

11 maps in Special Collection.

**THE GEOLOGY AND GEOCHEMISTRY OF THE
DAISY CREEK PROSPECT,
A STRATABOUND COPPER-SILVER OCCURRENCE
IN WESTERN MONTANA**

by

Clifford R. Stanley

A.B., Dartmouth College, 1980.

**A thesis submitted in partial fulfillment
of the requirements of the degree of
Master of Science**

in

The Faculty of Graduate Studies

Department of Geological Sciences

**We accept this thesis as conforming
to the required standard:**

**University of British Columbia
November, 1984**

© Clifford R. Stanley, 1984

86

In presenting this thesis in partial fulfilment of the requirements for an advanced degree at the University of British Columbia, I agree that the Library shall make it freely available for reference and study. I further agree that permission for extensive copying of this thesis for scholarly purposes may be granted by the head of my department or by his or her representatives. It is understood that copying or publication of this thesis for financial gain shall not be allowed without my written permission.

Department of Geological Sciences

The University of British Columbia
1956 Main Mall
Vancouver, Canada
V6T 1Y3

Date October 12, 1984

Frontispiece - View of the South Daisy Creek
drainage basin. The occurrence lies on the
left hand fork near the head of the stream.
Stratigraphy dips obliquely to the left
and toward the foreground at
approximately 45°.



ABSTRACT

Stratabound copper, lead, and silver minerals at the Daisy Creek prospect are hosted by crossbedded feldspathic quartzites and orthoquartzites of the Bonner Formation, Missoula Group, Belt Supergroup. Detailed float rock-chip mapping of a 60,000 m² soil grid indicates that mineralization occurs in reduced, coarse grained quartzites, over- and under-lain by fine grained, hematitic quartzites and siltites. Sixty km² regional mapping shows that the prospect lies on the eastern limb of a major broad, open syncline, and that two phases of deformation have folded the strata.

Paragenetic relationships in float rock-chips indicate that early diagenetic pyrite was replaced successively by galena, chalcopyrite, and argentiferous bornite and chalcocite. Later, middle diagenetic oxidation resulted in the formation of cupriferous goethite, barite, and acanthite. Holocene supergene oxidation formed abundant malachite, cerrusite, and pyromorphite from these sulfides and earlier oxidation products.

Soil grid geochemical anomalies of Cu, Pb, Zn, Ag, Ba, and Hg occur as overlapping zones over the prospect. These geochemical zones overlie and correlate well with areas containing different gossan types. The types of different gossan morphologies and anomalous geochemical concentrations suggest that the zones are mineralogically related and that they are characterized by the presence of galena, chalcopyrite, bornite, and chalcocite.

The paragenetic relationships, geochemical soil zoning

geometries, gossan morphologies, and geochemical and mineralogical constraints on the mineralization suggest that sulfide deposition occurred before the onset of silica cementation. Furthermore, it involved the introduction of metal-chloride complexed, oxidized ground waters along an aquifer into locally reduced sediments. Diffusion of copper, and later oxygen, from the oxidizing ground waters into the reduced zone resulted in the formation of Cu-sulfides and their subsequent oxidation.

A new method of statistical analysis involving the regression of stream sediment data from around the Daisy Creek prospect reveals that the prospect can be precisely located using this technique, and that numerous other 'false anomalies' are recognized. Traditional methods for anomaly recognition fall short of providing the information derived from this technique.

TABLE OF CONTENTS

TITLE PAGE	i
FRONTISPIECE	ii
ABSTRACT	iii
TABLE OF CONTENTS	v
LIST OF FIGURES	vii
LIST OF PLATES	ix
LIST OF TABLES	xii
ACKNOWLEDGMENTS	xiii
CHAPTER 1: INTRODUCTION	1
CHAPTER 2: REGIONAL GEOLOGY	7
Structural Geology	8
Stratigraphy and Sedimentology	21
CHAPTER 3: DETAILED GEOLOGY	37
Soil Grid Geology	37
Mineralogy	51
CHAPTER 4: EXPLORATION GEOCHEMISTRY	81
Stream Sediment Geochemistry	81
Soil Geochemistry	105
CHAPTER 5: CONCLUSION	126
BIBLIOGRAPHY	129
APPENDICES: 1 - Description of Stratigraphy	144
2 - Stream Sediment Survey - Geochemical Data ..	161
3 - Stream Sediment Relative Error Plots	170

TABLE OF CONTENTS (CONT)

4 - Stream Sediment Probability Plots	176
5 - Stream Sediment Regression Output	217
6 - Soil Survey - Geochemical Data	231
7 - Soil Relative Error Plots	235
8 - Soil Survey Probability Plots	241

LIST OF FIGURES

	<u>Page</u>
Figure 1. Daisy Creek Area Location Map	2
Figure 2. Fishtrap Creek Syncline	9
Figure 3. Poles to A-C Joints in Bonner and Libby Formations	10
Figure 4. Poles to Bedding in the Bonner and Libby Formations and Cambrian Strata	12
Figure 5. Poles to Bedding for Phase 1 Folds	13
Figure 6. Plunge Orientations for Phase 1 Folds Before and After Phase 2 Deformation	15
Figure 7. Regional Angular Unconformity Between Cambrian and Proterozoic Strata	20
Figure 8. Paleocurrent and Shoreline Directions, Ybn ₄ Member, Bonner Formation	30
Figure 9. Soil Textures From the Daisy Creek Prospect	50
Figure 10. SEM/EDS Spectral Image of Monazite	54
Figure 11. Line Diagram of Paragenetic Mineral Relationships	59
Figure 12. Van der Veer Diagram of Paragenetic Mineral Relationships	60
Figure 13. SEM/EDS Spectral Image of Cupriferous Goethite ..	66
Figure 14. Acanthite Rim/Cupriferous Goethite Rim Thickness Comparison	70
Figure 15. Thresholds for Histogram Distributions	90

LIST OF FIGURES (CONT)

	<u>page</u>
Figure 16. Thresholds for Frequency Distributions	91
Figure 17. Regression Predicted vs. Observed Copper Concentrations - Log Values	97
Figure 18. Regression Predicted vs. Observed Copper Concentrations - Arithmetic Values	98
Figure 19. Soil Grid Geochemical Concentrations: Cu	114
Figure 20. Soil Grid Geochemical Concentrations: Pb	115
Figure 21. Soil Grid Geochemical Concentrations: Zn	116
Figure 22. Soil Grid Geochemical Concentrations: Ag	117
Figure 23. Soil Grid Geochemical Concentrations: Ba	118
Figure 24. Soil Grid Geochemical Concentrations: Hg	119
Figure 25. Soil Grid Geochemical Concentrations: As	120
Figure 26. Soil Grid Geochemical Concentrations: Co	121
Figure 27. Zoning Relationships for Soil Anomalies at the Daisy Creek Prospect	123

LIST OF PLATES

	<u>page</u>
Frontispiece	i
Plate 1. Strong A-C Joints in Ybn ₃ Member of the Bonner Formation	11
Plate 2. Surface Geology of the Daisy Creek Area	<i>in Special (pocket) collections</i>
Plate 3. Generalized Stratigraphy of the Daisy Creek Area	<i>in Special (pocket) collections</i>
Plate 4. Trough Crossbeds in the Ybn ₄ Member of the Bonner Formation	26
Plate 5. Generalized Stratigraphic Cross-Section of the Bonner Formation in the Daisy Creek Area .	<i>in Special (pocket) collections</i>
Plate 6. Generalized Cross-Section of the Bonner and Mount Shields Stratigraphy	<i>in Special (pocket) collections</i>
Plate 7. Outcrop of the Upper Member of the Libby Formation	31
Plate 8. Outcrop of the Lower Member of the Libby Formation	32
Plate 9. Cambrian Flathead Sandstone	34
Plate 10. "Outcrop" of the Cambrian Gordon Shale	35
Plate 11. Redox Facies and Grain Size Distributions of the Daisy Creek Prospect	<i>in Special Collections (pocket)</i>
Plate 12. Sedimentary Facies and Copper Mineralization Distributions of the Daisy Creek Prospect	<i>in Special Collections (pocket)</i>

LIST OF PLATES (CONT)

page

Plate 13. Generalized Bedrock Geology of the Daisy Creek Prospect	<i>in Special Collections</i> (pocket)	
Plate 14. Gossan Type Distributions of the Daisy Creek Prospect	<i>in Special Collections</i> (pocket)	
Plate 15. Cut Surface of Coarse Grained Quartzite from Type D Gossan (Ybn ₄ Member, Bonner Formation		42
Plate 16. Cut Surface of Coarse Grained Quartzite from Type E Gossan (Ybn ₄ Member, Bonner Formation		43
Plate 17. Cut Surface of Coarse Grained Quartzite from Type G Gossan (Ybn ₄ Member, Bonner Formation		45
Plate 18. Soil Color Distributions of the Daisy Creek Prospect	<i>in Special Collections</i> (pocket)	
Plate 19. Phosphatic Mud Chip from Ybn ₄ Member, Bonner Formation		53
Plate 20. Detrital Monazite/Zircon/Bastnaesite		55
Plate 21. Altered and Replaced Detrital Feldspar		57
Plate 22. Chalcopyrite Replacing Cubic Pyrite Form		61
Plate 23. Cupriferous Goethite Replacing Chalcopyrite ..		64
Plate 24. Acanthite Reaction Rim and Silver Element Map		68

LIST OF PLATES (CONT)

page

Plate 25. Cupriferous Goethite Replacing Chalcopyrite with Various Thicknesses of Acanthite Reaction Rims	72
Plate 26. Cupriferous Goethite Replacing Chalcopyrite with Radial Cracks	73
Plate 27. Barite Occurring in Cracks and Rimming Cupriferous Goethite	74
Plate 28. Hinsdalite in Vacant Cavity in Quartzite	76
Plate 29. Hinsdalite	77
Plate 30. Myrmekitic Intergrowth of Limonite and Malachite	79
Plate 31. Regional Stream Sediment Survey, Standard Threshold Selection Techniques	<i>in Special Collection</i> (pocket)
Plate 32. Regional Stream Sediment Survey, Standard Regression/Threshold Selection Technique	<i>in Special Collection</i> (pocket)

LIST OF TABLES

	<u>page</u>
Table I. Detection Limits for Stream Sediment Data	82
Table II. Replicate Samples from Stream Sediment Survey ...	84
Table III. Detection Limits for Soil Grid Data	106
Table IV. Replicate Samples from Soil Grid Survey	108
Table V. Thresholds for Soil Grid Data	113

ACKNOWLEDGMENTS

The author gratefully acknowledges the financial support, technical assistance, and encouragement of Anaconda Minerals Company. Several employees deserve special mention: Mr. John Miller, Dr. Walter C. Reise, Dr. Alan Reid, and Dr. Werner Raab, for their insights regarding the geochemistry data; Mr. Charles Brannon, Mr. Keith Rhea, and Mr. Steve Carlsen, for their assistance in evaluating the geology of the field area; Ms. Julie Flammang, Mr. David Miller, Mr. Gardiner Cross, and Mr. Kevin Hayner, as well as all of the field crew (who are too numerous to mention) for their contributions to sample collection, data acquisition, and sample preparation.

Thanks are also due to Asger Bentzen, for his aid in expediting the learning process of FORTRAN and the UBC computer system. John Newlands drafted and photographed all of the visual aids, and his help is gratefully appreciated.

Special thanks goes to Dr. W.C. Barnes, for his constructive comments and advice regarding the regional structure and sedimentology, and the patience and generosity he showed in visiting the field area.

Finally, appreciation and thanks are extended to Dr. A.J. Sinclair, for his questions, comments, suggestions, and pragmatic philosophy regarding statistical data reduction and the interpretation of geochemical data, for aid in understanding the ore mineralogy, for his ability to play the devil's advocate, and for his guidance and encouragement throughout the study.

CHAPTER 1

INTRODUCTION

Location and History

The Daisy Creek prospect, located in the Cabinet Mountains of Lolo National Forest, northwestern Montana, (Figure 1; 115° 06' W, 47° 48' N; R 27 W, T 24 N, S 30), consists of stratabound copper, silver, and lead minerals contained in the Bonner Formation of the Belt Supergroup. The property lies at 1,470 m elevation and can be reached by driving north out of Thompson Falls, Montana along the Thompson River Road for 23 km, then the Fishtrap Creek Road for 13 km, and finally the Daisy Creek Road for 2 km to the Richard's Peak Trailhead. The property lies 200 m south of the 0.5 km mark on the Richard's Peak Trail.

Potential for significant mineralization in the area was indicated by anomalies from a regional stream sediment geochemistry program conducted by Anaconda Minerals Company during the summer of 1982. Geologic investigation later that fall discovered malachite mineralization in float rock-chip samples with grades of up to 2.0% Cu, 0.3% Pb, and 70 grams Ag per tonne.

Aside from several small, barely recognizable pits located over part of the mineralized zone, and probably dug by prospectors during the 1930's, no evidence of mining or prospecting activity was observed in the area. As a result, a pedogeochemical grid survey was conducted during the fall of 1982

Daisy Creek Area Location Map

Sanders County, Montana

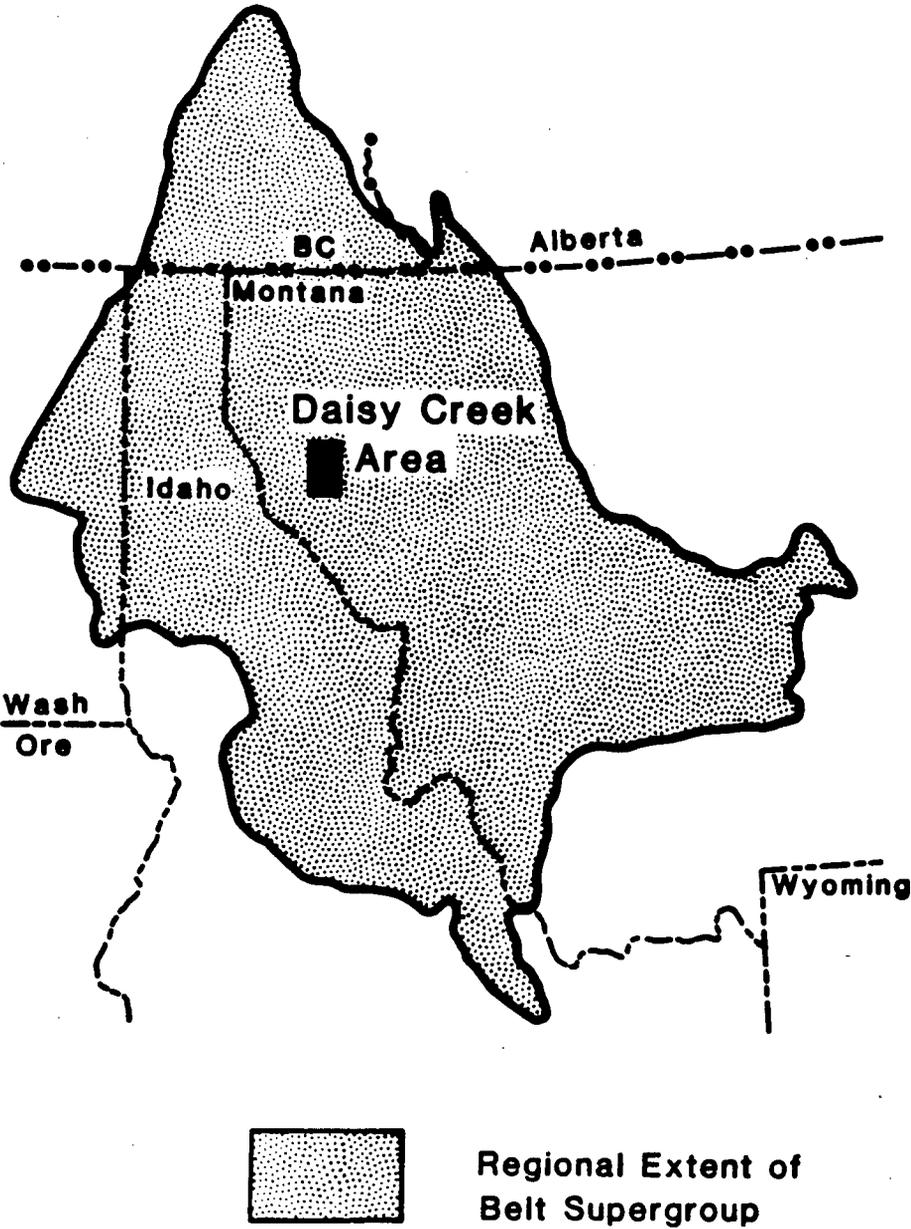


Figure 1

which demonstrated anomalous concentrations of Cu, Ag, and Pb, but failed to bound these anomalies.

Work in 1983, therefore, was devoted to defining the extent of the pedogeochemical anomalies and determining the nature of the mineral occurrence. This work consisted of 60 km² of regional (1:24,000) geologic mapping, 60,000 m² of detailed (1:1200) soil and float rock-chip mapping, and the collection of 736 additional soil samples.

Surface Features

Topography in the Daisy Creek area is rugged, with elevations ranging from 900 to 1,700 m. Slopes are steep (up to 45°) and densely forested. Bedrock generally weathers recessively and outcrop comprises less than 1% of the field area. No outcrops are present within 0.5 km of the mineral showing at Daisy Creek.

Vegetation in the Daisy Creek area is abundant. South facing slopes support mature canopies of ponderosa pine (Pinus ponderosa) at lower elevations, and younger stands of douglas fir (Pseudotsuga taxifolia) at higher elevations. Lodgepole pine (Pinus contorta latifolia) and subalpine fir (Abies lasiocarpa) are found in dense immature stands on north facing slopes, and occur at all elevations. Understory consists of grasses and sedges beneath ponderosa pine, and deadfall and low shrubs beneath douglas fir, subalpine fir, and lodgepole pine. Creek bottoms commonly are choked with deadfall.

On a regional scale, soils generally consist of mature

brunisols to immature luvisols. Profiles consist of a >7 cm thick Lh horizon overlying a >50 cm Bm horizon which, in turn, overlies bedrock. Ae and Bt soil horizons occur sporadically, developing over argillaceous bedrock and alluvial deposits, under grassy parklands, and on shallow to flat slopes. Bedrock lies within 1 to 3 m of the surface except where alluvium is present. No evidence of past glacial activity or glacial deposits occur on the prospect or the 60 km² regional map area.

Due to high snow accumulation during the winter, spring runoff is nearly torrential and most fine detritus is flushed from the higher order stream beds. Runoff generally drops to a trickle by the end of June, though, and silt size material accumulates, making stream sediment sampling possible by July.

Since soils are well drained, low order streams are ephemeral during most of the summer, and country rocks are largely unreactive, physical weathering dominates over chemical weathering, and stream detritus is immature. This is especially true in the 1st and 2nd order streams within 1 km of the head of the drainage, where stream sediment is texturally and compositionally identical to adjacent soils except for a small amount of sorting and winnowing. Higher order streams contain sorted but generally unweathered sands and silts. Clays and Fe- and Mn-oxides are generally not abundant.

General Geology

The mineralized zone lies on a south facing slope and occurs continuously over 200 m of strike length and 15 m of true width. Strata dip 40° SW and strike approximately 165° .

Bedrock consists of moderately deformed, terrigenous, very fine to coarse grained, feldspathic quartzites and ortho-quartzites of the Bonner Formation (Belt Supergroup, Helikian Era) which lie on the eastern limb of a large north-south trending syncline. These sediments were deposited as prograding sands on a deltaic mud flat near the center of the Belt Basin. The Cu-Ag-Pb showing appears to be stratabound and occurs within a coarse grained, crossbedded facies of the Bonner Formation, which generally is more reduced than the surrounding strata.

Minerals of economic interest consist of galena, chalcopyrite, bornite, chalcocite, cerussite, pyromorphite, hinsdalite, cupriferous goethite, acanthite, malachite, chalcantite, chrysocolla, native silver, barite, and limonite. Unweathered sulfide minerals occur in the quartzite, cementing and often replacing detrital grains. Secondary oxide phases replace these sulfides and occur as coatings on weathered surfaces.

Paragenetic mineral relationships, sedimentary evidence, offset soil geochemical anomalies of Cu, Pb, Zn, Ag, Ba, and Hg, and gossan types which form zones across the prospect area support the hypothesis of the sequential replacement of early diagenetic pyrite by galena, Cu-sulfides, and argentiferous

Cu-sulfides. Geochemical evidence suggests Cu- and Pb-sulfide deposition occurred as a result of the interaction of oxidized, metal-bearing ground water with reduced pore water-bearing and/or pyrite-bearing, unconsolidated sediments, some time during diagenesis. As a result, the mineral showing at Daisy Creek is interpreted to be a stratabound sedimentary copper-type occurrence, and to have formed in a manner similar to other Cu-Ag deposits within the Revett Formation in the Western Montana Copper Sulfide Belt.

CHAPTER 2

REGIONAL GEOLOGY

The Daisy Creek prospect, located in the overthrust belt of western Montana, is hosted by Helikian age sedimentary rocks of the Belt Supergroup. In the Daisy Creek area, no intrusive, volcanic, or high grade metamorphic rocks are present. Low grade load metamorphism has reached lower greenschist facies (Winston, 1978). Chlorite and muscovite, which are the major metamorphic minerals, were probably derived from recrystallized detrital micas, clays, and altered feldspars. Metamorphism had little effect on the other minerals in the strata, which are predominantly quartz and hematite.

The following structure, stratigraphy, and sedimentology sections describe the major features of the host rock in the Daisy Creek area.

Structural Geology

The Daisy Creek prospect occurs in the Libby Thrust Zone, one of several east verging imbricate thrust packages which account for the bulk of the structural shortening and displacement in the overthrust region of Western Montana (Price, 1984). Numerous minor imbricate, and probably listric, detachment surfaces, and several major open anticlines and synclines comprise most of the structure in the Libby Thrust Zone (Reynolds, 1984a, b). Fold axis orientations and thrust directions are consistent with a major stress axis oriented at a 070° to 080° azimuth.

The Daisy Creek prospect lies on the eastern limb of one of these major open synclines (Figure 2; from Harrison et al., 1981) with an axis which trends 164° and plunges 12° , based on ac joint orientations (Figure 3; Plate 1). Bedding strikes 162° and dips vary from 65° SW in the northern third of the field area, to 25° SW in the southern third (Figure 4; Plate 2).

Phase 1 Structures

Two phases of deformation are evident in the Daisy Creek area. Phase 1 structures are minor with respect to the overall deformation, taking the form of two tightly folded anticline-syncline pairs which are located at the Snowslip/Wallace Formation contact, and within the Mount Shields Formation (Plate 2). Axes of these folds trend 286° and 281° , and plunge 26° and 31° SE respectively (Figure 5). The steep

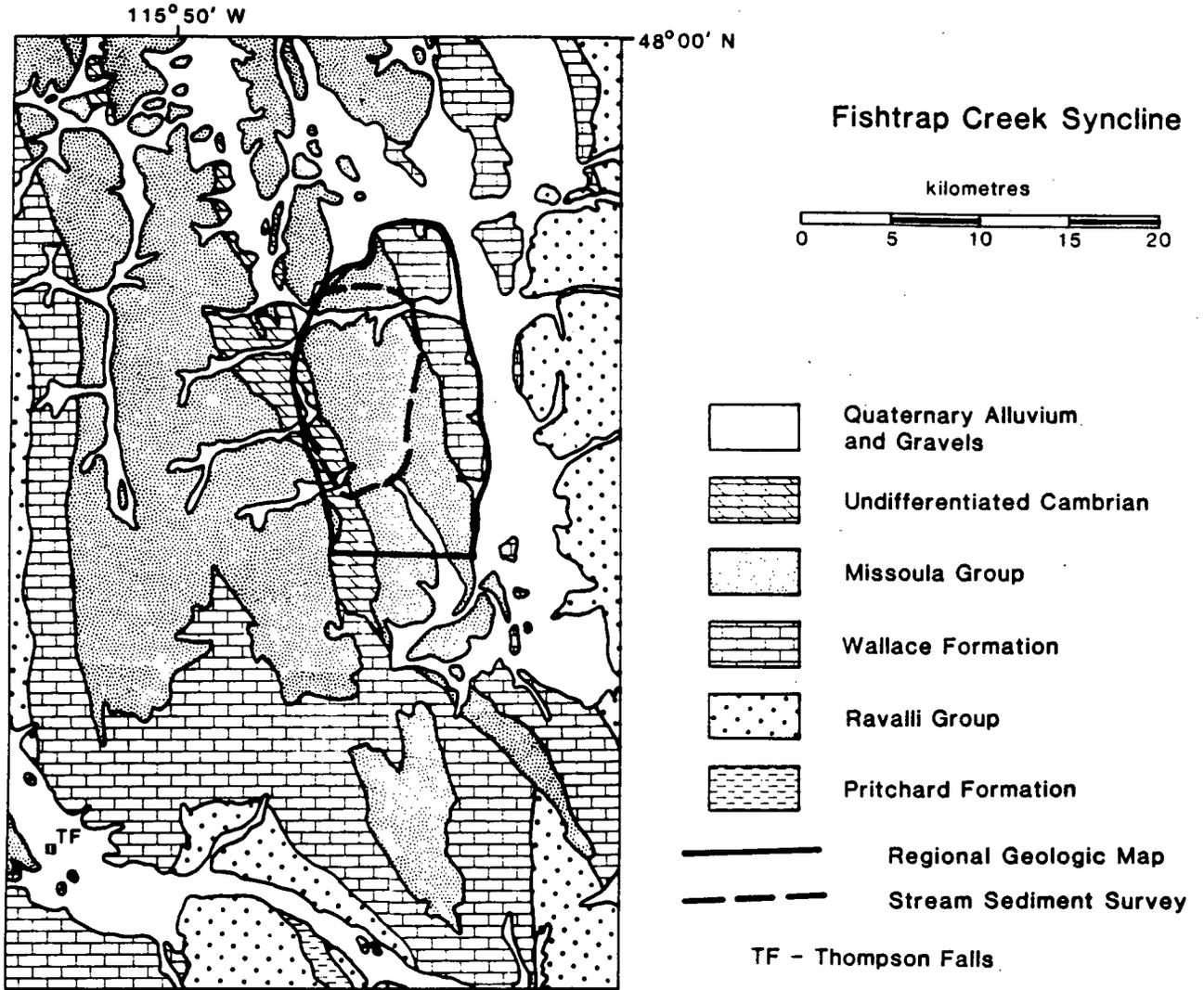
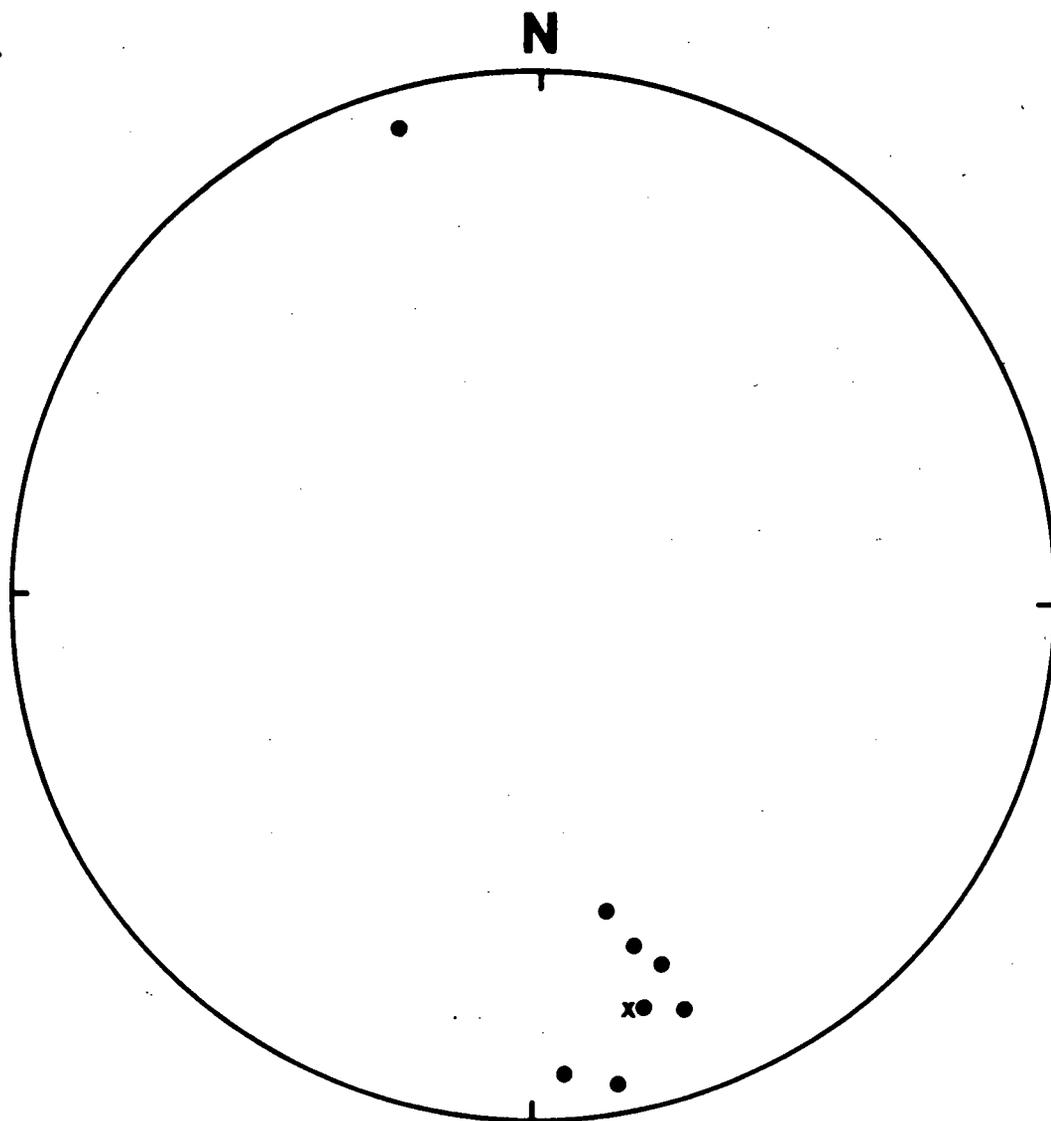


Figure 2

Poles to A-C Joints in Bonner and Libby Formations for Major 164° Trending Syncline



- Poles to A-C Joints (n=8)
- x mean plunge of Synclinal Axis

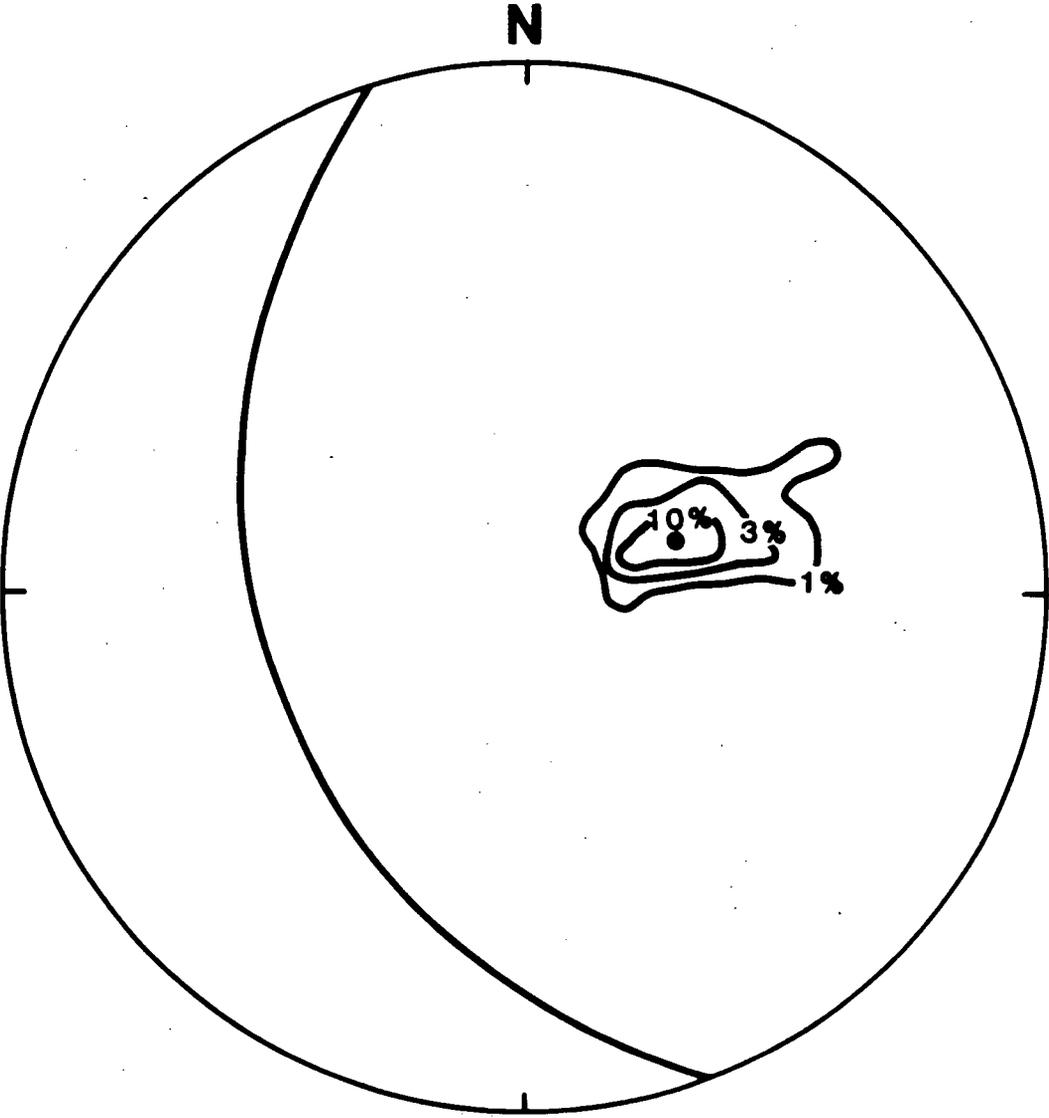
Figure 3

Plate 1

Strong ac joints in Ybn₃ member
of the Bonner Formation, a hematitic,
flat laminated, fine grained quartzite.
(outcrop is approximately 5 m wide)



