50m grid between -1000m and -2150m east, and -150m and -650m north.

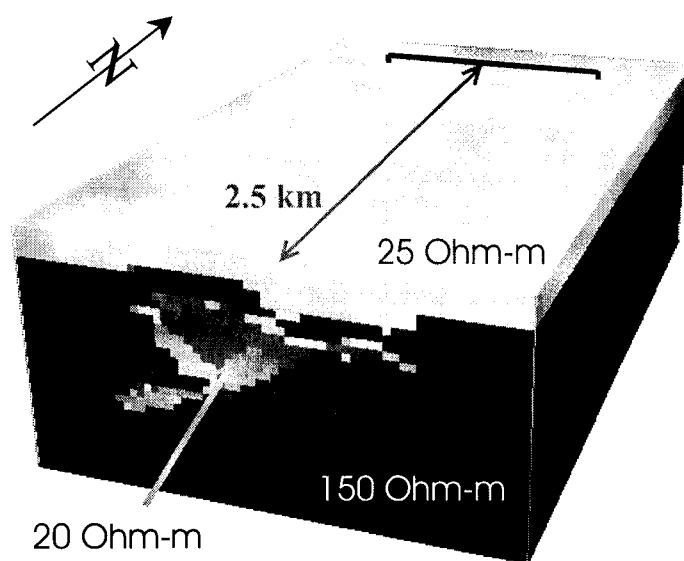


Figure 6.10: A perspective view of conductivity rock-model (including keel), cut to show the conductivity structure at -400m south. The approximate resistivity values of the model are shown, as is the location of the transmitter that was used in the field CSAMT survey.

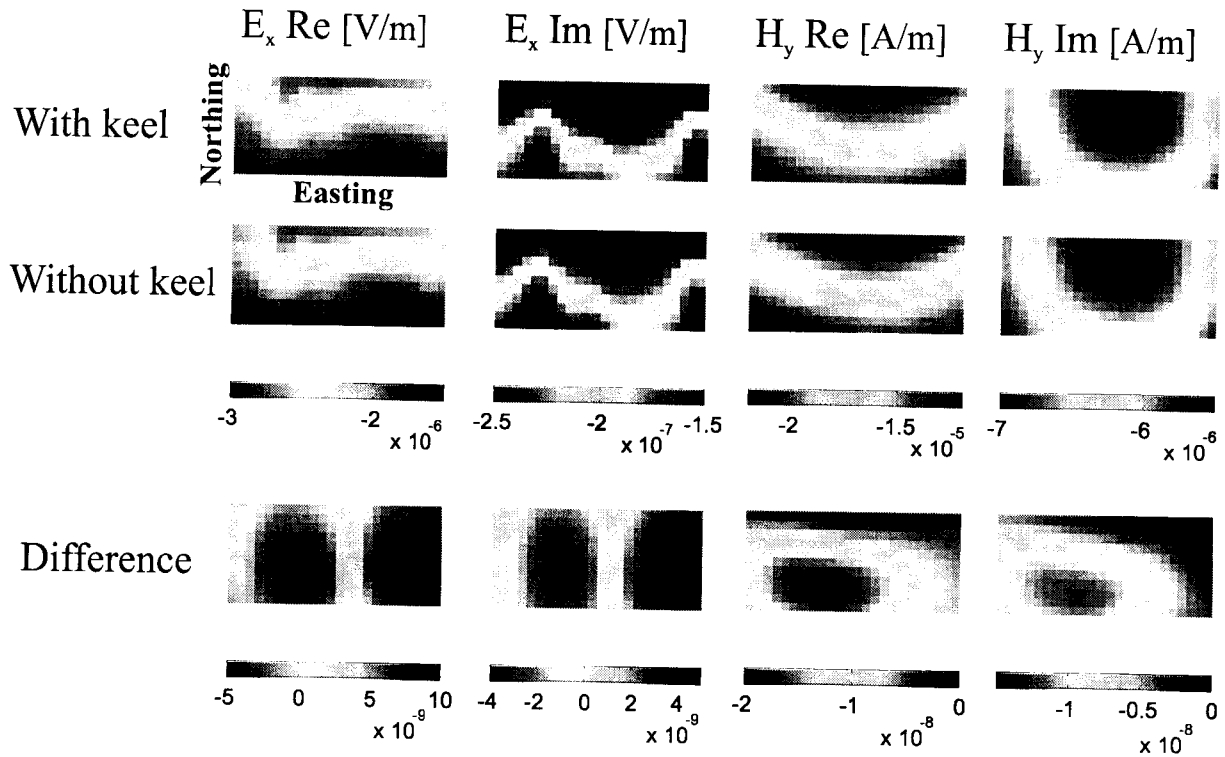


Figure 6.11: Color plot of synthetic real and imaginary E_x and H_y fields, at the surface over the San Nicolás deposit, that emanate from a far, 16Hz, grounded transmitter. The first and second rows show the field components that would be produced over conductivity models with, and without the keel, respectively (these are plotted on the same scales). The third row of plots show the difference between fields produced with and without the keel for each component. Each plot covers the area over the deposit, as described previously, although the northing and easting values are not shown.

Plan-view, color plots of real (Re) and imaginary (Im) parts of E_x and H_y , for all stations, are shown in Figure 6.11. The first row of plots shows the fields that would be produced over a conductivity structure of the deposit that includes the keel, the second row of plots shows those that would be produced if the keel were absent, and the third row of plots displays the difference. The difference the keel makes to the real part of the E_x field ranges from -5×10^{-9} to 10×10^{-9} . When compared to the actual measured values, we can see this is less than one percent of the total signal. This signal level will be most likely saturated by noise, and therefore not useful for keel detection. Similar observations are made for the remaining field component measurements, where the effect of the keel is about two percent of the imaginary part of the E_x field, and less than one percent of both the real and imaginary parts of the H_y component.

This forward modeling suggests that, for a grounded current source oscillating at 16Hz, noise levels much smaller than five percent would be needed for the keel to be modeled even if the information contained in the E_x and H_y field components can be extracted through 3D inverse modeling.

The analysis for other frequencies, or for other E and H components has not been carried out, so it is difficult to draw general conclusions. Furthermore, only relative signal magnitudes and noise levels have been considered, however, absolute signal-to-noise levels are also important — if more current were to enter the ground the magnitudes of the electric and magnetic fields would be larger and, assuming that noise levels remain the same, and a coherent response from the keel above the noise levels could be measured.

It will probably be difficult to detect the keel using a plane wave source, that is a transmitter a large distance away from the target. Greater signal is needed for detection and this can come about through a change in survey design. The next section looks at bringing the source in closer to the target in order to increase the signal.

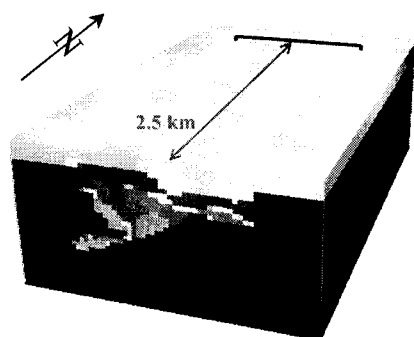
6.4.2 CSAMT/CSEM survey design

Another way to increase the signal from the keel would be to change the design of the survey. In contrast to potential field methods, whose sources are fixed, active-source methods, such as CSAMT, allow for the repositioning of sources as well as receivers⁴. In addition, many more components of the resulting fields can be measured, including real and imaginary parts of x-, y-, and z-components of electric and magnetic fields, and different source frequencies can also be employed. Accordingly, there are many more variables to consider when determining survey design, and optimization methods can be used to design the appropriate surveys given a specific target. Glenn and Ward (1976), Boerner and West (1989), and Maurer and Boerner (1998) address optimization of electromagnetic survey designs. For this study, however, only simple survey-design changes are considered, for a single frequency of 16Hz, in an attempt to increase the response of the keel by increasing the current-density within the keel⁵. The forward modeling is performed in order to demonstrate the types of surveys that could be conducted in order to detect the keel, rather than conclusively show the best survey design for the job.

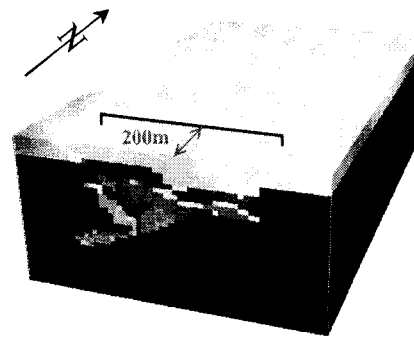
By bringing the current source, or transmitter, closer to the deposit, it is expected that more current will enter the deposit, and therefore larger signals can be measured. Consequently, the signal from the keel should also be larger and the absolute response from the keel may become higher than the noise levels. Figure 6.12 shows the four transmitter locations that are used for the forward modeling. The characteristics of each transmitter setup is described below.:

⁴When sources are used that generate primary electric and magnetic fields that are not planar in nature at the receiver locations, the survey is considered to be a controlled source electromagnetic survey (CSEM), and no longer a controlled source audio magneto-telluric survey (CSAMT).

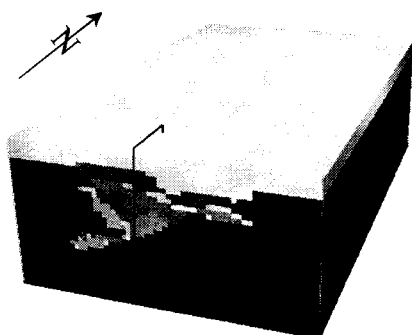
⁵An increase in current density within the keel is expected to result in larger magnitude secondary fields emanating from the keel.



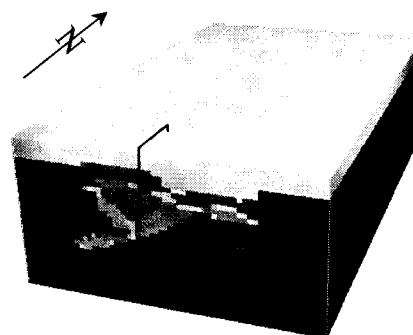
(a) *Far* location of transmitter; as used in the field survey.



(b) *Near* location of transmitter.



(c) Transmitter with one end below the deposit down a drill-hole.



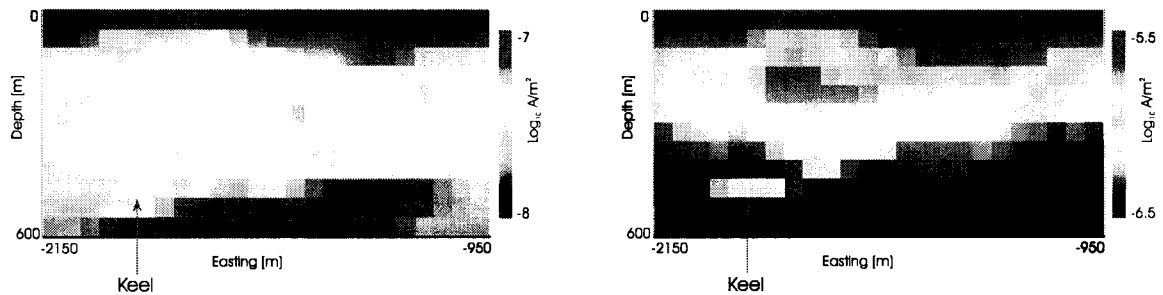
(d) Transmitter located with one end within the deposit down a drill-hole.

Figure 6.12: Perspective views of conductivity rock-model showing locations of the transmitters used for forward modeling of electromagnetic data.

1. Figure 6.12(a) shows the *far* grounded transmitter location that was used in the field survey. This is 1.6 km long, parallel to the perceived geologic strike, and is 2.5 km north of the center of the deposit (-400m N).
2. For the near transmitter location, the source remains grounded on the surface and is brought closer to the deposit (Figure 6.12(b)). The transmitter is now located along line -200m N (very close to the deposit), and is still 1.6 km long.
3. Figure 6.12(c) shows a very different transmitter location. One end of the transmitter is located at the surface just north of the deposit (-1600m E, -200m N) and one end is located in a drill-hole below the deposit at a depth of 525m at -1600m E, -400m N.
4. The final transmitter location is shown in 6.12(d). This is also located with one electrode at the surface and one down a drill-hole, but now the electrode is grounded within the deposit at a depth of 325m.

A plot of current density magnitudes for models cells on a north-facing cross-section through the deposit is shown in Figure 6.13. An increase of current densities in the keel of more than one order of magnitude results from decreasing the distance to the transmitter. The current density within the keel increases in magnitude to over 1×10^{-5} A/m² for the down-hole source locations (not shown), which is an increase of more than two orders of magnitude over the original field survey.

The new transmitter locations have increased the currents within the keel as well as the rest of the deposit. Synthetic measurements of E_x and H_y are now made at the surface for all three transmitter locations on conductivity structures with, and without, the keel. The following figures show the measurements made with these transmitter locations and compare them with the results of the simulated field observations (Section 6.4.1).



(a) Magnitude of current density through a north-facing cross-section of the rock-model at -400m north for the *far* transmitter location (2500m north of the deposit) used in the field survey.

(b) Magnitude of current density through a north-facing cross-section of the rock-model at -400m north for the *near* transmitter (at the northern edge of the deposit) location.

Figure 6.13: North facing cross-sections through a rock-model, which includes the keel, showing the magnitude of current density in each cell for two, (a) far and (b) near, transmitter locations. The cross-section is located at -400m north and shows an increase in current density within the keel (the small anomaly to the lower left of each section) from about 5×10^{-8} to 6×10^{-7} A/m².

Four figures are presented; one for each part (real and imaginary) of the E_x and H_y components. Within each figure there are four columns of plots; each column represents one transmitter location. Each column of plots contain synthetic readings calculated for: (1) a conductivity structure with the keel, (2) a conductivity structure without the keel, and (3) the difference between the two. Each plot is for receivers at the surface and covers an area over the deposit which extends from -2150m to -1000m east, and -150m to -650m north (the northing and eastings are not shown). The plots shown for observations with all of the transmitter locations, with and without the keel, are presented at the same scale (given by a color bar to the left), while the difference plots, for all transmitter locations, are presented at a different color-scale.

As shown in Figure 6.14, the real part of the E_x component of the electric field measured at the surface increases by two orders of magnitude when the transmitter

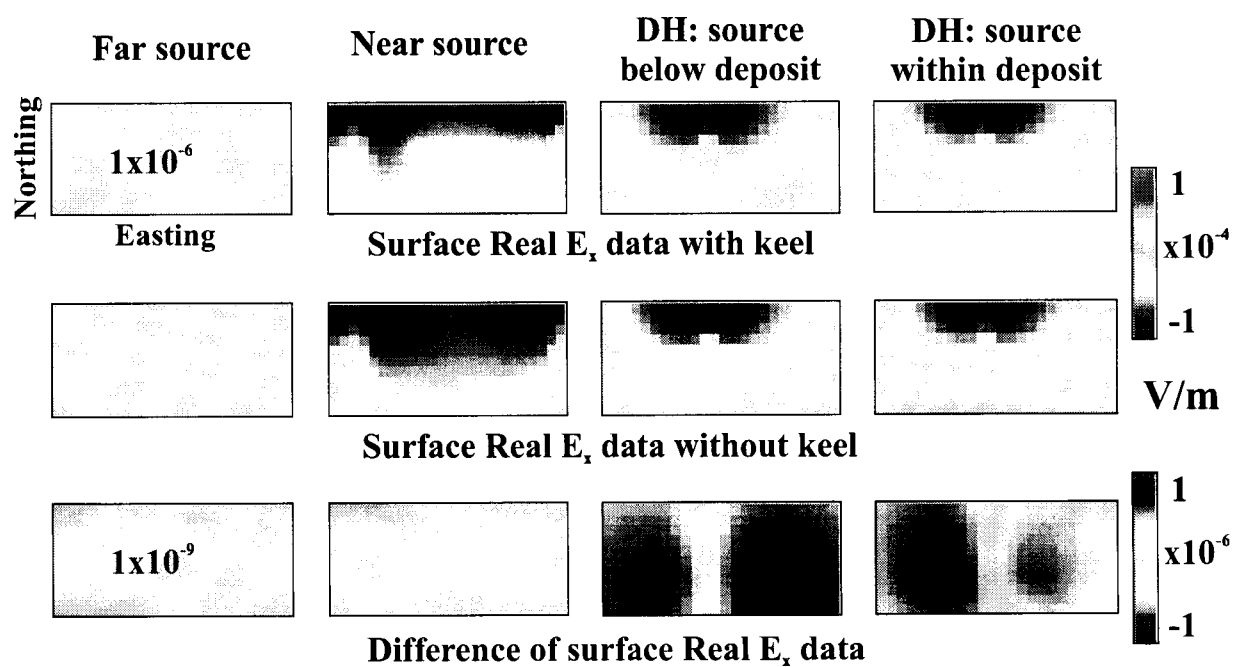


Figure 6.14: Real part of E_x surface measurements made for four different transmitter source locations over conductivity rock-models of the San Nicolás deposit with and without the keel.

location is brought close to the deposit. The difference, between measurements made on conductivity structures with and without the keel, increases by just under two orders of magnitude when the transmitter is brought near to the deposit, whereas it increases by nearly three orders of magnitude when one end of the transmitter is placed near the keel, down a drill-hole. Also, the response from the keel seems larger when the transmitter is placed below the deposit, as opposed to within the deposit.

A similar result is shown for the imaginary part of the x-component of the electric field at surface receiver locations. While the observed signal increases by about two orders of magnitude when the distance to the transmitter is decreased, the difference between measurements made with and measurements made without the keel is only significant when one end of the transmitter is located below the surface.

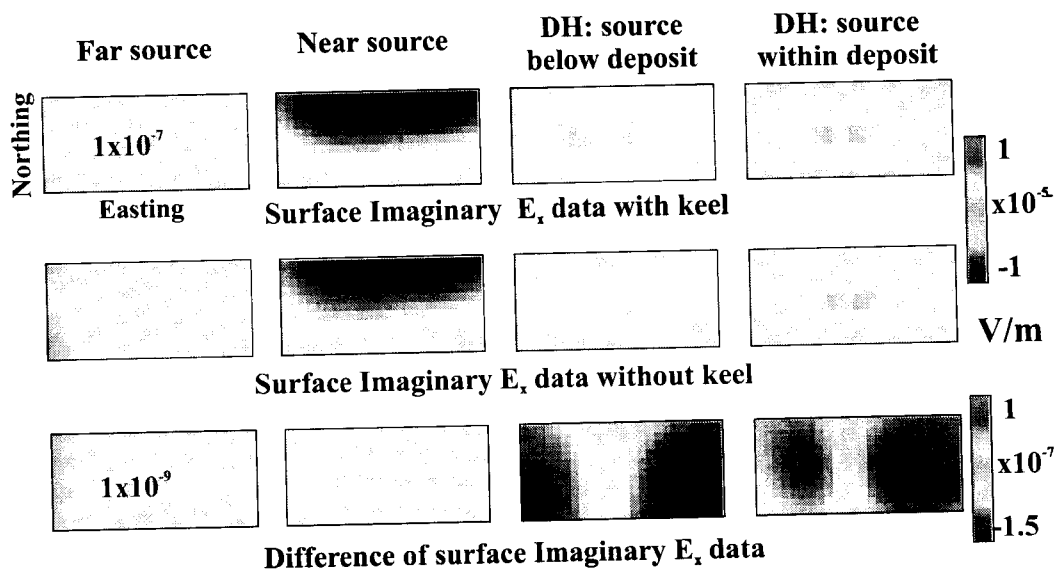


Figure 6.15: Imaginary part of E_x surface measurements made for four different transmitter source locations over conductivity rock-models of the San Nicolás deposit with and without the keel.