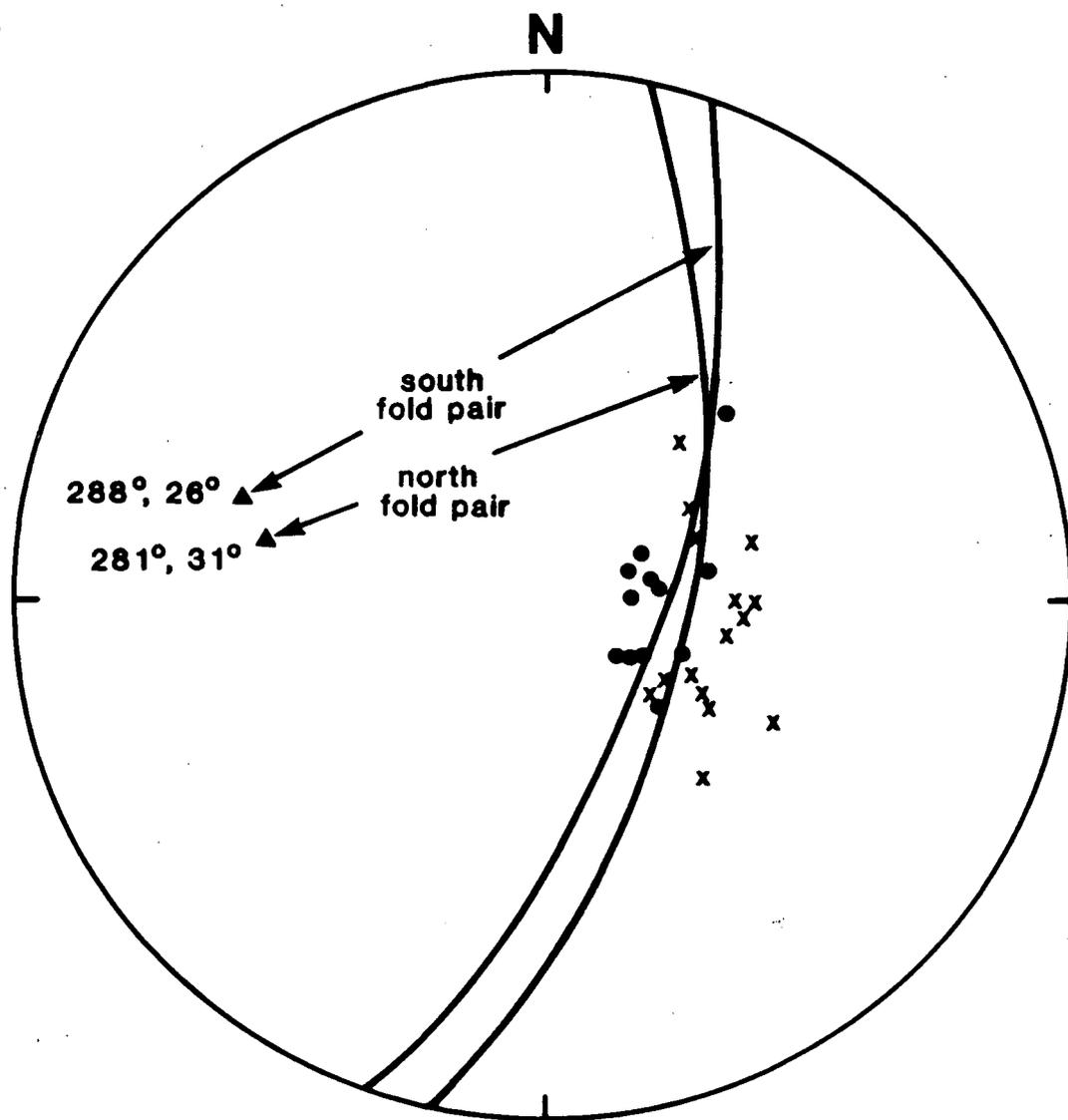
Poles to Bedding in the Bonner and Libby Formations and Cambrian Strata



(n = 148)

Figure 4

Poles to Bedding for Phase 1 Folds



- poles to bedding for the north fold pair (n = 12)
- x poles to bedding for the south fold pair (n = 16)
- ▲ fold plunge

Figure 5

plunge of these fold axes, which lie sub-parallel to the west-dipping bedding surfaces constituting the east limb of the major syncline described above, suggests that they have been folded. This probably occurred during phase 2 deformation, which caused the development of the 164° trending syncline. Prior to phase 2 deformation, the axes of these phase 1 folds were sub-horizontal and had trends of 283° and 277°, respectively (Figure 6).

Minor parasitic fold geometries on the limbs of phase 1 anticline-syncline pairs indicate SW vergence. These minor folds are observed only in carbonate cemented beds of the Wallace Formation and the Mixed Clastic Unit of the Mount Shields Siltite Member. This phenomenon may be related to the fact that, under stress, carbonate cement can recrystallize and twin more readily than silica cement. This would allow carbonate cemented beds to deform ductily, exhibiting folds, whereas silica cemented beds deform brittly under the same stress conditions.

Phase 1 deformation has not been reported previously in this area of the Belt Basin; however, both the pre-thrusting orientation of phase 1 structures (280°), and the present trend of these fold axes (285°), are sub-parallel to the orientation of the Lewis and Clark Line, a major E-W dextral fault system which crosses Western Montana and Northern Idaho, located 50 to 70 km to the SW of the field area. No time constraint for phase 1 deformation is suggested, except that it occurred before phase 2 deformation, and thus is Mesozoic or older.

Plunge Orientations for Phase 1 Folds before and after Phase 2 Deformation

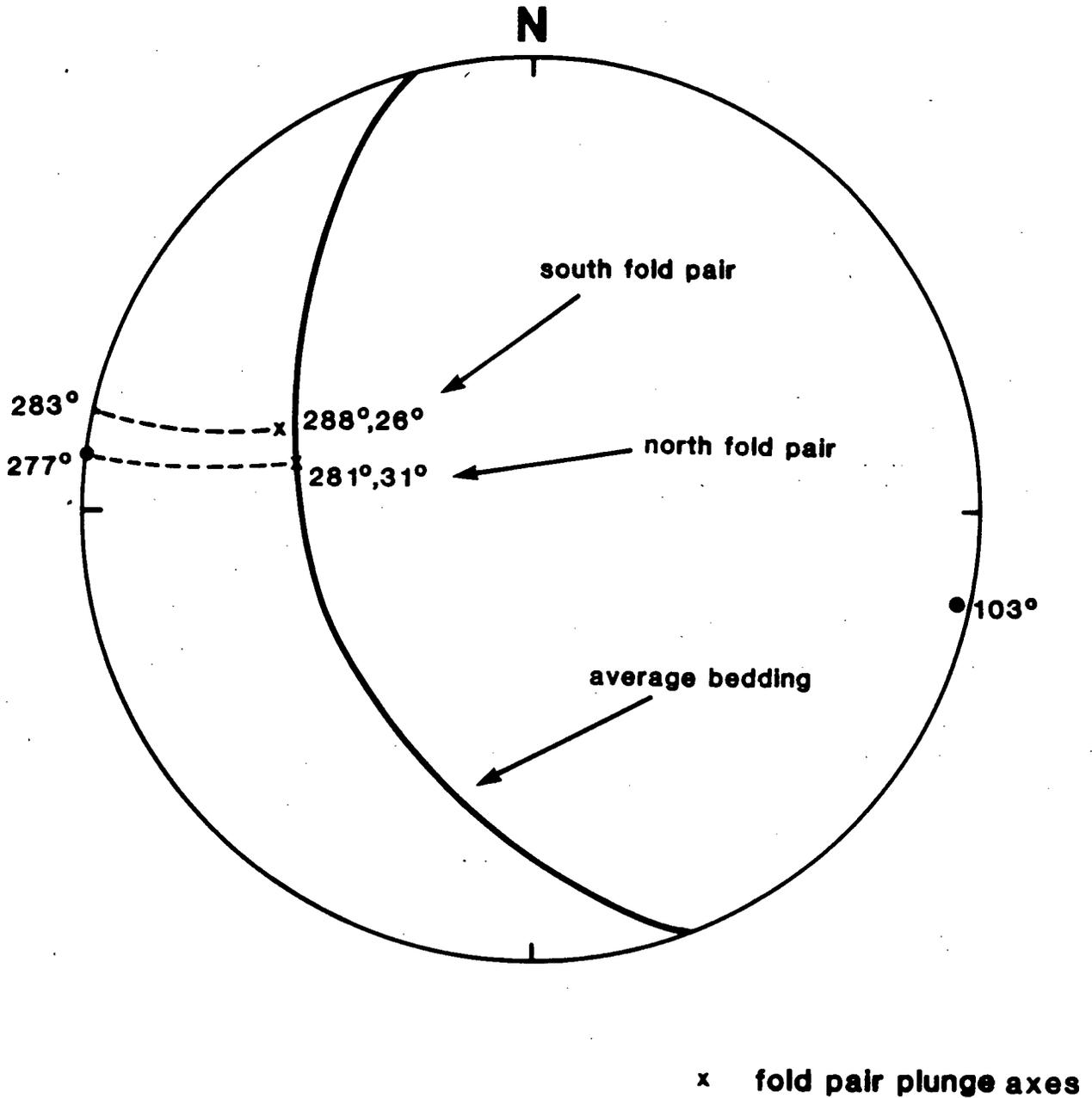


Figure 6

Phase 2 Structures

Phase 2 structures in the Daisy Creek area reflect the development of the Libby Thrust Zone. This phase of deformation is the most prominent in Western Montana and is related to overthrusting of Proterozoic (Belt) and early Paleozoic sediments over Mesozoic and Cenozoic strata (Wheeler, 1983). Phase 2 deformation was active from Jurassic through early Tertiary time. The structural style resulting from this phase of deformation is manifested locally in tight, closely spaced anticlines and synclines and shallow, low angle thrusts in the Daisy Creek area, and regionally, as open, broadly spaced anticlines and synclines and imbricate thrusts.

Structures in the Daisy Creek area related to phase 2 deformation include (Plate 2):

- 1) a tight, closely spaced anticline-syncline pair (trending 338° ; horizontally plunging) at the north end of the field area located in the Snowslip Formation;
- 2) a broad, widely spaced anticline-syncline pair (trending 327° ; doubly plunging) and a minor thrust fault (172° strike, 18° SW dip) located in the Mount Shields Formation, also at the north end of the field area, which constitutes a parasitic fold and thrust set on the eastern limb of the major syncline; this fold and thrust set indicates vergence to the NE; and

3) a homoclinal succession of west-dipping strata which constitutes the eastern limb of the major 164° trending syncline represented throughout the field area; the synclinal axis plunges 8° to the south, based on ac joint measurements.

All of these structures have orientations consistent with a major stress axis trending approximately 075° , which is comparable to that responsible for the development of the Libby Thrust Zone.

Other Structures

Numerous 100° striking, steeply dipping (commonly 80° to the north) quartz veins are recognized in the Daisy Creek area. These veins formed as syntaxial quartz infillings of either, expansion or radial joints during a post-phase 1 relaxation event (their strike is perpendicular to the phase 1 major stress axis), or, more likely, conjugate shear planes developed during phase 2 deformation (their strike is 25° from the phase 2 major stress axis). An obscure fibrous texture and minor brecciation with rotation of fragments is common in these veins. In addition, the veins contain hematite, chlorite, and barite, mimicing the mineralogy of the country rocks they cut (Appendix 1; Bonner and Libby Formations).

The quartz veins are observed cutting only the Libby and Bonner Formations. This may suggest that these formations have the highest tensile strength within the field area and thus deformed brittlely when one of the above deformation events

occurred. In contrast, closely spaced parasitic folds and thrusts formed during phase 2 deformation and closely spaced folds comprising phase 1 deformation occur only in the Mount Shields, Snowslip, and Wallace Formations. These formations probably have a lower overall strength than the overlying Bonner and Libby Formations, and thus deformed ductily under stresses which did not noticeably deform the Bonner and Libby Formations.

Finally, two major dip-slip faults cut the field area (Plate 2). A steeply dipping fault strikes 355° and is downdropped on the west approximately 210 m, based on marker horizon offset, regional bedding orientations, and the assumption of 100% dip-slip displacement. The fault occurs in the north part of the field area, and is accompanied by abundant deformation along a zone marginal to the fault. This deformation takes the form of tight, closely spaced antiforms and synforms, with attitude changes of over 90° in less than 10 m, numerous normal and reverse faults, and abundant bedding plane slips.

A second, steeply dipping fault occurs at the south end of the field area. It strikes 70° and is downdropped approximately 135 m (assuming 100% dip-slip displacement) on the south side. Distinct folding adjacent to the fault in the Libby and Bonner Formations is structurally consistent with dip-slip movement along the fault.

Minor malachite-stained reverse faults, located in the Bonner Formation adjacent to this dip-slip fault are oriented parallel to ag joints related to the phase 2 syncline. This suggests that this major fault is a reverse fault with a 70° northerly dip.

Neither of these faults can be related conclusively to either of the deformation phases which are observed in the Daisy Creek area; however, both appear to have undergone significant relative movement since the end of phase 2 deformation.

Regional Angular Unconformity

Separate analysis of bedding orientations for Cambrian and Proterozoic rocks indicates that a minor angular unconformity occurs between these rock packages. Rotation of the mean Cambrian bedding orientations in the west-central portion of the field area to horizontal demonstrates that, during Cambrian deposition, the underlying Proterozoic sediments had a strike of 162° with an 8° SW dip (Figure 7). Similar rotation of bedding orientations in the northwest and southwest parts of the field area yield similar results.

Regional Angular Unconformity between Proterozoic and Cambrian Strata

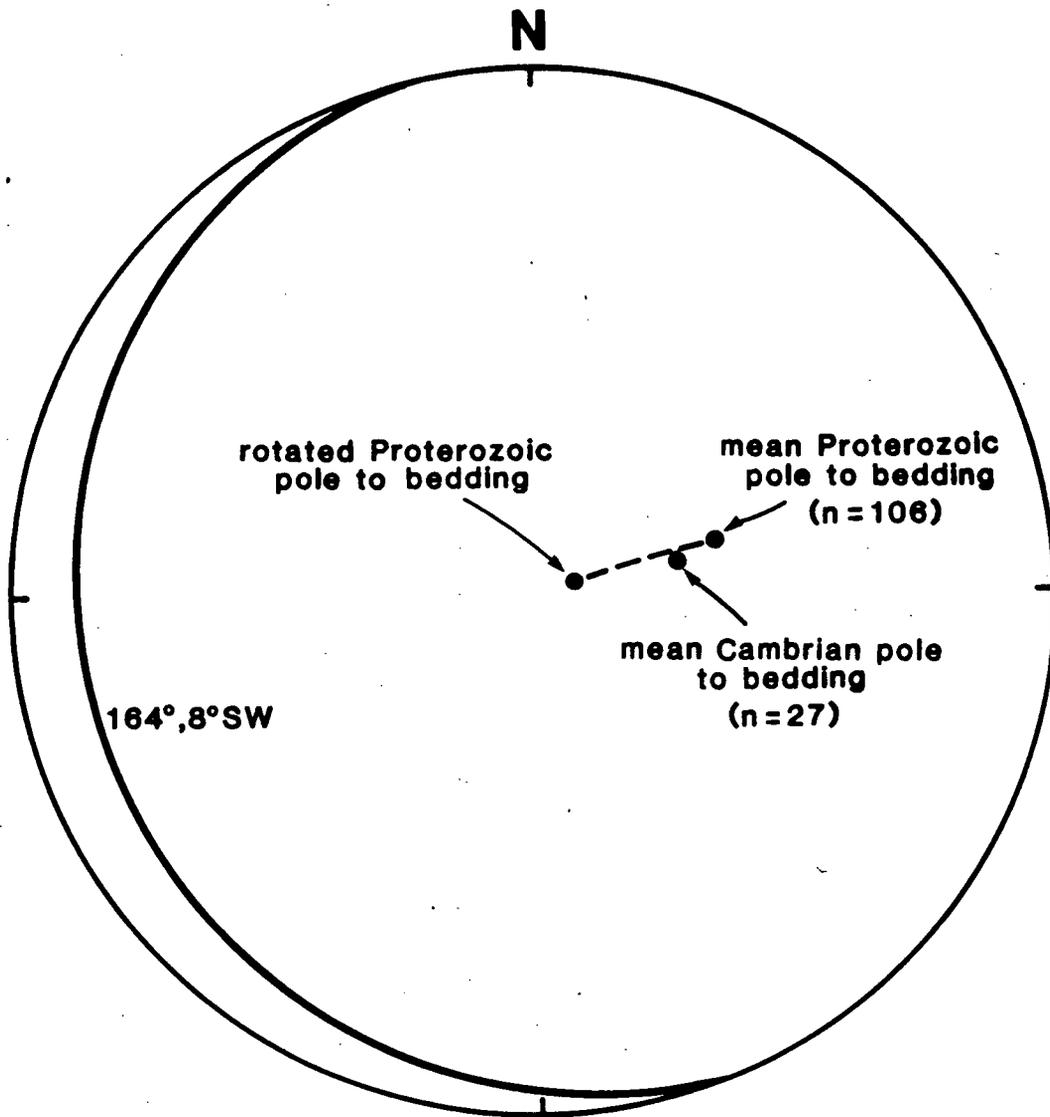


Figure 7

Stratigraphy and Sedimentology

Bedrock in the Daisy Creek area is entirely sedimentary. Two major lithologic packages are present. Helikian fluvial, deltaic, and basinal sediments over 4,200 m thick comprise the Missoula Group and underlying Middle Belt Carbonate of the Belt Supergroup (Johns, 1970; Croweley, 1972). These sediments occur below at least 550 m of a miogeoclinal Cambrian sequence (Lochman, 1957). A stratigraphic column is presented in Plate 3 and a referenced description of this stratigraphy is included in Appendix 1.

Belt Supergroup

In the Daisy Creek area, sedimentary deposition during Belt time appears to have occurred in largely intermittent pulses on a vast alluvial plain or fluvio-deltaic complex prograding into a major basin (Winston et al., 1977). This basin was probably surrounded by granitic and gneissic terranes of Archean and Hadrynian age. Detrital input apparently occurred during flash floods or seasonal monsoons because mud crack casts, mud chip conglomerates, and fining upward sedimentary sequences are common (Winston, 1984a). Most beds in the Daisy Creek area are fine grained and probably were located distal to flood sources; thus waters did not have enough energy to significantly scour previously deposited beds.

Belt sedimentation took place in a semi-arid environment (Winston et al., 1977) and, as paleo-magnetic data indicate, at latitudes less than 25° (Elston, 1983; Obradovich et al., 1984; Harrison, 1984b). Very little tectonic activity occurred during the bulk of Belt time, except at the edges of the basin, where boundary faults were active intermittently, perpetuating basinal subsidence and sedimentation. Syn-sedimentary faults also presumably occurred at the edge of the sedimentary prism.

Igneous activity was uncommon, consisting of only the Moyie Sills, which intrude the Pritchard/Aldridge Formation, and the Purcell Lavas which occur in the Shepard and Snowslip Formations (Harrison, 1972). Neither of these igneous units occurs in or near the field area. Units older than the Wallace Formation are not present in the Daisy Creek area and are not described in this report.

Middle Belt Carbonate

Wallace Formation

The only part of the Wallace Formation exposed in the Daisy Creek area is the upper part of the Middle Member. It is composed of dark grey to black, laminated, calcareous argillite and siltite, interbedded, in a very distinctive wavy form, with white dolomitic quartzite. The Upper Member, as described by Harrison (1984a), resembles the lower Green Argillite Member of the Snowslip Formation (Ysn₁; Appendix 1), but is not defined as

such in this report because of the lack of abundant carbonate cement.

The depositional environment for Wallace sedimentation consisted of quiet, shallow water carbonate mud-flats and shoals dominated by carbonate precipitation. These shoals contain convolute bedding, water escape structures, and ripple marks. Significant fine clastic input with a southern provenience also occurred during this time (Wallace et al., 1984).

The basin also contained minor local sub-basins (Wallace et al., 1976; Godlewski, 1977; Harrison, 1984a) which were sites for soft sediment slumping and sedimentary breccia development along the edges of these local depressions. No slumping features of this sort are recognized in the Daisy Creek area.

Missoula Group

The Missoula Group has been divided into three assemblages based on sedimentologic characteristics, lithofacies, source areas, and dispersal systems (Wallace et al., 1984). These consist, in the Daisy Creek area, of a lower assemblage composed of the Snowslip Formation, a middle assemblage composed of the Mount Shields and Bonner Formations, and an upper assemblage composed of the Libby Formation.

Snowslip Formation (Lower Assemblage)

The Snowslip Formation contains dominantly thin bedded, red and green argillites to very fine grained quartzites (Harrison and Campbell, 1963). Depositional environments for the upper Purple and Green Argillite Member (Ysn₂; Appendix 1; intercalated red and green beds) consisted of mud flats and shallow distributary channels on a vast sedimentary plain. Local lacustrine environments perpetuated sub-aqueous conditions and caused the formation of green beds due to the preservation of ferrous iron in chlorite, while sub-aerial exposure resulted in red beds, caused by the oxidation of iron and the development of hematite. Sub-aerial exposure of the sediment created desiccation cracks, and caused the development of mud chip conglomerates during later floods.

Green beds of the lower Green Argillite Member (Ysn₁; Appendix 1) were deposited in shallow water with rare sub-aerial exposure. Desiccation cracks did not develop because the sediment was always wet, and thus no mud chip conglomerates were formed.

Lithofacies changes across the basin suggest that the provenance for the Snowslip Formation was from the Hudsonian age Canadian Shield to the east and northeast (Wallace et al., 1984).

Mount Shields and Bonner Formations (Middle Assemblage)

The middle assemblage of the Missoula Group is significantly coarser grained than the underlying lower assemblage (Winston, 1984a, b). In the Daisy Creek area, the Bonner Formation and underlying Mount Shields Formation consist of two major packages of prograding sediments, one deposited in supra-tidal, fluvial and delta-plain mud flat environments (Bonner Formation), and the other deposited dominantly in inter-tidal, strandline and possibly sub-tidal, pro-delta environments (Quartzite Member, Mount Shields Formation). These packages are separated by a series of transgressive sediments comprising the Siltite Member of the Mount Shields Formation.

The Bonner Formation consists of trough and epsilon crossbedded, coarse grained quartzites (braided channels; Plate 4) and flat to ripple laminated, fine grained quartzites (inter-channel mud flats). These inter-channel mud flats were temporarily covered by shallow waters, as suggested by abundant ripple marks, but normally dried quickly and were exposed to atmospheric oxidation, because green beds are rare (Plate 1).

Facies and facies thicknesses change rapidly along strike (Plate 5), and many mappable units thin and thicken appreciably along the eight kilometers of strike length in the field area. These variations are common in fluvial environments due to channel migration during stream evolution, and due to local variations in topography.

The two crossbedded members of the Bonner Formation (Ybn₂

Plate 4

**Trough crossbeds in Ybn₄ member of
the Bonner Formation, a vitreous,
coarse grained quartzite.**



20 a

and Ybn₄; Appendix 1) represent two minor but separately recognizable periods of progradation, the first extending further to the north than the second (Plate 6; from Winston et al., 1977, and Newton, 1982). Each is correlatable with a probable time-equivalent conglomeratic facies which occurs in the southern part of the Belt Basin. These may represent periods of greater uplift in the highlands surrounding the Belt Basin and/or periods of larger clastic input from these highlands into the Belt Basin.

The Mount Shields Formation consists of two members (Winston et al., 1977). The upper Siltite Member (Ymss; Appendix 1) consists of purple and green transgressive siltites and argillites and overlies the lower Quartzite Member (Ymsq; Appendix 1), another sequence of prograding sediments which is slightly finer grained than the Bonner Formation.

Deposition of the Siltite Member (Ymss; Appendix 1) occurred in shallow water, which often dried or drained away. Occasionally, carbonate cemented the sediments. Green beds were deposited in more permanent waters. Red beds may have been deposited in shallow standing water that soon dried, resulting in desiccation cracks and mud chip conglomerates when later flooding occurred. The waters were not totally quiescent during sedimentation, because the existence of asymmetric tuning fork and undulatory ripple marks suggests that some sort of current existed. These ripple marks may have been caused by currents related to the waning floods which introduced the sediment, but more likely, were caused by the waters draining from the flood plain well after sediment deposition.

Deposition of the lower Quartzite Member (Ymsq; Appendix 1) appears to have been substantially different in form than the Bonner Formation. These differences are especially evident in the Pyritic Quartzite Unit (Ymsq₃; Appendix 1). No trough or epsilon crossbeds were noted in this unit, suggesting that it was not deposited in a classical fluvial environment. Instead, only planar and tabular-tangential crossbeds occur. In addition, abundant authigenic pyrite occurs in this unit. These features suggest that the quartzite was deposited as sub-aqueous sheet sands and bars and that it did not dry out significantly after sedimentation. Conditions of this type can only be found in inter-tidal strandline or sub-tidal pro-delta environments.

The Hellgate Quartzite Unit (Ymsq₁; Appendix 1), which is the facies equivalent of the Hellgate Quartzite Formation located further to the south (Harrison, 1977), also exhibits crossbeds similar to the Pyritic Quartzite Unit. However, the pink color of its beds suggests that it either dried out after subaqueous deposition, or went through redoxomorphic oxidation during diagenesis where Fe^{II} was oxidized to Fe^{III}.

The underlying Lower Green Siltite Unit (Ymsq₂; Appendix 1) of the Quartzite Member was also deposited in a largely subaqueous environment, with little sub-aerial exposure after sedimentation because no desiccation cracks or mud chip conglomerates occur. Thus, all units of the Quartzite Member of the Mount Shields Formation appear to have been deposited in a sub-aqueous inter-tidal strandline and/or sub-tidal pro-delta environment.

The fluvial systems which formed the Bonner and Mount Shields Formations had a granitic source located to the southwest, on 'Belt Island', and on the Dillon Block, south of the Willow Creek Fault Zone (Winston et al., 1977). Limited paleo-current measurements from the Daisy Creek area indicate paleo-transport directions toward 334° (Figure 8) with a shore line roughly perpendicular to this orientation (064°).

Libby Formation (Upper Assemblage)

The Libby Formation (upper assemblage) consists dominantly of mint blue to olive green chert-bearing argillite and siltite. This unit was deposited in shallow water but sediments rarely were exposed to the atmosphere. Mud chip conglomerates occur but these may have been transported into the shallow water from sub-aerial mud flats nearby. Most mud chips are not rounded, suggesting a very proximal source. Available evidence indicates that the Upper Member (Y1bu; Appendix 1; Plate 7) was deposited in shallower water than was the Lower Member (Y1bl; Appendix 1; Plate 8). Both members show some evidence of periodic desiccation; however, the more common desiccation cracks and red beds in the Upper Member indicate that drying was a more frequent phenomenon there than in the Lower Member.

During deposition of this upper assemblage, the sediment source was from the south (Wallace et al., 1984; and based on lithofacies data suggesting finer grained and thinner laminated sediments toward the north - from exposures in the Daisy Creek

Paleocurrent and Shoreline Directions Ybn 4 Member, Bonner Formation

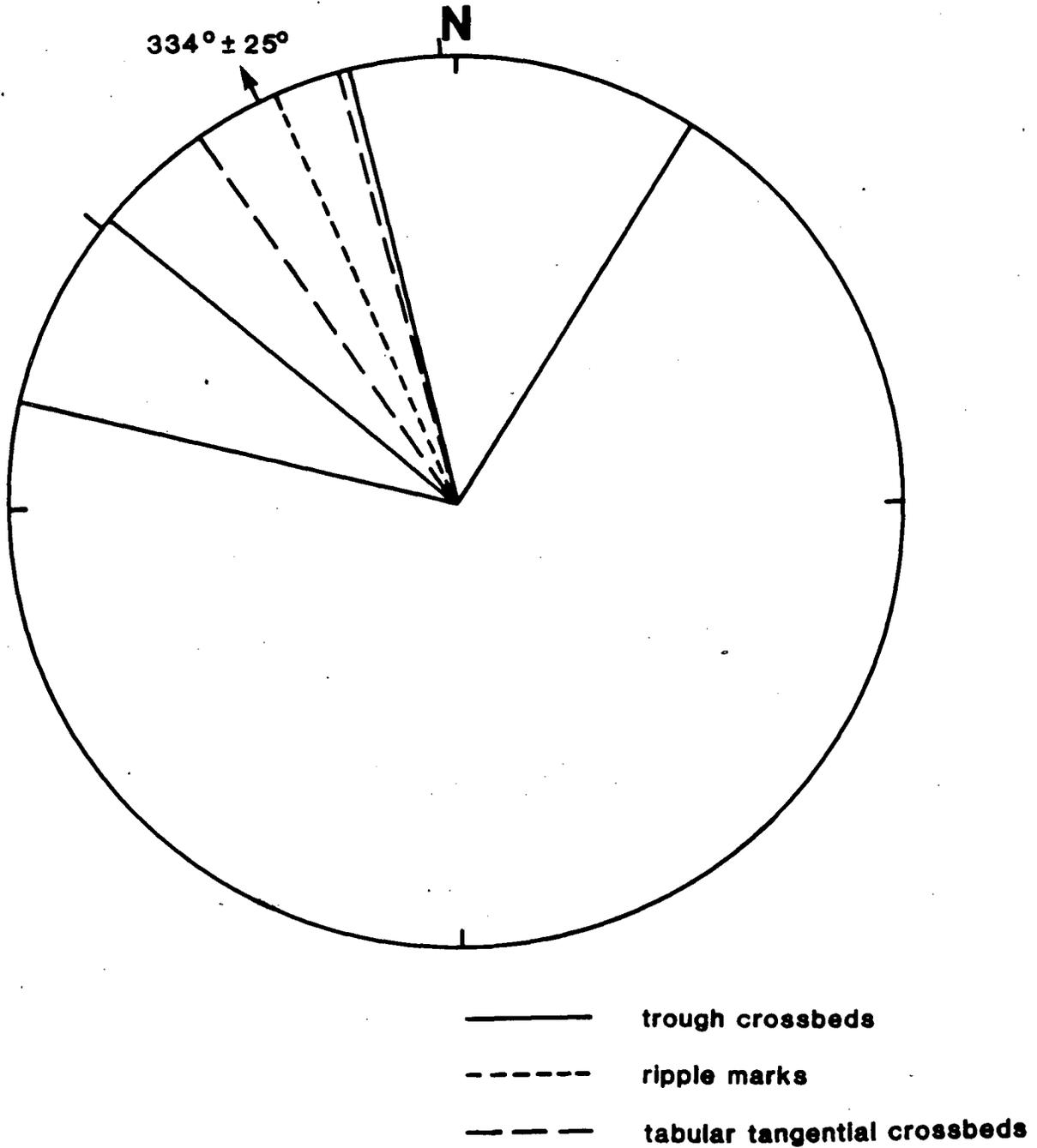


Figure 8

Plate 7

**Outcrop of upper member of the
Libby Formation (Y1bu). Note
swaley bedding and brownish black
Fe-oxide stains.**



Plate 8

Outcrop of lower member of the
Libby Formation (Y1b1). Note
strong ac joints and pale
tan weathering.



area and around Libby, Montana).

Middle Cambrian Sediments

Above the Libby Formation, and separated by a low angle regional unconformity, lies a transgressive sequence of Middle and Upper Cambrian clastic and carbonate rocks. These rocks occur on the west side of the field area. At the base of this sequence is the Flathead Sandstone, a trough and epsilon crossbedded, coarse grained, quartz- and jasper-rich sandstone overlying a flat laminated, fine grained, hematitic sandstone (Plate 9), deposited in braided channel and inter-channel mud flat environments, respectively, at the edge of a marine basin (Lochman, 1950). The Flathead Sandstone is similar to the Bonner Formation, except that no evidence exists suggesting that the Belt Basin was marine at the time of Bonner deposition.

Overlying the Flathead Sandstone lies the Gordon Shale, a fissile, fossiliferous shale deposited in a marine basin of intermediate depth (Plate 10; Lochman, 1950).

Finally, the Damnation and Dearborn Dolomites overlie the Gordon shale and cap the Middle Cambrian section in the field area. Both of these sandy dolomites contain shaly interbeds.

These formations underlie the rest of the Middle Cambrian, Upper Cambrian, and Devonian section, which is not exposed in the Daisy Creek area, and represent a major Paleozoic transgression which onlapped onto North America from the west, commencing during Middle Cambrian time. These sediments were deposited on

Plate 9

Cambrian Flathead Sandstone (Cfs).
Note maroon 'inter-channel mud flat'
facies interbedded with white, crossbedded
'channel facies'.



Plate 10

"Outcrop" of Cambrian Gordon Shale (Cgs).



- 35 a -

top of a partially eroded surface (regional angular unconformity with an 8° dip) of the Belt Supergroup, and are the youngest rocks exposed in the area.

CHAPTER 3

DETAILED GEOLOGY

Soil Grid Geology

The Daisy Creek prospect consists of galena, chalcopyrite, bornite, chalcocite, cupriferous goethite, malachite, acanthite, cerussite, pyromorphite, hinsdalite, chalcanthite, chrysocolla, native silver, and barite disseminated erratically as blebs in medium to very coarse grained, vitreous feldspathic quartzites and orthoquartzites of the Bonner Formation which strike 162° and dip 40° SW. A detailed float rock-chip map was generated by evaluating the average lithologic characteristics at soil sample sites located on a square grid 15 m on a side. The soil grid, covering over $60,000 \text{ m}^2$, is irregular in shape and extends beyond the known mineralized limits. A minimum of 15 rock-chips, greater than 5 cm on a side collected from each soil sample hole, were used to evaluate the geologic characteristics of each site.

Float Rock-Chip Maps

Soil float rock-chip maps are presented in Plates 11 and 12. Results demonstrate a close correlation between the mineral showing (Chapter 4) and several geologic features. These include grain size, sedimentary bed forms, and redox potential. These geologic characteristics allow the subdivision of the Bonner

Formation into several members, as described in Appendix 1, and depicted in Plate 3. A summary map of these members as they occur on the soil grid is presented in Plate 13. Copper minerals appears to be contained entirely within the Ybn₄ member of the Bonner Formation.

Two other stratabound Cu-Ag prospects in or immediately adjacent to the Ybn₄ member are known to occur in southwestern Montana. The Pacman prospect, located in the Southern Pioneer Mountains, and the Bluebird/Mona Lisa prospect, located at the west end of the Anaconda-Pintlar Range, consist of chalcocite and minor bornite which cement and partially replace detrital grains. Host lithologies for these prospects are similar to that at Daisy Creek, although they are significantly coarser grained (Newton, 1982).

In addition to the sedimentologic criteria at the Daisy Creek prospect described above, a distinct zoning pattern of gossan types can be observed across the deposit. Weathered disseminated authigenic sulfide and ferroan carbonate blebs, now converted to Fe-oxide phases (limonites), exhibit distinct color and form differences which are correlatable with original sulfide minerals. These are distributed in zones across the deposit and surrounding strata (Plate 14).

Disseminated limonite blebs occur, as gossans, in all members of the Bonner Formation (Ybn₁₋₅; Appendix 1), as well as the Purple and Green Argillite Unit (Ymss₄; Appendix 1) and Pyritic Quartzite Unit (Ymsq₃; Appendix 1) of the Mount Shield Formation. Descriptive terminology has been borrowed from

Blanchard (1968), Whitten (1965), Ney et al. (1976), and Anderson (1982). Limonites from the finer grained red beds, which are flat to ripple laminated and thin bedded (Ybn₁, Ybn₃, Ybn₅, and Ymss₄; Appendix 1; type A), are indigenous and only minor constituents of the host rock. They are orange in color, appear to be powdery, amorphous aggregates, and completely fill the up to 5.0 mm wide cavities in which they are found.

In contrast, limonite from the coarser grained crossbedded units of the Bonner Formation (Ybn₂ and Ybn₄; Appendix 1; type B) are abundant and red in color, indigenous, appear crystalline in form, and only dust the inside surfaces of the up to 1.0 cm wide cavities they occupy.

Limonite from the Pyritic Quartzite Unit (Ymsq₃; Appendix 1; type C) appears to be a weathering product of the 1.0 mm authigenic pyrite cubes found in this strata and is very different from the disseminated limonite found in the other units and members. It is orange in color, slightly porous, retains the original euhedral pyrite cubic form, is 1.0 to 3.0 mm in width, amorphous, indigenous, and generally is rimmed by red fringing and exotic limonite which is finely disseminated in all directions up to 1.0 cm away from the bleb.

The orange and red indigenous disseminated limonite blebs, found in both thin bedded, flat laminated, fine grained quartzites and crossbedded, coarse grained quartzites, occur ubiquitously throughout the Belt Basin and may be weathering products of authigenic ferroan calcite or ferroan dolomite which act to cement grains in the terrigenous clastics. The color

difference may be related to the permeability of the respective units; the coarser grained units have higher permeabilities and thus are exposed to supergene oxidation for a longer period of time during the erosion process than the finer grained beds because oxygen bearing waters can circulate to greater depths in these porous units. This would allow time for limonites from the coarser grained beds to mature through recrystallization and dehydration to more crystalline forms of Fe-oxide, specifically hematite, whereas limonites from the finer grained beds would be relatively immature and consist of poorly crystalline ferric hydroxide and goethite. In addition, the high permeability of the coarser grained rocks would allow the removal of most of the iron in the coarse grained beds, leaving only hematite dusts to coat cavity surfaces.

All of these types of limonite blebs appear to be unrelated to the Cu-bearing zones in the Ybn₄ member of the Bonner Formation. They occur throughout the entire strike length of these strata, and do not appear to be related spatially in abundance, occurrence, color, or crystallinity to the mineralization.

Limonite which occurs on the prospect itself has completely different characteristics than those described above. Four different types occur (Plate 14):

TYPE D - seal brown, indigenous and minor fringing, highly abundant (covers 10% to 50% of fresh rock surface), highly vitreous euhedral limonite up to 1.0 mm wide which dusts the inside surfaces of up to 1.5 cm

wide cavities in only the coarsest grained crossbedded quartzites (Plate 15). This limonite occurs at the base of the Ybn₄ member (Appendix 1; east side) at the north end of the prospect. Malachite occurs with this limonite as coatings on highly micaceous quartzites.

TYPE E - red, indigenous, abundant (covers 10% to 30% of the rock surface), vitreous euhedral limonite which thickly dusts to coats but does not fill 1.0 cm wide cavities in coarse to fine grained quartzites (Plate 16). It is gradational with type D limonite and usually indistinguishable from type B limonite (described above), except that they usually do not fill the host cavities. This limonite occurs in the Ybn₄ member (Appendix 1) from the top to the base of the unit and is located at the north end of the mineralized zone in a NW to SE trending band. Malachite occurs with this gossan as coatings in highly micaceous quartzites.

TYPE F - orange, indigenous and fringing, abundant (5% to 25%), subhedral to anhedral, non-vitreous, slightly porous limonite often intergrown with malachite which thickly coats and almost fills 0.5 to 1.0 cm wide cavities in medium to fine grained quartzites. Often coats cracks which cut bleb cavities and occurs in rocks containing black oxide coatings on fracture surfaces. This limonite occurs from the base to the top of the Ybn₄ member (Appendix 1) over the central portion of the mineralized zone in a NW to SE trending band.

Plate 15

Cut surface of coarse grained quartzite
from Type D Gossan (Ybn₄ Member, Bonner Formation).
Note hematite dusting vacant cavities
up to 5 mm wide (scale in mm).



- 42 a -

Plate 16

Cut surface of coarse grained quartzite
from Type E gossan (Ybn₄ Member, Bonner Formation).
Note limonite coating up to 5 mm wide cavities
(scale in mm).



- 43 a -

TYPE G - yellow to orange, indigenous, moderately abundant (1% to 15%) limonite completely filling 5.0 mm wide bleb cavities, as intergrowths of dense, yellow, adamantine, subhedral to anhedral crystals (Plate 17) with various amounts of orange, porous, powdery anhedral Fe-oxide. Occurs with black vitreous aggregates which coat grain boundaries in accumulations up to 1.0 cm in size. Occurs from the top to the base of the Ybn₄ member (Appendix 1) in medium to fine grained quartzites but is more prevalent near the top of the unit.

The four limonite types described above are zoned across the prospect. From NE to SW, the sequence of zones consists of gossan type D, E, F, and G. Blanchard's (1968) descriptions of gossans developed over different types of sulfide minerals correspond well with the different types of gossans developed over the prospect. Limonite types D and E, which together constitute a continuum, are similar in occurrence to those gossans which develop over bornite and/or chalcocite (mainly red to seal brown, indigenous limonite; Blanchard, 1968), which occur in country rocks with a large buffering capacity provided by high carbonate, mica, or feldspar contents, and which contain only small amounts of sulfuric acid- or ferric sulfate- producing sulfides (pyrite). The Ybn₄ member (Appendix 1) has a low pyrite content and thus could conceivably develop gossans of this type. Its limonite is restricted to indigenous forms because iron is not soluble in the reacting meteoric waters and thus remains in place. Abundant malachite would also be produced in this

Plate 17

Cut surface of fine grained quartzite
from Type G gossan (Ybn₄ Member, Bonner Formation).
Note yellow cerrusite and black Fe-oxides
disseminated throughout (scale in mm).



- 45a -

gossan, occurring in rocks with high buffering capacities caused by high carbonate, feldspar, or muscovite contents.

Limonite types D and E occur over zones which contain anomalous soil Ag concentrations (Chapter 4). Since Ag can be readily substituted in the bornite and chalcocite lattices up to significant levels (>1%), and malachite occurs in micaceous quartzites, the mineralogic zone underlying these limonites is interpreted to have consisted of chalcocite (under type D limonite) and bornite (under type E limonite) minerals.

Limonite type F is similar in occurrence to those developed over chalcopyrite mineralization (orange, porous, indigenous and fringing limonite myrmekitically intergrown with malachite; Blanchard, 1968) in the aforementioned country rock. It occurs over a zone with anomalous soil Cu concentrations (Chapter 4) where float hand samples contain recognizable chalcopyrite. For this reason, type F limonite appears to occur over a chalcopyrite mineralogic zone in the Daisy Creek prospect.

Finally, limonite type G is similar in occurrence to those gossans developed over galena mineralization (yellow to orange, indigenous limonite; Blanchard, 1968). This limonite occurs over a zone of anomalous soil Pb concentrations (Chapter 4) and corresponds spatially with a portion of the prospect containing numerous prospect pits, presumably dug by prospectors who were successfully panning the soils for galena and other insoluble galena oxidation products, such as anglesite, cerussite, or pyromorphite. Heavy mineral concentrates from soils over this gossan contain abundant amounts of both cerussite, the yellow,

adamantine mineral in these float rock-chips, and pyromorphite. For this reason, the G type limonite is interpreted to occur as a gossan over a galena zone in the Daisy Creek prospect.

The existence of these gossan types suggests that these rocks contained a significant buffering capacity. This probably consisted of carbonate, probably ferroan calcite or ferroan dolomite which has now released its carbonate to malachite, feldspars, now largely replaced by clay minerals, and detrital muscovite. Although very little carbonate is observed at the surface, except for malachite, it may have comprised a significant percentage of the host rock, based on the current porosity of the weathered mineralized quartzites, and the abundance of gossanous cavities after carbonate in several of the strata in the Daisy Creek area (gossan types A and B). The carbonate presumably has been dissolved and remobilized under acidic conditions created by the oxidation of the sulfides on the prospect to sulfate and sulfuric acid.

The zonation of gossan type representing chalcocite, bornite, chalcopyrite, and galena mineralogic zones is similar to the mineralogic zonations observed at many other sedimentary copper deposits (e.g., Kupferschiefer (Rentzsch, 1974), Zambian Copperbelt (Annels, 1974; Bartholome, 1974; Van Eden, 1974; Bowen and Gunatilaka, 1976), White Pine (Ensign et al., 1968; Brown, 1971, 1974), Troy (Hamilton and Balla, 1983; Hayes, 1984; Hughes, 1984), Northwest Territories (Watson, et al., 1975; Chartrand and Brown, 1978), Africa (Caia, 1976), Udokan Basin (Bakun, 1967), Adelaide (Rowlands, 1974), and in general: Renfro, 1974; Smith,

1976; Holcombe, 1977; Brown, 1978; Bjorlykke and Sangster, 1981; Gustafson, 1981). Careful mapping of gossan characteristics has thus given indications of the unweathered mineralogy of the Daisy Creek Prospect and could be a helpful exploration tool on stratabound-type copper deposits in environments similar to those at Daisy Creek.

Soil Map

In addition to the float rock-chip maps, a soil map was also obtained from soil color and grain size data (Plate 18). Distinct soil color zones can be recognized which correspond with different underlying lithologies.

Rusty-brown (10 YR 6/3 or 'Pale Brown' on the Munsell Soil Color Chart) soils occur over the Libby Formation rocks and tend to be fine grained. Pinkish-brown (5 YR 6/4 or 'Light Reddish Brown') soils occur over the Bonner Formation red-bed, fine grained quartzites and tend to be fine grained but commonly feel granular when rolled between the fingers. These soils occur over both members Ybn₅ and Ybn₃ (Plate 18; Appendix 1) on the Daisy Creek soil grid.

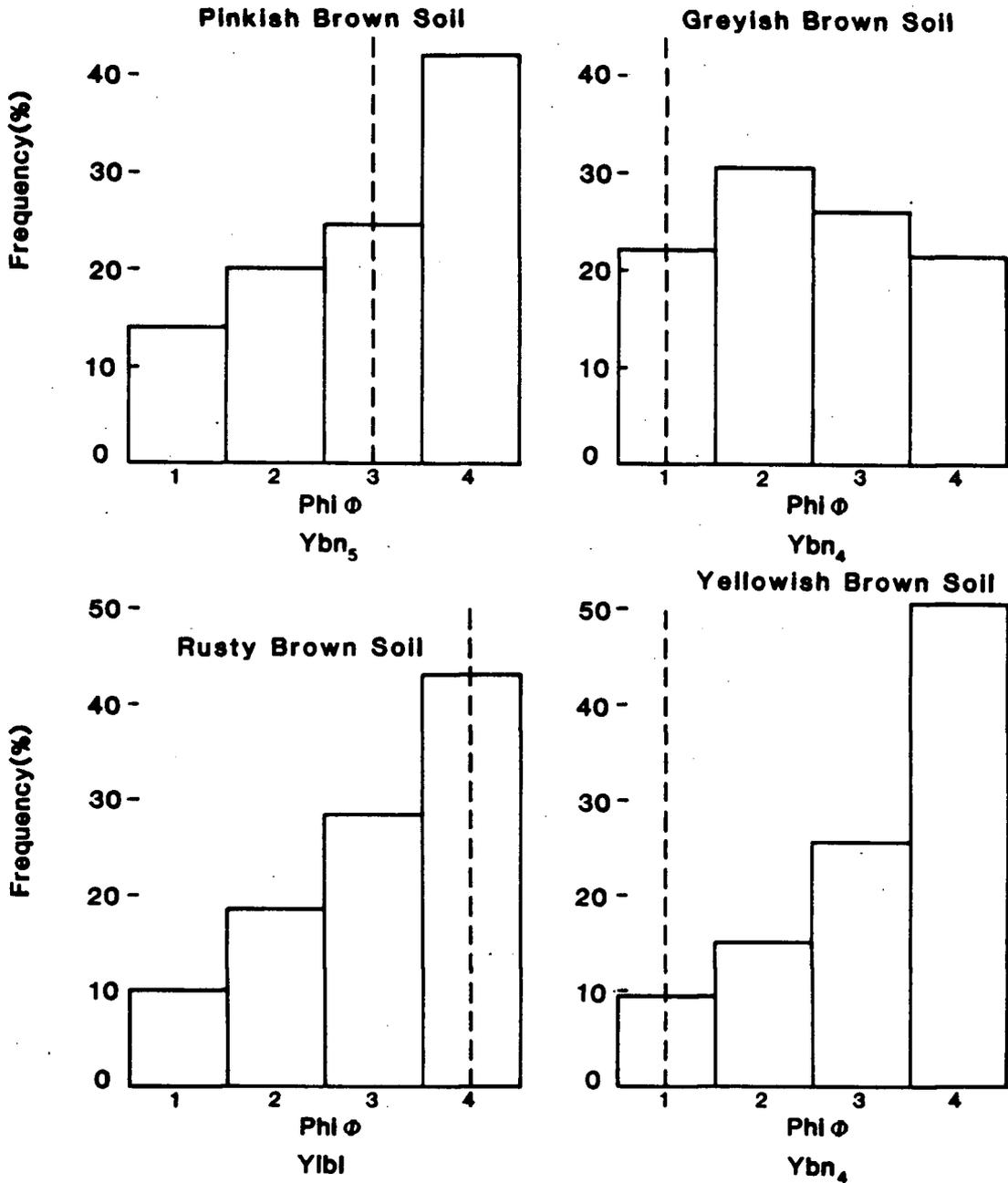
Over the Ybn₄ member of the Bonner Formation (Plate 18; Appendix 1), two soil types occur. The first, with a greyish-brown (7.5 YR 6/4 or 'Light Brown') soil, is very coarse grained and granular. It occurs over areas in this member which do not contain significant amounts of economic minerals. The second soil tends to be yellowish-brown (10 YR 7/5 or 'Very

Pale Yellowish Brown') and is very fine to powdery in texture. This soil type tends to occur over the Ybn₄ member (Plate 18; Appendix 1) in areas with significant mineralization and correlates well with zones of anomalous Cu and, to a lesser extent, Pb soil concentrations (Chapter 4).

Distinct differences in textural data for typical soils from each group are apparent in Figure 9. Yellowish-brown soils exhibit much finer grain sizes than the greyish-brown soils despite both occurring over the same lithology. This may be due to more acidic conditions in soils overlying mineralization, caused by the oxidation of sulfides and the resulting sulfuric acid produced. Destruction of much of the cement, feldspars, and non-quartz detrital grains would thus take place, reducing the overall grain size. The yellowish-brown color likely is due to the formation of immature ferric hydroxides from iron bearing phases in the Ybn₄ member (Appendix 1). More alkaline conditions associated with barren zones of the Ybn₄ member would not have produced substantial decreases in grain size or significant and rapid destruction of iron bearing phases, and thus remain neutral (i.e., grey) in color and coarse grained, like the unmineralized portions of the Ybn₄ member.

Results indicate that the limits of mineralized zones can be determined without geochemical analysis by using soil maps based on careful examination of soil color and grain size. In addition, in areas without outcrop, soils can be effective for indirect mapping of the underlying lithologies.

Soil Textures From The Daisy Creek Prospect



* dashed line represents average phi(Φ) value for underlying sedimentary rock

Figure 9

Mineralogy

Over 100 hand specimens of the soil float rocks occurring over the Daisy Creek prospect were collected for laboratory study. All were examined under a binocular microscope. In addition, 27 thin sections, 45 polished sections, and two polished thin sections were examined. Eleven polished sections were evaluated using a scanning electron microscope and energy dispersive x-ray emission spectrometer. Numerous x-ray diffraction scans were also made to identify individual minerals.

Rock Forming Minerals

The mineralized rocks at the Daisy Creek prospect consist of feldspathic quartzites and orthoquartzites. Quartz is the predominant mineral and comprises up to 85% of the rock. Feldspars are largely destroyed, and this probably occurred during both a diagenetic phyllo-morphic event, when they were altered to muscovite, and during a supergene oxidation event, when they were weathered to kaolinite. For this reason, and because the heavy minerals present in the country rock suggest that the source terrane for this detritus was granitic, the feldspars are probably largely potassium varieties.

Heavy minerals in these quartzites are (in order of decreasing abundance) magnetite, apatite, monazite, zircon, leucoxene, and specularite. The specularite may be recrystallized accumulations of Fe-oxides which formed during diagenesis from

the destruction of minor mafic Fe-bearing minerals. All other minerals are of detrital origin. Leucoxene appears to be an altered remanent of detrital ilmenite because it occurs mostly with indigenous and fringing limonite (Dimanche and Bartholome, 1976; Marod and Aldahan, 1982; Bartholome, 1974).

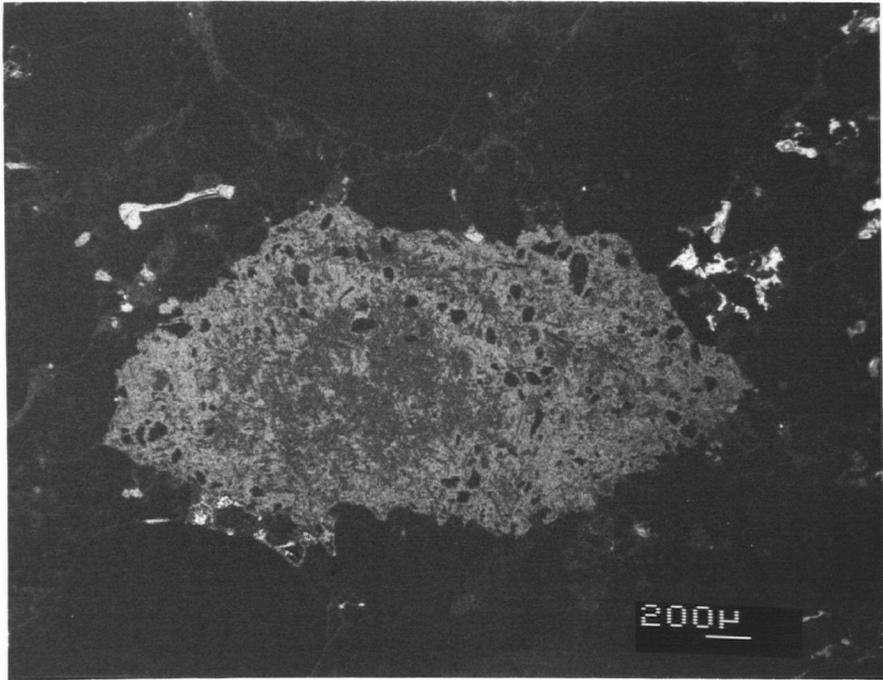
Apatite occurs as discrete detritus and as a major constituent in green mud chips. The Ybn₄ member contains very large amounts of apatite as both mud chips (Plate 19) and discrete grains (up to 1%), and heavy mineral concentrates of soils located over this member reveal that 15% of the apatite present in the soil occur as euhedral prisms of apatite which may be growing in situ, probably as a fluoro-carbonate apatite. This suggests that the Ybn₄ member has an extremely high phosphorus content.

The monazites in these quartzites occur with two distinctly different chemistries, one consisting of dominantly a cerium/neodymium phosphate, and the other a cerium/lanthanum phosphate (Figure 10). In addition, some of the Ce/La type monazites are rimmed by zircon (Plate 20), which have epitaxial overgrowths of bastnaesite, a yttrium/rare earth fluoro-carbonate, in this case containing gadolinium, dysprosium, and minor samarium and neodymium.

Texturally, the feldspathic quartzites and orthoquartzites of the Ybn₄ member exhibit several predictable trends. The finer the grain size, the more muscovite and less feldspar is present. This suggests that the feldspar was abraded in the fluvial environment at a very rapid rate and altered to clay minerals

Plate 19

SEM photomicrograph (backscattered electrons)
of phosphatic mud chip from the Ybn₄ member,
Bonner Formation.



SEM/EDS Spectral Image of Monazite

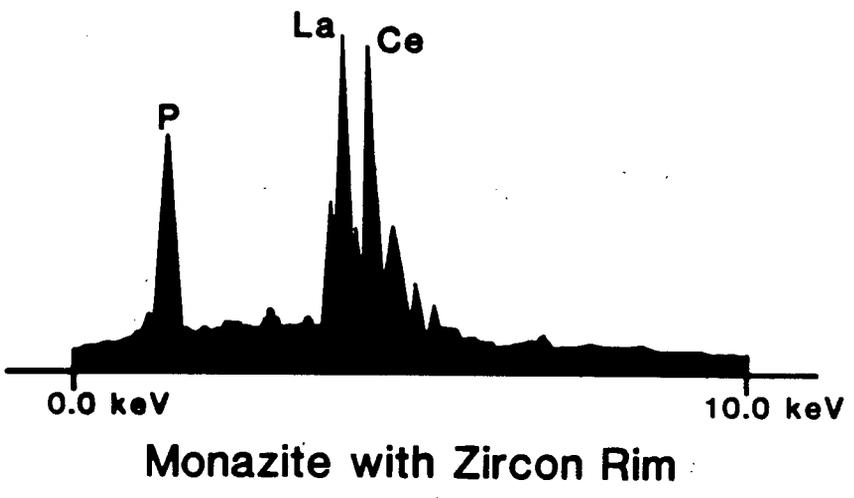
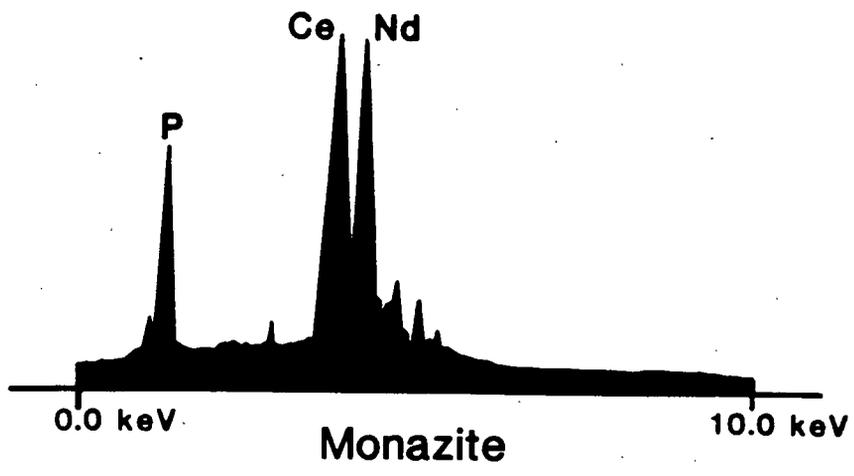
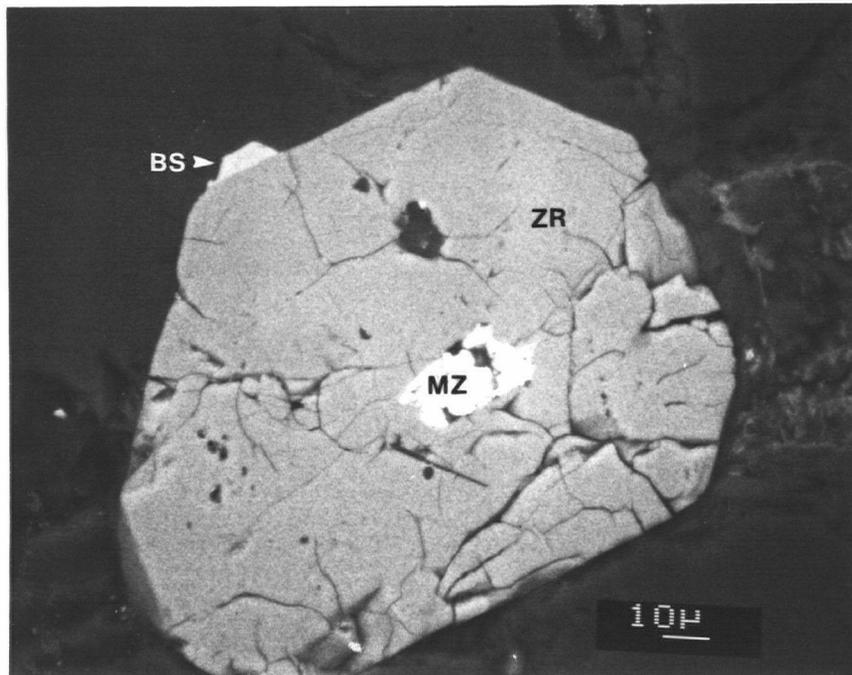


Figure 10

Plate 20

SEM photomicrograph (backscattered electrons)
of detrital heavy mineral. Core consists
of Ce/La monazite (MZ); Rim consists of zircon (ZR).
Note epitaxial overgrowth of Gd/Dy/Sm/Nd yttrian
bastnaesite (BS), a rare earth yttrio-fluoro-carbonate.



when reduced to very small grain sizes. These small grains, having large surface-to-volume ratios, would be chemically reactive and alter readily to muscovite. The larger grains of feldspars did not react chemically until they were abraded to smaller grain sizes because of their small surface-to-volume ratios.

Thus, muscovites in finer grained quartzites are generally replacing the destroyed feldspars (Plate 21), whereas in the coarser grained quartzites, those that occur are probably detrital. This detrital muscovite is very prevalent in the mineralized zone and is often coated with extensive amounts of secondary malachite.

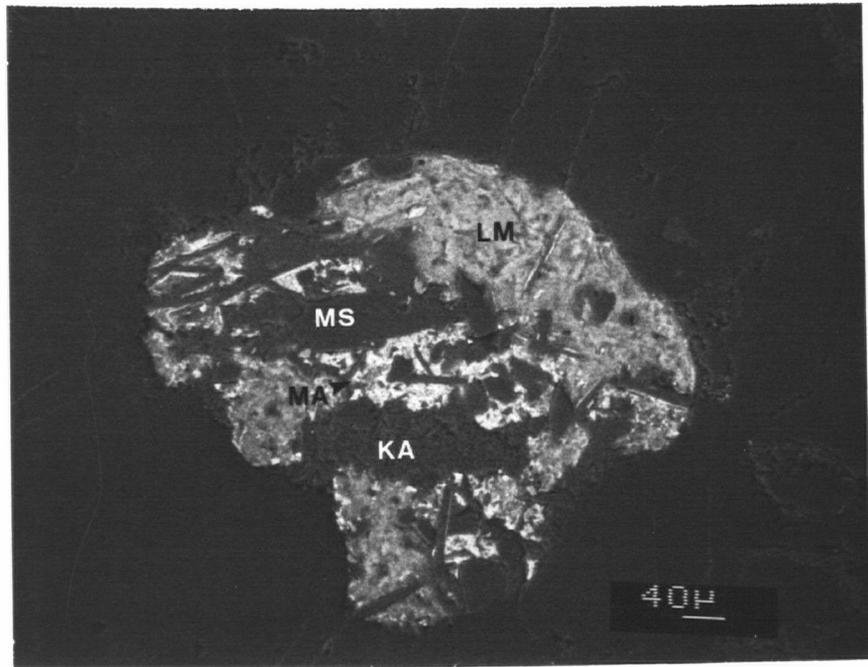
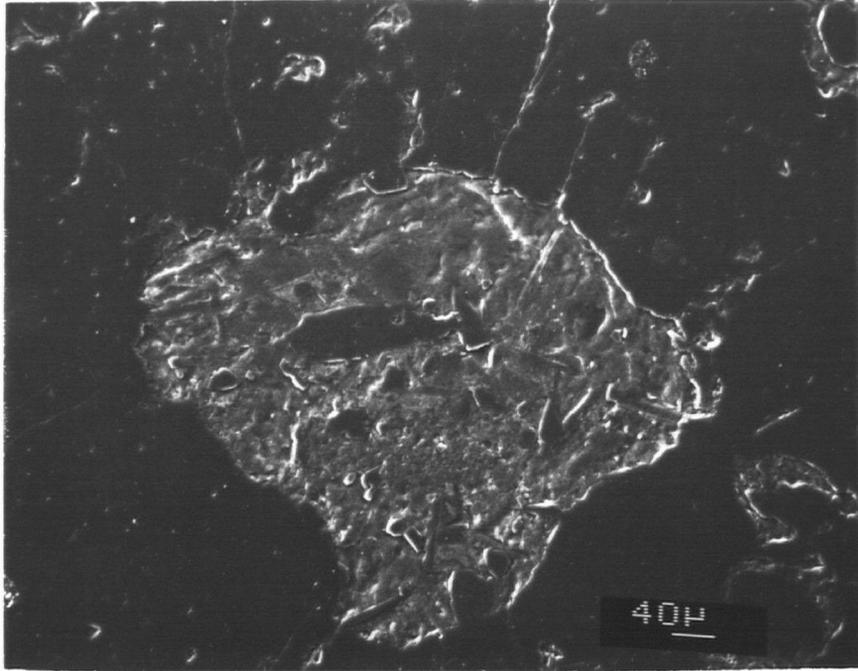
The matrix of the quartzites generally is composed of quartz, hematite, and carbonate, in order of decreasing abundance. Carbonate was not observed, but limonitic cavities, which probably contained a ferroan carbonate that has since weathered out, are abundant in many units and occur as disseminations up to 5.0 mm in size.

In general, the finer grained quartzites contain more hematite cement than the coarser grained quartzites. This cement is extremely prevalent but formed before the onset of quartz cementation and the development of quartz overgrowths.

Quartz overgrowths are more common in the cleaner, coarser quartzites. In the coarsest crossbedded units, the overgrowths have developed to such a point as to exhibit a true graphic texture. Furthermore, the quartz overgrowths surround all of the sulfide minerals at the prospect as well as the cupriferous

Plate 21

SEM photomicrographs (secondary electrons above,
backscattered electrons below) of remains of
detrital feldspar, now consisting of
malachite (MA), limonite (LM),
muscovite (MS), and kaolinite (KA).



goethite seen rimming and replacing many chalcopyrite grains. This indicates that economic minerals precipitated before the onset of quartz overgrowth development (i.e., before late diagenesis).

No mafic minerals were observed in either the coarse or fine grained quartzites of the Bonner Formation, probably because they were rare in the well sorted quartzites in the first place, and would have been unstable in the diagenetic environment, altering to hematite and limonite.