Very similar results are also observed for surface measurements of the magnetic field,

and are summarized below. For the three transmitter locations close to the deposit, the real part of the H_y field increases by about two orders of magnitude, while the imaginary part increases by just under two orders of magnitude. The difference between H_y measurements made on conductivity rock-models with and without the keel increases by two orders of magnitude only when one end of the transmitter is grounded close to the keel, down a drill-hole.

6.4.3 Summary of CSAMT/CSEM forward modeling

Forward modeling has shown that current densities, and the resulting electromagnetic fields, all increase when the transmitter is brought closer to the deposit. The response of the keel relative to the total signal observed also increases, but does not exceed one percent of the total signal. However, with these elevated field strengths, the noise levels may remain close to their original level, and the signal-to-noise could be improved.

The results presented are for two field components (E_x and H_y) at one frequency (16 Hz). These components are measured in CSAMT surveys because they have maximum coupling with the propagating plane-wave source with the transmitter geometry used, and therefore will produce the strongest signals. As other survey geometries are considered, the source does not approximate a plane wave, and the orientation of maximum signal is harder to determine. Therefore, other field components should be measured so all information is recorded. In addition, the aspect of placing the receiver locations down a drill-hole has not been presented. This would probably result in an increase in signal-to-noise response of the keel depending on the proximity of the drill-hole to the keel. Finally, the response of the keel at different frequencies might also show different results.

The preliminary study made on the design of CSAMT surveys, targeted at detecting the keel, clearly demonstrates the scope of this problem. As such, definite conclusions cannot be made regarding the detectability and definition of the keel with CSAMT

techniques, apart from stating that the range of survey designs available probably includes one that could provide data to sufficiently model the keel.

6.5 DC resistivity

Inversion of DC (direct-current) resistivity measurements, made using the 'Realsection' array configuration, has not been successful in identifying the San Nicolás deposit. This may be due to current channeling in the conductive overburden, or the small number of transmitter locations (more transmitter locations seem to have a greater constraint on conductivity distributions at depth). Even so, forward modeling is performed in order to quantify how sensitive the survey is to the keel, and to determine if other survey designs might provide more information about the keel.

In order to determine if the signal from the keel is above noise levels, the noise levels must be known, however, only estimates of the noise are available. Noise levels of five percent were assigned to the DC data for inverting, and, although the recovered models didn't seem very successful, noise levels between two and eight percent typical for such surveys. For a better understanding of the noise levels, instrument sensitivity and specific field conditions would have to be taken into consideration.

6.5.1 Simulation of Realsection resistivity observations

Two conductivity models are used for the forward modeling; one contains the structure of the keel and one does not. A synthetic Realsection⁶ survey was conducted over both of these models, the results of which are presented as apparent resistivity values in Figure 6.19. The survey, which has a layout as described in Appendix C.6, was conducted

⁶*Realsection* is the name given to a type of modified gradient array survey that was devised by Quantec Geoscience, and employs several transmitter dipoles of different lengths. When plotted on a section, the resulting picture merely represents the data and not the true electrical structure of the earth.

over the deposit on line -400m north and involves measuring the potential difference between two receiver electrodes. This potential difference is the actual measurement made in the field, but apparent resistivities are commonly calculated using the following equation, [Telford, Geldart, and Sheriff, 1990b].

$$\rho_a = \frac{2\pi V}{Ik}, \quad (6.1)$$

where k is a geometric factor that is determined by the position of the transmitter and receiver electrodes, V is the potential difference measured with the receiver dipoles, and I is the transmitted current.

The synthetic Realsection data are plotted in Figure 6.16 and show a response from the keel that is about one percent of the total signal. This amplitude is below the expected noise levels (5 %).

Another survey design is employed to see if the response from the keel can be increased. The dipole-dipole survey design is chosen. Dipole-dipole arrays are usually more sensitive to lateral variations in conductivity, and, although the depth penetration may not be as great as the Realsection survey, it is hoped that a greater sensitivity to the keel is observed. Dipole-dipole arrays also employ many more transmitter locations and therefore can provide better constraint of inversion models at depth.

6.5.2 Forward modeling of dipole-dipole resistivity observations.

A dipole-dipole array was used to collect synthetic resistivity data over two rock-models of the San Nicolás deposit; one that includes the keel, and one that does not. The dipole-dipole array employed has thirteen 100m transmitter dipole locations, each with five 50m, and five 100m, associated receiver dipoles. For display purposes, resistivity values are plotted at the intersection of two forty-five degree lines that lead down from the center

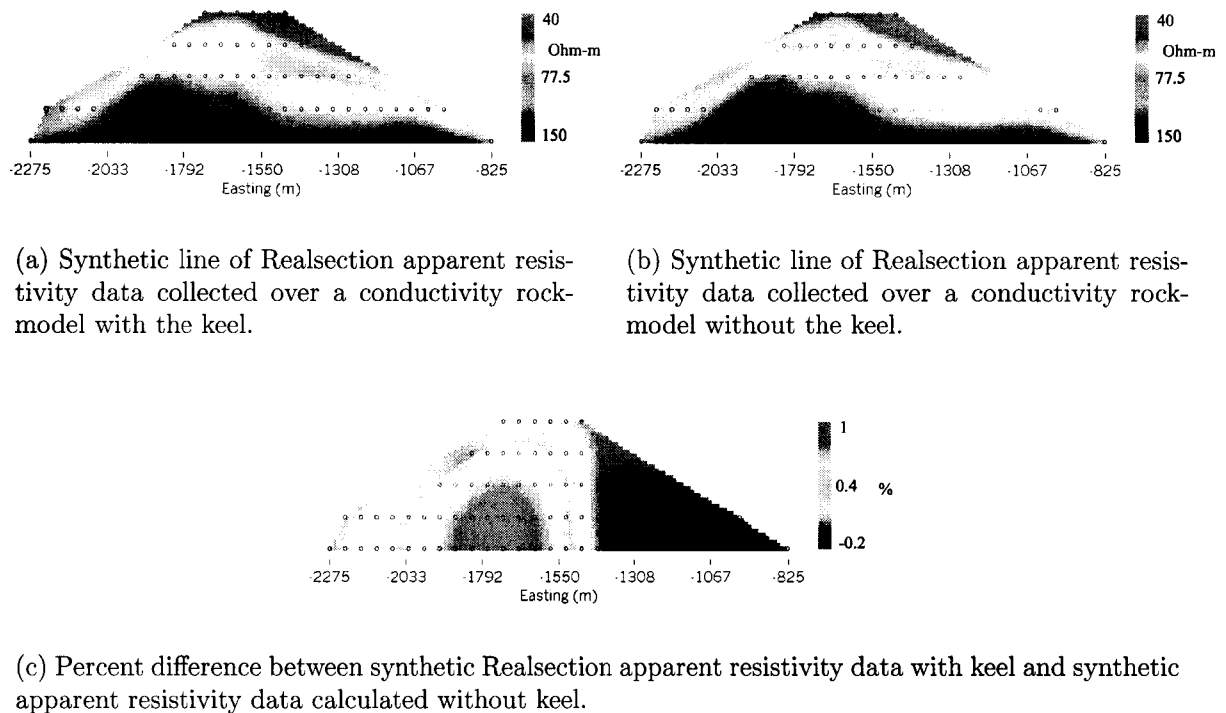


Figure 6.16: Synthetic, DC, apparent resistivity, Realsection data produced from conductivity rock-models of San Nicolás. The x-locations refer to the midpoint of the receiver dipoles, while the y-axis refers to different transmitter dipole positions; the bottom line corresponds to data collected with the longest transmitter, and the top line corresponds to data collected with the smallest transmitter length. The section is just a representation of the data and should not be taken as the true structure of the earth. The percent difference between data produced from models with, and models without the keel is also shown. Note: the apparent resistivity color-bar is on a log scale with low resistivities (high conductivities) displayed in red, while the percent difference is plotted on a linear scale.

of the transmitter (I) and receiver (v_x) dipoles, as shown in Figure 6.17. This *pseudo-section* is a convenient way of plotting the data but does not represent the structure of the earth. The values are interpolated and displayed in Figure 6.18.

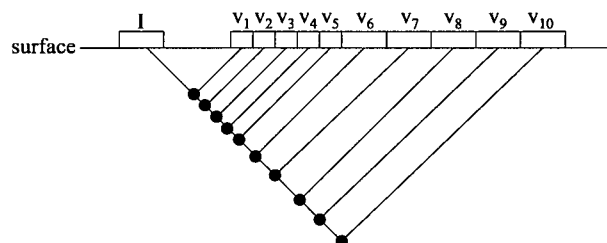
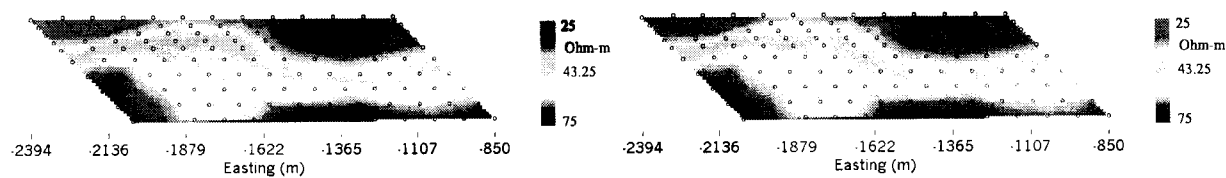


Figure 6.17: Schematic diagram of dipole-dipole array for one transmitter location, showing the pseudo-section plotting points relative to the electrode geometry.

The synthetic dipole-dipole data collected using the two conductivity rock-models show apparent resistivity values that range from 25 to 75 Ohm-m. The signal from the keel, as determined by calculating the difference between the data collected over the two models, has resistivity contrast values ranging from -2 to 0.3 Ohm-m. This response is up to 2.75 % of the magnitude of the total signal, which suggests that this dipole-dipole array configuration is more sensitive to the keel than the Realsection array, however, it is still below the assumed noise level of 5%.

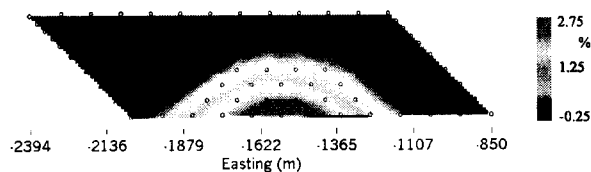
6.5.3 Summary of DC resistivity forward modeling

While inverse modeling of DC Realsection resistivity data hasn't resulted in an appropriate conductivity model of San Nicolás, the data have been shown to be sensitive to the presence of a conductive keel. Unfortunately, the response from the keel is below expected noise values and so other survey designs are sought in order to increase the signal-to-noise.



(a) Pseudo-section of synthetic, dipole-dipole, resistivity data collected over a rock-model that includes the keel.

(b) Pseudo-section of synthetic, dipole-dipole, resistivity data collected over a rock-model without the keel.



(c) Percent difference between synthetic dipole-dipole resistivity data with keel and synthetic dipole-dipole data without keel.

Figure 6.18: Synthetic apparent resistivity dipole-dipole pseudo-sections produced from conductivity rock-models of San Nicolás. Pseudo-sections of data produced over conductivity rock models with and without the keel are plotted on a log scale along with the percent difference between the two, which is plotted on a linear scale. The plotting convention use is shown in Figure 6.17.

A traditional dipole-dipole survey design was chosen as an improvement over the Realsection array in detecting the keel. In changing from a Realsection array to a dipole-dipole array, the response from the keel has been shown to increase by almost three-fold (from 1% of the total signal to 2.75%). Furthermore, if this data were to be inverted it is expected that the dipole-dipole data would constrain the model better at depth than the Realsection data. Unfortunately, the response from the keel is still below estimated noise levels so other survey designs are required, and down-hole methods may prove necessary in order to receive a coherent signal from the keel.

The increase in response from the keel demonstrates an improvement in survey design, but comes at the expense of more field work. Five transmitter and eighty-five receiver dipoles were used in the Realsection survey, versus thirteen transmitter and one hundred and thirty receiver dipoles for the dipole-dipole survey.

Optimization of DC resistivity surveys for a given target is a very broad subject due to the many different transmitter-receiver geometries that can be employed. The work presented in this section introduces the subject of DC resistivity survey design with two-dimensional surveys over a three-dimensional object, but investigations into three-dimensional designs would provide more complete results.

6.6 Chargeability

Realsection IP data collected at San Nicolás shown in Figure D.15, show apparent chargeability values up to 24 msec. Physical property measurements suggest the deposit as a likely source of the signal, and inversion of this data, along with gradient array IP data, has shown a distribution of anomalously chargeable material that corresponds well with the deposit (Section 4.5). Although the keel was not modeled through inversion of the IP data, a concentration of chargeable material is observed along the fault-zone that bounds

the deposit to the east. It is possible that these chargeability values persist within the keel and IP surveys could be used to detect the keel. As a result, forward modeling of IP data is used to determine the sensitivity of IP data to the presence of the keel.

Synthetic IP data are calculated using DCIPF3D — a forward modeling routine that first calculates the primary voltages over a conductivity distribution, and then calculates the secondary voltages based on the chargeability of the material. An overview of the forward modeling method is given in Appendix C.6.

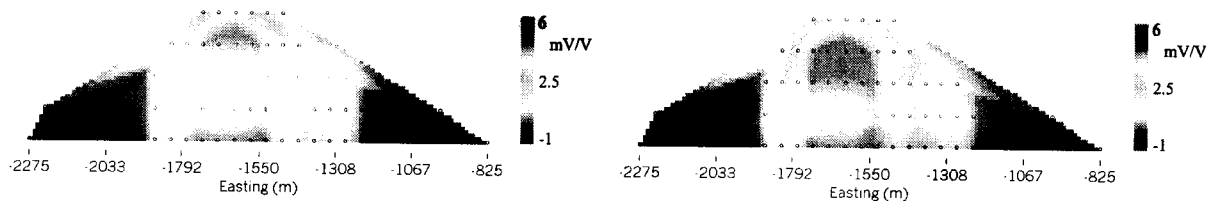
A chargeability rock-model is created from digitizing geologic cross-sections and assigning values to the rock units based on physical property values. A maximum value of seventy-five milli-Volts per Volt⁷ is assigned to the massive-sulphide. As before, synthetic observations are made over two chargeability models; one with the keel and one without. The difference between the two sets of observations will indicate the sensitivity of the survey technique to the presence of the keel. Two survey designs are examined for sensitivity to the keel: the Realsection survey design that was used in the field, and the dipole-dipole survey design that was used in Section 6.5.2.

The success of a survey in providing information about the keel depends on the strength of the signal from the keel and the noise levels encountered during the survey. Standard deviations of about 3 percent of the apparent chargeability plus a base level of 1 mV/V were assumed for the inversion of apparent chargeability data, and resulted in a reasonable model with some structure, but without too much noise. Therefore, a signal from the keel would have to be above this level in order for it to add information to the inversion model.

⁷Here, the apparent chargeability is calculated by taking the ratio of secondary voltages to primary voltages.

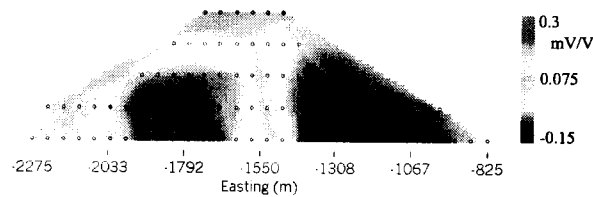
6.6.1 Simulating Realsection IP field observations.

Synthetic apparent chargeability data are first calculated for the survey layout that was used in the field.



(a) Synthetic line of Realsection apparent chargeability data collected over chargeability rock-model with the keel.

(b) Synthetic line of Realsection apparent chargeability data collected over chargeability rock-model without the keel.



(c) Difference between synthetic realsection apparent chargeability data from a chargeability rock-model including the keel and synthetic apparent chargeability data from a chargeability rock-model without the keel.

Figure 6.19: Synthetic apparent chargeability Realsection data produced from chargeability rock-models of San Nicolás. The x-locations refer to the midpoint of the receiving dipoles, while the y-axis refers to different transmitter dipole positions; the bottom line of data being collected with the longest transmitter, and the top line of data being collected with the smallest transmitter length. The difference between the two data sets is also plotted.

The synthetic Real-section array apparent chargeability data are displayed in Figure 6.19. The difference between data collected on models with and without the keel range between -0.15 and 0.3 mV/V. The difference is below the prescribed noise level and so, if the noise levels are appropriate, this survey design would not be useful in identifying the keel.

In an attempt to increase the chargeability signal from the keel, a dipole-dipole array configuration is examined using forward modeling methods. The results from this study are presented below.

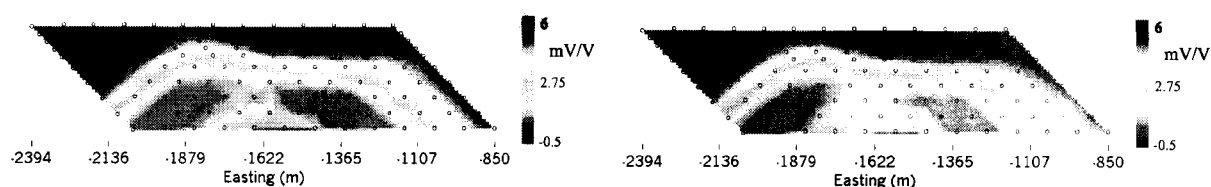
6.6.2 Synthetic dipole-dipole chargeability observations.

Synthetic dipole-dipole IP surveys, conducted on two chargeability models (one with the keel and one without), employ the same array configuration used in the DC resistivity survey described in Section 6.5.2. The data, displayed as apparent chargeabilities in Figure 6.20, show values that range from -0.5 to 6 mV/V. The response of the keel is shown in Figure 6.20(c) as the difference between data calculated using a chargeability model that includes the keel, and one that does not. This difference has values that range from -0.5 to 1.1 mV/V. These data do not show a response from the keel that is above the noise level, however, they are larger than the response from the keel that was calculated for the Realsection IP data.

6.6.3 Summary of chargeability forward modeling

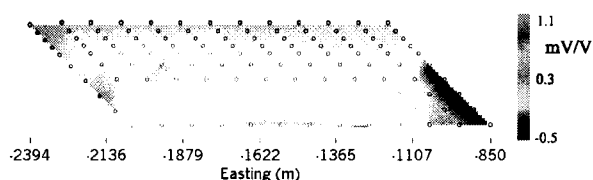
Forward modeling of apparent chargeability data over a chargeability rock-model, which is thought to represent that present at San Nicolás, shows that it is unlikely Realsection IP data contain information about the keel that is above the expected noise levels. Therefore, insight into the presence of the keel is not expected from inversion of this data.

Forward modeling of dipole-dipole chargeability data shows that an increase in response from the keel can be achieved, however, the amplitude is still below the expected noise levels. If the chargeability values have been appropriately assigned to the rock-model and the expected noise levels are also approximately correct, then detecting the keel using IP methods may have to involve reducing the distance from the keel to either



(a) Pseudo-section of synthetic dipole-dipole apparent chargeability data collected over a chargeability rock-model with the keel.