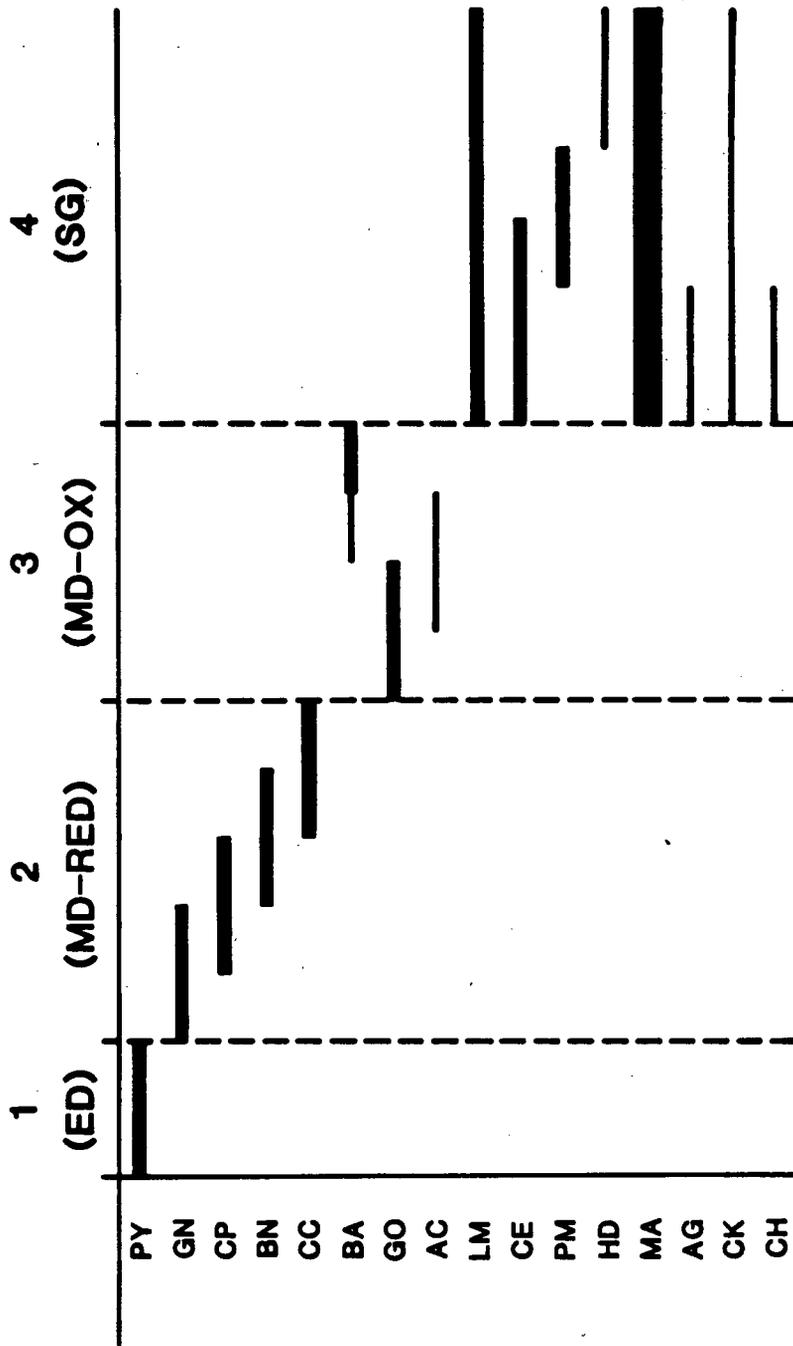
Economic Minerals

The paragenetic relationships and occurrence of the copper bearing and associated minerals at the Daisy Creek prospect are complex. Four different episodes of mineralization are indicated, as summarized in the line diagram (Figure 11) and the Van der Veer diagram (Figure 12).

Early Diagenesis

The first episode is characterized by what appears to have been the authigenic development of euhedral pyrite which grew at the expense of both feldspar and quartz grains. Detrital quartz grains adjacent to pyrite cubes commonly are truncated by the flat sides of the pyrite (Plate 22), whereas feldspars appear to have been replaced preferentially, but generally are not replaced completely.

LINE DIAGRAM OF PARAGENETIC MINERAL RELATIONSHIPS



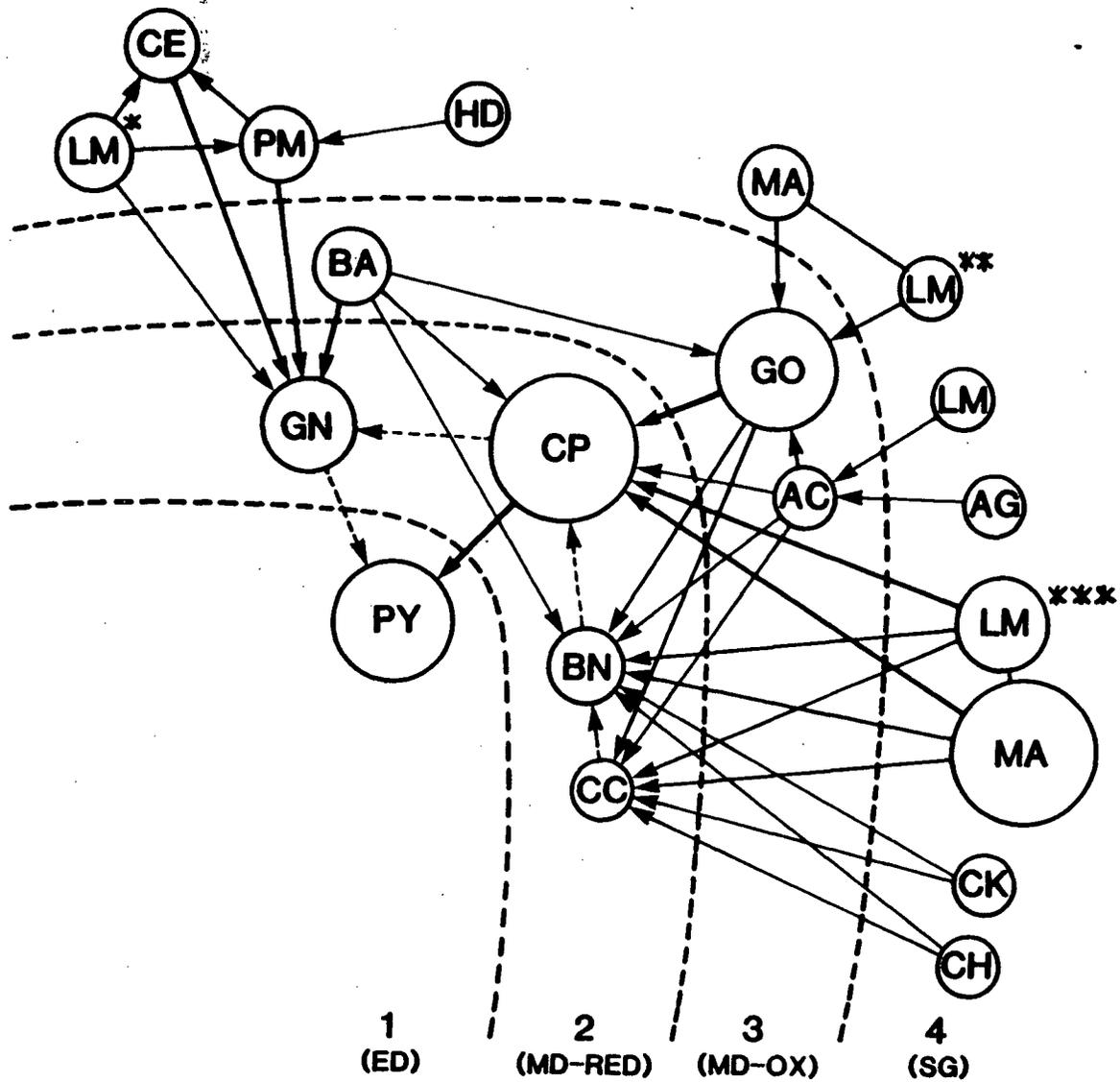
Abbreviations :

ED early diagenesis
 MD-RED middle diagenesis (reducing)
 MD-OX middle diagenesis (oxidizing)
 SG supergene oxidation

PY pyrite
 GN galena
 CP chalcopyrite
 BN bornite
 CC chalcocite
 BA barite
 GO Cu-goethite
 AC acanthite

LM limonite
 CE cerussite
 PM pyromorphite
 HD hinsdalite
 MA malachite
 AG native Ag
 CK chalcantite
 CH chrysocolla

Figure 11



**Van der Veer Diagram of Mineral Paragenesis
at the Daisy Creek Prospect**

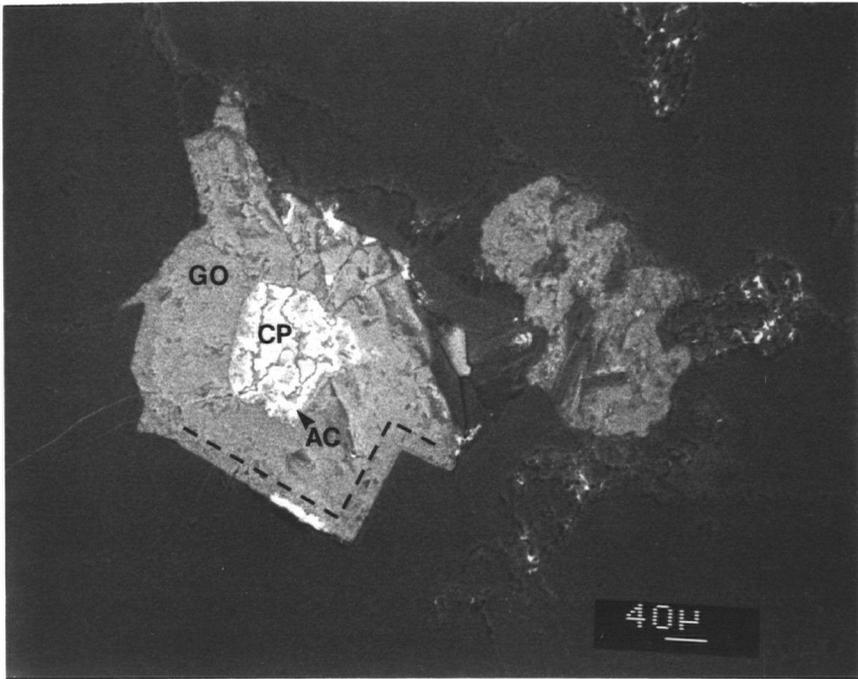
(see Figure 11 for abbreviations)

- * limonite from Type G gossan
- ** limonite from Type F gossan
- *** limonite from Type D and E gossans

Figure 12

Plate 22

SEM photomicrograph (backscattered electrons) of chalcopyrite (CP) completely replacing cubic form of pyrite. Chalcopyrite is, in turn being replaced by cupriferous goethite (GO).



Precipitation of pyrite in the Ybn₄ member probably took place while the sediment was unconsolidated (i.e., early diagenesis) as a result of bacterial activity which decomposed organic material buried in the sediment. Authigenic pyrite does not occur in beds which contain large accumulations of hematite cement. This may suggest that these sediments probably were either buried at too rapid a rate, or remained too wet to allow the introduction of oxygen into the sediment and the oxidation of iron and sulfur in the sediment to occur.

The iron in the sediment contained in both pyrite and hematite, was probably supplied by the destruction of mafic minerals during early diagenesis. Under reducing conditions this iron was fixed as pyrite, and under oxidizing conditions, it formed as ferric hydroxide, which later dehydrated and recrystallized to goethite and then hematite.

The high phosphorus content of the Ybn₄ member may be related indirectly to the development of pyrite in the sediment. The phosphate may have acted as a nutrient for algae, allowing a major algal bloom to occur in that unit. The algae, which grew at the surface, would be buried with the sediment and could have provided the organic material necessary to allow sulfur reducing bacteria (Desulfovibrio or Desulfotomaculum) to reduce sulfate to sulfide. Coincidentally, the Ybn₄ member contains evidence of early diagenetic pyrite development only in the Daisy Creek prospect area, and this also is the place where the largest amount of phosphatic mud chips occur.

Reducing Middle Diagenesis

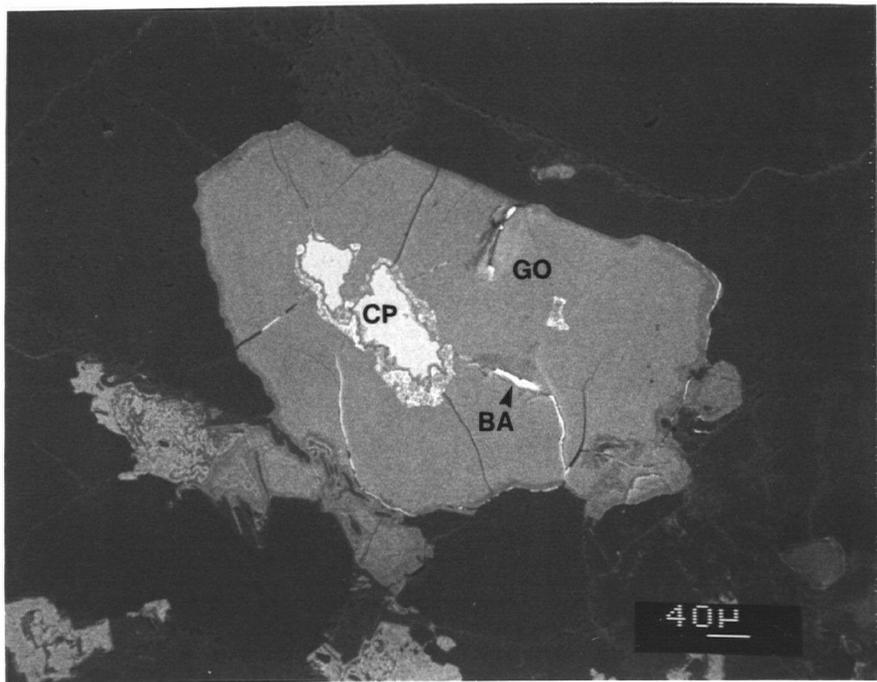
Additional mineral reactions appear to have occurred after the development of early diagenetic pyrite, but before the development of quartz overgrowths, which mark the start of late diagenesis. Most obvious is the replacement of pyrite by chalcopyrite (Plate 22). Replacement of pyrite by galena also probably occurred, but evidence for this is less conclusive, because both minerals have cubic forms. Thus the replacement of galena by chalcopyrite may also have occurred, but no evidence for this has been observed.

The replacement of pyrite by Cu-bearing phases probably occurred under acidic conditions, where copper would be mobile as a chloride complex. The precipitation of pyrite during early diagenesis probably occurred under alkaline conditions, because pyrite replaces the quartz, which would be unstable under these conditions. Thus, the formation of Cu-bearing sulfides represents a distinctly different mineral precipitation event, which differed in its chemical environment from that during early diagenesis.

Indirect evidence for the successive replacement of chalcopyrite by bornite and then chalcocite during the Cu-sulfide precipitation event is also present. Plate 23 depicts a typical chalcopyrite grain in one of the highly phosphatic, mud chip bearing, coarse grained, crossbedded quartzites of the Ybn₄ member. Chalcopyrite comprises the core of the crystal and appears to replace pyrite as a pseudomorph. The chalcopyrite is

Plate 23

SEM photomicrograph (backscattered electrons)
of cupriferous goethite (GO) replacing
chalcopyrite (CP). Note rim on cupriferous
goethite with higher Cu concentration.



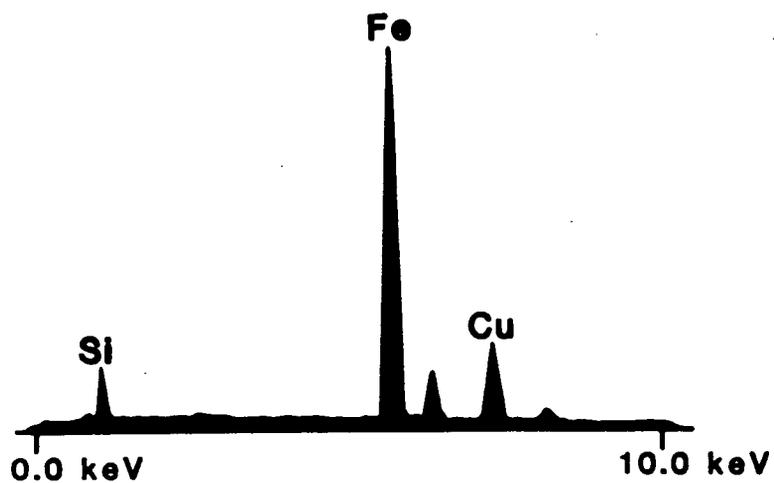
replaced marginally by cupriferous and silica bearing goethite (Figure 13). Very consistent ratios of Cu and Fe occur across the cupriferous goethite, based on a qualitative EDS scan, and may reflect the stoichiometry of the original phase. This suggests that the Cu was present in this phase which was replaced by goethite, and probably consisted of chalcopyrite.

In addition, a rim can be seen which surrounds the cupriferous goethite. This is also a goethite phase, but with a distinctly higher Cu to Fe ratio (i.e., a higher copper concentration; Figure 13). This high Cu goethite may represent an entirely different original phase which has been oxidized, but which originally contained both Cu and Fe. Since bornite, as well as chalcocite, has been observed in heavy mineral concentrates from soils over the prospect, it is the obvious choice and it is reasonable to assume that bornite replaced chalcopyrite.

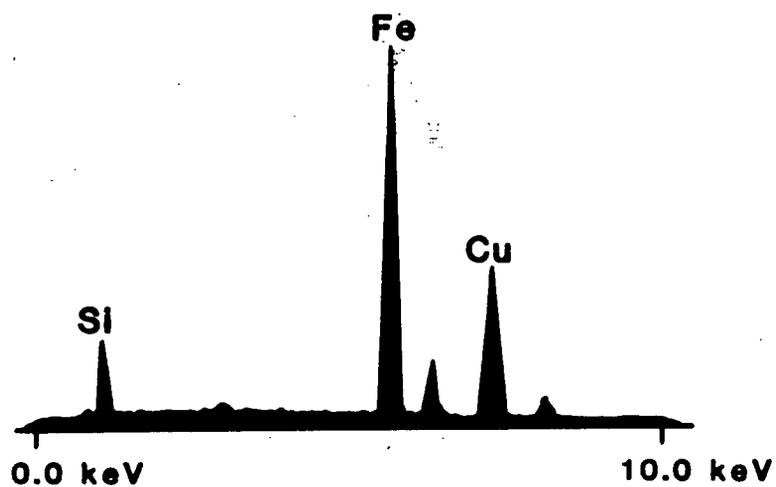
Paragenetic relationships observed in detrital grains from heavy mineral concentrates suggest that chalcocite subsequently may have replaced bornite. In any case, neither bornite nor chalcocite comprise major amounts of the sulfide in the deposit.

This scenario is consistent with the paragenetic relationships observed at many other stratabound copper deposits (Troy (Hamilton and Balla, 1983; Hayes, 1984), White Pine (Ensign et al., 1968; Brown, 1971, 1974), Kupferschiefer (Rentzsch, 1974), Zambian Copper Belt (Annels, 1974; Bartholome, 1974; Van Eden, 1974), etc.).

SEM/EDS Spectral Image of Cupriferous Goethite



Cupriferous Goethite Core



Cupriferous Goethite Rim

Figure 13

Oxidizing Middle Diagenesis

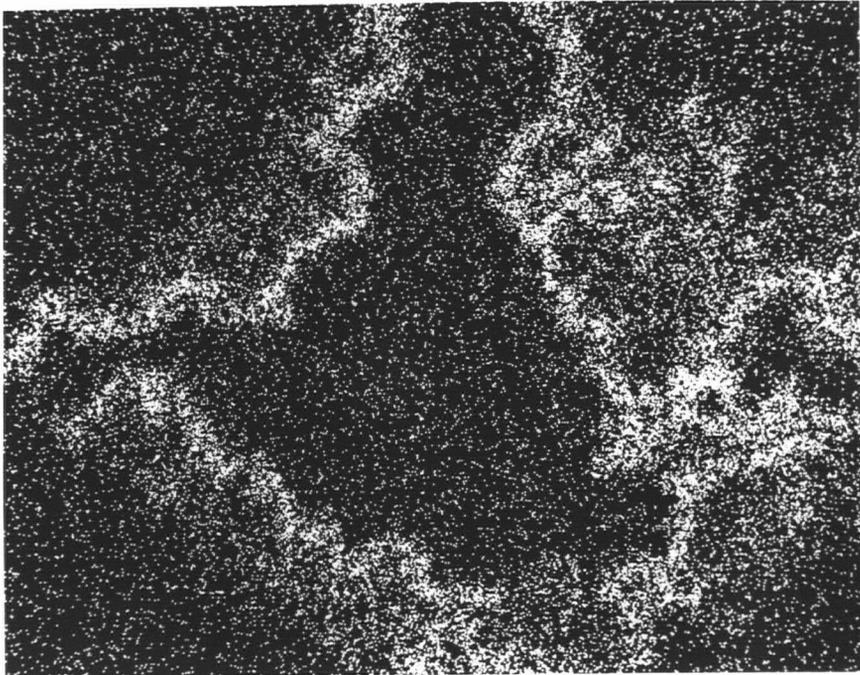
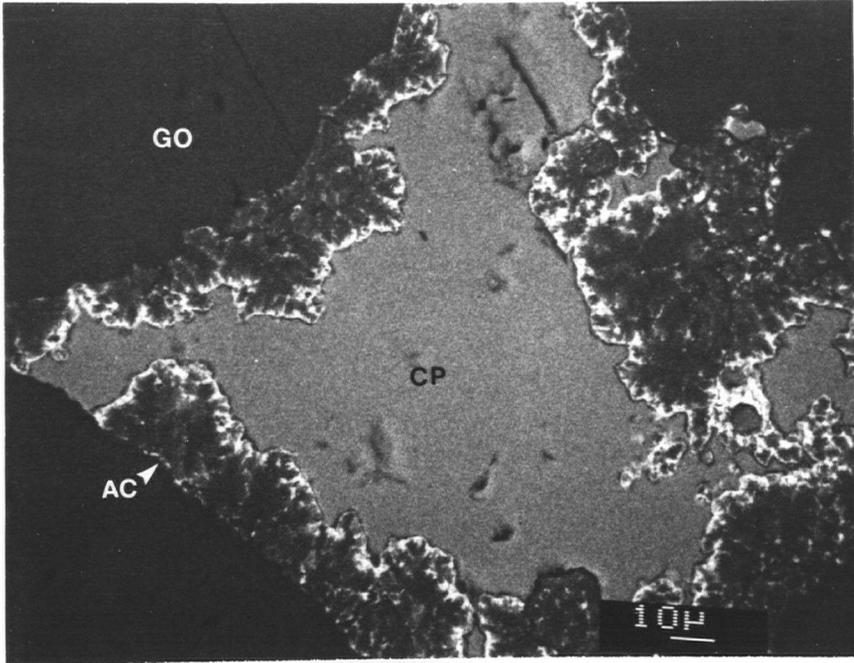
After the successive replacement of pyrite by galena (?), chalcopyrite, bornite, and possibly chalcocite (?), many of the sulfides at the Daisy Creek prospect underwent oxidation, probably during middle diagenesis (i.e., before silica cementation). All copper sulfides underwent at least partial oxidation. Plates 22 and 23 show a chalcopyrite crystal partly replaced by cupriferous goethite. In addition, a silver bearing sulfide, probably acanthite, occurs just outside the unaltered chalcopyrite remanent within the confines of the cupriferous goethite (Plate 24). This 'acanthite' appears to be an oxidation product of the sulfide(s) present, yet neither the chalcopyrite nor the cupriferous goethite are argentiferous.

Stratabound copper deposits in the Belt Supergroup commonly contain significant concentrations of silver (Clark, 1971; Harrison, 1974; Harrison et al., 1977; Lindsey, 1977; Lange and Eby, 1981; Newton, 1982; Lange and Sherry, 1983). This silver is contained largely in the crystal lattice of bornite and chalcocite in these deposits (Hurlbut and Klein, 1977).

The chalcopyrite in Plates 22 and 23 is partly oxidized, but in this case, two oxidation processes seem to have occurred simultaneously. The first involves the oxidation of ferrous (+2) to ferric (+3) iron. The second involves the oxidation of sulfide (-2) to sulfate (+6). Iron oxidation involves the loss of only one electron whereas sulfide oxidation involves the loss of eight electrons. Ignoring kinetic controls, which are generally very

Plate 24

SEM photomicrograph (backscattered electrons; above)
of cupriferous goethite (GO) replacing
chalcopyrite (CP), with acanthite (AC) zone
surrounding the chalcopyrite.
Below is backscattered element map for silver.



rapid in redox processes, it seems reasonable to assume that the oxidation of iron would take place faster than the oxidation of sulfide. This is because the ionization potential for the oxidation of ferrous iron is less than that for sulfide. Thus, precipitation of goethite would occur in the presence of coexisting sulfide. This sulfide would be available to react with any metal released in the oxidation process, in this case silver from the bornite lattice, and would immediately precipitate, forming acanthite.

Thus, as the oxidation process continued silver was released and, while some diffused outward and was lost, some diffused inward, encountered sulfide, and precipitated in voids and replaced goethite. The silver-rich zone adjacent to the chalcopyrite is thus a result of silver migrating with the oxidizing front in the chalcopyrite/goethite crystal inward from the bornite rim. As further oxidation occurred, sulfide eventually converted to sulfate, releasing silver to be lost outward and captured inward by newly released sulfide. The lack of any copper sulfide phase could be due to the favored stability of acanthite over any possible Cu phase (chalcocite or covellite), or because Cu fits readily into the goethite lattice.

Figure 14 demonstrates the association of the thickness of the high Cu goethite rim, and presumably related to the amount of bornite and, thus, the amount of silver present before oxidation, with the thickness of zone where acanthite is present outside of the chalcopyrite, and presumably related to the amount of silver now present after oxidation.

Acanthite Rim/Cupriferous Goethite Rim Comparison

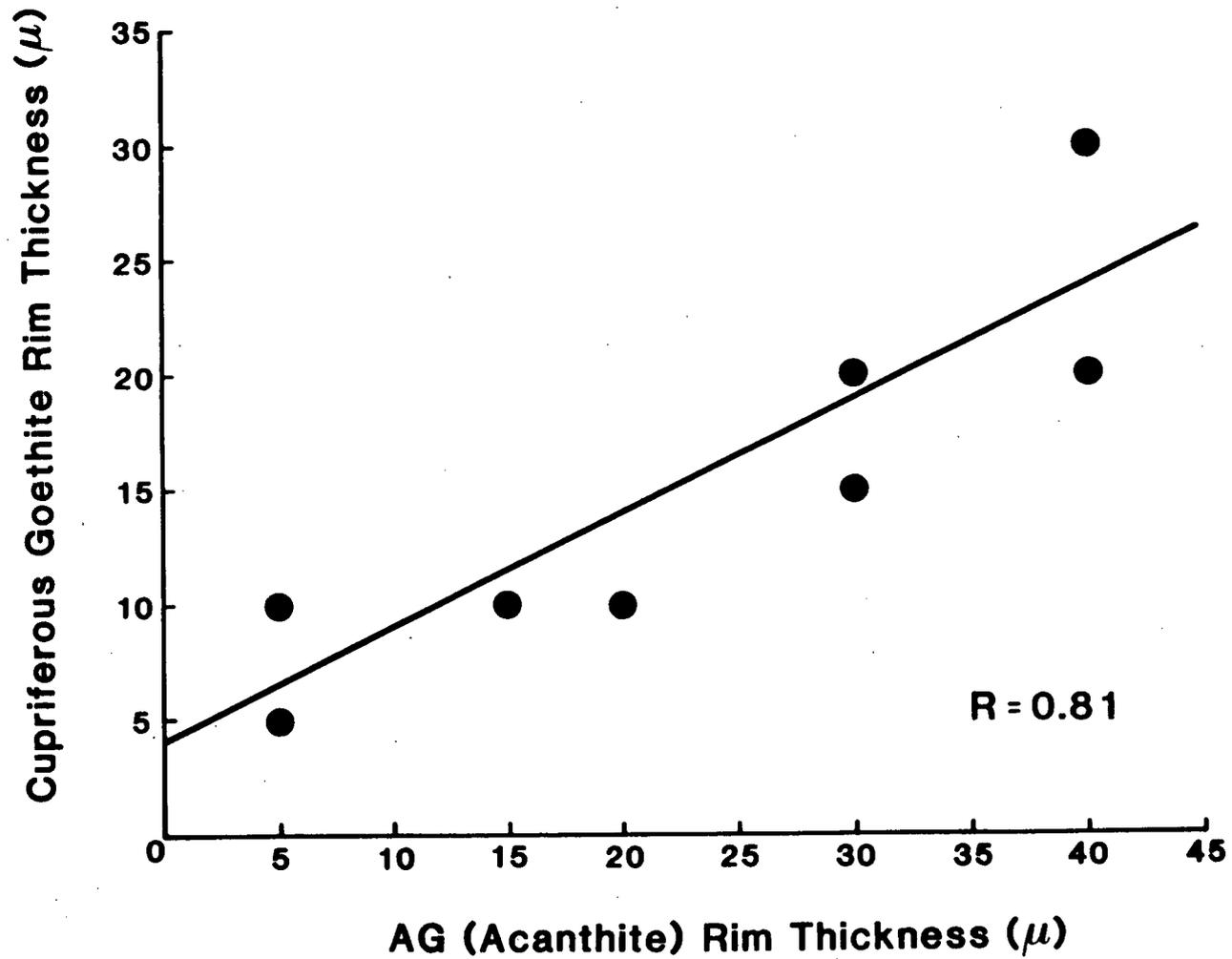


Figure 14

In one case, two discrete cores of chalcopyrite are connected by rims of cupriferous goethite (Plate 25). One chalcopyrite grain contains acanthite around it and has a Cu-rich goethite rim, whereas the other has neither. The high correlation (0.81; >99th%ile) of these two thickness variables (Figure 14) suggests that the silver was, in fact derived from the high Cu goethite rim, which probably consisted of bornite originally. If the original sulfide had been chalcocite, iron would have had to have been introduced, resulting in highly variable Cu/Fe EDS ratios, which are not observed.

Radial cracks are observed in the cupriferous goethite, but do not extend into the chalcopyrite (Plate 26). These cracks (Plates 22, 23, 25, and 26) are probably the result of a recrystallization of the goethite to a more crystalline form which resulted in a net loss in volume or due to volume loss in the chalcopyrite to goethite reaction. If the cracks were due to the dehydration of ferric hydroxide to goethite, a well known diagenetic reaction, a larger volume loss would occur, resulting in much larger cracks.

Barite commonly fills these cracks and also occurs sporadically along the rims of cupriferous goethite (Plate 27). These barite 'veinlets' formed after the formation of goethite, and are probably the last oxidation product, in that they derived their sulfate from the oxidation of chalcopyrite and acanthite.

The formation of the barite probably occurred during middle diagenesis because of the low solubility of barium in the supergene environment. Supergene waters would not be able to

Plate 25

SEM photomicrographs (backscattered electrons) of cupriferous goethite (GO) replacing chalcopyrite (CP). Note various thicknesses of high Cu cupriferous goethite rims, and various thicknesses of acanthite (AC) zones.

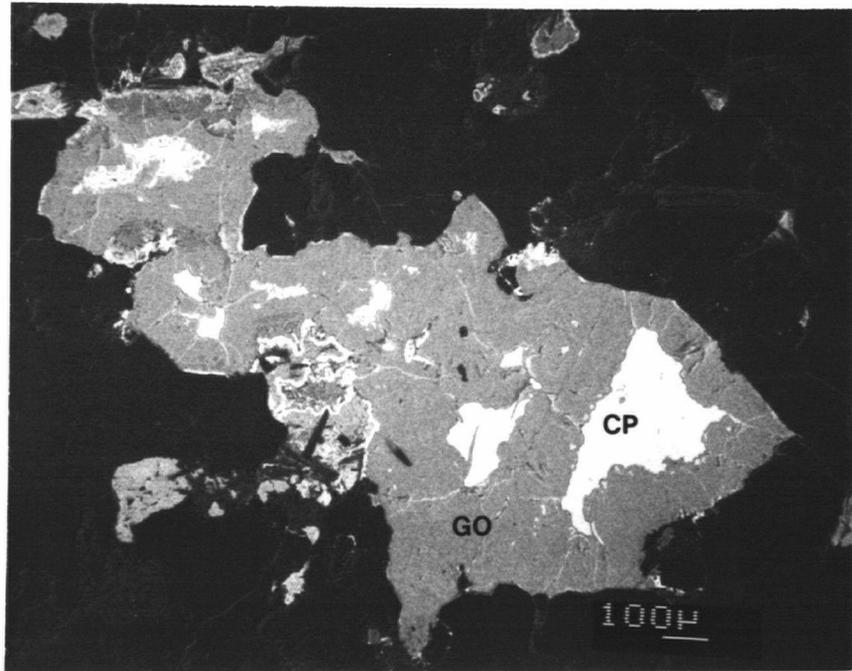
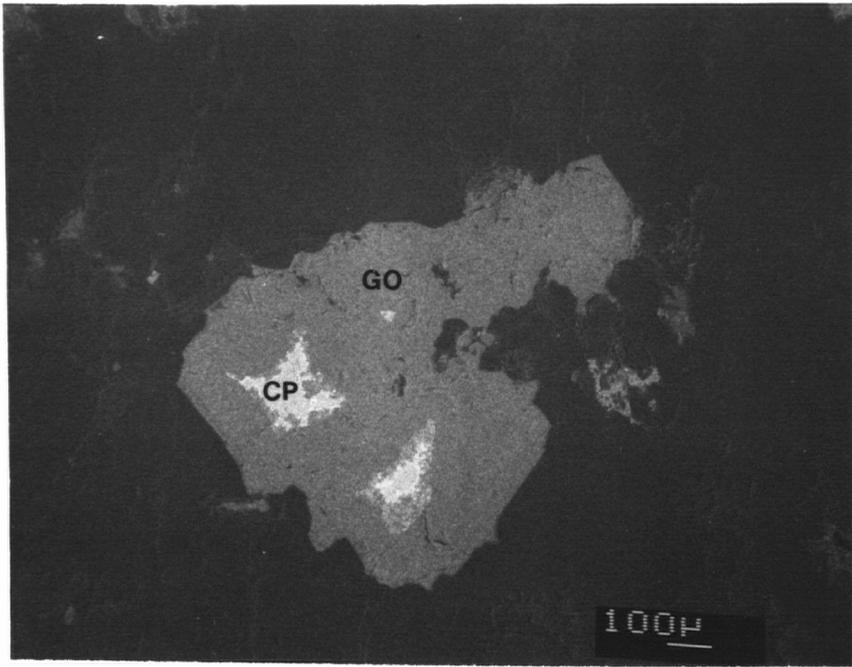


Plate 26

SEM photomicrograph (backscattered electrons) of cupriferous goethite (GO) replacing chalcopyrite (CP). Note cracks in cupriferous goethite do not penetrate chalcopyrite.