(b) Pseudo-section of synthetic dipole-dipole apparent chargeability data collected over a chargeability rock-model without the keel.



(c) Difference between synthetic dipole-dipole apparent chargeability data from a chargeability rock-model including the keel and synthetic apparent chargeability data from a chargeability rock-model without the keel.

Figure 6.20: Synthetic apparent chargeability dipole-dipole data produced from chargeability rock-models of San Nicolás. The plotting convention used is shown in Figure 6.17. The difference between the two data sets is also plotted.

the transmitter, the receivers, or both. This could be achieved by placing electrodes (transmitter or receiver electrodes) down a drill-hole.

It should also be noted that the secondary voltages measured in the dipole-dipole survey are about one tenth the magnitude of those measure in the Realsection survey. This could have implications regarding the signal-to-noise of the response of the keel as lower magnitude observations may be closer to ambient noise and instrument sensitivity levels.

The examination of two survey layouts hardly constitutes a comprehensive analysis of how to detect the keel using IP methods, however, it has been shown that different types of surveys have different sensitivity to the keel and, should the rock-model values be correct, down-hole methods may be necessary.

6.7 Conclusions

Forward modeling, using physical property rock-models at the deposit scale, has shown that all of the data collected at San Nicolás could be sensitive to the presence of the keel, yet only the gravity data would show a response that is significantly above the noise level. This is consistent with the inversion results presented in Chapter 4, in which the gravity data produced a density distribution of the deposit that had an extension in the vicinity of the keel, whereas the inversion of other data did not.

The surveys employed in the field were designed to detect anomalous structure and bodies of unknown location and geometry, and they were successful in finding San Nicolás. However, changes to the survey design are needed when deep, small ore-bodies such as the keel are the target. To this end, different survey designs have been examined in an attempt to increase the observable response of the keel. Insight has been provided about what data should be measured in order to see the keel.

The target of the investigation in a potential-field survey is also the source of the signal. For this reason there is a limit to the signal from the keel that can be measured at the surface. Consequently, down-hole methods may be needed. On the other hand, changing transmitter-receiver geometries on the surface can improve the magnitude of the response from the keel, when active-source methods are used. However, down-hole methods also prove useful for active source measurements.

The use of drill-holes in geophysical surveys at San Nicolás is a logical progression in the search for deeper ore as many drill-hole already exist. This ties in with the over-all detailed exploration work that has been the motivation for this chapter. Information from drill-holes can be used to constrain inversion models in the form of a priori information (Chapter 5), or geophysical data can be collected down-hole. I have shown how both of these methods can increase the level of understanding of the subsurface, and therefore, increase the efficiency of a deposit-scale exploration program.

The issue of increasing the signal-to-noise of the response from the keel has been addressed by assuming noise levels and trying to increase the signal based on survey design. For a complete study in this area the noise should also be investigated. Different sources of noise, and how absolute noise levels, and noise levels relative to the signal, change in response to different survey designs could also be addressed.

As well as considering different survey designs of the methods used to collect data at San Nicolás, other methods may be more successful in defining the keel, and investigation into such methods as MMR (magneto-metric resistivity) could prove fruitful.

Chapter 7

Conclusions and Suggestions for Further Work

The work completed at San Nicolás has successfully demonstrated how geophysical inversion can be used at various stages of an integrated exploration program. This work has been presented within the context of an exploration program that is summarized by Figure 7.1.

Geologic and geophysical exploration models were formulated and provided direction for the subsequent work at different stages. Regional-scale inverse modeling allowed conclusions to be made about physical property distributions of the earth. These models were given a geologic context through interpretation using physical property measurements. The results provide important information for finding mineralization, and could have been used to direct follow-up work.

Local-scale, forward and inverse modeling, along with physical property analysis, has shown how geophysics can be used to define mineralization well enough to successfully spot drill-holes. This leads to further iterations at the local scale that include using drill-hole information directly in an inversion, and modifying surveys in order to help delineation.

The work completed has demonstrated how geophysical inversion can be used within general exploration programs by providing information specific to San Nicolás.

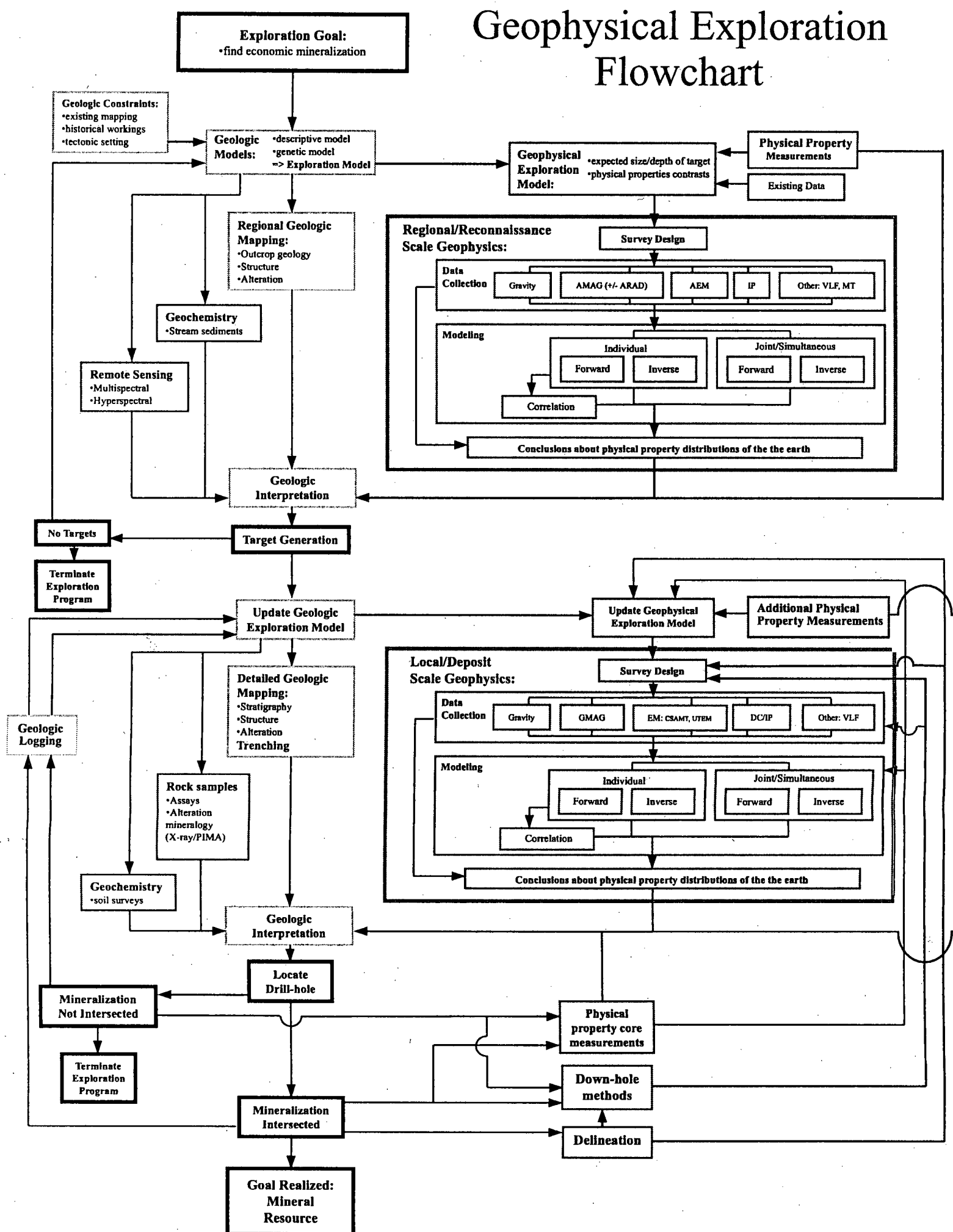


Figure 7.1: Geophysical Exploration Methodology Flowchart.

7.1 Summary of results

Regional-scale geophysical programs can cover a large amount of land in a small amount of time. The information that they contain can often be more easily interpreted using geophysical inversion. Individually, the inversions can produce a number of targets for selection. In combination, the number of follow-up targets can be greatly reduced, thus focusing an exploration program faster than before.

At San Nicolás, regional-scale modeling has provided information about the extents of a horst, within which the deposit is contained. The location and geometry of deep intrusives has also been modeled. These intrusives might have been the source of mineral-rich hydrothermal fluids and are therefore important for future exploration efforts. In addition, faults have been defined that cross-cut the horst and contain tertiary volcanics. Discriminating between these faults and older faults within the horst may be useful in finding structures that control mineralization. Shallow dense and susceptible mafic intrusives have also been identified.

Combining the results from several physical property distribution models has significantly reduced the number of targets. Apart from the deposit, other areas of mineralization were also detected.

The potential of airborne EM methods to detect the deposit has also been investigated. With both the frequency-domain and time-domain methods only information regarding the upper 200m of the earth has been recovered. Forward modeling of the time-domain data, however, suggest that deeper conductivity information might be recoverable.

Local-scale inversions can provide detailed information about the subsurface. This can be used to guide drilling programs. Furthermore, feedback from drilling programs (in the form of physical property measurements on the core, or down-hole surveys) can be used

to improve physical property inversion models and thus increase the success of drilling. The use of drill-core information in such an iterative manner has been demonstrated.

Deposit-scale modeling at San Nicolás has produced physical property models of the subsurface that correspond well with the extents of the deposit. The results confirm the ability of geophysics to find the concealed deposit. The models produced also have geometric features that emulate those of the deposit. Details such as the extension of susceptibility material to the north and the concentration of chargeable material along the fault are also seen at this scale.

Including drill-hole information within inversions results in improved models being constructed. The physical property range and the extent of the deposit are much better constrained with the inclusion of one or more drill-holes. Physical property distributions away from the drill-holes also change, especially below the deposit. This allows the possibility of more sulphide below the deposit to be all but discounted.

The inclusion of drill-hole information in an inversion addresses the discrepancy between physical property measurements made on small drill-core samples and the physical property values recovered in inversion models. This discrepancy comes from inability of a large model cell to adequately represent fine-scale structure. Furthermore, the types of models being constructed have to be consistent with known information, that is smooth models should not be expected when sharp discontinuities have already been found. These issues have been addressed and possible solutions have been put forward, however much more understanding is required.

The use of forward modeling to evaluate the sensitivity of survey designs has also been demonstrated. The availability of forward modeling routines allows for quantitative evaluation of survey designs that can lead to more appropriate surveys being conducted in the future. While not addressing the topic of optimizing survey designs, ideas with which survey design can proceed have been introduced.

Forward modeling has shown that most of the geophysical survey designs employed at San Nicolás were probably inadequate to detect the keel. Simple survey design changes are suggested that, when evaluated using forward modeling techniques, show an increased sensitivity to the keel.

7.2 Further work

The work completed at San Nicolás has successfully demonstrated how geophysical inversion can be used at various stages of an integrated exploration program. In doing so, many issues have been raised that have not been conclusively resolved. Some of these issues form the basis of suggestions for further work. These are presented below:

1. The chargeable nature of the San Nicolás deposit and the forward modeling analysis that was performed, both suggest that information about the chargeability of the earth could be contained in the *Questem*, time-domain, airborne, electromagnetic data. If such information could be recovered from the data this would be prove valuable for two reasons. Firstly, the ability to search for chargeable bodies from the air could revolutionize regional exploration programs. Secondly, accounting for the chargeable material will allow for more accurate, and possibly more deeply resolved, conductivity models to be generated. So how should these data be inverted? Cole-Cole parameters along with a conductivity value could be recovered for each layer of a one-dimensional model. However, this would dramatically increase the number of unknowns in the problem and non-unique combinations of the Cole-Cole parameters may be able to produce the same data. Additional information may be required to recover unique solutions. The data could provide additional information if all three components of the magnetic field were to be inverted simultaneously. Furthermore, three-dimensional inversion of these data could provide more information and result

in much more easily interpreted models. Some of these issues are already being addressed by the UBC-GIF group.

2. More accurate modeling of the deposit by inverting surface CSAMT data might have resulted in less-drilling needed for the mineralization at San Nicolás to be defined. In order to define the deposit more accurately, three-dimensional inversion of CSAMT might prove useful. Forward modeling of CSAMT data suggested that only half of the recovered model accurately reflects the information encoded in the data, so a full 3D inversion might better define the deposit.
3. For external information to be used to maximum benefit within an inversion, methods of representing fine-scale physical property features in a large model must be developed. But, how can this information be used to constrain an inversion? Is bounding the range of values a model cell can assume the best method, and how much detail should be accounted for? These are questions that certainly need much more investigation before we can produce models that are truly consistent with the earth.
4. The subject of survey design has only been briefly addressed in this thesis. The advent of widespread inversion modeling allows for data to be collected that don't have to provide information at face-value. We can start to depend in inversions to decode the desired information about the subsurface. Optimization techniques could be used to design surveys specific to the task-at-hand that will be able to adequately resolve the earth where needed. This will hopefully reduce ambiguities in the resulting models. Consequently, survey design is an important topic that could receive increased attention in the future.

What does it mean for a survey to be well designed? Does it mean that the data are

sensitive to the target, or does it mean that the data can adequately constrain an inversion that will model the target? These are questions that need to be addressed at the survey design stage of each exploration program, and the answers will depend on the type of subsurface information being sought.

Other obvious questions that come to mind with respect to survey design are the following: (1) How efficient are current survey designs at defining the subsurface? and, (2) How might current survey designs be improved? The answers to these questions will depend on the application of the method, and may be contained within the context of an exploration program, even within this thesis. As better survey design methods are developed, surveys could become much more tailored to the exploration stage at which they are implemented. This would result in an even more integrated and efficient exploration program.

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Appendix A

Coordinate System

Many of the data at San Nicolás were collected using different coordinate systems. In order to compare and analyze different data sets and models with one another, a standard coordinate system was chosen to which all other data were transformed prior to modeling. The system chosen was a local coordinate system that has a north orientated about 23 degrees to the east of true north. This coordinate system was chosen because the geologic sections and much of the data provided by Teck were already in this coordinate system. Drill-hole fences were also orientated with the local grid. The airborne data was provided in the UTM (Universal Transverse Mercator) Zone 13, NAD (North American Datum) 27 coordinate system. As a result a transformation is presented that was used for converting from UTM(meters) to the local grid system, also in meters. This is accurate to within a few meters at the deposit.

$$X_{LOC} = (\cos(-0.36824) * (X_{UTM} - 192432)) + (\sin(-0.36824) * (Y_{UTM} - 2500845)) - 2000 \quad (A.1)$$

$$Y_{LOC} = -(\sin(-0.36824) * (X_{UTM} - 192432)) + (\cos(-0.36824) * (Y_{UTM} - 2500845)) - 400 \quad (A.2)$$

In the above equation, X_{LOC} is the local x-coordinate, X_{UTM} is the UTM x-coordinate, Y_{LOC} is the local y-coordinate, and Y_{UTM} is the UTM y-coordinate.

Appendix B

Physical Property Analysis

B.1 Introduction

Physical property measurements are essential in order to choose which survey to use and to interpret the results of a survey. This section presents the physical property values that were used in the work carried out at San Nicolás. The physical property *rock-models* are constructed using these physical property values. These rock-models were used for forward-modeling performed in the thesis.

B.2 Density

The density values that were used for in this thesis come from density measurements performed by *Quantec* on a variety of rock-types at San Nicolás, from density measurements that were made mainly on the ore for mine tonnage estimates, and, where specific rock-type information was insufficient, from physical property values found in *Applied Geophysics*, [Telford et al., 1990]. Density measurements are summarized below.

B.3 Magnetic Susceptibility

The magnetic susceptibility (MS) of a volume of rock is a function of the amount of magnetic minerals, (mainly magnetite and pyrrhotite), contained within the rock. These measurements can be interpreted to reflect lithological changes and the presence of alteration zones.

Rock-Type	Density (g/cm ³)
Quaternary Alluvium	1.8 - 2.2
Tertiary Breccia	2.1 - 2.5
Volcanoclastic sediments, tuffs, and flows	2.2 - 2.6
Mafic flows and breccias	2.5 - 2.9
Mafic flows and sediments	2.5 - 3.0
Massive sulphide	3.5 - 4.3
Semi-massive sulphide	3.0 - 3.5
Sulphide stringer zone	2.8 - 3.3
Rhyolite flows and breccias	2.2 - 2.6
Quartz feldspar rhyolite	2.3 - 2.6
Mafic volcanics and siliceous sediments	2.6 - 3.0
Graphitic Mudstone	2.2 - 2.6

Table B.1: Range of densities for rock-types encountered at San Nicolás.

B.3.1 Magnetic Susceptibility core measurements

Over one thousand magnetic susceptibility measurements were collected on core samples from thirteen drill-holes. These measurements are summarized in Figure B.1 for each rock type.

Reliability of measurements

Susceptibility measurements were collected mostly on three to four meter intervals along the drill-core (except in areas of high susceptibility variation). Tests were conducted on four different sections of core to see how well these coarse measurements sampled the detailed susceptibility features. Results for two test sections of core are presented in the following plots (Figures B.2 and B.3). The graphs show the original coarsely-sampled measurements and measurements made at finer intervals. The results show that the coarse measurements are probably adequately representative of the true susceptibility

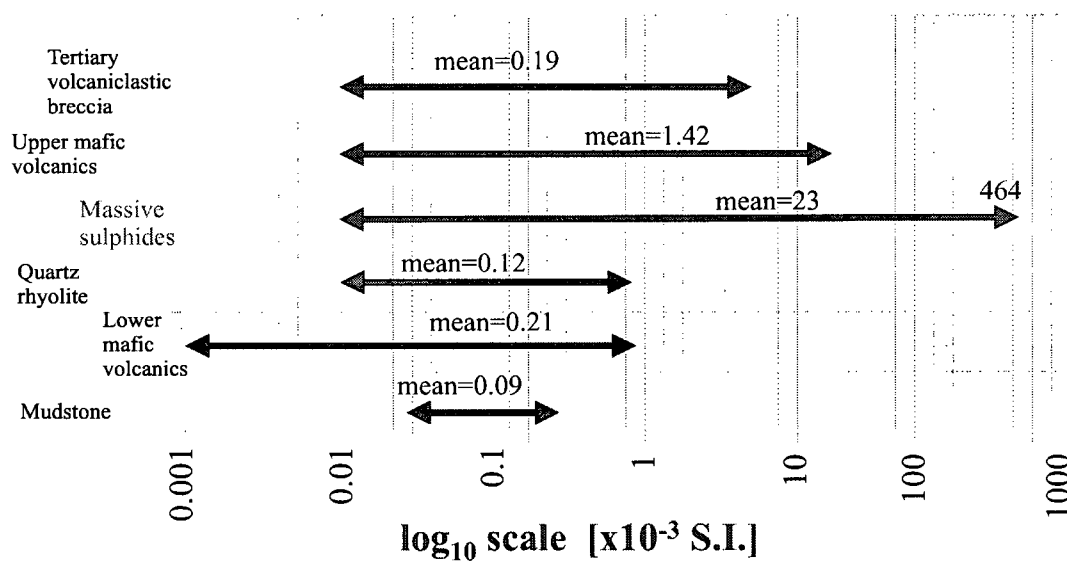


Figure B.1: The range of susceptibility values for different rock types at San Nicolás.

distribution.