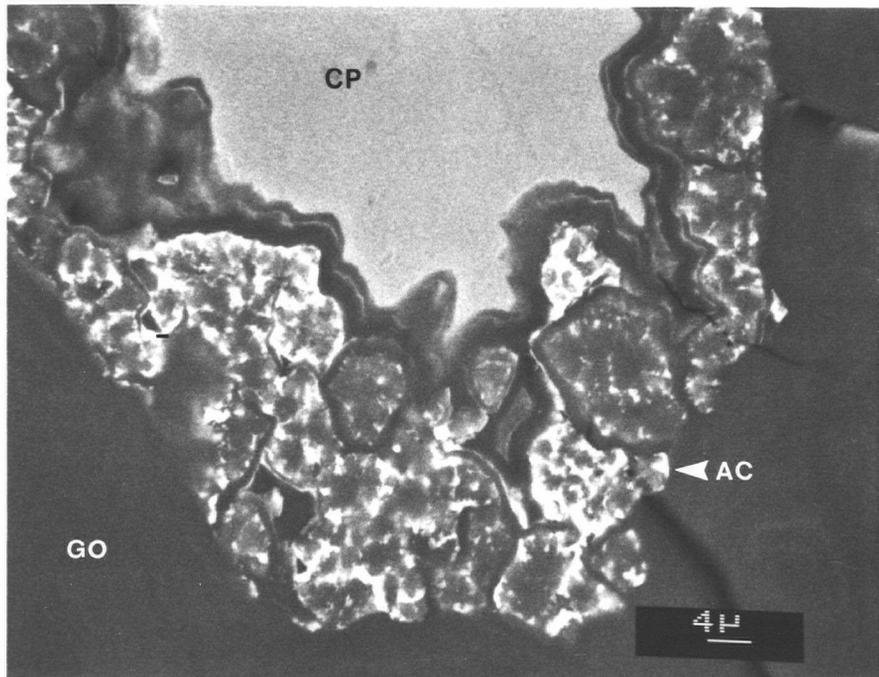
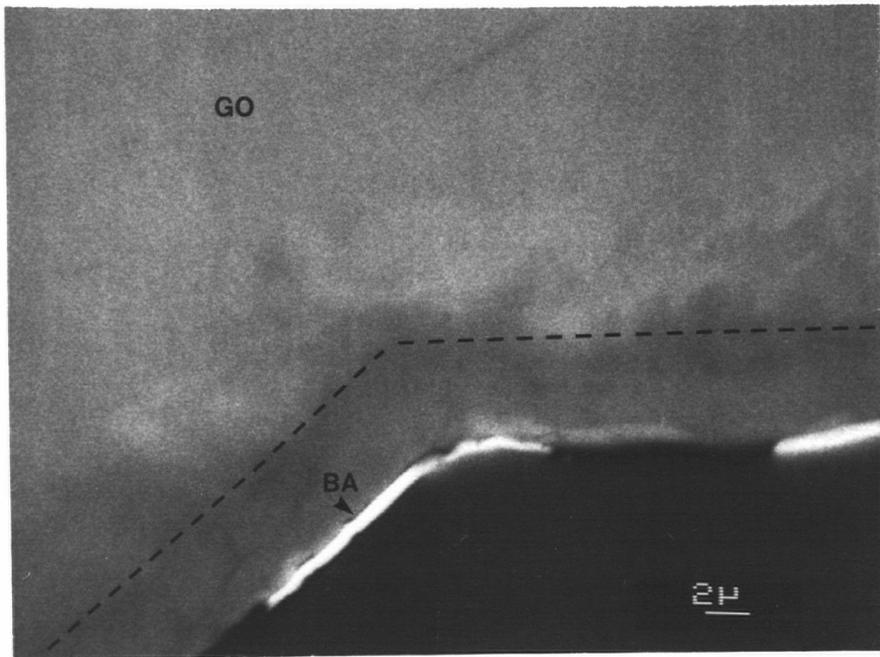
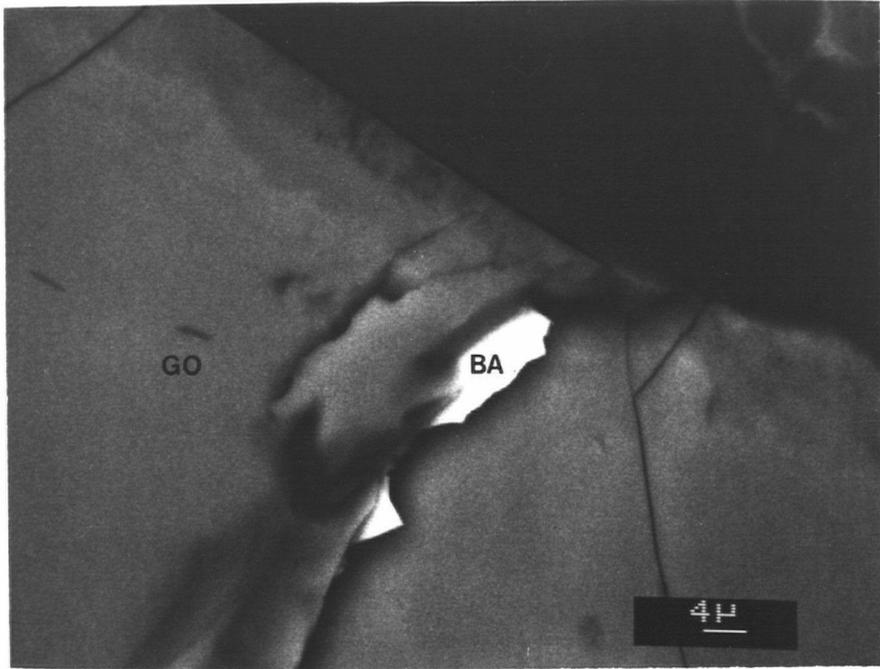


Plate 27

SEM photomicrographs (backscattered electrons)
of barite (BA) occurring in cracks in
cupriferous goethite (GO), and rimming
cupriferous goethite.



carry barium in solution to react with the sulfate diffusing out of the cupriferous goethite crystals; however, chloride brines possibly responsible for the sulfide minerals (Rickand, 1974; Rose, 1976; Rye et al., 1984) and later oxidation could carry it as chloride complexes, especially if it were being released from altering feldspars, where it is a trace to minor constituent.

In addition, major amounts of barite are observed as a replacement of galena. This reaction is probably also a result of oxidation, and, coincidentally, is suggested by an overlap in both Pb and Ba soil anomalies (Chapter 4).

Thus the precipitation of barite could only occur during middle diagenesis and represents the last middle diagenetic oxidation reaction. Quartz overgrowths surround many of these barite crystals.

Supergene Oxidation

Extensive Holocene supergene oxidation has occurred on all minerals in the Daisy Creek prospect. Galena has altered to cerussite and pyromorphite. Carbonate was probably derived from unstable ferroan carbonates which are also weathering out, and phosphate was supplied by the apatites in the Ybn₄ member (Appendix 1). A rare mineral, hinsdalite [PbAl₃(SO₄)(PO₄)(OH)₆] (Plates 27 and 28), occurs coating pyromorphite, and has been described at only three other localities in the world (Birch, 1977; Wilkinson, 1980).

Cupriferous goethite largely is altered to a myrmekitic

Plate 28

SEM photomicrographs (secondary electrons above;
backscattered electrons below) of
hinsdalite (HD). Note above hinsdalite
occurs in vacant cavity in quartzite.

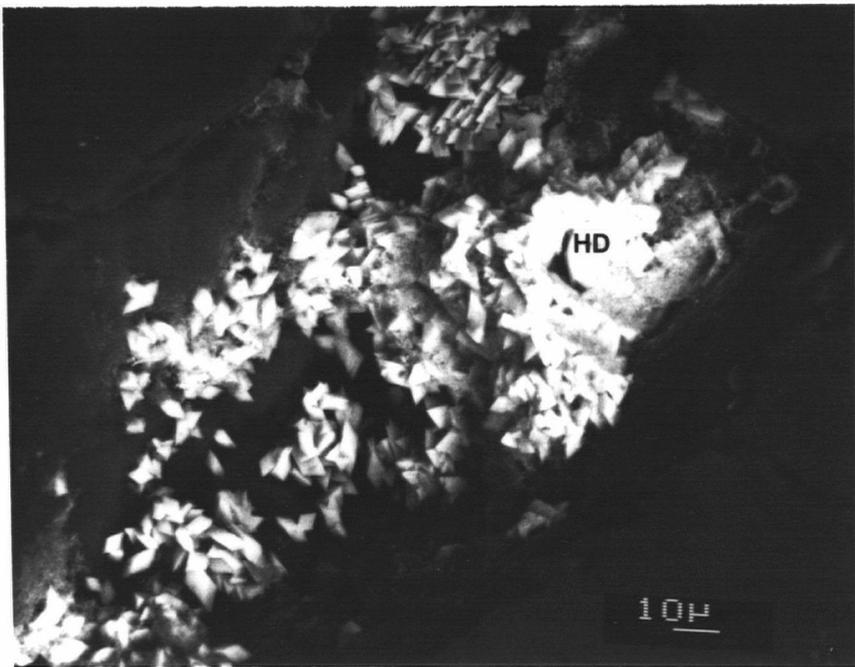
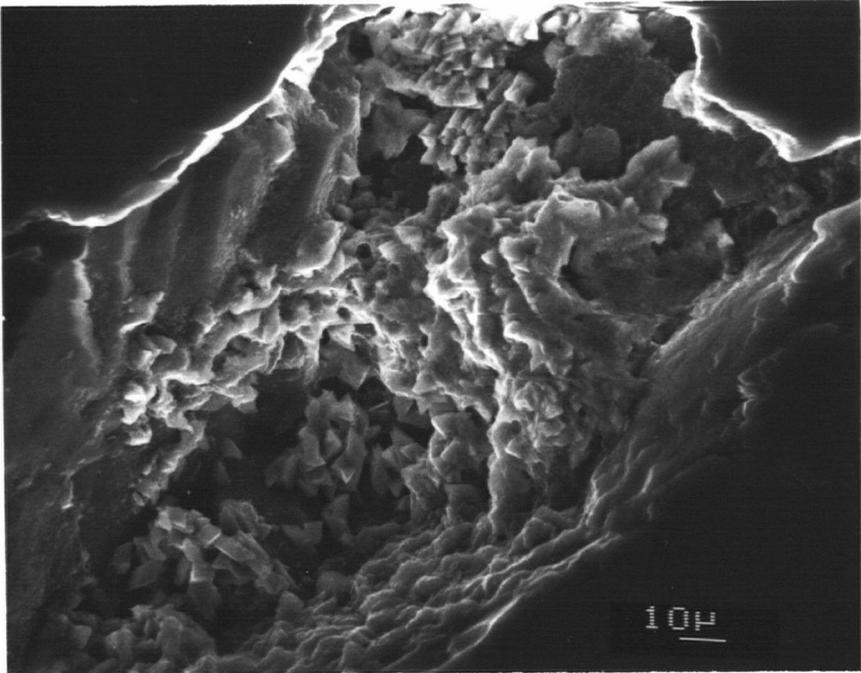
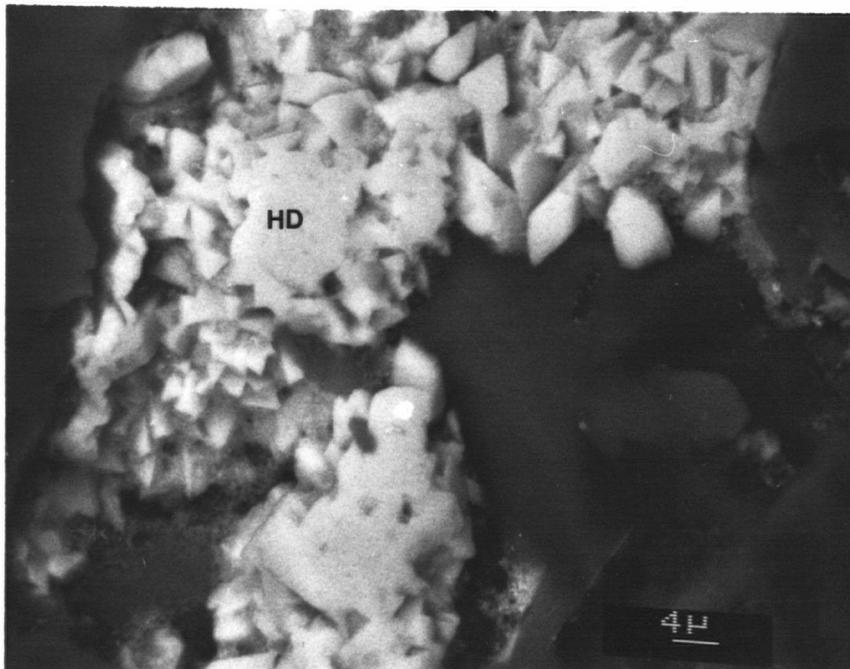
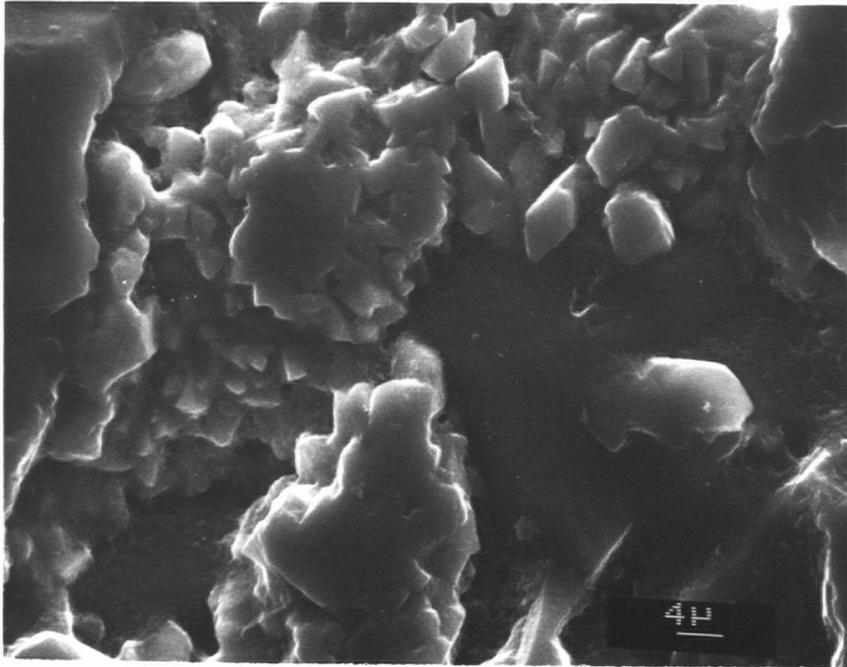


Plate 29

SEM photomicrographs (secondary electrons above;
backscattered electrons below) of
hinsdalite (HD).

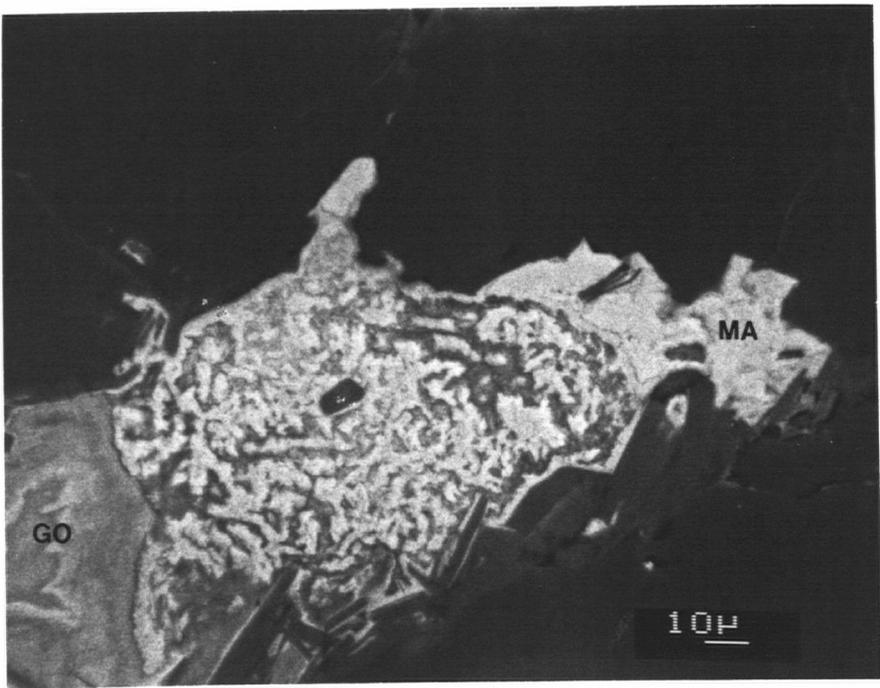


intergrowth of malachite and limonite (limonite type F) only where grain boundaries intersect the cupriferous goethite crystal (Plate 29). Some (<10%) of the limonite is manganiferous and these are the only limonites which contain intergrown native silver and malachite. Abundant malachite occurs within quartzites containing detrital muscovite and probably formed from the oxidation of chalcocite, and, where associated with limonite, bornite. Minor chrysocolla and chalcantite have replaced both bornite and chalcocite. Limonite appears as a replacement of galena, chalcopyrite, bornite, chalcocite, cupriferous goethite, acanthite, and possibly ferroan carbonate). Minor native wire silver has been observed in panned concentrates of soils on the prospect.

Thus four separate episodes of reactions involving metals in the Daisy Creek prospect can be identified. Each can be separated from the other by distinct mineralogic criteria and paragenetic position. Although the first episode (early diagenesis), and the last episode (supergene oxidation), occur separately in time, the middle two (during middle diagenesis) probably represent a continuum. The reducing event probably evolved from a replacement reaction under reducing conditions, to a replacement reaction under oxidizing conditions. A possible reaction mechanism for this includes an oxidizing metal-chloride complexed ground water (Shipulin, 1971; Renfro, 1974; Rickand, 1974; Rose, 1976; Whipple, 1984) interacting, through the diffusion of metals from the oxidizing to reducing conditions, with an early diagenetic pyrite-bearing permeable host rock containing reduced pore

Plate 30

SEM photomicrograph (backscattered electrons) of myrmekitic intergrowth of malachite (MA; light) and limonite (LM; dark). Note at left the liesegang diffusion rings in the cupriferous goethite (GO).



waters. Mineralization may have been a result of the infiltration of metalliferous ground waters adjacent to this reducing environment, and involved the initial replacement of pyrite by Cu-sulfide phases, probably as a result of copper diffusion from the oxidizing to reducing environment. Subsequent diffusion of free oxygen into the reducing environment resulted in the replacement of the Cu-sulfides by Fe-oxides.

CHAPTER 4

EXPLORATION GEOCHEMISTRY

Stream Sediment Geochemistry

During the summer of 1982, a stream sediment geochemistry survey was conducted by Anaconda Minerals Company. A data set comprising 416 samples surrounding the Daisy Creek Prospect from a portion of this survey was made available for interpretation.

This data set consists of sample number, location, and raw element concentration variables. Elements determined include Cu, Pb, Zn, Mo, Ag, Co, Mn, Fe, Ba, and F. Concentrations reported for all elements are in ppm, except for Fe, which is reported in percent. Detection limits are presented in Table I. A listing of the raw data is included in Appendix 2.

Each stream sediment sample consisted of a 1 kg composite of 10 subsamples collected along 30 m of the stream. Every attempt was made to collect non-organic fine stream detritus. Samples were collected every 750 m along the streams and above all stream confluences on each fork. This resulted in a sample density of approximately two samples per km² and a total coverage of over 210 km².

Samples were air dried, sieved to -80 mesh, and submitted for analysis by the Anaconda Minerals Company Sample Preparation Laboratory located in Monte Vista, Colorado. Analysis of the samples was made using atomic absorption spectrometry for all

Table I

Detection Limits for Stream Sediment Data
(values in ppm)

<u>Element</u>	<u>Detection Limit</u>	<u>Recode Value</u>
Cu	2.0	*
Pb	2.0	*
Zn	5.0	*
Mo	1.0	0.5
Ag	0.2	0.2
Mn	20.0	*
Fe	5000.0	*
F	20.0	*
Ba	20.0	*
Co	2.0	1.0

* denotes no samples at or below detection

elements except F, which was analyzed by the specific ion electrode technique. Dissolution was accomplished using aqua regia. All geochemical analyses were carried out by Bondar-Clegg Laboratories, Vancouver, B.C.

Control samples, collected in conjunction with the exploration samples, consist of 16 pairs of field replicates. These were samples which were collected at the same field site as the exploration samples, but were analyzed independently. These samples were thus used to determine the amount of combined sampling, preparation, and analytical error for each analysis.

Error Analysis

Samples which were collected as replicates were evaluated to determine the precision of the geochemical analyses (Meisch, 1967; Plant et al., 1975; Garrels and Goss, 1978). A listing of the replicate data is presented in Table II.

Statistical analysis of this data was made by comparing the median concentration of a replicate pair against the medial coefficient of variation for that pair ($CV_m\%$). The medial coefficient of variation is the medial standard deviation, derived using the median in the standard deviation formula instead of the mean, divided by the median. This evaluation technique is analogous to that described by Thompson and Howarth (1973, 1976a, b, 1978) for duplicate analytical data. The median $CV_m\%$ value is the best estimate of the relative error for a set of replicate analyses because it is a non-parametric

Table II
Replicate Samples from Stream Sediment Survey

SAMP NUM	REP NUM	Cu ppm	Pb ppm	Zn ppm	Mo ppm	Ag ppm	Mn ppm	Fe pct	Ba ppm	F ppm	Co ppm
33783	1	9	11	37	2.0	0.2	660	1.70	510	230	8
33784	2	8	11	37	2.0	0.2	615	1.70	490	300	6
33805	1	33	22	66	2.0	0.3	1000	2.10	950	350	9
33807	2	35	22	70	1.0	0.4	1150	2.20	970	430	9
33844	1	20	19	81	2.0	0.2	800	2.50	730	520	14
33845	2	24	19	84	2.0	0.2	820	2.50	710	430	14
38061	1	40	45	99	1.0	0.2	835	2.10	510	390	9
30638	2	48	52	102	1.0	0.4	960	2.20	530	250	13
46526	1	4	11	66	2.0	0.2	575	1.20	740	270	5
47400	2	6	10	63	1.0	0.2	530	1.20	720	220	4
46539	1	39	18	61	2.0	0.4	840	2.20	730	550	8
46540	2	41	18	61	3.0	0.6	770	2.15	700	480	8
46620	1	37	37	90	2.0	0.2	1100	2.25	860	630	9
46621	2	48	44	107	3.0	0.2	1050	2.40	840	420	9
46646	1	38	22	77	3.0	0.4	1200	2.10	870	540	7
46647	2	36	22	78	2.0	0.2	1250	2.10	860	770	8
46693	1	22	16	64	2.0	0.2	865	1.55	1010	420	8
46694	2	17	18	60	1.0	0.2	665	1.50	890	570	8
46836	1	15	14	45	2.0	0.2	540	1.25	780	270	7
46837	2	14	12	47	1.0	0.2	495	1.20	770	300	6
46842	1	10	11	75	3.0	0.2	258	1.65	770	400	6
46843	2	10	10	71	2.0	0.2	232	1.70	780	480	6
46844	1	37	15	62	2.0	0.3	1200	2.40	970	460	11
46847	2	37	15	60	3.0	0.2	1200	2.50	950	550	12
47417	1	64	21	65	2.0	0.2	1000	2.05	930	460	9
47378	2	50	18	72	3.0	0.2	830	1.90	880	390	8
47420	1	55	40	57	2.0	0.2	1400	1.80	750	360	10
47421	2	52	36	56	2.0	0.2	1200	1.75	740	400	10
61129	1	19	26	56	1.0	0.2	660	1.80	860	570	9
61130	2	20	23	57	0.5	0.2	750	1.95	820	600	8
68663	1	12	13	72	1.0	0.2	770	1.35	950	330	7
68664	2	15	13	64	0.5	0.2	600	1.45	920	270	5

approximation of the error. Results are presented Appendix 3.

The replicates demonstrate excellent precision for geochemical analyses. This may be due in part to the fact that most of the replicate samples were analyzed sequentially; thus, contamination and analytical drift were minimized. Relative errors for most elements are below 10%.

Many of the calculated $CV_m\%$ values for Ag and Mo are extremely high due to the lack of analytical precision at the near-detection limit concentrations (Table I) of the analyzed replicate samples. In the case of F, the only element determined by a technique other than atomic absorption spectrometry, the higher relative error of 9.7% may simply be related to a lower level of precision inherent in the specific ion electrode technique.

Statistical Analysis

Statistical evaluation of the geochemical data consisted of an initial probability plot analysis for each element to evaluate whether multiple populations were represented in the data (c.f., Sinclair, 1976). No obvious multimodal distributions were described (Appendix 4); however, all elements were lognormally distributed so further statistical analysis was performed on the logarithm of the element concentration.

In order to simplify the interpretation of the geochemical data, additional statistical evaluation was limited to operations on only the Cu concentration variable. This variable was chosen

because it is a major pathfinder element and ore constituent in stratabound copper deposits of the Belt Basin and the most probable economic target in the survey area.

Since classification of anomalous Cu concentrations based on multimodal distribution discrimination was impossible, arbitrary thresholds were chosen to define the anomalous samples. Two different, but often used, threshold selection techniques are presented.

The first involves the selection of a threshold at the mean plus two standard deviations concentration level. For Cu, this corresponds to 97 ppm, based on log-transformed data. Samples above this concentration are dotted in Plate 31. Four anomalous samples are recognized.

An alternative yet similar approach would be to select the threshold at the 95th percentile. This threshold corresponds to 63 ppm Cu and samples greater than this threshold are stippled on Plate 31. Numerous single point and several multi-point anomalies are recognized.

These two methods vary only in the number of samples which are recognized as anomalous. Classification is basically a boolean operation: anomalous, or not anomalous. No other information is provided. Anomalies might be ranked for follow-up based on the concentration level of the different anomalous samples, or based on the number of anomalous samples which are contiguous and thus constitute multiple point anomalies. Unfortunately, these criteria may be completely unrelated to criteria which define where mineralization might be located or

how it is expressed in the surface geochemical environment. This may be because the sample density is not appropriate for the mineralization type being sought, or because different lithologic units in the survey area have different background levels and true anomalous samples may not have the highest concentrations.

One would be forced, using this type of procedure, to follow up each of the anomalous samples to see if they are related to mineralization. Utilization of a geologic map might add some information which could be used to eliminate anomalous samples from consideration (e.g., those which occur over unfavorable host rocks), but no insight is added which allows an understanding of why the anomalies occur where they do, until they are actually field checked. No information is available which might rank these anomalies in terms of their potential for being related to mineralization or some other geochemical phenomenon such as contamination.

As a result, an alternative approach is presented which allows further interpretation of the geochemical data at this level of investigation and utilizes the multivariate nature of the available data.

Alternative Statistical Technique

The following statistical technique assumes that all samples fall into one of two categories; anomalous or background. Samples from a background category, or population, constitute those unrelated to any type of irregular geochemical concentration with

respect to the population as a whole, either from a single or multi-element perspective. Samples which are anomalous constitute those which are irregular for some reason.

Unfortunately, in most surveys, several different types of anomalous samples may be present. Different types of mineralization may exist which have different geochemical and mineralogical signatures. Vein-, skarn-, and stratabound-type mineralization each could have different suites of anomalous elements and produce different types of anomalous geochemical signatures. Likewise, different types of weathering processes may have occurred which create different types of geochemical anomalies. Elements which are anomalous in a displaced seepage anomaly may be entirely different from those in a clastically borne dispersion train anomaly, yet both may derive their anomalous metal concentrations from the same mineral deposit.

In a similar manner, there may be more than one type of background sample. The background geochemical category (or population) may actually be composed of several individual populations. These multiple background populations are common in geochemical datasets, and generally are related to fundamental geologic or geochemical factors or processes, such as the lithology which underlies or provides the source material for the sample, or the type of sample collected such as dry streams versus wet streams or B versus C soil horizons. It is up to the geochemist to separate the data into groups which correspond to these criteria and to treat each group separately. The groups can be processed statistically and results from each group can then

be interpreted with a minimum of ambiguity. In many cases, individual groups will demonstrate multimodal distributions not obvious in the data as a whole, while the combination of several multimodal distributions may result in an apparent unimodal distribution (a corollary of the central limit theorem).

The following statistical technique is recommended in cases where a single background population exists, either because the data are homogeneous or because they have already been divided into individual groups.

Threshold Selection

Two thresholds can be chosen to divide the data into three groups if a bimodal distribution is present. These groups constitute an anomalous population, a background population, and an intermediate zone in between (Figures 15 and 16). These thresholds are selected using a probability plot (c.f., Sinclair, 1976), and should be chosen to eliminate background samples from being assigned to the anomalous population and anomalous samples from being assigned to the background population. In this way, both high and low groups are 'pure', in that no other samples, or at least an insignificant number, would be misclassified. The intermediate zone consists of samples where classification into high and low groups is ambiguous because it represents the range of overlap of the background and possible anomalous populations.