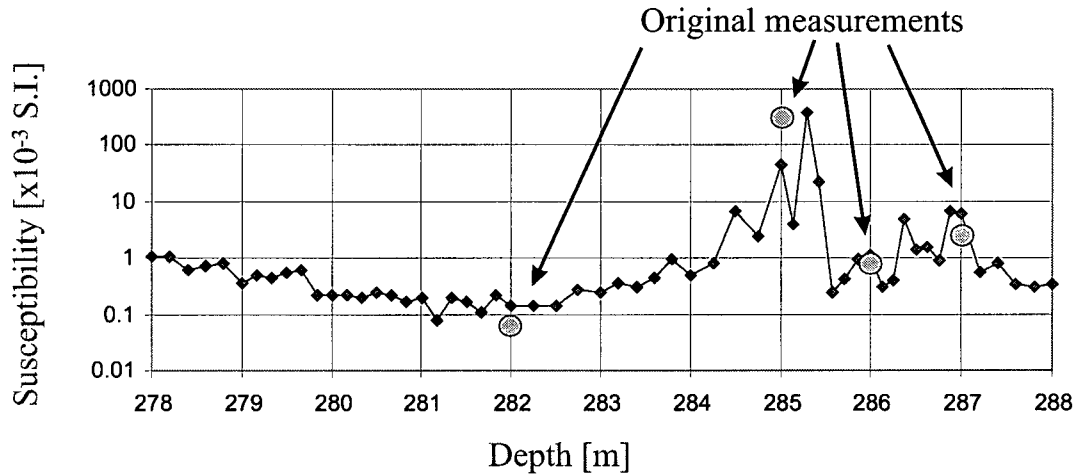


Figure B.2: Coarse and fine magnetic susceptibility measurements on a section of core from drill-hole SAL 70.

B.3.2 Composite models for bounding inversions

For magnetic susceptibility measurements that were not evenly sampled, the mean would not be a representative value of several readings as it would be biased towards the areas where more readings were taken. When calculating values to assign large cells from twenty or so unevenly sampled measurements, each measurement was weighted based on the amount of core that it might represent. A depth interval was determined for each measurement that corresponded to the distance between two midpoints. The midpoints were half way between the location of the measurement in question and the locations of the two measurements immediately above and below. Each measurement was weighted based on the length of this interval. It is hoped that this is an improvement over the mean but it is realized that this method is still somewhat ad hoc. The composite model

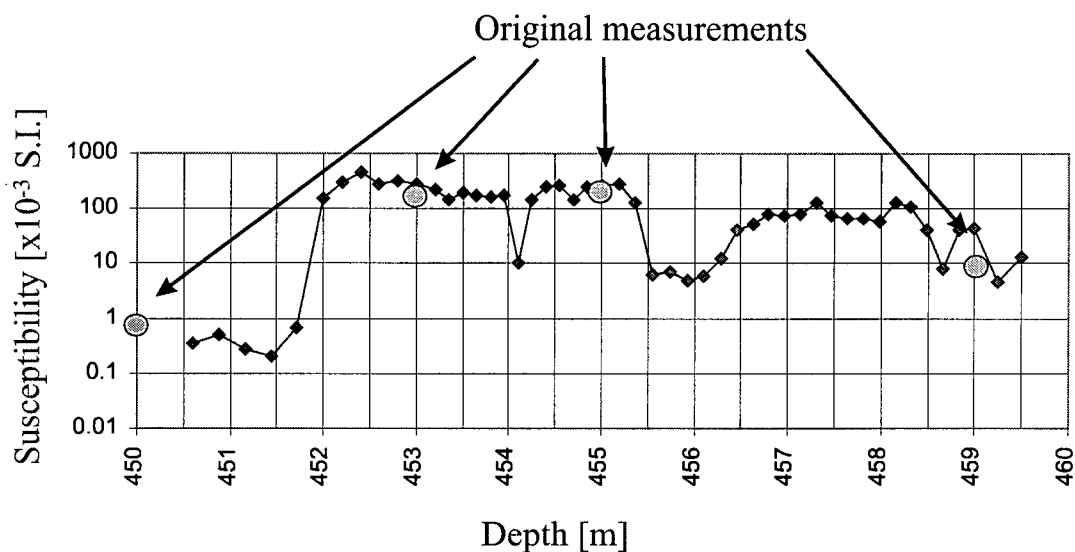


Figure B.3: Coarse and fine magnetic susceptibility measurements on a section of core from drill-hole SAL 34.

produced for SAL 25 using this method is shown in Figure B.4 along with the actual measurements.

B.4 Resistivity

Only twenty-one resistivity measurements were available for different rock-types at San Nicolás. The measurements are summerize below but it should be noted that they may not be representative of the values for each rock-type.

B.5 Chargeability

As with the resistivity values only twenty-one chargeability values were available. These did, however, sample all of the rock-types in the area.

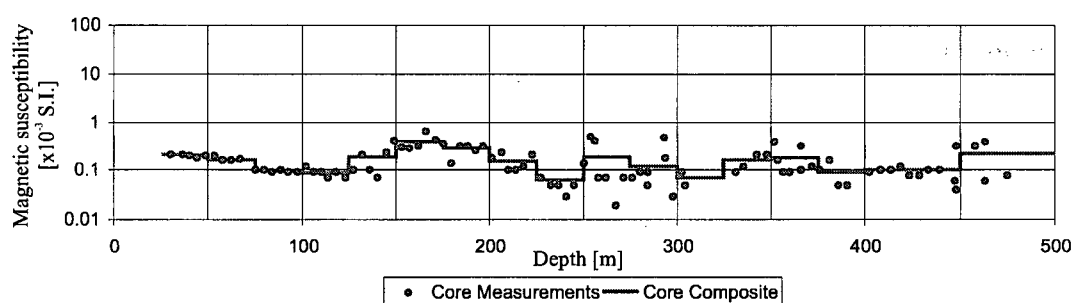


Figure B.4: Magnetic susceptibility measurements made on core from drill-hole SAL 25. A composite model was produced that was used to constrain inversion of ground magnetic data.

Rock-Type	Resistivity (Ohm-m)
Quarternary Alluvium	30?
Tertiary Breccia	15-30
Volcanoclastic sediments, tuffs, and flows	80
Mafic flows and breccias	80
Mafic flows and sediments	80
Massive sulphide	20
Semi-massive sulphide	20
Sulphide stringer zone	30
Rhyolite flows and breccias	100
Quartz feldspar rhyolite	100
Mafic volcanics and siliceous sediments	80
Graphitic Mudstone	100-200

Table B.2: Range of resistivities for rock-types encountered at San Nicolás.

Rock-Type	Chargeability (msec)
Quaternary Alluvium	5 - 10
Tertiary Breccia	10 - 30
Volcanoclastic sediments, tuffs, and flows	10 - 100
Mafic flows and breccias	30 - 80
Mafic flows and sediments	30 - 50
Massive sulphide	500
Semi-massive sulphide	400
Sulphide stringer zone	500
Rhyolite flows and breccias	10 - 20
Quartz feldspar rhyolite	100
Mafic volcanics and siliceous sediments	50 - 80
Graphitic Mudstone	30 - 70

Table B.3: Range of chargeabilities for rock-types encountered at San Nicolás.

B.6 *Physical property rock-models*

A geologic rock-model was constructed by digitizing five geologic sections (four east-west sections and one north-south section). The sections were digitized on a 25m x25m grid and each grid cell was assigned a rock-type based on the predominant content of each cell. In order to construct the physical property rock-models the digitized geology was assigned values based of the physical property data base presented above. Then an routine was used to interpolate between sections and construct a three-dimensional model. The interpolation routine used triangulation between cells at the same depth. For density and chargeability values a linear interpolation was used, for magnetic susceptibility and resistivity the Log_{10} of the value was interpolated. The resulting models were used for the keel sensitivity and survey design analysis presented in Chapter 6.

Figure B.5 shows the conductivity physical property rock-model that was used to forward model Questem, CSAMT, and DC resistivity data.

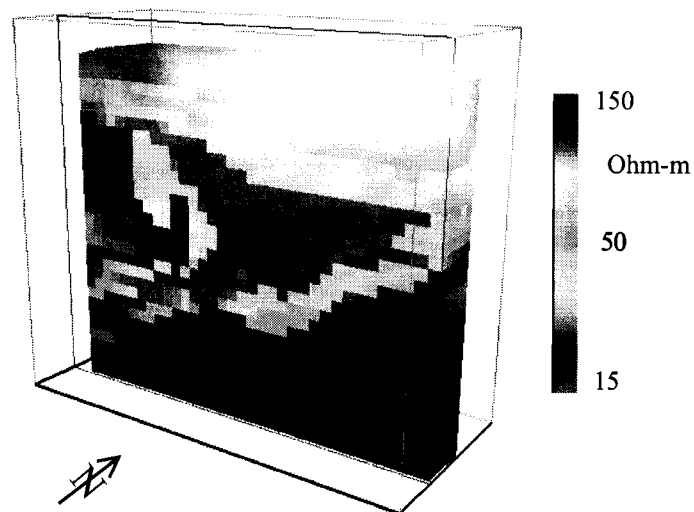


Figure B.5: Perspective view of the physical property conductivity model used for test forward modeling in Section 4.4.1. The model has been sliced at -400m N to expose a north-facing cross-section.

Appendix C

Geophysical methods

Many different geophysical methods were used in the collection of data at San Nicolás. This section gives a brief overview of each method and the underlying physical principles, and also defines the data collected in each case.

C.1 The gravity method

Gravity surveys involve measuring local irregularities in the earth's gravitational field with the aim of using these measurements to determine subsurface density variations. Newton's law of gravitation describes the force of attraction that acts at a distance between two masses. In the example shown in Figure C.1, the force exerted by mass m_2 on m_1 (\vec{F}_{21}) is expressed as:

$$\vec{F}_{21} = -\frac{\gamma m_1 m_2 \hat{r}}{|\vec{r}_2 - \vec{r}_1|^2} \quad (\text{C.1})$$

where G is the universal gravitational constant ($6.672 \times 10^{-11} \text{Nm}^2/\text{kg}^2$), m_1 to m_2 are masses, \hat{r} is the unit vector from m_1 to m_2 , and \vec{r}_1 and \vec{r}_2 are displacement vectors from the origin of m_1 to m_2 respectively. Furthermore the acceleration, or force per unit mass, due to m_2 at a point, defined by \vec{r}_1 can be stated as:

$$\vec{g}(\vec{r}_1) = -\frac{\gamma m_2 \hat{r}}{|\vec{r}_2 - \vec{r}_1|^2} \quad (\text{C.2})$$

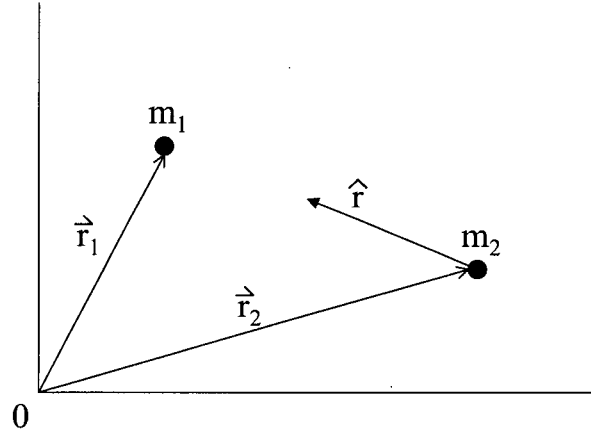


Figure C.1: The force exerted on mass m_1 by mass m_2 is in the direction of \hat{r} .

By rewriting \hat{r} as $\frac{(\vec{r}_2 - \vec{r}_1)}{|\vec{r}_2 - \vec{r}_1|}$, equation C.2 becomes:

$$\vec{g}(\vec{r}_1) = -\frac{\gamma m_2 (\vec{r}_2 - \vec{r}_1)}{|\vec{r}_2 - \vec{r}_1|^3} \quad (\text{C.3})$$

Due to superposition, for a number of point masses the gravitational acceleration can be written as:

$$\vec{g}(\vec{r}_1) = -\gamma \frac{\sum_{i=1}^N m_i (\vec{r}_i - \vec{r}_1)}{|\vec{r}_i - \vec{r}_1|^3}, \quad (\text{C.4})$$

and, for a distributed density function $\rho(\vec{r}_2)$, equation C.4 can be stated as the following volume integral:

$$\vec{g}(\vec{r}_1) = -\gamma \int_V \frac{\rho(\vec{r}_2) (\vec{r}_2 - \vec{r}_1) dv}{|\vec{r}_2 - \vec{r}_1|^3} \quad (\text{C.5})$$

Gravimeters use a mass on a zero-length spring to measure the vertical component of the gravitational field as stated below. The usual unit of measurement of the gravitational field is mGal, and most ground gravimeters are accurate to within about 0.01 mGal.

$$g_z(\vec{r}_1) = -\gamma \int_V \frac{\rho(\vec{r}_2)(z_2 - z_1)dv}{|\vec{r}_2 - \vec{r}_1|^3} \quad (\text{C.6})$$

Through the removal of regional fields (the field due to a background density distribution) we can focus on the gravitational field due to density variations within our depth of exploration.

$$\Delta g_z(\vec{r}_1) = -\gamma \int_V \frac{\Delta \rho(\vec{r}_2)(z_2 - z_1)dv}{|\vec{r}_2 - \vec{r}_1|^3} \quad (\text{C.7})$$

These data are typically collected at the surface of the earth on a 2D grid, and GPS is often used to accurately locate each gravity measurement to within a few centimeters, thus providing information about \vec{r}_1 and z_1 . The forward modeling operator used in the inversion procedure comes from the evaluation of Equation C.7 through discretization of the density distribution.

C.2 The magnetic method

Local variations in the earth's magnetic field are measured at the surface in order to gain information about subsurface magnetic susceptibility distributions. When a magnetizable body is placed in an external magnetic field it is magnetized by induction [Telford, Geldart, and Sheriff, 1990a]. The amount that the body is magnetized is measured by the magnetization vector \vec{J} , which is the dipole moment per unit volume. For this work we neglect the possible effects of remanent magnetization and demagnetization. We assume relatively small inducing magnetic fields, in which case the magnetization vector is proportional to, and is in the same direction as, the inducing field \vec{H} , thus:

$$\vec{J} = \kappa \vec{H}, \quad (\text{C.8})$$

where $\vec{H} = \vec{B}/\mu_0$ (μ_0 being the magnetic permeability of free space), and κ is the degree to which a body is magnetized by the inducing field or *magnetic susceptibility*. Most of the minerals which exhibit this type of magnetism are described as *ferrimagnetic*. These minerals have positive and relatively large susceptibilities.

In the magnetic experiment the inducing field is the earth's main magnetic field (\vec{H}_0), which is produced mainly by processes occurring in the outer core. The intensity of the total magnetic field resulting from the superposition of the main field and the secondary field is measured during magnetic surveys (Equation C.2).

$$\vec{B} = \vec{B}_0 + \vec{B}_S \quad (\text{C.9})$$

Through processing, the main field is removed and the secondary, or anomalous field is derived. The anomalous magnetic field due to a distribution of magnetizable materials can be stated as [Telford, Geldart, and Sheriff, 1990a]:

$$\vec{B}_S(\vec{r}_0) = \frac{\mu_0 \vec{H}_0}{4\pi} \int_V \nabla \nabla \frac{\kappa(\vec{r}) dV}{|\vec{r}_0 - \vec{r}|} \quad (\text{C.10})$$

It is this field (measured in Teslas, [T]) that is inverted in order to obtain a subsurface susceptibility distribution, $\kappa(\vec{r})$. Data are usually presented in nano-Tesla (nT).

The strength and orientation of the inducing field (\vec{H}_0) must be known for inversion of this data. This is obtained by using the program GEOMAG, written by John M.

Quinn of the USGS. GEOMAG used numerical models of the earth's magnetic field to calculate inclination, declination, total intensity, and their associated rates of change given a location and a date [Quinn et al., 1995].

C.3 Inductive electromagnetic methods

Inductive electromagnetic methods actively energize the ground using a time varying current source. The current source produces a magnetic field that interacts with conductive material in the earth, causing electric fields and inducing currents in the ground. The currents in the earth also produce a magnetic field. The magnetic fields emanating from the current in the transmitter are termed primary fields, while the magnetic fields emanating from the current in the ground are termed secondary fields. Both of these magnetic fields induce a voltage in the conductive receiver loop but it is the secondary fields that provide information about the conductivity of the subsurface.

C.3.1 Frequency-domain electromagnetic airborne surveys.

The Dighem frequency-domain airborne system produces an alternating current waveform in the transmitter loop and operates at several different frequencies. The sources and receivers are small loops that are towed behind the aeroplane (away from electromagnetic interference). In this system the loops are orientated in two configurations: horizontal co-planar, and co-axial, as shown in Figure C.2. The position of the receiver relative to the transmitter is kept constant throughout the survey.

Ward and Hohmann [Ward and Hohmann, 1987], derive an expression for the voltage generated in the receiving loop when a sinusoidal current waveform is produced in the transmitter loop and the system is in the presence of a conductive body:

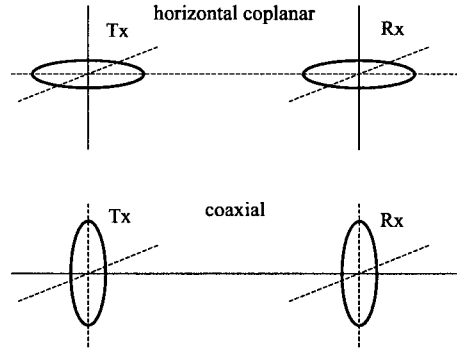


Figure C.2: Orientation of transmitter-receiver (Tx-Rx) pairs in the Dighem frequency-domain airborne electromagnetic system.

$$V = -i\omega\mu_0 NA(H_z^P(s) + H_z^S(\omega, \sigma, h, s)) \quad (\text{C.11})$$

where ω is the frequency, μ_0 is the magnetic permeability of free space, N is the number of turns in the receiver coil, A is the area of the receiver coil, $H_z^P(s)$ is the vertical component of the primary magnetic field at the center of the receiver coil (and is a function of the separation, s , between the transmitter and receiver coils), and $H_z^S(\omega, \sigma, h, s)$ is the vertical component of the secondary magnetic field at the center of the receiver coil, which is a function of frequency, conductivity (ω), height above the ground (h), and transmitter-receiver separation.

Normalization of this voltage by the transmitter current, and the representation of the signal as a ratio of the voltage generated in the presence of a conducting body to that generated in free space, allows the measured value to be written as:

$$\frac{H_z^S(\omega, \sigma, h, s)}{H_z^P(s)} \quad (\text{C.12})$$

In order to remove the effect of the primary field a *bucking coil* is used. This is an additional coil that measures the primary field at the receiver and allows the secondary field to be calculated. Because the secondary magnetic field is not in phase with the primary magnetic field, equation C.12 can be represented as inphase and quadrature components, and are usually given in parts per million of the primary field.

C.3.2 Time-domain electromagnetic airborne surveys

In the time-domain method, a current is passed through the transmitter loop and abruptly interrupted. According to Faraday's law, this change in primary field will induce currents in nearby conductors. As time progresses the currents diffuse from the surface of the conductors inwards [Nabighian and Macnae, 1987] and, for a conducting half space, downward and outward in a *smoke-ring* effect [Nabighian, 1979]. Currents will persist and propagate slower in conductors, and diminish and propagate faster in resistors. The secondary magnetic fields emanating from the induced currents generate an electric field and induce a voltage in a conductive receiver loop. These voltages are measured as a function of time. As a result, a slower decay the voltage will indicate the presence of a conductor. This is the classic use of time-domain electromagnetic systems.

Measurements are made by integrating the voltage over time-windows following the turn-off of the current source. The absence of a primary field during recording time (current off-time) allows for small signals to be measured, although some systems also measure during the on-time. The data can be presented as: $\partial B / \partial t$ for different times, the secondary magnetic field as a function of time normalized by the transmitter current, or as voltages induced in the receiver loops.

The transmitter is six turns of wire making a 186 square-meter loop, which is slung around an aeroplane (Figure C.3). When the peak current of about 500 Amps is reached the transmitter has a magnetic moment of half a million Amps per square-meter. The

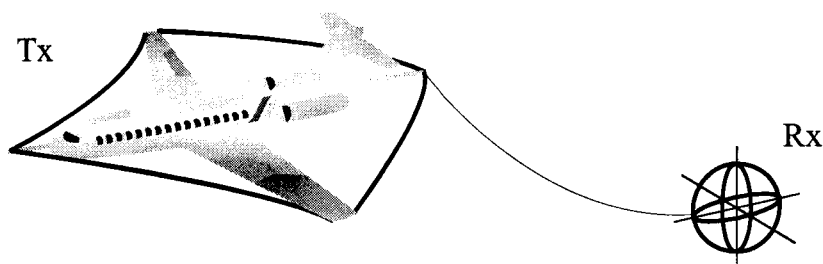


Figure C.3: Diagrammatic sketch of the Questem time-domain electromagnetic system. The transmitter (Tx) is placed around the aeroplane while the three-component receiver (Rx) is towed in a bird behind the plane. GPS positioning is used to measure the position of the receiver relative to the bird.

current waveform refers to the behavior of the current in the transmitter as a function of time. The Questem system uses a periodic half-sine waveform alternating in polarity to produce the primary magnetic field (shown in Figure C.4). The current is turned on for 4 msec and recordings are made by the three-component receiver over 15 subsequent time-windows during the 16 msec following the termination of the current pulse. The time-windows increase in width over the recording period so as to improve signal-to-noise in later times.

C.4 CSAMT surveys

The controlled source audio-frequency magnetotelluric (CSAMT) method is an electromagnetic technique that uses a grounded horizontal electric dipole source to directly *inject* currents into the ground. These currents produce magnetic fields that induce electric fields in the ground as the signal propagates from the source. It is assumed that the fields are measured far from the source (far-field¹) in which case the incident fields should

¹The far-field zone is defined as the region in which the source-receiver separation is much greater than the skin depth. The skin depth is approximately stated as $\delta = 503 \sqrt{\frac{\rho}{f}}$ [m], where ρ is the resistivity, and f is the frequency.

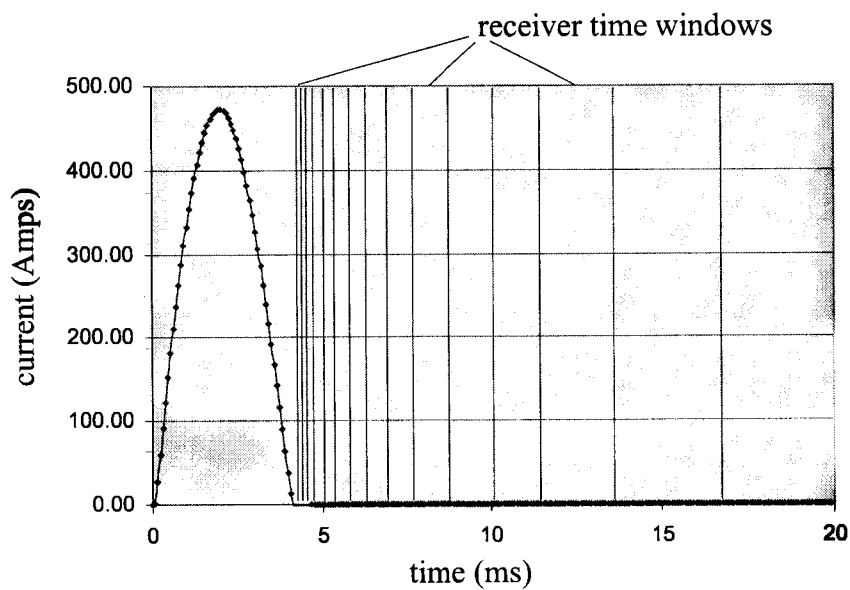


Figure C.4: The current in the transmitter as a function of time for 50% of one cycle. Receiver measurements are made over 15 time-windows during the 16 msec following the current on-time.

be planar in nature. Components of the electric and magnetic field are measured at a number of frequencies in the audio range (0.1 Hz to 10kHz) (Figure C.5). Perpendicular, horizontal electric and magnetic field values are used to calculate apparent resistivity and phase change at different frequencies. Zonge and Hughes, [Zonge and Hughes, 1987], provide a detailed description of the CSAMT method and derive an expression for far-field apparent resistivity (Equation C.13).

$$\rho_a = \frac{1}{2\pi\mu f} \left| \frac{E_x}{H_y} \right|^2 [\Omega m] \quad (C.13)$$

The phase difference between the electric and magnetic fields in Equation C.13 is also calculated and is measured in milli-radians (mrads).

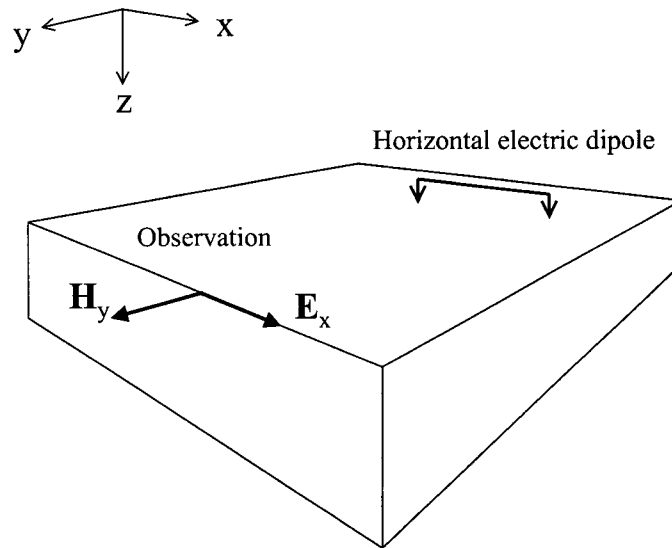


Figure C.5: Diagrammatic sketch of the CSAMT survey layout (modified from Zonge and Hughes).

C.5 DC resistivity methods

Resistivity methods involve injecting a current into the ground using grounded electrodes, and measuring the voltage between electrodes in the current flow (in Volts). The current, the voltage, and the geometry of the survey are measured so apparent resistivity values can be calculated.

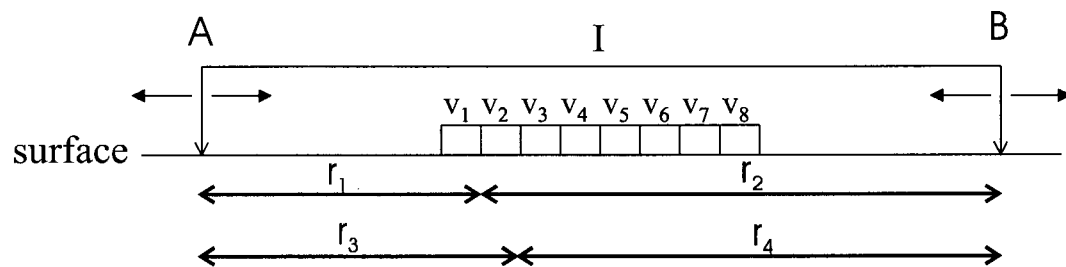


Figure C.6: “Real-section” array geometry used for collection of DC resistivity and IP data in Section 4.5.

For the array configuration show in Figure C.6 the apparent resistivity calculated from measurement V_2 would be:

$$\rho_a = \frac{2\pi V_2}{I} \frac{1}{\left[\left(\frac{1}{r_1} - \frac{1}{r_2}\right) - \left(\frac{1}{r_3} - \frac{1}{r_4}\right)\right]}$$

where V_2 is the measured potential, I is the current applied using electrodes A and B, and r_1 through r_4 are the distances shown on the diagram.

In formulating the forward modeling, we start with Ohm’s law:

$$\mathbf{J} = \frac{1}{\rho} \mathbf{E} \quad (\text{C.14})$$

The electric field, \mathbf{E} , is the gradient of the scalar potential V so we can write:

$$\mathbf{E} = -\nabla V \quad (\text{C.15})$$

With the above equation, and assuming that no charge accumulates ($\nabla \cdot \mathbf{J} = 0$), we can write an equation that relates the resistivity, the potential and the current sources for a homogeneous medium:

$$\nabla \cdot \left(\frac{1}{\rho} \nabla V \right) = I \quad (\text{C.16})$$

This equation is the basis for the DC resistivity forward modeling that is used in the inverse modeling — *DCINV3D* (Section E.7). *DCINV3D* constructs a resistivity distribution that will fit the potential data to a specified degree, given the source locations.

C.6 IP methods

Induced polarization occurs when a current is applied to the earth and there is an accumulation of charged particles in the pore fluid due to either the presence of metallic minerals, clay minerals or graphite, or restrictions in the pore itself. The ability for a material to accumulate these charges is summarized by its chargeability.

As a current is applied, the voltage measured across a pair of electrodes increases very fast initially (V_P in Figure C.7) and then more slowly (logarithmically) until a maximum voltage (V_T) is reached at a finite time [Telford, Geldart, and Sheriff, 1990c]². This slow increase in *overvoltage* corresponds to a build up of charges either side of a pore restriction, much the same way a capacitor charges up. This is known as electrode polarization and occurs in the presence of metallic minerals.

²When the applied current is interrupted, a slow decay is also observed after an initial rapid voltage reduction.

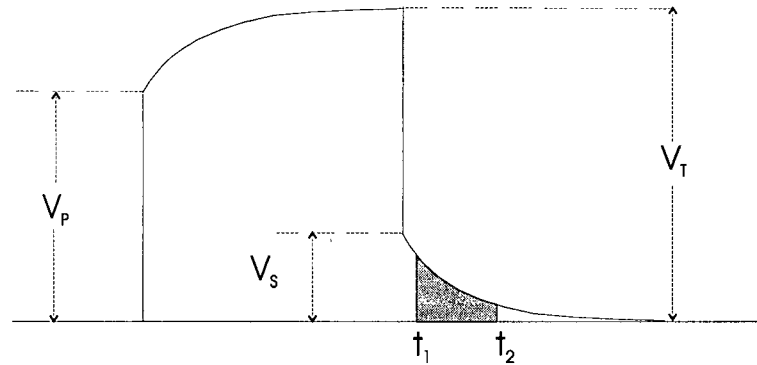


Figure C.7: The voltage response as a function of time, $V(t)$, in the presence of polarizable material, when a current is suddenly applied and then interrupted. $V_T = V_P + V_S$.

Measurements can be made in the time-domain or frequency domain. Time-domain measurements made at San Nicolás involved calculating the area under the decay curve between two times (t_1 and t_2 in Figure C.7) and normalizing it by the total voltage (V_T). In this case chargeability is defined in Equation C.17, and is measured in units of time (usually milli-seconds).

$$M = \frac{1}{V_T} \int_{t_1}^{t_2} V(t) dt \quad (\text{C.17})$$

The data are converted to secondary voltages (V_S) and forward modeling is carried out using the DC resistivity forward modeling (Equation C.16), but with the conductivity (σ) replaced with $\sigma(1 - \eta)$ [Seigel, 1959b]. Where η is the chargeability³. This produces large secondary voltage for high chargeabilities and small secondary voltages when η is small.

³The chargeability, η for this formulation is dimensionless and less than one. The apparent chargeability data η_a can be the ratio of V_S to V_P (also dimensionless and usually small). Alternatively M in Equation C.17 could be used (in units of seconds), in which case the modeled values of η are also shown in units of time.

Appendix D

Data

Data used for the inversions are presented in this section. The survey specifications associated with each data set are given in Table D.1. When inversions are performed the predicted data does resemble the observed data to the specified requirement, and often look identical upon first inspection. Data misfits are presented in Table E.1, however, the predicted data are not displayed themselves.

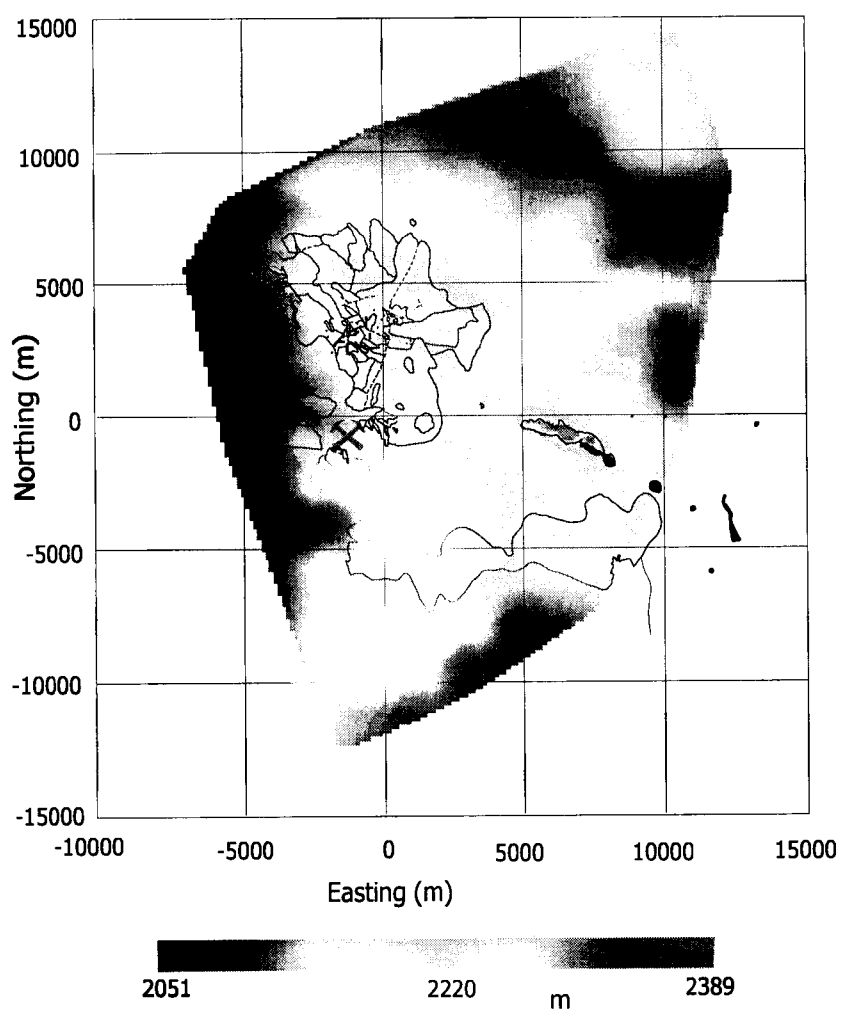


Figure D.1: Topographic elevations at San Nicolás. Apart from a hill at 6000m east, -1000m north, most of the area exhibits small topographic elevation changes. Outlines of outcropping geology are also shown, along with a hammer and pick symbol that locates San Nicolás at about -1700m east, -400m north.

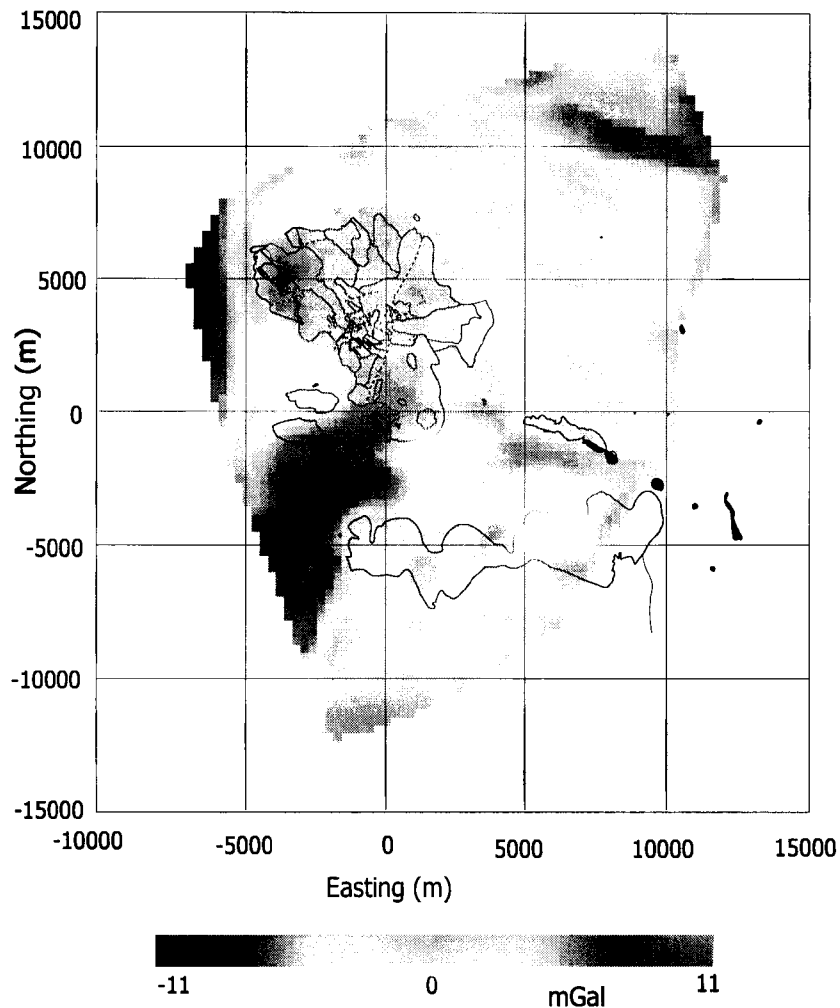


Figure D.2: Ground gravity observations. The data are the original data supplied by Quantec but with observations on the hill removed. The topographic map shown in Figure D.1 indicates a hill with over 100m of elevation gain. Initial inversions produced a large negative anomaly coincident with the hill. Observations made on and close to the hill were subsequently discarded, although the affects of the hill probably persist in some of the remaining observations in this area. A constant was also removed from the data in order to remove a first order regional field. An outline of outcrop geology is overlain.