Note that because we know that there is only one background population in the data, since we would have subdivided it into

THRESHOLDS FOR HISTOGRAM DISTRIBUTIONS

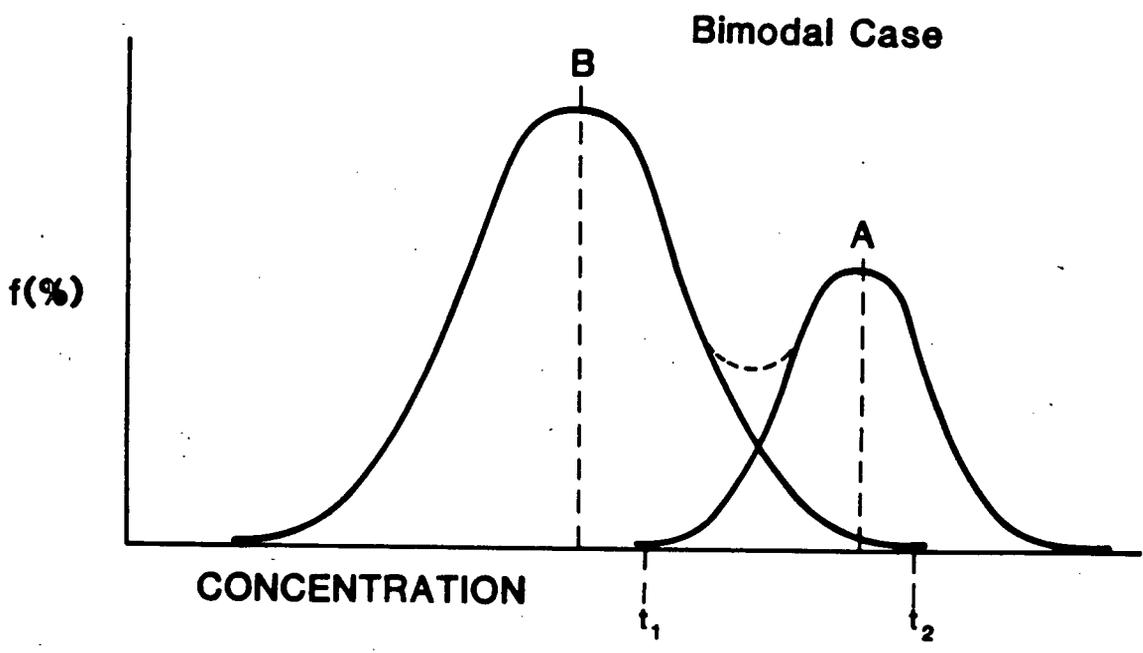
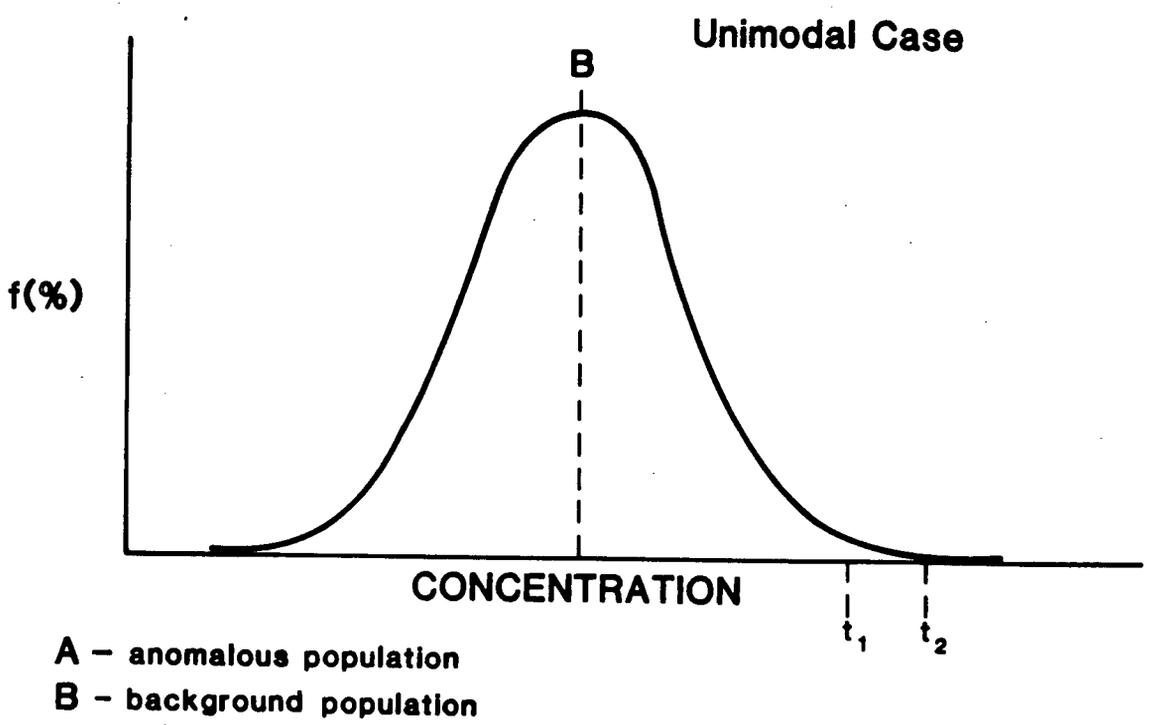


Figure 15

Thresholds for Frequency Distribution

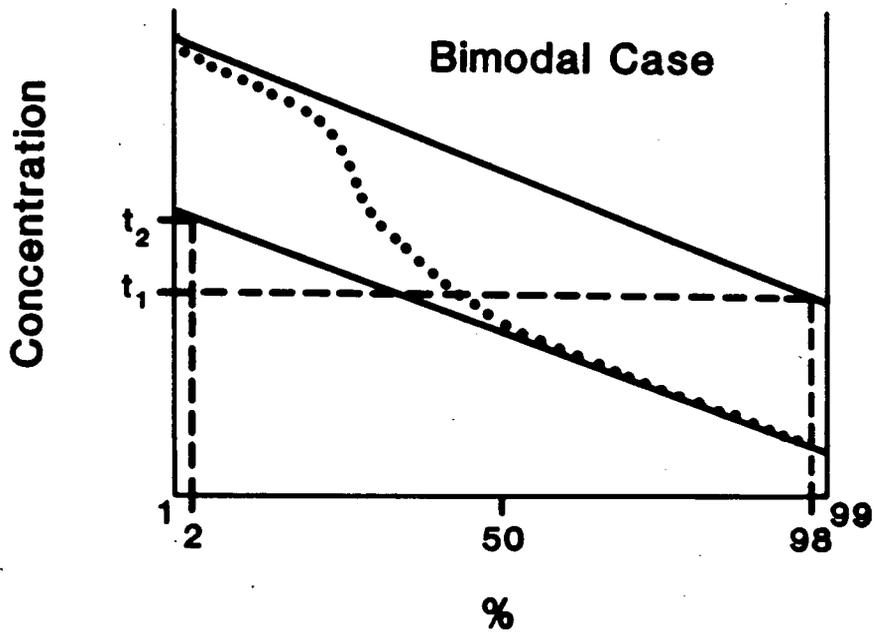
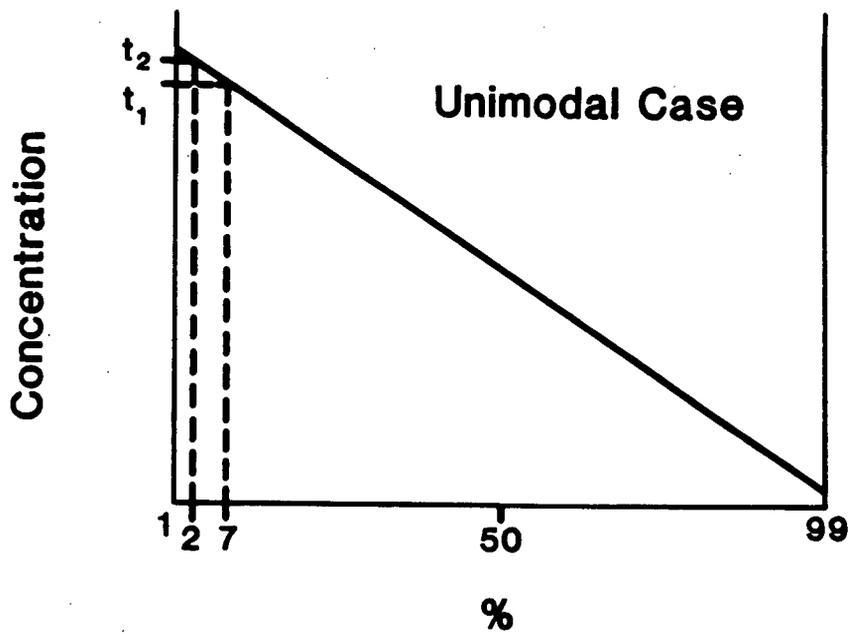


Figure 16

fundamental groups if there is more than one, only one type of background sample is present in the background population. This may not be the case in the anomalous population, because several different types of anomalies may be present.

An arbitrary choice of thresholds can be made if a unimodal population distribution exists. We assume that anomalous samples are present in the data, but they may not be sufficiently different from the background samples, at least on a single element basis, or do not constitute a large enough proportion of the data to be recognizable on the probability plot. Thus we choose a threshold which we are certain will be low enough to prevent the anomalous samples from being classified as background samples. This provides us with a 'pure' background dataset. In the case presented, a threshold was chosen at the 93rd percentile (mean plus 1.5 standard deviations in a perfect normal distribution), which corresponds to 60 ppm Cu.

In addition, a higher threshold at the 98th percentile has also been chosen (the mean plus two standard deviations in a perfect normal distribution), which corresponds to 90 ppm. This threshold may separate the zone of overlap from the anomalous population. In this case, since we do not have any evidence to support a bimodal distribution, and thus a second population, this threshold may be of little consequence, but it has been chosen to demonstrate the statistical technique in a manner consistent with the case where two thresholds are essential to the procedure.

Whether we now have a bimodal or unimodal population, the

data have been divided into three different groups. In both cases the lower, background group is 'pure', that is, all contained samples have a high probability of representing background.

Samples from the high group, with concentrations greater than the upper threshold, generally would be considered to be anomalous if a bimodal distribution existed. Thus, no further evaluation need be performed on these samples at this stage. Such is not necessarily the case for a unimodal distribution.

Samples from the low group, with concentrations less than the lower threshold, are considered to have too low a concentration to be of interest and are automatically categorized as background samples.

Samples from the third group, with concentrations between the two thresholds, are those whose classification as anomalous or background is indeterminable based solely on the concentration level, given the amount of analytical error in the data and because of population overlap. Such values need to be classified as either part of the anomalous population or part of the background population. The additional statistical analyses presented below are directed toward this end.

One approach to recognizing anomalous samples in the zone of population overlap is to determine what type of multielement geochemical signature an anomalous sample might have, and to recognize other samples which have a similar signature. Unfortunately, in many practical cases too few anomalous samples are available to determine the multielement characteristics of a typical anomalous sample. Additionally, several types of

anomalous geochemical signatures may be present in the data because of different mineralization sources or because several processes may be acting to form different types of geochemical anomalies.

As a result, an alternative approach is suggested which is designed to recognize the anomalous samples using criteria from information provided by the background population. This involves doing the opposite operation, that of determining the multielement geochemical signature for background samples and recognizing those samples which are not similar to this signature. These samples are then classified as anomalous. In this way, all possible anomalous samples are recognized and can be grouped and interpreted as to their origin and significance with further statistical techniques. A characterization of the typical background sample can be made because these samples come from only one population.

Regression Analysis

The method chosen to discriminate between anomalous and background samples in the concentration range of population overlap is a backwards stepwise multiple linear regression technique. A dependent variable, in this example, copper, is regressed against all other elements, the independent variables, for all of the samples which were classified as background using the probability plot technique (i.e., Cu concentrations <60 ppm). The resulting regression function was then applied to the samples

from the zone of overlap. In this case we will apply it to both anomalous and zone of overlap samples because we do not know if the upper group is truly anomalous. A scatter plot of the actual concentration versus the concentration predicted from the regression function is evaluated to determine which samples have a background multielement geochemical signature and which do not.

This technique is similar to that described by Matysek et al. (1982); however, by utilizing only background samples in the determination of the regression function, a high multiple regression correlation is possible because no samples from other populations are included.

Unfortunately, one must remember that every variable in a geochemical data set is subject to error. As a result, normal least-squares regression is not valid because the independent variables are not known precisely. This is an assumption in least-squares regression. An alternative approach must be utilized which accounts for errors in each variable. This technique, known as the 'reduced major axis regression technique', has been described by Dent (1933), Kermack and Haldane (1950), Till (1973), and Mark and Church (1977); however, the derivation of this form of regression has only been worked out for two variables. A description of the derivation for this technique which determines a regression function for more than two independent variables is in preparation.

The reduced major axis technique allows the construction of a regression line which corresponds to the general trend of the raw data where the expectation of the residual is 0 across the full

range of the data; this is not necessarily the case in a normal least-squares regression situation, where the average expectation of the residual is zero (Lindley, 1950; Ezekial, 1961; Kendall and Stuart, 1963; Jones, 1972, 1979; Spiegelman, 1982).

The advantage of this empirical technique is that it allows one to extrapolate the regression line outside the bounds of the data used to define the line from <60 ppm (the background population) into the zone of mixing of the anomalous and background populations (60 to 90 ppm). This enables one to evaluate the similarity of these samples with respect to the general trend of the background population. In this way, quantitative comparisons of residuals for the samples can be made whereby the residuals may be analyzed using probability plots to separate the populations and samples which have multielement geochemical characteristics which are similar to those of the background population can be classified as such.

Interpretation

Results of the reduced major axis backwards stepwise multiple linear regression technique for Cu are presented in Appendix 4. Elements with large contributions to the regression equation include Fe, Pb, Mn, and Ag (in decreasing order; Appendix 5). Figures 17 and 18 are plots of the predicted Cu concentration from the regression function versus the actual Cu concentration. Several samples from the zone of population overlap fall on or near the predicted regression line (slope = 1) and thus are

REGRESSION PREDICTED VS. OBSERVED COPPER CONCENTRATIONS

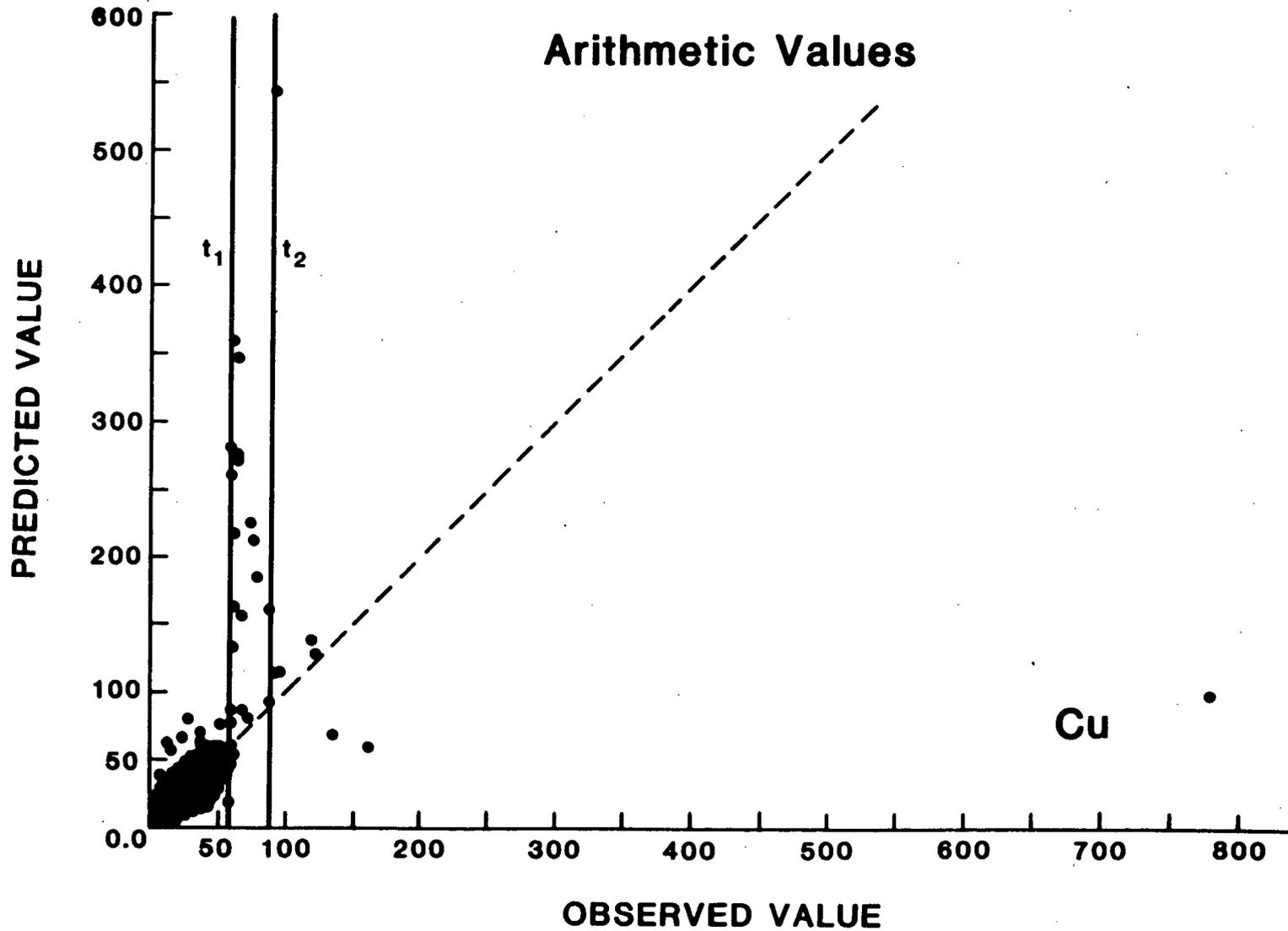


Figure 17

REGRESSION PREDICTED VS. OBSERVED COPPER CONCENTRATIONS

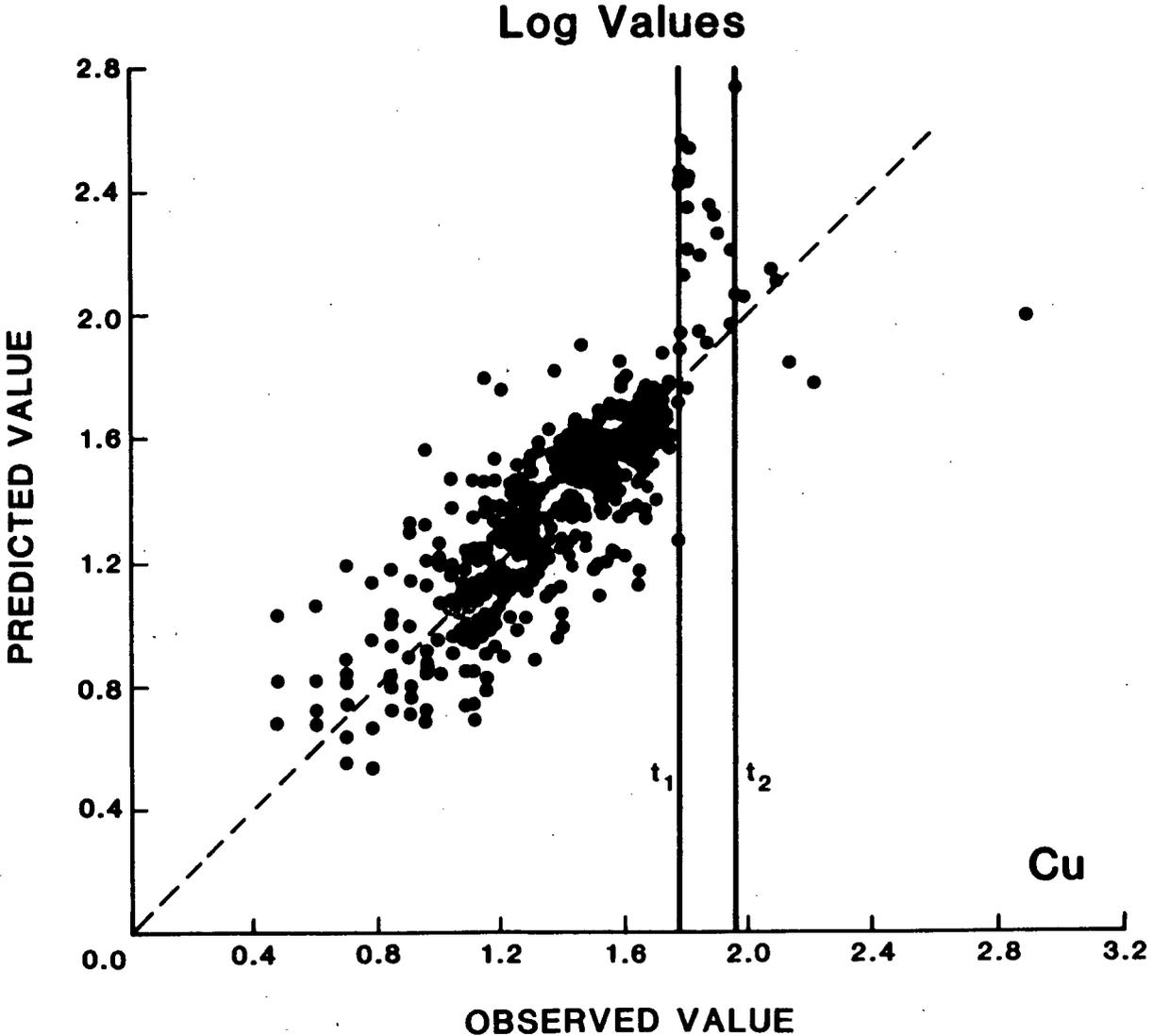


Figure 18

considered to be background samples.

In addition, three obvious trends for samples greater than 60 ppm are observed. One trend, which corresponds to samples with an actual Cu concentration which is greater than that predicted by the regression function (shallow slope), consists of three samples. Four samples, comprising a second trend (intermediate slope), which are deemed anomalous because their Cu concentration is greater than 90 ppm occur along the multiple regression line within the 95th percentile error bounds for the background population. Since we are only assuming that there is an upper anomalous population, these samples may or may not be truly anomalous. Finally, a third trend, which has predicted Cu concentrations which are greater than the actual concentrations (steep slope), comprises the remaining samples which do not fall on or near the regression line.

Plate 32 is a drainage basin map of the area covered by the geochemical survey around Daisy Creek. Drainage basins which contain samples deemed initially anomalous in Cu (>90 ppm) are stippled. Those drainage basins containing samples which are part of the shallow slope trend are lined vertically, those which are part of the steep trend are lined horizontally, and those which lie along the regression line are lined at an angle. Drainage basins containing samples deemed to be background are blank.

Evaluation of the anomalies with the use of the geologic map (Plate 2) reveals several important conclusions:

- 1) a two point anomaly occurs in the South Daisy Creek drainage basin and appears to be related to the

- mineralization at the Daisy Creek prospect; the sample at the head of the drainage basin (sample 46726; Appendix 2) may represent the actual mineralization; it is anomalous in Cu only and visible malachite and Cu-sulfides have been observed in heavy mineral concentrates from the stream; the other sample (46727) may be related to downslope dispersion, adsorbed concentrations of Cu on Fe- and Mn-oxides (because of its high Mn and Fe concentrations), or to contamination by a previous anomalous Cu concentration (46726);
- 2) a two point anomaly (46736 and 47424) occurs in the north end of the survey area at the Meadow Creek anomaly, a region known to contain malachite as discrete grains and coatings on stream detritus and which is underlain by the Snowslip Formation, a unit which contains abundant copper occurrences of the green bed type; samples from this anomaly have extremely high Cu concentrations;
 - 3) two single point anomalies (46699 and 68157) occur over the Bonner Formation, the same unit which hosts the Daisy Creek prospect; these may be related to adsorbed Cu on Mn- or Fe-oxides, because of their high Mn or Fe concentrations, or to other minor stratabound Cu-Ag occurrences in the Bonner Formation;
 - 4) two multiple point anomalies occur on the east side of the survey area and are underlain by the Wallace Formation (30739 and 38062; to the north) and the Snowslip Formation (33886, 33887, and 33888; to the south); both

occur adjacent to occupied dwellings and may represent contamination from these sources; the anomaly to the north has high Pb, Fe, and Mn concentrations, possibly suggesting that Cu, as well as these elements, were fixed by high pH's created by the carbonate bedrock. The anomaly to the south has high Fe and Mn concentrations, which may suggest Cu adsorption. Alternatively, these samples are sequentially ordered, and thus may represent contamination during analysis;

- 5) a semi-continuous set of anomalous concentrations occur along the 355° striking high angle fault at the north end of the field area; these anomalies (46827, 46833, and 47417) occur downhill from the Meadow Creek anomaly along this structure and may represent displaced anomalies created by ground water circulation and hydromorphic redistribution of metals along the fault to regions separate from the mineralization source; anomalous samples along this trend have high Fe and Mn concentrations and thus may contain large amounts of adsorbed Cu on accompanying Fe- and Mn-oxides;
- 6) a four point anomaly located at the head of the Daisy Creek drainage basin (46678, 46679, 47410, and 68665) is anomalous in Fe, and may be a seepage anomaly from the mineralization at the Daisy Creek prospect; these samples are on the opposite side of the valley as the Daisy Creek Prospect at about the same elevation; consequently, this interpretation is not likely. These

samples may be related to green bed-type mineralization occurring in the underlying Mount Shields Formation; The Purple and Green Argillite Unit of the Siltite Member contains minor disseminated limonite blebs which are similar to those at the Daisy Creek prospect; alternatively, these samples are those which are greater than 90 ppm but which fall along the multiple regression line (intermediate slope) and thus have a similar multi-element geochemical signature as the background population; they may not, in fact, be truly anomalous; the large area covered by this anomaly may not be related to a point source, representing mineralization, as at the Daisy Creek prospect and the Meadow Creek anomaly, but rather to a large zone related to high background concentrations in that area.

Both the Daisy Creek prospect and the Meadow Creek anomaly contain anomalous samples which fall along the shallow slope trend on the regression scatterplot (Figure 17). These anomalous samples are all at the heads of their respective drainage basins and are at fairly high elevations. Areas such as these generally have immature stream sediments which have not undergone extensive chemical weathering. At both locations, malachite has been observed coating rocks in the stream bed (c.f., Cazes, 1981). As a result, the shallow slope trend is considered to be related to detritus containing mineralogically borne metal concentrations in malachite, cupriferous Fe-oxides, and Cu-sulfides, which are all seen in heavy mineral concentrates from these areas.

In contrast, the samples comprising the steep slope trend of the regression scatterplot (Figure 17) are located in sites where hydromorphic dispersion and seepage anomalies seem reasonable given the geologic environment. Some occur downstream from known mineralization, adjacent to dwellings, and along major faults. In addition, they generally contain high concentrations of Fe and/or Mn. For this reason, the steep slope trend of anomalous samples is interpreted as representing anomalous Cu concentrations which resulted from hydromorphic metal transport, and which are now adsorbed to oxides.

Finally, samples which are part of the anomalous group which clusters along the regression line are interpreted to be either not anomalous at all, or a seepage anomaly related to minor green bed occurrences in the Mount Shields Formation.

Note that samples which originally were classified as anomalous are not necessarily related to the two areas of mineralization and that samples from the zone of population overlap are, in many cases, as important to the overall interpretation as those samples initially classified as anomalous.

Using this anomaly recognition technique, a determination can be made as to whether several types of anomalies are present in the data. These anomaly types can be recognized and distinguished based on the multi-element geochemical trends of the anomalous samples. Ranking of the anomalies can be made based on combined input from the statistical evaluation and geologic criteria. Finally, since a rigid threshold value is not applied, samples

which comprise the higher concentration tail of the background population can be recognized and eliminated from consideration.

Conversely, for the standard anomaly ranking techniques described previously, a rigid arbitrary threshold is assigned, and, thus, samples from the high concentration tail of the background population cannot be eliminated. This results in an abundance of uninterpretable single point anomalies which still have to be field checked.

The most prominent anomaly in the 95th percentile threshold example is located at the head of the Daisy Creek drainage basin, a location which is thought to be unimportant based on the probability plot/regression technique and where no mineralization is known to exist. The mean plus two standard deviation technique recognizes only four samples and incorrect priority might be assigned because of the very high concentrations at the Meadow Creek anomaly (785 ppm Cu), and because it consists of two contiguous anomalous samples.

Thus, the probability plot/regression technique described above offers significant advantages over the standard 95th percentile and mean plus 2 standard deviations techniques, and represents a better analytical tool in the evaluation of geochemical data. Anomaly ranking, elimination of random scatter, and multiple anomalous population recognition provide additional interpretive criteria for the exploration geochemist and allow the ready recognition of significant geochemical anomalies.

Soil Geochemistry

Pedogeochemical data from the 1982 soil grid survey was also evaluated statistically and interpreted. A geochemical data set of 116 samples was provided by Anaconda Minerals Company. Information for each sample includes sample number, location coordinates, and element concentrations for Cu, Pb, Zn, Mo, Ag, Co, As, Hg, and Ba. A listing of these data is presented in Appendix 6. Detection limits for the geochemical concentration data are given in Table III.

All samples consisted of 1 kg of material from a single soil pit. Depths ranged from 10 to 35 cm but all samples were collected from the top 15 cm of the Bm soil horizon. These soils were generally less mature than the average soil underlying the regional map area and are classified as orthic-dystric to degraded-dystric brunisols (Miller, 1982). No development of any Ae or Bt soil horizon occurs.

The samples were air dried, sieved to -80 mesh and submitted for analysis by the Anaconda Minerals Company Sample Preparation Laboratory located in Monte Vista, Colorado. All geochemical determinations were done using atomic absorption spectrometry, except for As (colorimetric), and Hg (gold foil). These analyses were carried out at Skyline Laboratories, Golden, Colorado.

Sample sites are located on a N-S oriented square grid which is 15 m on a side. The total soil grid measures 180 m by 180 m.

The survey consisted of 107 routine soil samples and 9 site replicate samples collected at three different grid locations.

Table III

Detection Limits for Soil Grid Data
(values in ppm)

<u>Element</u>	<u>Detection Limit</u>	<u>Recode Value</u>
Cu	5.0	*
Pb	5.0	2.5
Zn	5.0	*
Mo	1.0	**
Ag	0.2	0.1
Co	5.0	2.5
As	5.0	5.0
Hg	0.005	*
Ba	20.0	*

* denotes no samples at or below detection

** denotes no samples above lower detection limit

These site replicate samples were located at points approximating the vertices of an equilateral triangle three meters on a side, which was centered around a routine exploration sample. All were collected in non-anomalous terrane.

In addition, 26 samples were resubmitted for a second analysis of all elements except Hg. Three of these sample replicates consist of known and well blended standards. These sample replicates, along with the site replicates, were considered together and used to evaluate the sampling, preparation, and analytical error for the geochemical analyses.

Error Analysis

Replicate sample data (Table IV) were processed statistically in the same manner as the stream sediment replicate samples. Plotted results are presented in Appendix 7. Relative errors are low and acceptable for most elements (approximately 5%). Copper, Pb, Zn, Ag, and As all exhibit higher levels of relative error at lower concentrations, near their respective detection limits. One replicate analysis for Ba is highly imprecise, but the remainder are acceptable. Molybdenum, Co, and to a lesser degree, Ag and Hg all have most of their replicate values at or near their respective detection limits, so the relative error determined is not significant. More realistic relative error estimates, obtained by inspection based on the average CV%, for these elements would be 10.0% for Mo, 3.0% for Ag, 6.0% for As, and 5.0% for Hg. All are within acceptable levels for geochemical

Table IV

Replicate Samples From Soil Grid Survey

SAMP NUM	REPL NUM	Ag (ppm)	Cu (ppm)	Pb (ppm)	Zn (ppm)	Co (ppm)	Mo (ppm)	As (ppm)	Hg (ppb)	Ba (ppm)
88858-V	1	0.1	25	2.5	50	2.5	1	5	30	670
88860-V	2	0.1	25	2.5	45	10.0	1	5	20	850
88862-V	3	0.1	20	2.5	45	10.0	1	5	30	650
88864-V	4	0.1	25	2.5	45	10.0	1	5	20	680
88886-V	1	0.1	30	2.5	45	10.0	1	5	50	780
88888-V	2	0.1	75	2.5	80	2.5	1	5	60	750
88890-V	3	0.1	15	10.0	70	10.0	1	5	40	870
88812-V	4	0.8	20	5.0	55	10.0	1	5	30	690
88972-V	1	0.1	30	15.0	45	5.0	1	5	10	570
88974-V	2	0.1	30	10.0	70	10.0	1	5	20	660
88976-V	3	0.1	25	20.0	45	5.0	1	10	30	470
88978-V	4	0.1	50	15.0	90	5.0	1	5	20	650
10001-S	1	18.0	110	335.0	165	2.5	4	1100	300	1050
10002-S	2	17.0	110	335.0	160	5.0	6	1050	300	1000
10003-S	3	18.0	115	330.0	170	5.0	6	1100	320	920
10004-S	4	19.0	110	375.0	160	5.0	6	960	300	1050
10011-S	1	5.9	70	75.0	170	10.0	1	390	100	1000
10012-S	2	5.4	70	70.0	140	10.0	4	430	100	1000
10013-S	3	5.6	75	70.0	155	10.0	4	370	90	1050
10021-S	1	0.1	60	20.0	235	15.0	1	30	20	830
10022-S	2	0.1	65	15.0	240	15.0	2	20	20	890
10023-S	3	0.1	50	20.0	225	5.0	4	20	20	870
88956	1	0.1	35	10.0	95	10.0	1	5		**470
	2	0.1	35	10.0	45	5.0	1	5		530
88028	1	0.1	60	20.0	235	2.5	1	5		610
	2	0.1	50	10.0	35	5.0	1	5		590
88972	1	0.1	35	15.0	50	2.5	1	5		540
	2	0.1	30	15.0	45	5.0	1	5		570
88984	1	0.1	125	15.0	75	10.0	1	5		720
	2	0.1	130	15.0	70	5.0	1	5		780
88928	1	0.1	5	5.0	25	5.0	1	5		400
	2	0.1	15	2.5	20	2.5	1	5		440
88810	1	0.1	135	2.5	70	5.0	1	5		800
	2	0.4	125	2.5	70	10.0	1	5		850

Table IV Cont.

SAMP NUM	REPL NUM	Ag (ppm)	Cu (ppm)	Pb (ppm)	Zn (ppm)	Co (ppm)	Mo (ppm)	As (ppm)	Hg (ppb)	Ba (ppm)
88842	1	0.1	35	5.0	50	5.0	1	10		610
	2	0.1	30	5.0	50	5.0	1	5		240
88878	1	0.1	30	10.0	45	10.0	1	10		570
	2	0.8	30	5.0	50	2.5	1	5		620
88002	1	0.1	40	20.0	35	5.0	1	10		470
	2	0.1	35	20.0	35	5.0	1	10		440
88832	1	1.7	1050	15.0	65	5.0	1	10		790
	2	1.7	950	20.0	65	2.5	1	10		710
88898	1	0.1	40	5.0	60	5.0	1	5		710
	2	0.1	25	5.0	50	2.5	1	30		620
88940	1	0.2	90	205.0	60	10.0	1	5		750
	2	0.2	80	240.0	55	15.0	1	5		750
88061	1	0.4	90	5.0	80	10.0	1	5		620
	2	0.4	85	10.0	75	2.5	1	5		610
88078	1	0.7	50	5.0	75	5.0	1	5		590
	2	0.6	55	10.0	75	10.0	1	5		520
88019	1	0.1	30	5.0	40	2.5	1	5		620
	2	0.1	30	2.5	35	10.0	1	5		540
88064	1	0.4	40	15.0	120	5.0	1	5		620
	2	0.4	35	15.0	120	2.5	1	5		580
88094	1	1.0	35	5.0	85	5.0	1	5		600
	2	0.8	50	10.0	80	10.0	1	5		510
34739	1	0.1	35	5.0	105	10.0	1	5		690
	2	0.4	25	2.5	95	10.0	1	10		630
88870	1	0.1	15	2.5	35	5.0	1	5		570
	2	0.1	15	2.5	35	5.0	1	5		570
88910	1	0.1	15	2.5	35	2.5	1	5		530
	2	0.1	15	2.5	30	2.5	1	5		520
88998	1	0.7	30	10.0	85	5.0	1	5		530
	2	0.6	35	15.0	105	10.0	1	5		550
70001-R	1	0.1	15	20.0	65	5.0	1	10		610
	2	0.1	15	20.0	65	10.0	1	10		580

Table IV Cont.

SAMP NUM	REPL NUM	Ag (ppm)	Cu (ppm)	Pb (ppm)	Zn (ppm)	Co (ppm)	Mo (ppm)	As (ppm)	Hg (ppb)	Ba (ppm)
70002-R	1	0.1	5	2.5	20	2.5	1	5		240
	2	0.1	15	2.5	20	2.5	1	5		280

V - represents site replicate samples

S - represents standards

R - represents rock chip sample

* - All non-designated samples are duplicate soil samples.

** - All standards, rock chip, and duplicate soil samples are sample replicates (the crushed material was submitted a second time for analysis); site replicates are actually a different sample collected at the same site.

analyses.