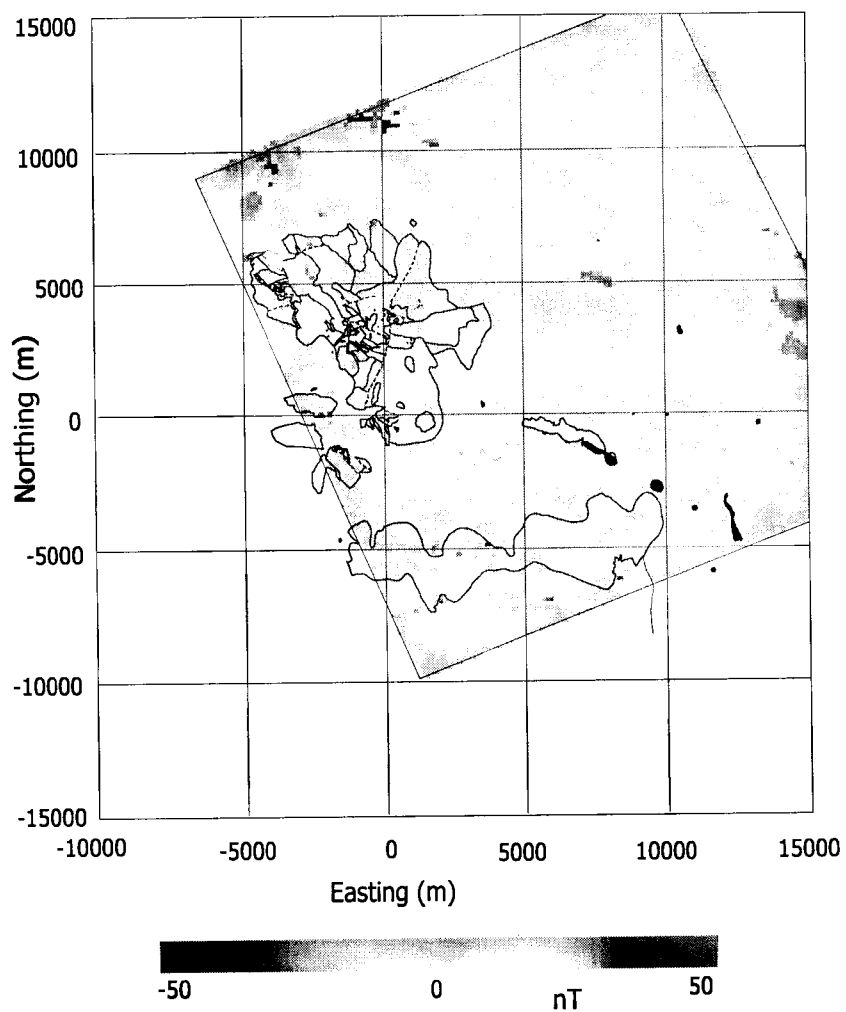


Figure D.3: Dighem airborne total-field magnetic data. The deposit is identified by a subtle anomaly at -1600m east, -400m north.

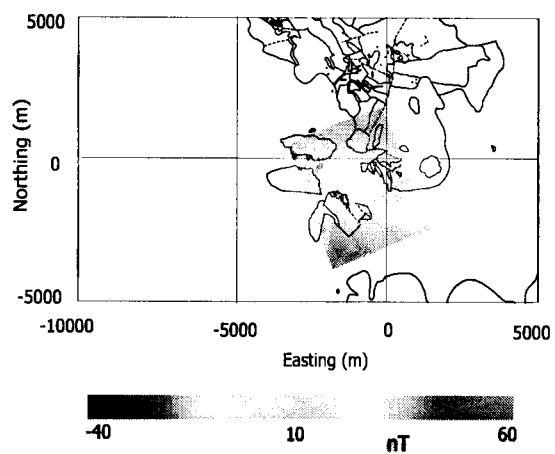


Figure D.4: Subset of Dighem airborne total-field magnetic data used for inversion.

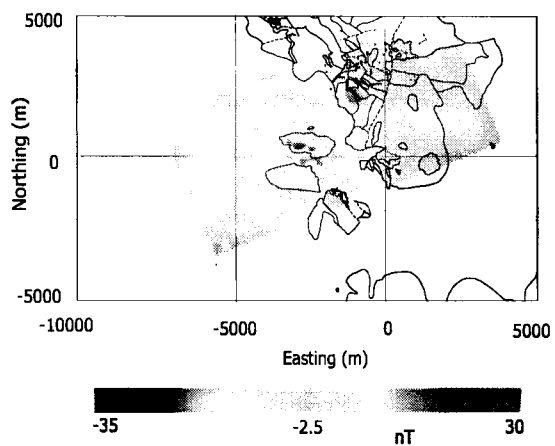


Figure D.5: Falcon airborne total field magnetic data.

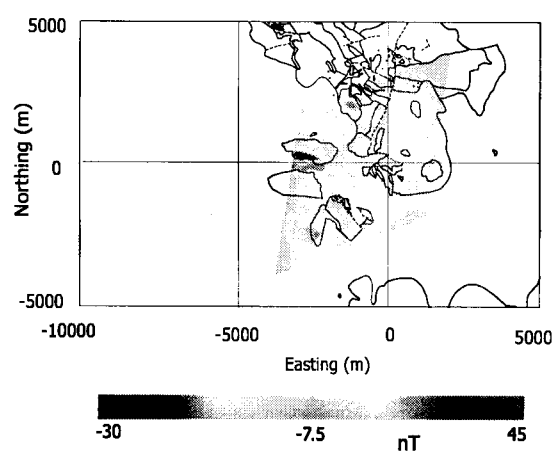
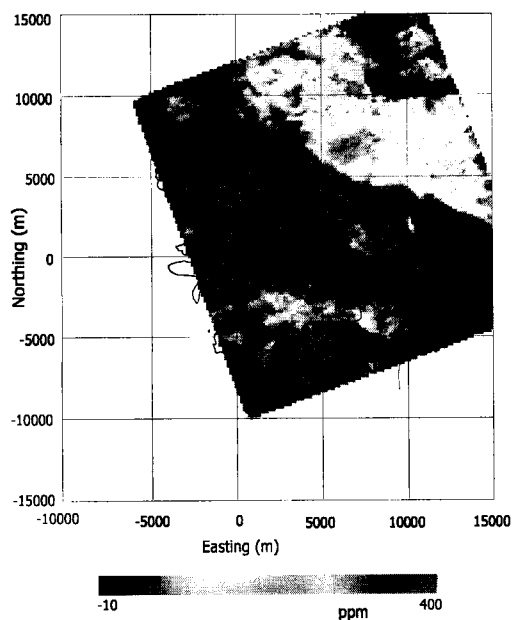
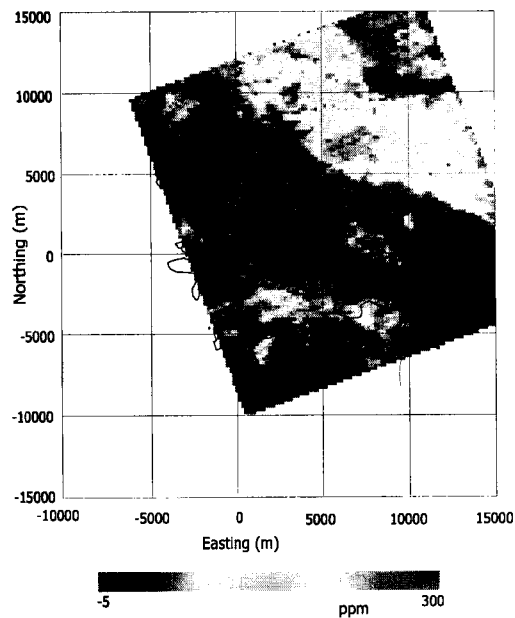


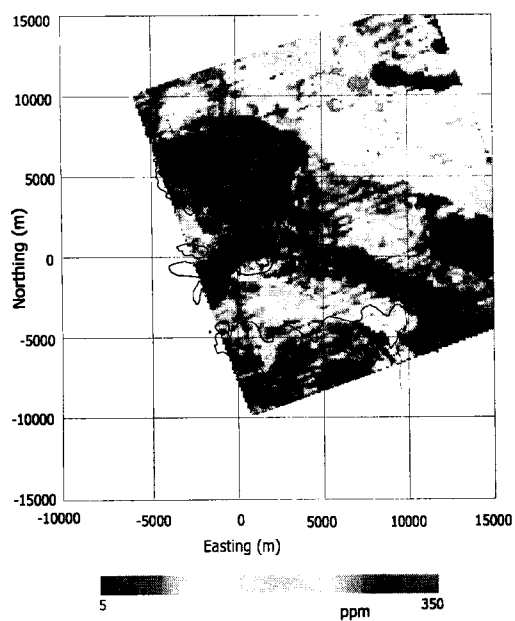
Figure D.6: Questem airborne total field magnetic data.



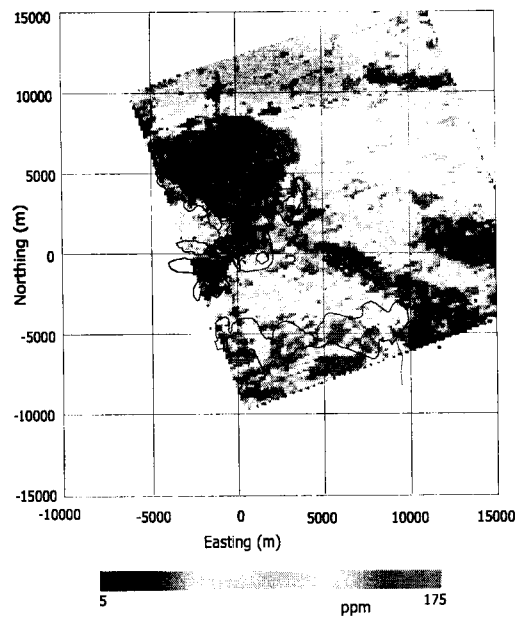
(a) Inphase coplanar 466Hz



(b) Inphase coaxial 1046Hz

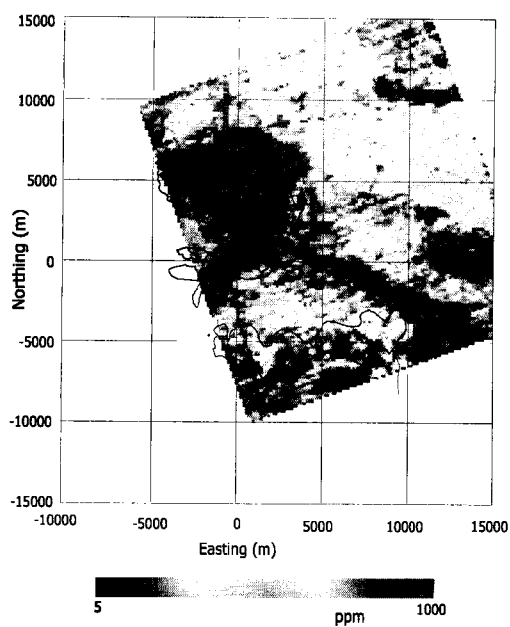


(c) Quadrature coplanar 466Hz

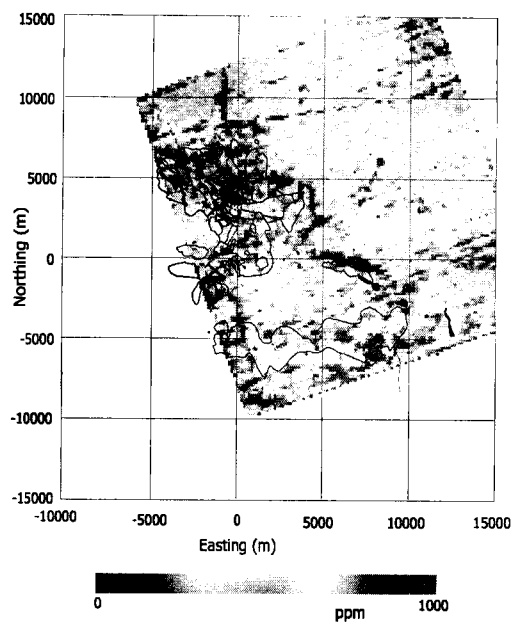


(d) Quadrature coaxial 1046Hz

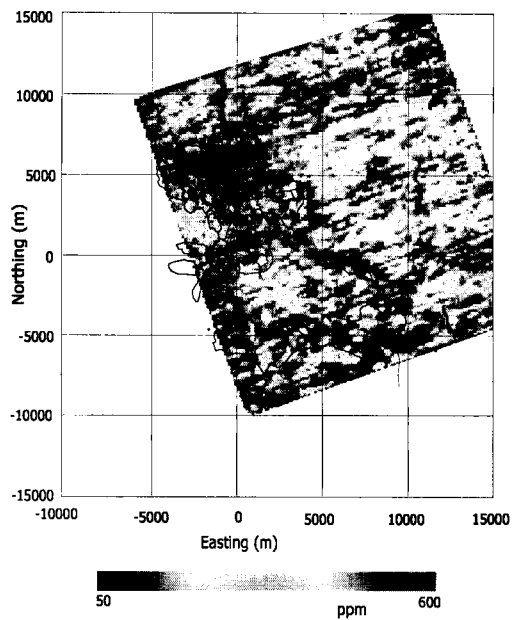
Figure D.7: In-phase and quadrature frequency-domain electromagnetic data from the Digheem system: 466Hz, 1046Hz, and 6495Hz



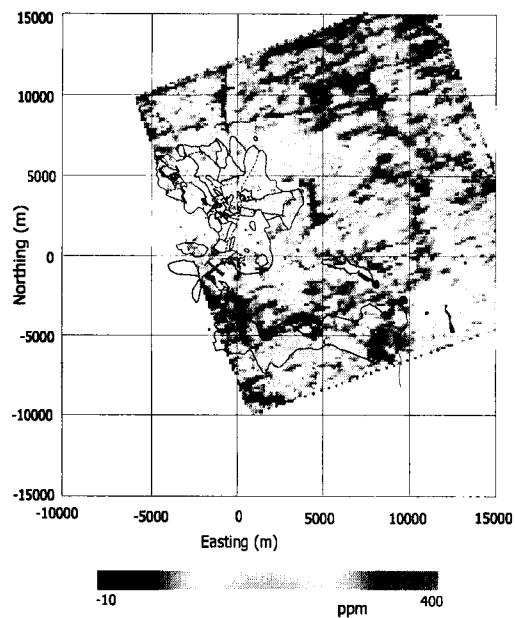
(a) Inphase coplanar 7206Hz



(b) Inphase coplanar 56kHz



(c) Quadrature coplanar 7206Hz



(d) Quadrature coplanar 56kHz

Figure D.8: In-phase and quadrature frequency-domain electromagnetic data from the Dighem system: 7206Hz and 56kHz

Questem Z-component data along flight line over the deposit.

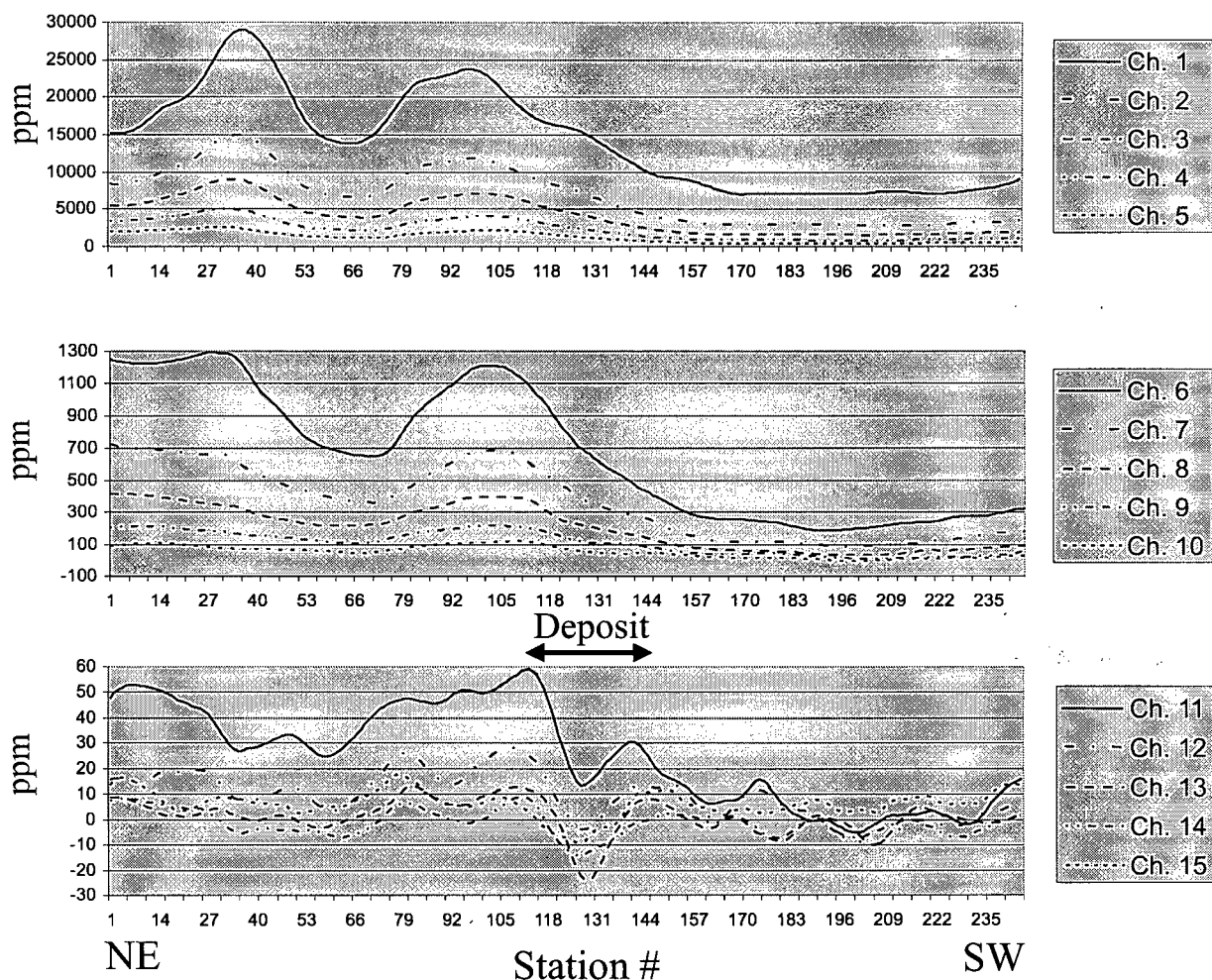


Figure D.9: Profiles of Z-component dB/dT data for the fifteen time-channels (0 to 18 msec after the transmitter current turn-off) along line 10. The upper panel shows data for the first five time channels, the center plot shows data for the next five time channels, and the lower panel shows data for the last five time channels. The deposit corresponds to the strong negative response in the later time channels. Thirty-one lines of data were collected and different orientations but this is the only line that has been presented.

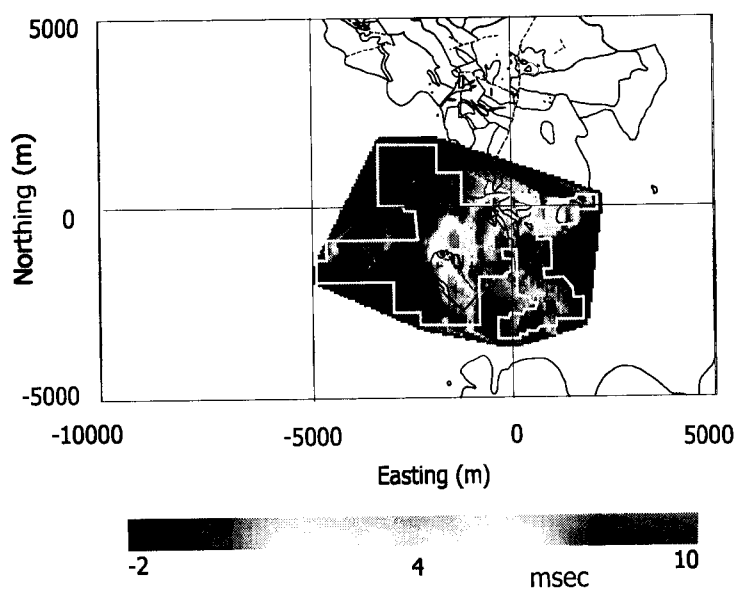


Figure D.10: Gradient Array IP data. The boundary shows the extension of the survey measurements and an outline of outcropping geology is also shown.