At one replicate site, the central exploration sample and two of the three replicate samples are in good agreement; however, the third replicate sample is different. This is because it was collected in a gully which drains an area of anomalous Cu, Pb, Zn, and Ba soil concentrations. This sample is more of a stream sediment sample than a soil sample and its higher metal concentrations appear to be due to clastic and hydromorphic dispersion downhill from the corresponding soil anomalies. The other three samples at this replicate site were collected over routine non-anomalous soil locations. This discrepancy imposed an artificially high relative error measurement for this replicate group but has not increased the overall relative error determination significantly.

Statistical Analysis

The average concentration of most of the elements analyzed in the soil survey are extremely low, especially for residual soils. This may be due to the inability of the soils to fix many of the soluble metals of interest. Because the underlying bedrock is composed primarily of quartz with hematite cement, resulting soils have similar compositions, and these soil constituents have little potential for adsorbing and concentrating metals. In addition, bedrock beneath the soil grid contains low amounts of mafic minerals and other impurities, such as volcanoclastics, which normally contain large amounts of base metals. Thus,

background concentrations for these metals are very low. Nevertheless, rocks containing economic minerals do have anomalous soil concentrations developed over them.

Statistical evaluation of the pedogeochemical data was limited to a probability plot analysis because this technique effectively separated distinct populations within the data; hence, additional processing was deemed unnecessary. Most elements display obvious multimodal probability density functions, and have easily interpretable anomalous populations which correspond spatially with geologic criteria.

Probability plots are included in Appendix 8 and population distribution (anomaly) maps are presented in Figures 19 to 26. Thresholds used to obtain the maps were derived from the probability plots and represent the optimal concentrations for discriminating anomalous concentrations from the background population (c.f., Sinclair, 1976). Table V contains a list of these thresholds for each element.

Note that Mo has been omitted from consideration because no Mo concentrations are greater than the 1 ppm detection limit. In addition, As and Co anomaly maps are presented (Figures 25 and 26), but they do not exhibit any recognizable trend and have a very low concentration range and contrast.

Interpretation

The soil grid maps show several features. First, Cu defines the most obvious anomaly. Silver, Pb, and to a lesser degree,

Table V

Thresholds for Soil Grid Data
(values in ppm)

<u>Element</u>	<u>Modes</u>	<u>Threshold</u>	<u>Percentile</u>
Cu	3	100.0	22
		220.0	12
Pb	3	15.0	28
		90.0	7
Zn	3	65.0	21
		80.0	7
Ag	3	0.2	4
		0.4	2
Hg	3	0.030	32
		0.050	2
Ba	2	800.0	9
Co	2	10.0	7
As	3	5.0	9
		10.0	2

* Mo is omitted because all values were equal to the detection limit.

Soil Grid Geochemical Concentrations

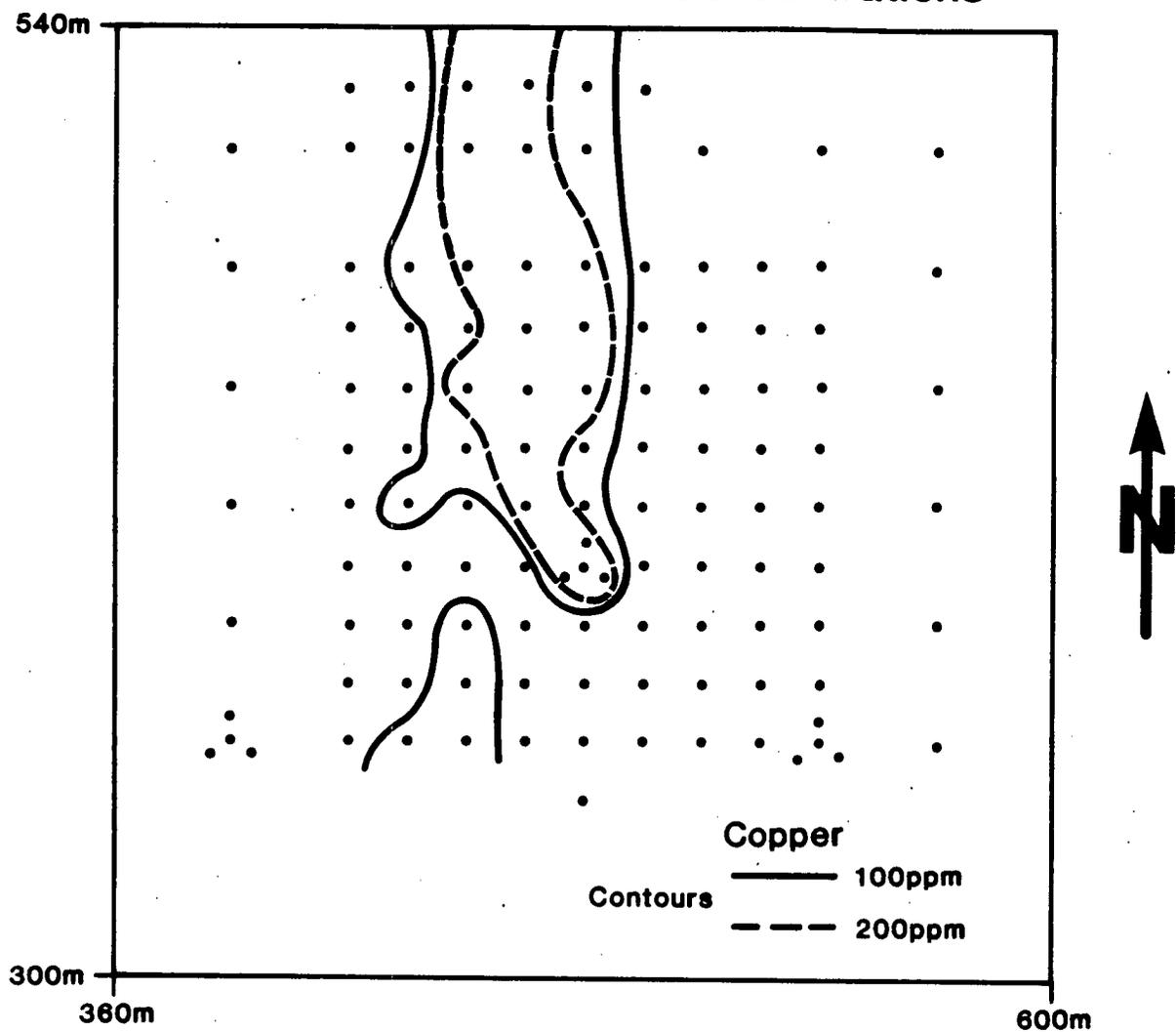


Figure 19

Soil Grid Geochemical Concentrations

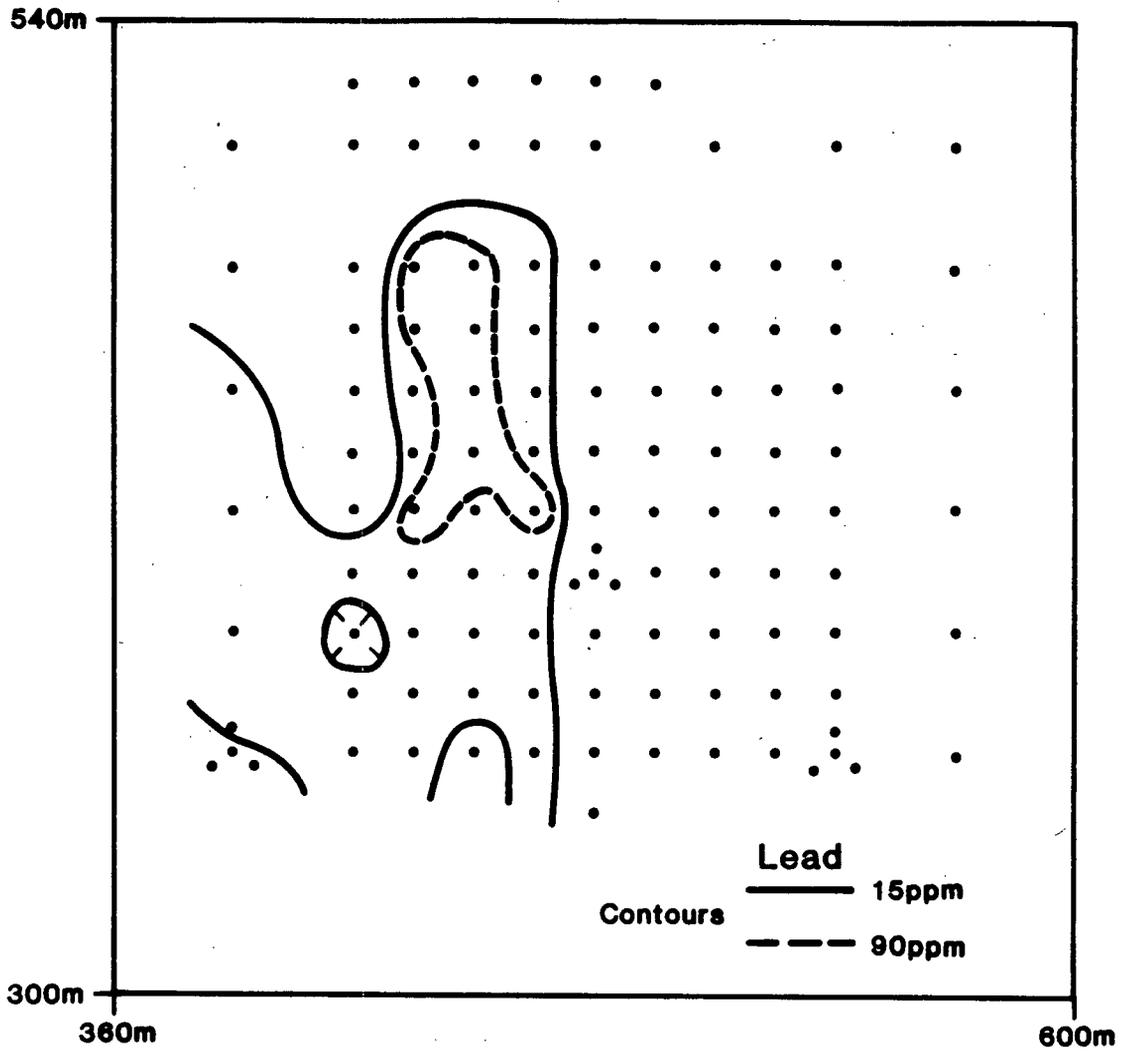


Figure 20

Soil Grid Geochemical Concentrations

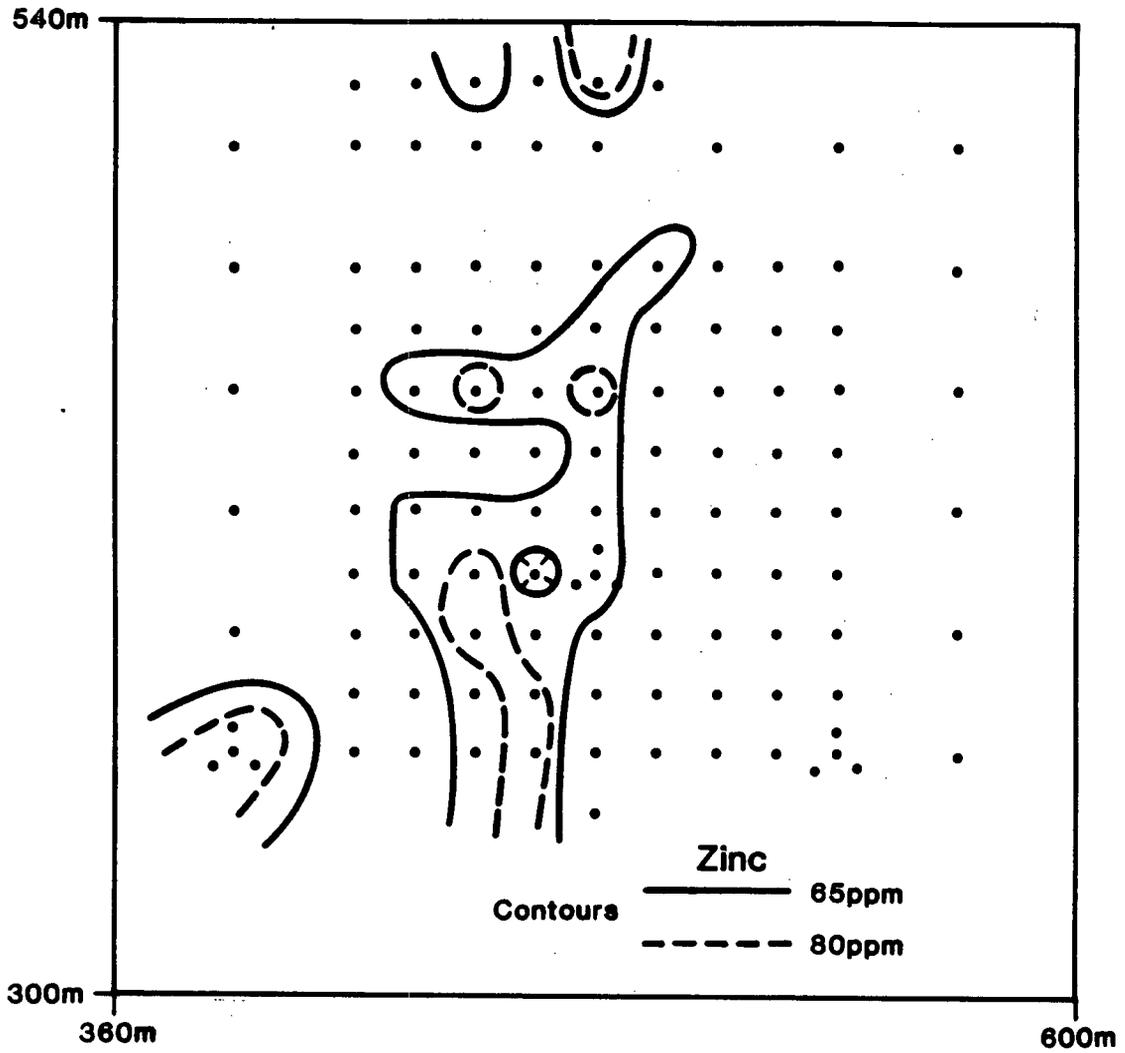


Figure 21

Soil Grid Geochemical Concentrations

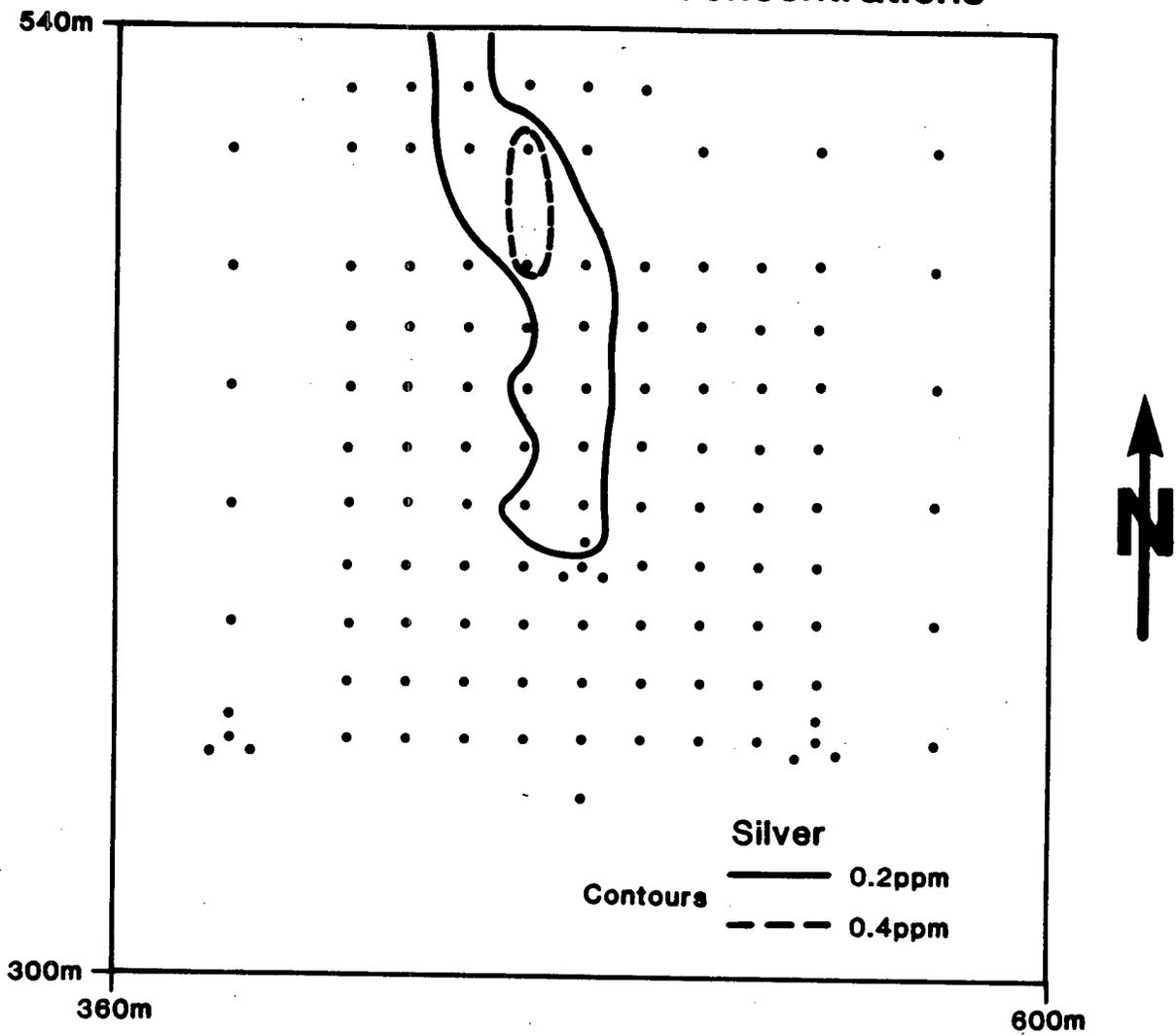


Figure 22

Soil Grid Geochemical Concentrations

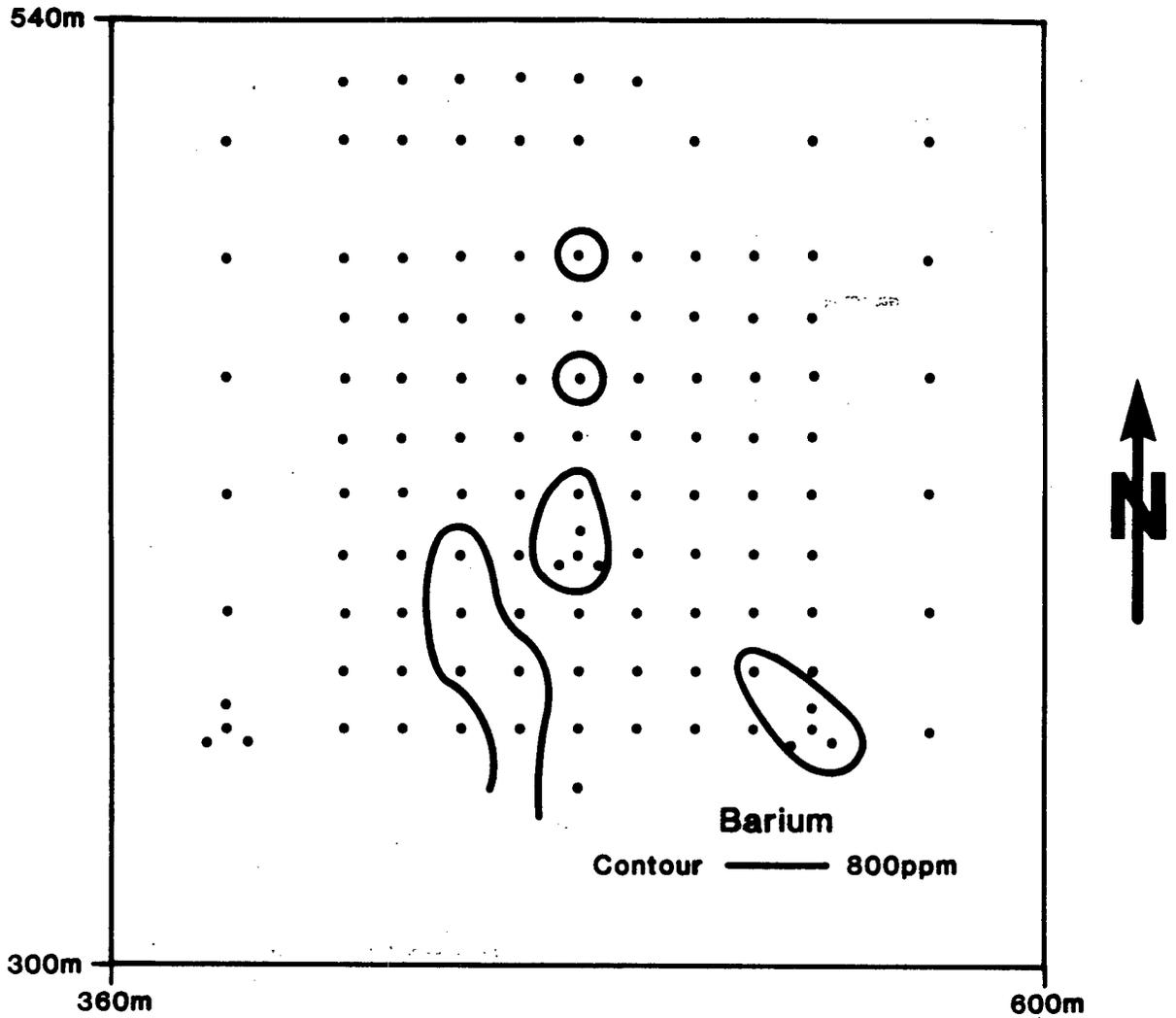


Figure 23

Soil Grid Geochemical Concentrations

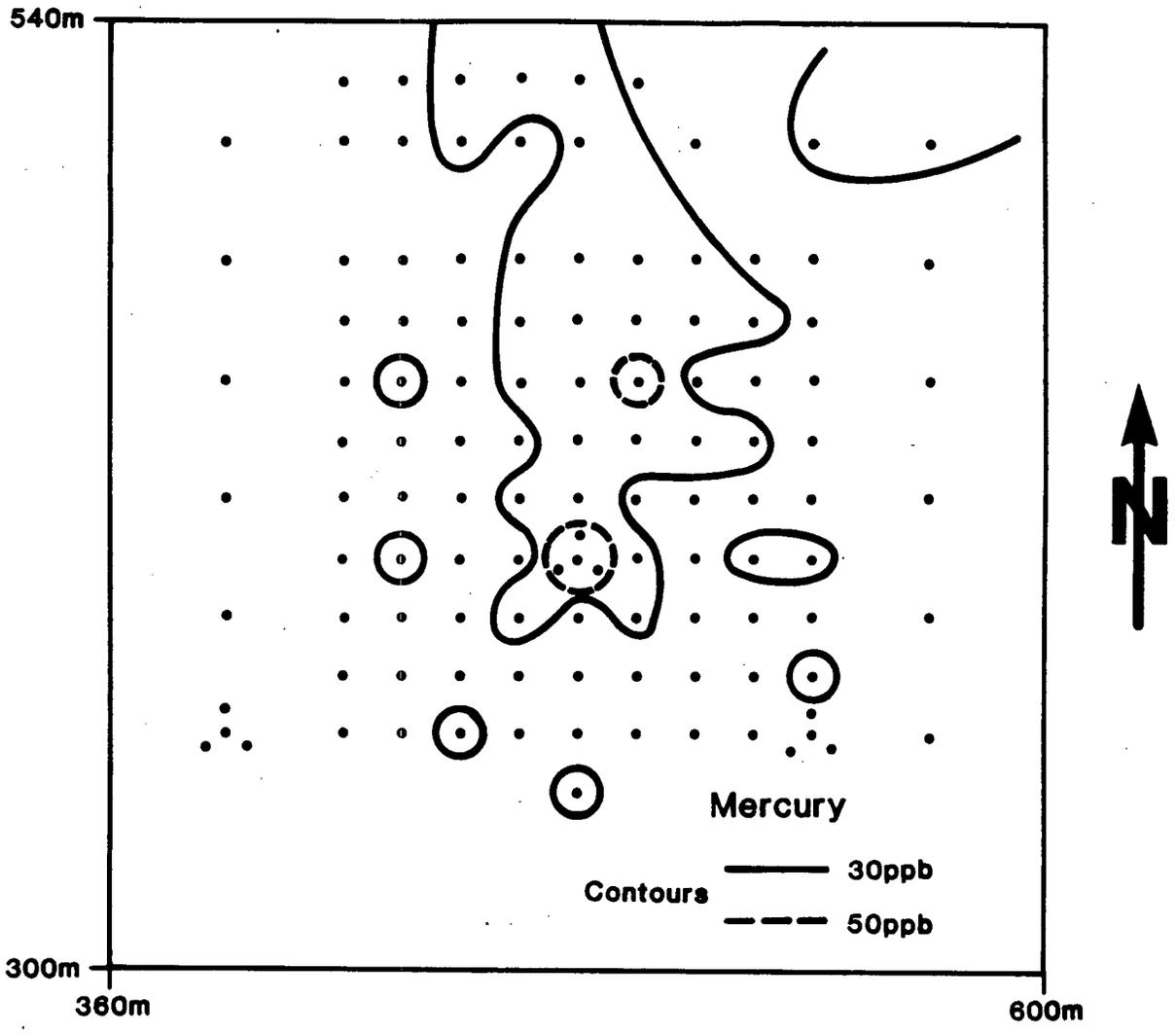


Figure 24

Soil Grid Geochemical Concentrations

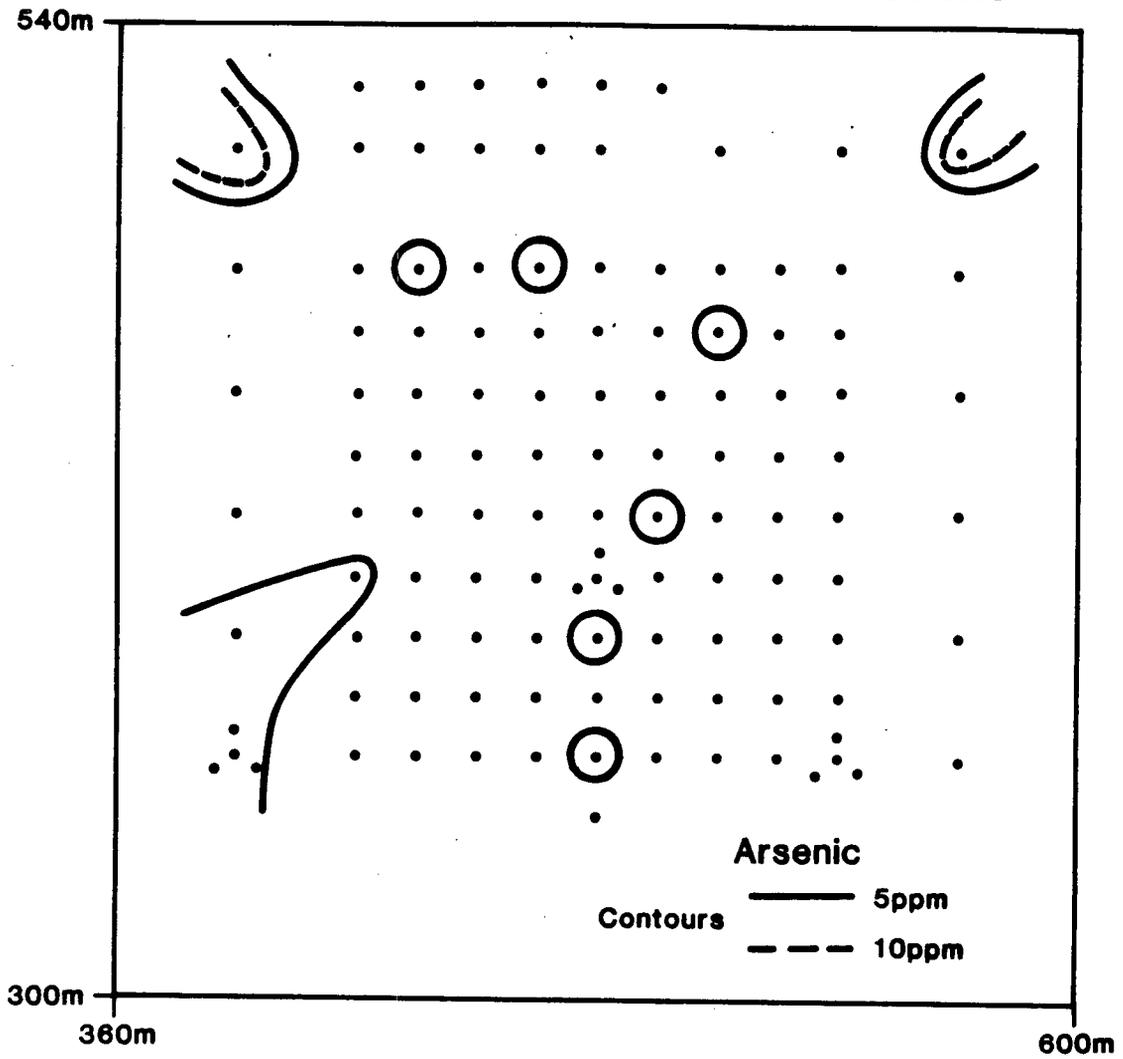


Figure 25

Soil Grid Geochemical Concentrations

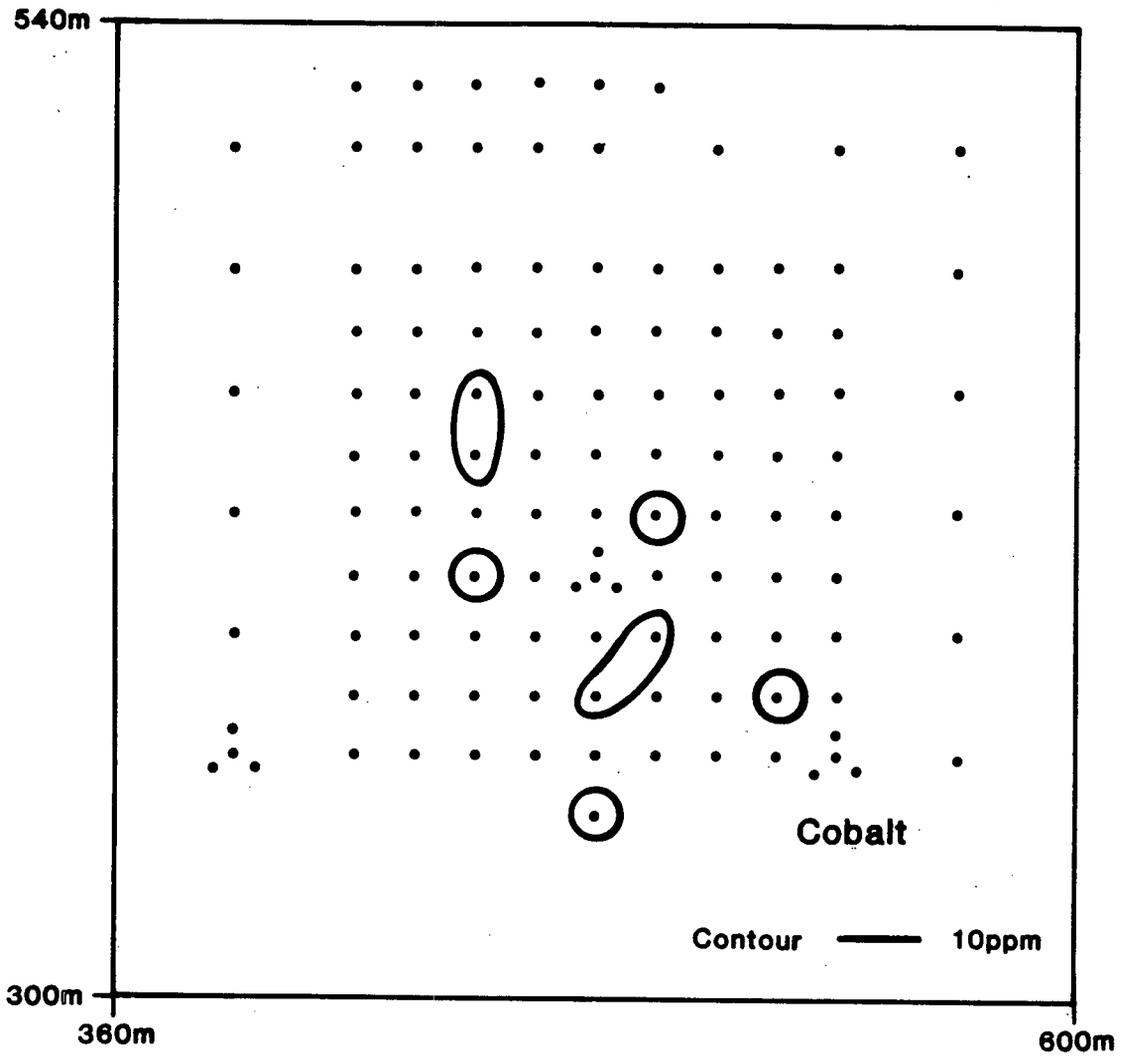


Figure 26

Zn and Ba all have anomalous values which roughly overlie the Cu anomaly but are less coherent spatially. Careful inspection reveals that these anomalies are actually slightly displaced from one another. Moving from southwest to northeast, the sequence of overlapping anomalies consists of Pb and Ba, Zn, Cu, and finally, Ag (Figure 27). Also note that most of the anomalous Hg concentrations occur to the east of these anomalies and, thus, may also be part of this zonation.