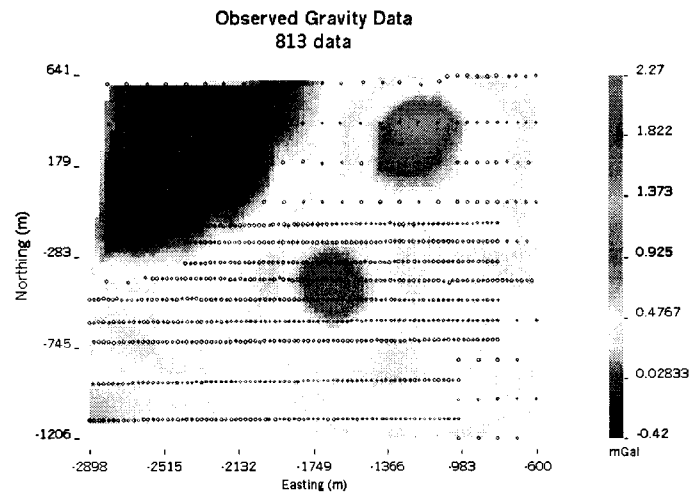


Figure D.11: Reduced gravity data. This gravity data has had a regional signal removed. The resulting gravity data were used for producing the density-contrast models in Chapters four and five.

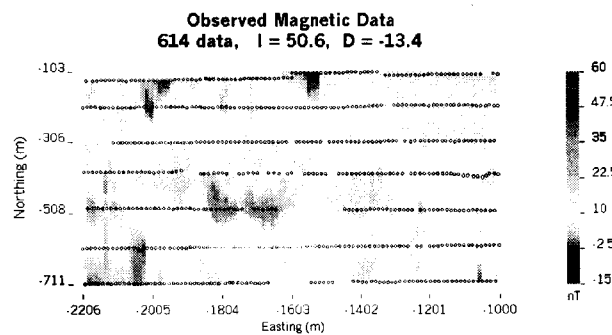


Figure D.12: Total-field ground magnetic data. The observation locations are displayed and the gaps where the data have been discarded are clearly visible.

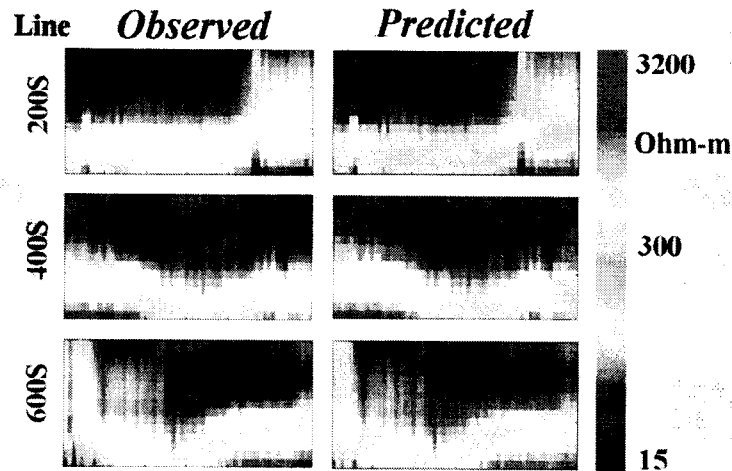


Figure D.13: Observed and predicted CSAMT apparent resistivity data. The data was collected along three lines, -200m N, -400m N, and -600m N. Each data plot is plotted with frequency in the vertical axis (ranging from 8192Hz at the top to 0.5Hz at the bottom), and easting along the horizontal axis (-2200m E to -800m E).

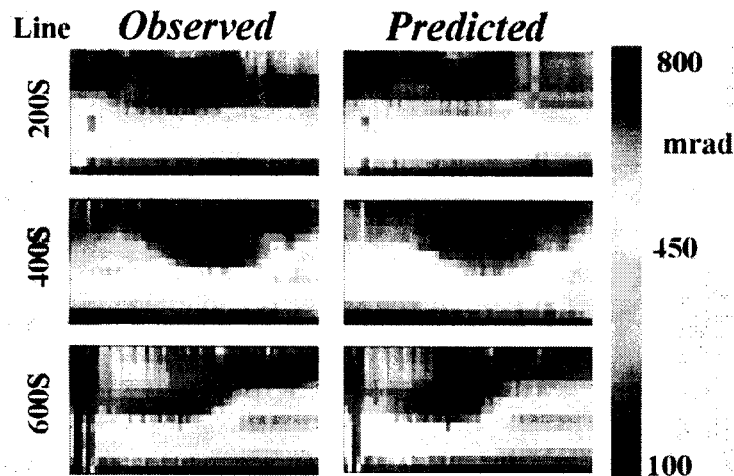


Figure D.14: Observed and predicted CSAMT phase data. Each data plot is plotted with frequency in the vertical axis (ranging from 8192Hz at the top to 0.5Hz at the bottom), and easting along the horizontal axis (-2200m E to -800m E).

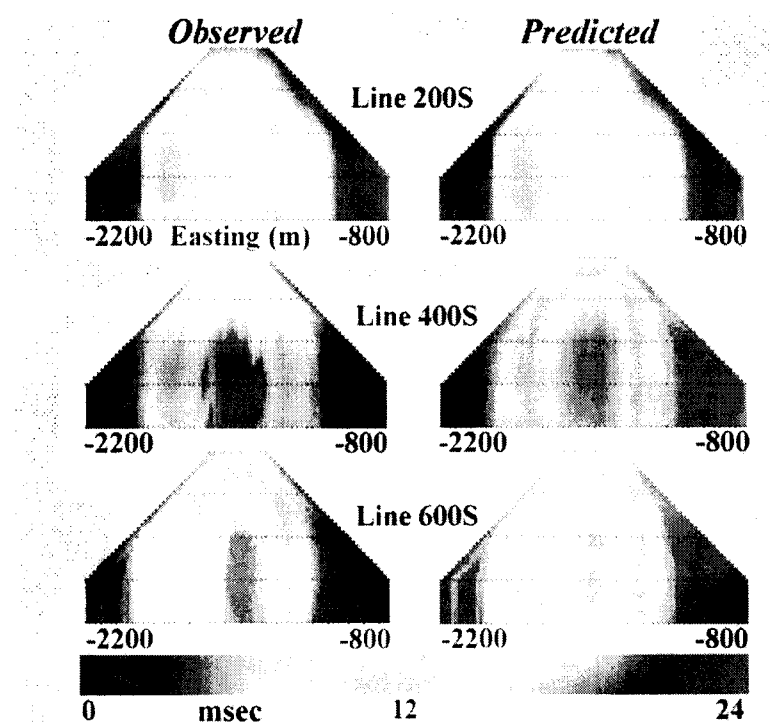


Figure D.15: Observed and predicted “Realsection” IP data. Five rows of data are plotted in each panel. Each row corresponds to a different transmitting dipole length.

Survey Type	Configuration	Number of Data	Line Spacing	Station Spacing	Preprocessing	Acquired by:	Year collected	Other Specifications
Gravity	East-West lines	17671	100m-200m	25m-100m	Instrument, Tidal, Free-air, Bouguer, Regional	Quantec de Mexico and Geociencias Consultores	1998	GPS positioning
Airborne Magnetic(1)	East-West lines (UTM)	281107	200m	8m	Diurnal corrections, leveling, removal of inducing field	Digheem	1997	40m height
Airborne Magnetic(2)	East-West lines (UTM)	3302	250m	60m	Diurnal corrections, leveling, removal of inducing field	BHP	2001	100m height
Airborne Magnetic(3)	N030E lines (UTM)	10115	200m	14m	Diurnal corrections, leveling, removal of inducing field	Fugro	1998	90m height
Ground Magnetics	East-West lines	614	100m	12.5m	Diurnal corrections, contaminated data discarded	Quantec	1998	
Airborne Electromagnetics (Frequency)	East-West lines (UTM), co-axial and coplanar	56258	200m	40m	primary field removed	Digheem	1997	40m height, 8m Tx-Rx separation, Freq: 466, 1046, 6495, 7206, 56627Hz
Airborne Electromagnetics (Time)	N030E lines (UTM)	10115	200m	14m	primary field removed	Fugro (Questem 450)	1998	90m height, 60m Tx-Rx separation, 15 time-windows
C.S.A.M.T.	East-West lines	10800 (60/stn)	200m	25m	Apparent resistivity and phase derived from E_x and H_y field measurements	Quantec	1998	15 frequencies (0.5-8192Hz), 1.7km Tx dipole length, Mean Tx-RX separation: 3.7km
Gradient Array I.P.	East-West lines (local)	6631	100m and 50m	25m	NA	Quantec	1998	20 and 25m dipoles
'Realsection' I.P.	East-West lines (local)	564	200m	25m	NA	Quantec	1998	25m dipoles, 5 transmitter dipole separations: 500m to 2500m
Core susceptibility measurements	Point measurements made on drill-hole core samples	1124	NA	0.5 - 5 m	NA	Nigel Phillips, UBC-GIF	2001	KT-9 magnetic susceptibility meter

Table D.1: Geophysical survey specifications

Appendix E

Inversion

In the body of the thesis the models have been presented without much discussion. Here we provide a brief description of the inversion codes used and the parameters that were used to invert each data set.

E.1 Introduction

This section gives a brief overview of the algorithms used for the inversions performed in this thesis. All of the algorithms were developed by the UBC Geophysical Inversion Facility under the auspices of the JACI (Joint and Cooperative Inversion of Geophysical and Geologic Data) research consortium.

E.2 *GRAV3D*

GRAV3D is a collection of algorithms that are used for forward modeling, inversion, and sensitivity calculations of gravity data. The developers of this code give a description of the method in [Li and Oldenburg, 1998b]. As described in Chapter 2 the earth model is discretized into a large number of cuboidal cells of constant density contrast. The user can design a range of model objective functions to be minimized, subject to fitting the data, in order to recover a density contrast distribution. Functionality allows for incorporating *a priori* information through the weighting of reference models, and bounding the range of values that cells can assume. A depth weighting scheme that counteracts the natural

decay of the gravitational source is needed to invert data that has no inherent depth information. This allows all parts of the model to play a role in the solution, and therefore introduces depth information into the model.

E.3 *MAG3D*

MAG3D is a suite of algorithms used to invert surface and borehole magnetic responses due to a three-dimensional susceptibility distribution ([Li and Oldenburg, 1997]). When susceptible material is magnetized by an inducing field, an anomalous magnetic field is produced that is superimposed on the inducing field. The anomalous, or induced, magnetic field is inverted to recover the susceptibility distribution. The algorithms assume that susceptible values are small (a reasonable assumption for most geologic scenarios with paramagnetic materials), and that no demagnetized or remnantly magnetized material is present. Residual magnetic field observations, locations and error estimates, the inducing field orientation and strength is required for inversions to proceed. The objective function minimized in the inversion allows various types of a priori information to be included. In addition weighting is incorporated to distribute the susceptibility in depth. Wavelet transforms are used for compression of the sensitivity matrix, allowing for faster run-times. The logarithmic barrier method ensures that positive magnetic susceptibility values can be enforced and bounds can be applied.

E.4 *CSAMT1D*

CSAMT1D is a near- and far-field 1-D code for recovering conductivity structures from E and H field components or the derived quantities of apparent resistivity and phase. The source electric field is a grounded loop and measurements are made in the frequency domain.

E.5 *EM1DFM*

EM1DFM inverts airborne or ground EM frequency-domain data to recover simultaneously, or separately, 1-D conductivity and magnetic susceptibility models. The source and receivers are represented by magnetic dipoles. This is a fast inversion routine and can be applied to regional data sets ([Farquharson and Oldenburg, 2000]).

E.6 *TEM1D*

TEM1D is used to invert the vertical component of the time-domain electromagnetic data to recover a layered earth model. The forward modeling is done in the frequency domain and a frequency-to-time transform performed at each iteration.

E.6.1 Complex forward modeling

A modified version of the *TEM1D* forward modeling code was used to produce the response over a polarizable layered earth. The Cole-Cole parameters are used to define the frequency-dependence of the model.

E.7 *DCIP3D*

DCIP3D is a suite of algorithms that are used to invert DC resistivity and IP data. The algorithms can be used in sequence. First, *DCINV3D* is used to invert DC resistivity data (primary potentials) and produce a resistivity distribution of the earth. Secondly, the resistivity model can be used to calculate sensitivities for the IP inversion (*IPSEN3D*). Thirdly IP data (secondary potentials) can be inverted using *IPINV3D*. The model objective function is designed to allow for a priori information to be included.

E.8 *EH3D*

EH3D is a forward modeling algorithm that computes frequency-domain electromagnetic responses from 3D conductivity and/or susceptibility distributions. The conductivity and susceptibility distributions can exhibit large discontinuities. The source can be either a grounded electric source, a loop source, or a plane-wave. Real and imaginary parts of electric and magnetic fields due to the source are computed over the whole model for different frequencies.

E.9 *EH3DTD*

EH3DTD is a forward modeling algorithm that computes time-domain electromagnetic responses from a given 3D conductivity structure of the earth. As with the frequency-domain equivalent a number of different sources can be used and the electric and magnetic fields are calculated throughout the whole model region. Different time-dependent waveforms can be used as sources.

Inversion Description	Inversion Program	Number of Data	Recovered model	Noise estimation	Number of cells	Cell Size	Reference Model	Achieved Misfit	Inversion Time	Other parameters
Regional Gravity	GRAV3D	17671	3D density-contrast	0.01 mGal	707625	250 x 250 x 100m	0 g/cm ³	482646	33.2hr	depth weighting, topography included
Dighem airborne magnetics	MAG3D	2566	3D magnetic susceptibility	4nT	167040	100 x 100 x 50m	0 S.I. x10 ⁻³	253	2hr	depth weighting, topography, inclination=50.638, declination=13.47 (local grid), inducing field=44000nT
BHP airborne magnetics	MAG3D	3318	3D magnetic susceptibility	4nT	862580	100 x 100 x 50m	0 S.I. x10 ⁻³	329	11.5hr	depth weighting, topography, inclination=50.638, declination=13.47 (local grid), inducing field=44000nT
Fugro airborne magnetics	MAG3D	10115	3D magnetic susceptibility	4nT	338640	100 x 100 x 50m	0 S.I. x10 ⁻³	241	38.5hr	depth weighting, topography, inclination=50.638, declination=13.47 (local grid), inducing field=44000nT
Gradient-array IP	MAG3D	6631	3D chargeability	1 msec	426400	50 x 50 x 50m	0 msec	4642	17.75hr	Inducing current orientation: inclination 0°, declination 90°
Local scale gravity	GRAV3D	813	3D density-contrast	1 mGal	446684	25 x 25 x 25m	0 g/cm ³	2654	3.2hr	depth weighting, topography
Ground magnetics	MAG3D	614	3D magnetic susceptibility	2 nT	222880	25 x 50 x 12.5m	0 S.I. x10 ⁻³	425	1hr	depth weighting, topography, inclination=50.638, declination=13.47 (local grid), inducing field=44000nT
CSAMT	CSAMT1D	60/stn	1D layered-earth resistivity	2-5% apparent resistivity and 1-2 degrees phase	40	1 - 10,000m	100 Ohm-m	60	6min/stn	
"Real-section" and gradient-array IP	DCIP3D	1182	3D chargeability	5% + 1 msec	117600	25 x 25 x 25m	0 msec	241	9hr	Electrode weighting

Table E.1: Inversion parameters

Appendix F

Gradient IP data modeled using *MAG3D*

Gradient-array data has been modeled using *MAG3D*, an inverse modeling package for inverting magnetic field data. Similarities between the nature of data allow this. The advantage of inverting gradient-array IP data in this way is that depth information can be introduced in the model. As with potential field data, depth information is not contained in these data. Here we show why this method is valid, and test it with some synthetic examples.

MAG3D generates 3-D susceptibility distributions based on the following forward modeling of the anomalous magnetic field from susceptibility distributions, κ , in an inducing field [Li and Oldenburg, 1997]:

$$\vec{B}(\vec{r}) = \frac{\mu_0 \vec{H}_0}{4\pi} \cdot \int_V \nabla \nabla \frac{1}{|\vec{r} - \vec{r}'|} \kappa dv, \quad (\text{F.1})$$

where \vec{r} is the point of observation of the magnetic field. This conservative field is derivable from the scalar magnetic potential [Grant and West, 1965], which is written as:

$$\Psi(\vec{r}) = -\frac{\mu_0 \vec{H}_0}{4\pi} \cdot \int_V \nabla \frac{1}{|\vec{r} - \vec{r}'|} \kappa dv, \quad (\text{F.2})$$

Gradient-array IP data can be interpreted and modeled as magnetic potential field data because of the dipole source of both fields. Seigel [Seigel, 1959a] describes the over-voltage effect associated with a metallic conducting sphere in an electrolyte as equivalent

to a dipolar source whose moment is proportional to the current density vector (\vec{j}). The dipole source has a volume current moment of $\vec{M} = -m\vec{j}$, where m is the chargeability. The potential due to a distribution of current dipoles can be written as:

$$\phi(\vec{r}) = \frac{1}{4\pi\sigma} \int_V \nabla \frac{1}{|\vec{r} - \vec{r}'|} \vec{M} dv \quad (\text{F.3})$$

where σ is the conductivity of the medium. Expressions of potentials in equations F.2 and F.3 have the same form if the conductivity and electric field are uniform. The negative of the data is used, as the dipoles would be aligned in antiparallel with the inducing current, so the inducing field is set antiparallel with respect to the position of the current electrodes.

F.1 Four synthetic examples

A synthetic example was used to test the use of *MAG3D* in this manner. It is expected that the conductivity of the deposit is anomalous with respect to its surroundings, so several synthetic experiments were carried out with different conductivity distributions in order to determine if this method is valid for use at San Nicolás.

A 100 msec chargeable cuboid was generated from 20m x 20m x 10m cells positioned between -60m and 60m east, -60m and 60m north, and 80m and 120m depth (Figure F.1). The body is hosted in a background of 0 msec chargeability.

A current of 1 Amp was transmitted through grounded electrodes at two locations straddling the target (-600m east, 0m north and 600m east, 0m north). Eighty-four gradient IP measurements were made using 12, 40m dipole locations on seven lines centered around the chargeable body. *DCIPF3D*, which is a forward part of the *DCIP3D* inversion package, was used to generate these measurements. Four different conductivity models were used to model this data (Figure F.2):

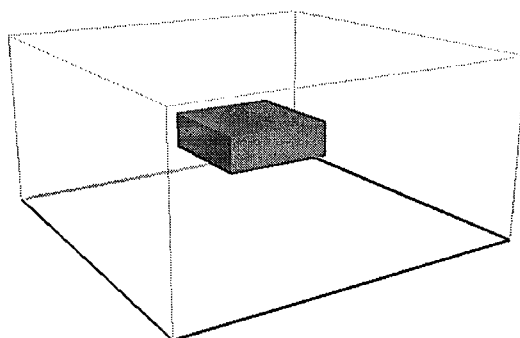


Figure F.1: Synthetic chargeable body.

1. a half-space of 0.01 S/m
2. a 0.1 S/m conductive body of the same size and location as the chargeable body in a background of 0.01 S/m
3. a 1 S/m conductive body of the same size and location as the chargeable body in a background of 0.01 S/m
4. a 0.1 S/m conductive body of the same size and location as the chargeable body in a background of 0.01 S/m with a 40m thick conductive overburden of 0.1 S/m (This model is thought to most closely resemble the conductivity distribution at San Nicolás.)

The four synthetic data sets (one for each conductivity model) were contaminated with random noise with a mean magnitude equal to 3% of value of the data. Each datum was assigned an appropriate standard deviation (3% of each datum) and the data sets were inverted using *MAG3D*. An incident field with 0 degrees inclination, 90 degrees declination, and 1 nT field strength was used.

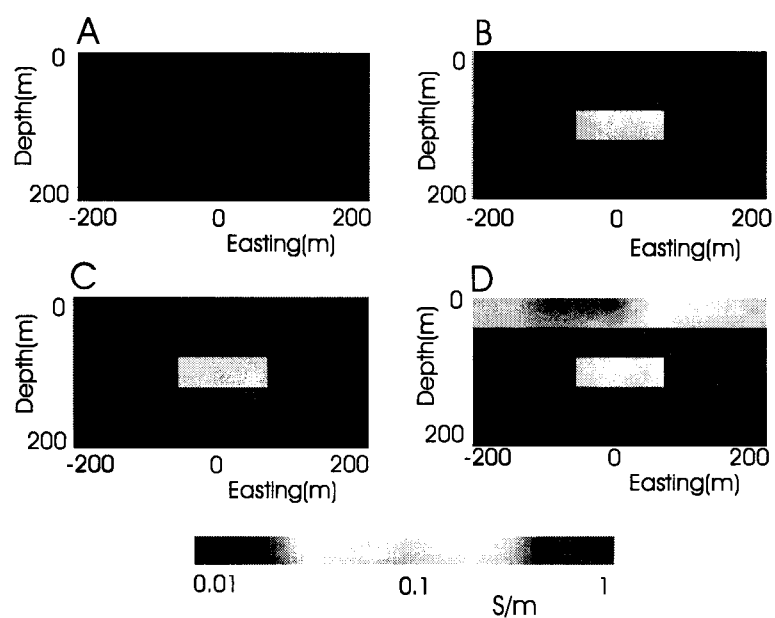


Figure F.2: North-facing cross-sections through the conductivity models: A) 0.01 S/m half-space, B) 0.1 S/m block in a 0.01 S/m background, C) 1 S/m conductive block in a 0.01 S/m background, D) 0.1 S/m block and a 40m thick 0.1 S/m overburden in a 0.01 S/m background.

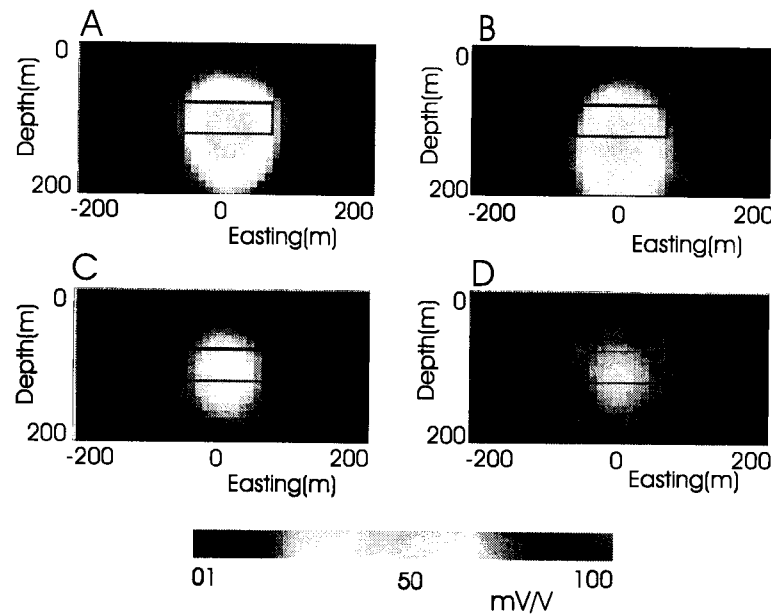


Figure F.3: North-facing cross-section through recovered chargeability models.

Figure F.3 shows the recovered chargeability models for the four models used. As expected, the half-space conductivity model gives the best results; the recovered values are closest to the true model value of 100 mV/V. As the conductivity increases, the recovered model deteriorates. The model with the conductive overburden is the worst. This is possibly because only a small amount of current was able to penetrate to the depth of the target. (A smaller current density will produce a smaller dipole source). In all cases there is good localization of the anomaly.

F.2 Modeling the electric field

As well as being affected by the conductivity structure, the electric field also must also be uniform in order for the treatment of gradient array IP data as potential field data. *EH3D* was used to forward model the electric field generated by the transmitter for each

of the conductivity models. The results are presented in Figure F.4, which shows the vector representation of the real part of the electric field for the north-facing cross-section at -400m north. The colors represent the magnitude of the field, and the line is orientated in the direction of the field vector with the white tip being forward.

The electric field is distorted by interactions with the conductivity structure. It is not surprising to see the best recovered chargeability results coming from the uniform conductivity and electric field distribution, while the less successful results are obtained from conductivity structures that distort the electric fields. Although the electric field and conductivity structure isn't too anomalous in the case where a conductive overburden is added, the current is concentrated near the surface (Figure F.5). This current channeling reduces the signal relative to the noise, which is reflected in smaller amplitude data, and results in a recovered model that is reduced in magnitude.

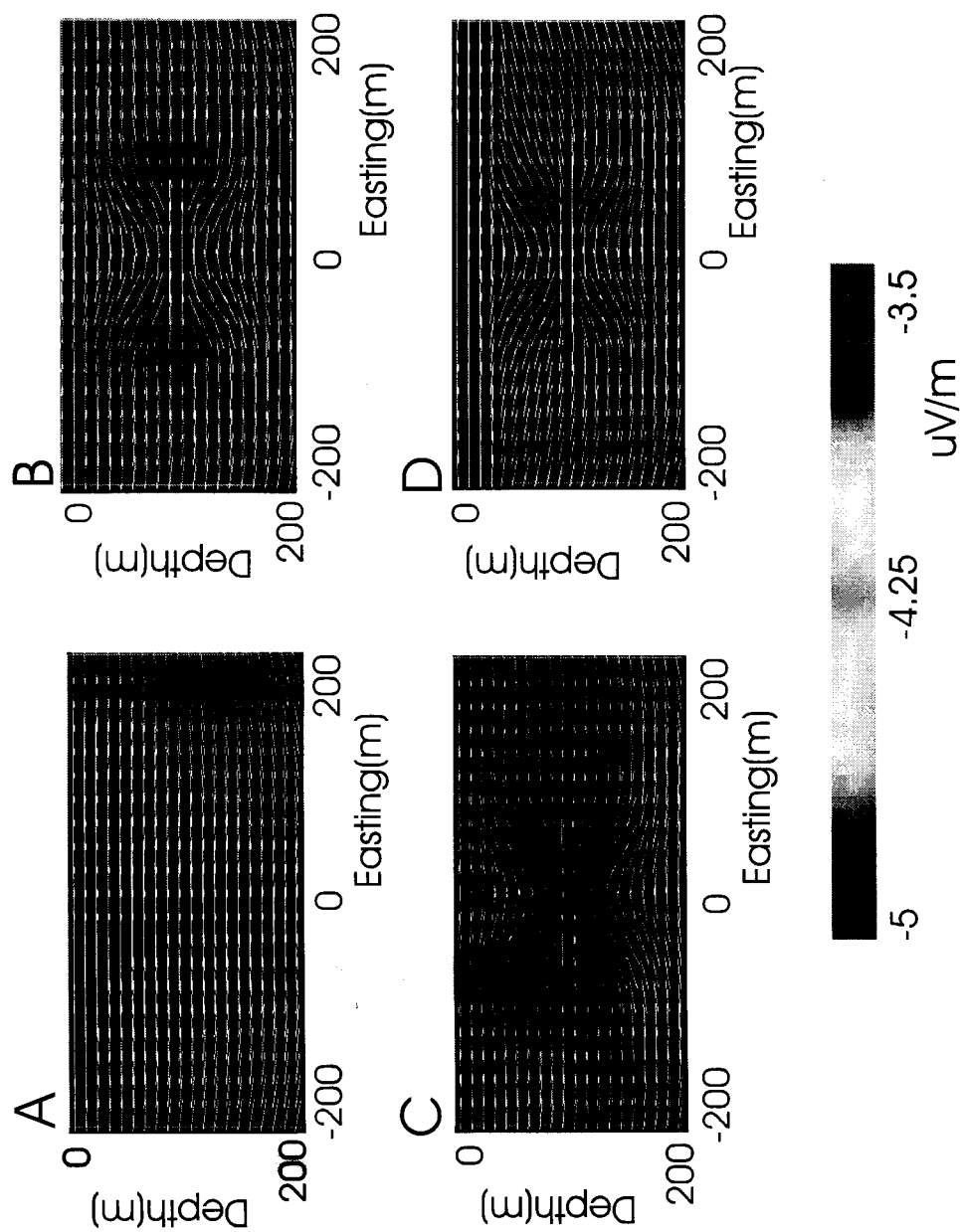


Figure F.4: North-facing cross-section through vector model of real component of the electric field.

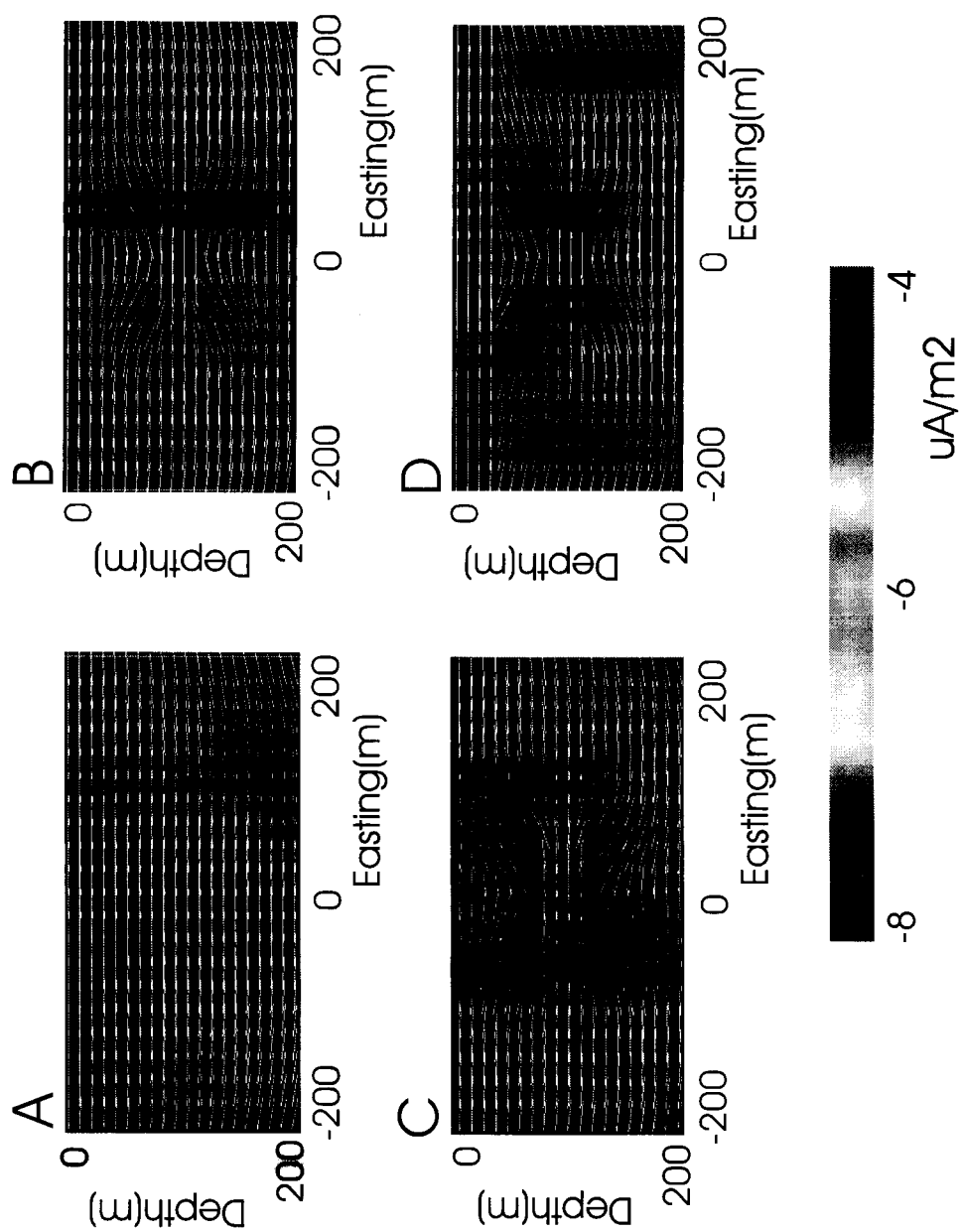


Figure F.5: North-facing cross-section through vector model of real component of the current density.

F.3 Summary

Gradient array IP data can be regarded as a potential field, and can be treated as pseudo-magnetic data if the conductivity structure over which the data are collected isn't anomalous enough to distort the electric field. Conductivity distributions that have contrast of two orders of magnitude still recover a useful chargeability model. Inversion of this data using *MAG3D* can produce good chargeability models as the data is weighted in order to provide depth information. In the case of San Nicolás the conductivity isn't too anomalous although the overburden has a detrimental effect on the data.

Appendix G

IP electrode weighting

When inverting IP data using *DCIP3D*, accumulations of high chargeability values are sometimes observed coincident with the location of the transmitting electrodes. This *electrode-effect* is even more significant when the data have a small number of transmitter electrodes. The data collected at San Nicolás used the gradient-array, and “Real-section” array electrode configurations; both of which have a small number of transmitter electrodes. As a result, the electrode-effect is observed in the inversion models.

This section introduces an ad-hoc weighting procedure that can be used to reduce the electrode-effect. The weighting is invoked using the w_s weighting matrix in Equation 2.2. The value of these weights are simply increased in the vicinity of the electrode. The cells around the electrode will be preferentially weighted closer to the reference model, which is set as zero.

A weighting distribution is sought that smoothly decreases from a high value at the electrodes to a background weighting value of one away from the electrodes. The electrode weighting must effectively suppress charge accumulations around the electrodes but must also decay fast enough so as not to influence the true chargeability distribution.

The following set of conditions was used to empirically develop a weighting function.

$$w_m = \begin{cases} \sum_{n=1}^N ar_{m,n}^{-b} & \text{when } w_m \geq c, \\ 1 & \text{when } w_m < c. \end{cases} \quad (\text{G.1})$$

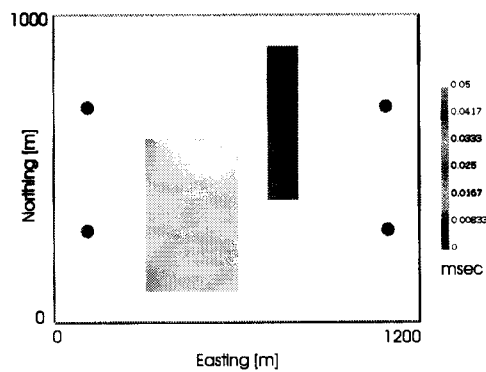
where, w_m is the weighting value assigned to the m^{th} component of the diagonals of the matrix w_s (Equation 2.2), n refers to the n^{th} transmitter electrode, N is the number of transmitter electrodes, $r_{m,n}$ is the distance from the n^{th} electrode to the center of the m^{th} cell, and a , b , and c are all empirically derived constants. a increases the weighting value, b determines the rate at which the weight decays away from the electrode location, and c determines the value the weighting function can obtain before the default weights are used.

The weighting around each electrode is superimposed to form a weighting function that

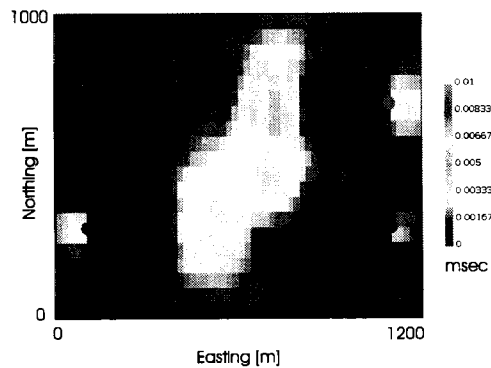
exhibits hemispheres of high values below each transmitter electrode. a 's of one, two, and three were applied to the synthetic model. A value of 100,000 was used for a , and a background value of one was assigned once the weights were less than 5 (the values assigned to c) (this cut-off was used to reduce the size of the volume that is effected by the electrode weights).

Weighting functions were tested for a number of simple synthetic chargeability models with different values of b . A value of 2 produced results that seemed to have the least artifacts. The results of using this weighting function is shown in Figure G.

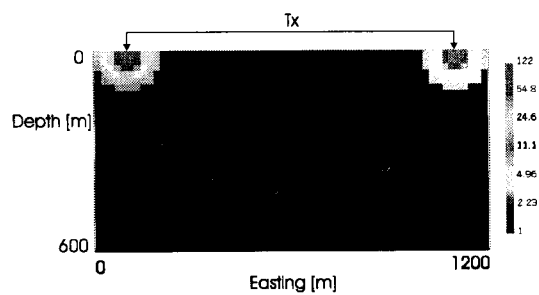
The effect of implementing the weights not only reduces the high values around the electrodes but also increases the magnitude of the anomaly. This increase helps the model achieve values closer to that of the true model. Limitations of this weighting function are obvious. This weighting method might not be appropriate if the electrode were to be placed in chargeable material. Also when the electrode is very close to the center of the cell the method might produce undesired effects. An alternative would be to construct a sensitivity based weighting function. This would make parts of the model less sensitive to the data and therefore allow them to be closer to the reference model.



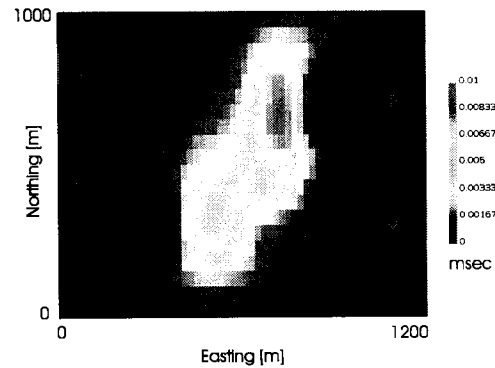
(a) Synthetic chargeability model.



(b) Inversion model



(c) Weights.



(d) Weighted inversion model.

Figure G.1: Removal of electrode-effect via weighting.