The progressive displacement of geochemical anomalies seen in the soil data described above, may be a spatial manifestation of the paragenetic relationships and the zonation of gossan types described in Chapter 3. The zonation is similar to geochemical and mineralogical zonations observed at other stratabound Cu deposits (Kupferschiefer, White Pine, Zambian Copperbelt, Troy; Rentzsch, 1974; Brown, 1971, 1974; Annels, 1974, Bartholome, 1974, Van Eden, 1974, Holcombe, 1977; Hamilton and Balla, 1983, Hayes, 1984). The existence of Hg in the zonation sequence has not been reported previously, except for Conner (1980; referring to the influence of nearby intrusions on the geochemical concentrations of Hg, As, and Sb in mineral occurrences in the Spokane Formation, Belt Supergroup). Because there are no intrusions recognized in the Daisy Creek area, the Hg zonation may support the concept of chloride complexation of metals as a transporting species for the ore constituents, since mercury, like Ag and Cu, has a strong chloride complex (Rose, 1976).

If the anomaly zonation observed at the Daisy Creek prospect is actually a manifestation of a migrating redox boundary which

Zoning Relationships for Soil Anomalies at the Daisy Creek Prospect

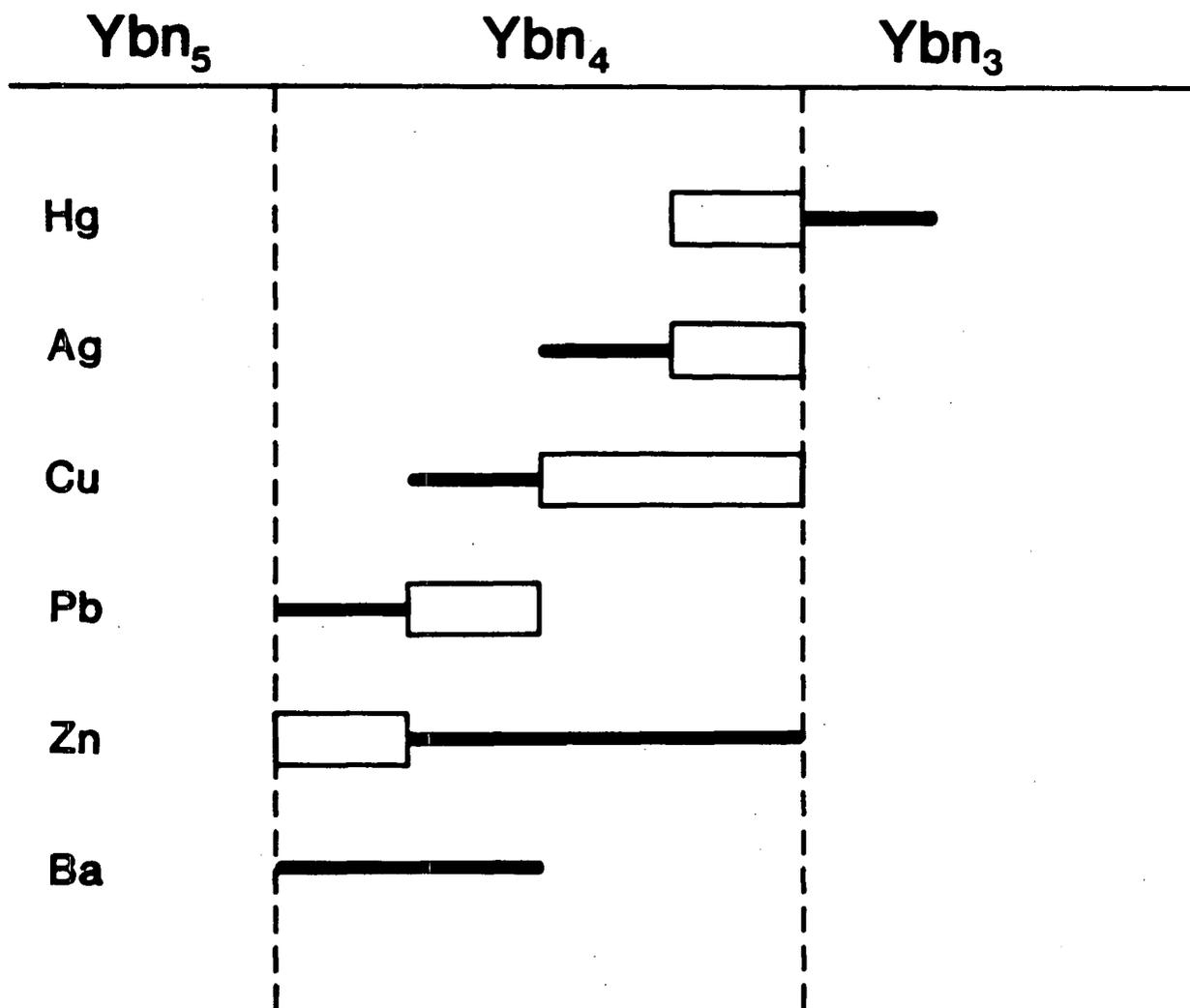


Figure 27

was responsible for the mineralization (Hayes, 1984; Holcombe, 1977), the apparent direction of movement of the oxidized waters would be from NE to SW, based on flow directions recognized at other stratabound Cu deposits referred to above (Brown, 1971; Hayes, 1984). This would correspond to fluids migrating up stratigraphy and toward the south, from the base of the Upper Crossbedded Quartzite Member of the Bonner Formation (Ybn₄; Appendix 1; in the prospect area composed of dominantly coarse to very coarse grained, crossbedded vitreous quartzites), into the top of the member (composed of finer grained, flat laminated quartzites). The base of the Upper Crossbedded Quartzite Member may actually have been an aquifer which allowed the introduction of the oxidized metal bearing waters to diffuse (Brown, 1971) into reducing conditions in the overlying finer grained quartzites.

The zonation observed is also consistent with the zonation of the gossans observed across the deposit (Chapter 3). Red and brown limonite occurs at the north end of the prospect and near the stratigraphic 'base' of the Cu anomaly in the Upper Crossbedded Quartzite Member of the Bonner Formation, along with anomalous Ag concentrations (Figure 22). Orange limonite occurs centered on the Cu anomaly, commonly containing intergrown supergene malachite. Yellow to orange limonite occurs over the Pb and Ba soil anomalies, at the south end of the prospect and at the stratigraphic 'top' of the Cu anomaly.

Therefore, the pedogeochemical data support observations regarding the paragenetic sequence and mineralogic zonation at

the Daisy Creek Prospect. In addition, these data give indications of the hydrodynamic and/or diffusion gradient during the time of mineralization and a further understanding of the geochemical constraints which helped form the deposit. Finally, the soil data precisely outline the location of the mineral occurrence and the various constituent mineralogical zones.

CHAPTER 5

CONCLUSION

The Daisy Creek prospect consists of stratabound Cu-Pb-Ag minerals within quartzites of the Bonner Formation, part of the Missoula Group, the uppermost member of the Belt Supergroup. The sedimentary host rocks consist dominantly of quartz- and hematite-cemented feldspathic quartzites and orthoquartzites deposited during the second period of fluvio-deltaic progradation which occurred during Bonner deposition. Copper bearing mineral assemblages are hosted by a reduced, medium to coarse grained, crossbedded member of the Bonner Formation (Ybn₄; Appendix 1) which contains no hematite cement but high amounts of phosphate. The depositional environment for this quartzite was fluvial.

Regional mapping reveals that two phases of deformation affected the area, one with fold axes oriented at 285° with a SSW vergence, and a second, later event (with a fold axis orientation of 164°) related to overthrusting of Precambrian and Paleozoic rocks to the ENE during Mesozoic and Cenozoic time.

Mineral concentrations of economic interest consist of galena, chalcopyrite, bornite, chalcocite, cupriferous goethite, acanthite, barite, cerussite, pyromorphite, hinsdalite, malachite, native silver, and limonite. These minerals are distributed across the prospect in four different mineralogic zones characterized by galena, chalcopyrite, bornite, and chalcocite which are zoned from SW to NE. These zones are

recognizable based on the different types of gossans developed over them.

Paragenetic relationships of the Cu-bearing mineral assemblages suggest that the present day mineralogy is a result of four different episodes of mineral reactions. The first occurred during early diagenesis and consisted of the precipitation of authigenic pyrite, probably by sulfur reducing bacteria. The second episode consisted of the sequential replacement of pyrite by more Cu rich sulfides: chalcopyrite, bornite, and chalcocite. The third episode involved the oxidation of these sulfide minerals, forming cupriferous goethite and barite, and releasing silver from the bornite and chalcocite lattice to form acanthite. Both episode two and episode three occurred during middle diagenesis and probably represent a continuum of chemical reactions between oxidized, metal-bearing, migrating ground waters and reducing pore waters or pyrite bearing sediments. Lastly, Holocene supergene oxidation resulted in the precipitation of malachite, cerussite, pyromorphite, and limonite from the metal-bearing phases. The paragenetic relationships are consistent with the mineralogic zones observed across the prospect.

Stream sediment geochemical data was seen to precisely locate the Daisy Creek prospect as well as another stratabound copper prospect located in the field area. Statistical analysis led to the recognition of displaced hydromorphic anomalies and anomalies caused by contamination as different in character from those related to mineralization.

Finally, a soil grid survey over the mineralization clearly defines the limits of the mineralization and correlates well with the different mineralogic zones located at the prospect. Concentrations are lower than those observed in float rock chip samples, probably due to surficial leaching of the economic and pathfinder metals during supergene oxidation and soil development. Careful mapping of soil color and grain size successfully predict the underlying lithologic units and the extent of the mineralization.

Paragenetic mineral relationships, gossan zoning, soil geochemical anomalies, and geochemical evidence suggest that the Daisy Creek prospect formed as a result of the interaction of pyrite-bearing quartzites with reduced pore waters, with migrating, oxygenated, metalliferous, briny ground waters. Diffusion of metals and free oxygen resulted in the replacement of pyrite by Pb- and Cu-sulfides, and then their subsequent oxidation. The apparent direction of flow for this migrating diffusion front appears to have been from the NE to the SW, suggesting that diffusion took place toward the south and upwards stratigraphically.

Bibliography

- Anderson, J.A., 1982, "Characteristics of Leached Capping and Techniques Appraisal", in Advances in Geology of Porphyry Copper Deposits, Southwestern North America", ed. by Titley, S.R., University of Arizona Press, Tucson, Ch. 12, pp. 275-298.
- Annels, A.E., 1974, "Some Aspects of the Stratiform Ore Deposits of the Zambian Copperbelt and Their Genetic Significance", in Gisements Stratiformes et Provinces Cuprifères, ed. by Bartholome, P., Centenaire de la Societe Geologique de Belgique, Leige, Belgique, pp. 235-254.
- Bakun, N.N., Volodin, R.N., and Krendel'ev, F.P., 1967, "Genesis of Udokansk Cupriferous Sandstone Deposit (Chitinsk Oblast)", International Geology Review, Vol. 8, No. 4, pp. 455-466.
- Bartholome, P., 1974, "On the Diagenetic Formation of Ores in Sedimentary Beds With Special Reference to Kamoto, Shaba, Zaire", in Gisements Stratiformes et Provinces Cuprifères, ed. by Bartholome, P., Centenaire de la Societe Geologique de Belgique, Leige, Belgique, pp. 203-217.
- Birch, B., 1977, "Hinsdalite from the Zeehan Area, Tasmania", Mineralogical Record, Vol. 8, No. 5, pp. 28-29.
- Bjorlykke, A., and Sangster, D.F., 1981, "An Overview of Sandstone

Lead Deposits and Their Relation to Red-Bed Copper and Carbonate-Hosted Lead-Zinc Deposits", Economic Geology, 75th Anniversary Volume, pp. 179-213.

Blanchard, R., 1968, "The Interpretation of Leached Outcrops", Nevada Bureau of Mines Bulletin, No. 66, 196 pp.

Bowen, R., and Gunatilaka, A., 1976, "The Zambian Copperbelt", Science Progress, Oxford, Vol. 63, pp. 87-109.

Brown, A.C., 1971, "Zoning at the White Pine Copper Deposit, Ontonogon County, Michigan", Economic Geology, Vol. 66, pp. 543-573.

Brown, A.C., 1974, "The Copper Province of Northern Michigan, U.S.A.", in Gisements Stratiformes et Provinces Cupriferes, ed. by Bartholome, P., Centenaire de la Societe Geologique de Belgique, Leige, Belgique, pp. 317-330.

Brown, A.C., 1978, "Stratiform Copper Deposits - Evidence for their Post-Sedimentary Origin", Minerals Science and Engineering, Vol. 10, No. 3, pp. 172-183.

Caia, J., 1976, "Paleogeographical and Sedimentological Controls of Copper, Lead, and Zinc Mineralization in the Lower Cretaceous Sandstones of Africa", Economic Geology, Vol. 71, No. 2, pp. 409-422.

Cazes, D.K., 1981, "Geochemical Analyses of Stream Sediment and Heavy Mineral Concentrates Collected Near a Stratabound Copper/Silver Occurrence in the Cabinet Mountain Wilderness Area, Montana", M.Sc. Thesis, University of Idaho, Moscow, Idaho, 114 p.

Chartrand, F., and Brown, A.C., 1978, "Diagenetic Origin of Stratiform Copper Mineralization at Coates Lake, Redstone Copperbelt, N.W.T., Canada", extended abstract, Ecole Polytechnique (Genie Mineral), Montreal, 3 p.

Clark, A.L., 1971, "Strata-bound Copper Sulfides in the Precambrian Belt Supergroup, Northern Idaho and Northwestern Montana", Society of Mining Geology of Japan, IMA-IAGOD Meeting, 1970, Special Issue No. 3, pp. 261-267.

Conner, J.J., 1980, "Possible Hydrothermal Signature in Belt Stratabound Copper Deposits, Montana", U.S. Geological Survey Professional Paper, No. 1175, p. 12.

Crowley, F.A., 1972, "Mines and Mineral Deposits (Except Fuels) Sanders County, Montana", Montana Bureau of Mines and Geology, Bulletin, No. 81, 88 p.

Dent, B.M., 1933, "On Observations of Points Connected by a Linear Relation", Proceedings of the Physical Society of London, Vol. 47, Part 1, pp. 92-108.

- Dimanche, F., and Bartholome, P., 1976, "The Alteration of Ilmenite in Sediments", *Minerals Science and Engineering*, Vol. 8, No. 3, pp. 187-201.
- Elston, D.P., 1984, "Magnetostratigraphy of the Belt Supergroup, A Synopsis", in The Belt, ed. by Hobbs, S.W., Montana Bureau of Mines and Geology, Special Publication No. 90, pp. 88-90.
- Ensign, C.O. Jr., White, W.S., Wright, J.C., Patrick, J.L., Leone, R.J., Hathaway, D.J., Trammell, J.W., Fritts, J.J., and Wright, T.L., 1968, "Copper Deposits in the Nonesuch Shale, White Pine Michigan", in Ore Deposits of the United States, Graton and Sales, American Institute of Mining, Metallurgical, and Petroleum Engineers, Maple Press, York, Pa., Vol. 1, pp. 460-488.
- Ezekial, M., 1961, "Methods of Correlation and Regression Analysis", John Wiley and Sons, New York, pp. 311-317.
- Garrett, R.G., and Goss, T.I., 1978, "The Evaluation of Sampling and Analytical Variation in Regional Geochemical Surveys", in Geochemical Exploration, ed. by Theobald, P.K., and Watterson, D., DREGS, pp. 371-383.
- Godlewski, D.W., 1977, "The Origin of the Middle Wallace Breccias", Geological Society of America, Abstracts with Programs, Vol. 9, No. 6, p. 727.

- Gustafson, L.B., and Williams, N., 1981, "Sediment-Hosted Stratiform Deposits of Copper, Lead, and Zinc", *Economic Geology*, 75th Anniversary Volume, pp. 139-181.
- Hamilton, J.M., and Balla, J., 1983, "Belt-Purcell Supergroup; Troy (Spar Lake) Mine", *Field Trip Guidebook, Geological Association of Canada, Annual Meeting, Victoria, B.C.*, pp. 25-31.
- Harrison, J.E., 1972, "Precambrian Belt Basin of Northwestern United States: Its Geometry, Sedimentation, and Mineral Occurrences", *Geological Society of America Bulletin*, Vol. 83, pp. 1215-1240.
- Harrison, J.E., 1974, "Copper Mineralization in Miogeosynclinal Clastics of the Belt Supergroup, Northwestern U.S.", in Gisements Stratiformes et Provinces Cuprifères, ed. by Bartholome, P., *Centenaire de la Societe Geologique de Belgique, Leige, Belgique*, pp. 353-366.
- Harrison, J.E., Griggs, A.B., and Wells, J.D., 1981, "Generalized Geologic Map of the Wallace 1° X 2° Quadrangle, Montana and Idaho", *U.S. Geological Survey*, 1:250,000.
- Harrison, J.E., 1984a, "Stratigraphy and Lithocorrelation of the Wallace Formation", in The Belt, ed. by Hobbs, S.W., *Montana Bureau of Mines and Geology, Special Publication No. 90*, pp. 24-26.

- Harrison, J.E., 1984b, "Session on Geochronology and Geophysics", in The Belt, ed. by Hobbs, S.W., Montana Bureau of Mines and Geology, Special Publication No. 90, pp. 97-100.
- Harrison, J.E., and Campbell, A.B., 1963, "Correlations and Problems in Belt Series Stratigraphy, Northern Idaho and Western Montana", Geological Society of America Bulliten, Vol. 74, pp. 1413-1428.
- Harrison, J.E., Lange, I.M., and Harrison, J.P., 1977, "Stratabound Copper Occurrences in Green Beds of the Belt Supergroup, Western Montana", Field Trip Guide No. 3, Rocky Mountain Section, 30th Annual Meeting, Geological Society of America, 12 p.
- Hayes, T.S., 1984, "The Relation between Stratabound Copper-Silver Ore and Revett Formation Sedimentary Facies at Spar Lake, Montana", in The Belt, ed. by Hobbs, S.W., Montana Bureau of Mines and Geology, Special Publication No. 90, pp. 63-64.
- Holcombe, C.J., 1977, "Sedimentary Coppers (Copper in Sedimentary Formations): Report on the 1976-1977 World Study Group", 1976-77 World Study Group Report, 94 p.
- Hughes, G.J. Jr., 1984, "Session on Strata-Related Mineral Resources", in The Belt, ed. by Hobbs, S.W., Montana Bureau of Mines and Geology, Special Publication No. 90, pp. 69-80.
- Hurlbut, C.S. Jr., and Klein, C., 1977, "Manual of Mineralogy, 19th

Edition", John Wiley and Sons, New York, pp. 235-244.

- Johns, W.M., 1970, "Geology and Mineral Deposits of Lincoln and Flathead Counties, Montana", Montana Bureau of Mines and Geology, Bulletin, No. 79, 79 p.
- Jones, T. A., 1972, "Multiple Regression With Correlated Independent Variables", Mathematical Geology, Vol. 4, No. 3, pp. 203-218.
- Jones, T.A., 1979, "Fitting Straight Lines When Both Variables Are Subject to Error - I. Maximum Likelihood and Least-Squares Estimation", Mathematical Geology, Vol. 11, No. 1, pp. 1-25.
- Kendall, M.G., and Stuart, A., 1963, "The Advanced Theory of Statistics, Vol. 2", Griffin, New York, pp. 375-418.
- Kermack, K.A., and Haldane, J.B.S., 1950, "Organic Correlation and Allometry", Biometrika, Vol. 37, No. 1, pp. 30-41.
- Lange, I.M., and Eby, D.E., 1981, "Stratabound Copper-Silver-Bearing Oolitic Deposits of the Belt Supergroup of Montana", Economic Geology, Vol. 76, No. 4, pp. 933-937.
- Lange, I.M., and Sherry, R.A., 1983, "Genesis of the Sandstone (Revett) Type of Copper-Silver Occurrences in the Belt Supergroup of Northwestern Montana and Northeastern Idaho", Geology, V. 11, pp. 643-646.

- Lindsey, T.R., 1977, "Mineralogy and Trace Metals in a Revett-Type Stratabound Sulfide Occurrence, Northwestern Montana", M.S. Thesis, University of Montana, Missoula, Montana, 141 p.
- Lindley, D.V., 1950, "Regression Lines and the Linear Functional Relationship", Journal of the Royal Statistical Society: Supplement, Vol. 9, pp. 218-244.
- Lochman, C., 1957, "Paleoecology of the Cambrian in Montana and Wyoming", Geological Society of America, Memoir 67, Treatise on Marine Ecology and Paleoecology: Paleoecology, Vol. 2, pp. 117-162.
- Mark, D.L., and Church, M., 1977, "On the Misuse of Regression in Earth Science", Mathematical Geology, Vol. 9, No. 1, pp. 63-75.
- Matysek, P., Sinclair, A.J., and Fletcher, W.K., 1982, "Rapid Anomaly Recognition and Ranking for Multi-element Regional Stream Sediment Surveys", Geological Fieldwork - 1981, British Columbia Ministry of Energy, Mines, and Petroleum Resources, Victoria, B.C., pp. 176-186.
- Meisch, A.T., 1967, "Theory of Error in Geochemical Data", U.S. Geological Survey Professional Paper No. 574-A, pp. 1-17.
- Miller, J.K., and Miller, D.E., "Interim Report on the Daisy Creek Prospect, Rocky Mountain District", Anaconda Minerals Company

Internal Report, p. 2.

- Morad, S., and Aldahan, A.A., 1982, "Authigenesis of Titanium Minerals in Two Proterozoic Sedimentary Rocks From Southern and Central Sweden", *Journal of Sedimentary Petrology*, Vol. 52, No. 4, pp. 1295-1305.
- Newton, E., 1982, "A Petrographic Study of Cu-Ag Mineralization in the Bonner Formation of the Precambrian Belt Supergroup of Western Montana", B.S. Thesis, Dartmouth College, Hanover, N.H., 89 p.
- Ney, C.S., Cathro, R.J., Panteleyev, A., and Rotherham, D.C., 1976, "Supergene Copper Mineralization", in Porphyry Deposits of the Canadian Cordillera, Special Vol. 15, Canadian Institute of Mining and Metallurgy, Ch. 9, pp. 72-78.
- Obradovich, J.D., Zartman, R.E., and Peterman, Z.E., 1984, "Update of the Geochronology of the Belt Supergroup", in The Belt, ed. by Hobbs, S.W., Montana Bureau of Mines and Geology, Special Publication No. 90, pp. 82-84.
- Ohle, E.L., 1968, "Copper Mineralization in the Upper Part of the Copper Harbor Conglomerate at White Pine, Michigan - Discussion", *Economic Geology*, Vol. 63, pp. 190-195.
- Plant, J., Jeffery, K., Gill, E., and Fage, C., 1975, "The Systematic

Determination of Accuracy and Precision in Geochemical Exploration Data", Journal of Geochemical Exploration, Vol. 4, pp. 467-486.

Price, R.A., 1984, "Tectonic Evolution of the Purcell (Belt) Rocks of the Southeastern Canadian Cordillera and Adjacent Parts of the United States", in The Belt, ed. by Hobbs, S.W., Montana Bureau of Mines and Geology, Special Publication No. 90, pp. 47-48.

Renfro, A.R., 1976, "Genesis of Evaporite-Associated Stratiform Metalliferous Deposits - A Sabhka Process", Economic Geology, Vol. 69, pp. 33-45.

Rentzsch, J., 1974, "The Kupferschiefer in Comparison with the Deposits of the Zambian Copperbelt", in Gisements Stratiformes et Provinces Cupriferes, ed. by Bartholome, P., Centenaire de la Societe Geologique de Belgique, Leige, Belgique, pp. 395-418.

Reynolds, M.W., 1984a, "Session on Structure and Tectonics", in The Belt, ed. by Hobbs, S.W., Montana Bureau of Mines and Geology, Special Publication No. 90, pp. 52-56.

Reynolds, M.W., 1984b, "Tectonic Setting and Development of the Belt Basin, Northwestern United States", in The Belt, ed. by Hobbs, S.W., Montana Bureau of Mines and Geology, Special Publication No. 90, pp. 44-46.

- Rickard, D.T., 1974, "Low Temperature Copper Geochemistry: Geological Aspects", in Gisements Stratiformes et Provinces Cupriferes, ed. by Bartholome, P., Centenaire de la Societe Geologique de Belgique, Leige, Belgique, pp. 1-34.
- Rose, A.J., 1976, "The Effect of Cuprous Chloride Complexes and the Origin of Red-bed Copper and Related Deposits", Economic Geology, Vol. 71, pp. 1036-1048.
- Rowlands, N.J., 1974, "The Geology of Some Adelaiddian Stratiform Copper Occurrences", in Gisements Stratiformes et Provinces Cupriferes, ed. by Bartholome, P., Centenaire de la Societe Geologique de Belgique, Leige, Belgique, pp. 419-427.
- Rye, R.O., Whelan, J.F., and Harrison, J.E., 1984, "The Origin of Copper-Silver Mineralization in the Ravalli Group as Indicated by Preliminary Stable Isotope Studies", in The Belt, ed. by Hobbs, S.W., Montana Bureau of Mines and Geology, Special Publication No. 90, pp. 104-107.
- Sinclair, A.J., 1976, "Applications of Probability Graphs in Mineral Exploration", Association of Exploration Geochemists, Special Publication #4, 95 p.
- Smith, G.E., 1976, "Sabhka and Tidal-Flat Facies Control of Stratiform Copper Deposits in North Texas", in Handbook of Stratabound and Stratiform Ore Deposits; II Regional Studies and

Specific Deposits, Vol. 6, ed. by Wolf, K.H., Elsevier, Amsterdam, pp. 408-446.

Spiegelman, C., 1982, "A Note on the Behavior of Least-Squares Regression Estimates When Both Errors Are Subject to Error", Journal of Research of the U.S. National Bureau of Standards, Vol. 87, No. 1, pp. 303-308.

Shipulin, P.K., 1971, "On Thermodiffusion of Components During the Formation of Stratabound Sulfide Ore Deposits", Society of Mining Geology of Japan, IMA-IAGOD Meeting, 1970, Special Issue No. 3, pp. 274-276.

Till, R., 1973, "The Use of Regression in Geomorphology", Area, Vol. 5, No. 4, pp. 303-308.

Thompson, M., and Howarth, R.J., 1973, "The Rapid Estimation and Control of Precision by Duplicate Determination", Analyst, Vol. 98, No. 1164, pp. 153-160.

Thompson, M., and Howarth, R.J., 1976a, "Duplicate Analysis in Geochemical Practice, Part 1: Theoretical Approach and Estimation of Analytical Reproducibility", Analyst, Vol. 101, pp. 690-698.

Thompson, M., and Howarth, R.J., 1976b, "Duplicate Analysis in Geochemical Practice, Part 2: Examination of Proposed Method and Examples of Its Use", Analyst, Vol. 101, pp. 699-709.

- Thompson, M., and Howarth, R.J., 1978, "A New Approach to the Estimation of Analytical Precision", *Journal of Geochemical Exploration*, Vol. 9, pp. 23-30.
- Van Eden, J.G., 1974, "Depositional and Diagenetic Environment Related to Sulfide Mineralization, Mufulira, Zambia", *Economic Geology*, Vol. 69, pp. 59-79.
- Wallace, C.A., Harrison, J.E., Klepper, M.R., and Wells, J.D., 1976, "Carbonate Sedimentary Breccias in the Wallace Formation (Belt Supergroup), Idaho and Montana, and Their Paleogeographical Significance", *Geological Society of America, Abstracts with Programs*, Vol. 8, No. 6, p. 1159.
- Wallace, C.A., Harrison, J.E., Whipple, J.W., Ruppel, E.T., and Schmidt, R.G., 1984, "A Summary of Stratigraphy of the Missoula Group of the Belt Supergroup and Preliminary Interpretation of Basin-Subsidence Characteristics, Western Montana, Idaho, and Eastern Washington", in *The Belt*, ed. by Hobbs, S.W., Montana Bureau of Mines and Geology, Special Publication No. 90, pp. 27-29.
- Watson, I.M., Mustard, D.K., and Blusson, S.L., 1975, "The Redstone Bedded Copper Deposit, District of MacKenzie, Northwest Territories", *Economic Geology*, Vol. 70, No. 1, pp. 251-252.
- Wheeler, J.O., 1984, "Columbian and Pacific Orogens Revisited",

Program with Abstracts, Symposium - Cordilleran Geology and Mineral Exploration: Status and Future Trends, Cordilleran Section - G.A.C., Vancouver, B.C., pp. 45-48.

Whipple, J.W., 1984, "Session on Geochemistry and Isotope Geology", in The Belt, ed. by Hobbs, S.W., Montana Bureau of Mines and Geology, Special Publication No. 90, pp. 108-111.

Whitten, G., 1965, "Leached Outcrops", in Exploration and Mining Geology, Vol. 2, Eighth Commonwealth Mining and Metallurgical Congress, Australia and New Zealand, ed. by Lawrence, L.J., Ch. 5, Guides to Ore, pp. 193-204.

Wilkinson, W.H. Jr., Roe, A., and Williams, C., 1980, "Some Unusual Secondary Minerals From the Mineral Park Mine, Mohave County, Arizona", Mineralogical Record, Vol. 11, No. 4, pp.42-45.

Winston, D., 1978, "Fluvial Systems of the Belt Supergroup, Montana and Idaho, U.S.A.", in Fluvial Sedimentology, ed. by Miall, A.D., Memoir #5, Canadian Society of Petroleum Geologists, pp. 343-359.

Winston, D., 1984a, "Sedimentology and Stratigraphy of the Missoula Group", in The Belt, ed. by Hobbs, S.W., Montana Bureau of Mines and Geology, Special Publication No. 90, pp. 30-32.

Winston, D., 1984b, "Sessions on Middle Belt Carbonate and Missoula Group and on Belt-Purcell Stratigraphy", in The Belt, ed. by

S.W., Montana Bureau of Mines and Geology, Special Publication
No. 90, pp. 39-42.

Winston, D., Lange, I.M., Bleiwas, D., and Godlewski, D., 1977,
"Alluvial Fan, Shallow Water, and Sub-Wave Base Deposits of the
Belt Supergroup Near Missoula, Montana", Field Trip Guide No. 5,
Rocky Mountain Section, 30th Annual Meeting, Geological Society
of America, 41 p.

APPENDIX 1

DESCRIPTION OF STRATIGRAPHY

Description of Stratigraphy

The following is a description of the stratigraphy from the Daisy Creek Area (Harrison, 1972; Harrison and Campbell, 1963; Winston, 1984-A, 1984-B). Nomenclature and symbols used here is referenced in the text. A stratigraphic column depicting the following description is located in the pocket as Plate ?.

CAMBRIAN SECTION

555 to ? m

Undifferentiated Dolomite - Middle to Upper Cambrian (Cdu) :

Dearborn Dolomite (Formation) occurs at base of this section and consists of a light to medium grey, thick bedded, sandy dolomite with minor shale interbeds and partings; beds commonly 5.0 to 15.0 cm thick; minor cross-cutting calcite veins; top not exposed.

210 to 555 m

Damnation Dolomite (Formation) - Middle Cambrian (Cdd) : light to medium grey, thin bedded, sandy dolomite with abundant shale interbeds and partings; beds commonly 2.0 to 5.0 cm thick; convolute bedding; abundant vuggy cavities are filled with calcite and minor quartz; abundant crosscutting calcite veins.

30 to 210 m

Gordon Shale (Formation) - Middle Cambrian (Cgs) : waxy, olive green, greyish green, and green, highly fissile, phyllitic shale; beds commonly 2.0 to 10.0 mm thick; minor interbedded planar crossbedded sandstone lenses and dolomite cemented beds; orange limonite after pyrite occurs along bedding planes and crossbeds in sandstone lenses.

0 to 30 m

Flathead Sandstone (Formation) - Middle Cambrian (Cfs) : thin bedded, red, reddish grey, and maroon, flat laminated, coarse grained siltstone to fine grained sandstone; abundant scour and fill structures; beds commonly 1.0 to 4.0 cm thick; interbedded and generally overlain by thick bedded, light pink, light orange, and light tan, very coarse grained sandstone, grit, and conglomerate; contains abundant rounded quartz and jasper grains; moderately well sorted; commonly contains trough and epsilon crossbeds as well as heavy mineral lag lamellae; beds commonly 4.0 to 20.0 cm thick; often light colored grit fills scours in the red siltstone.

----- Regional Angular Unconformity -----

HELIKIAN SECTION - MISSOULA GROUP

-2025 to 0 m

Libby Formation - Helikian (Ylb) : composed of two members;

-690 to 0 m

Upper Member (Ylbu) : thin to very thin bedded, dominantly olive green, medium to very fine grained, thin parting, chloritic siltite with minor medium brown, greyish purple, and dark grey beds; contains abundant muscovite along bedding planes and partings; muscovite also occurs in beds, which weather to medium brown; abundant mud chip breccias, rip up clast conglomerates, and imbricated mud chip lag deposits; mud chips are both siliceous (grey) and argillaceous (green); beds occur as fining upward couplets between 0.2 and 1.5 cm thick; desiccation crack casts are common; abundant oscillation ripple marks with minor asymmetrical tuning fork and undulatory ripple marks; bed forms are commonly nearly planar but scour and fill, pinch and swell, lenticular, and swaly bed forms are also present; exhibits poorly developed joint surfaces; black Fe- and Mn-oxides (as dendritic wad) occur commonly on joint, fracture, and weathering surfaces; lower section from -690 to -180 m contains common dark grey cherts and medium grey cherty siltites interbedded with olive green siltites; syneresis crack casts filled with milky quartz occur

commonly in the pure dark grey chert beds; the upper section from -180 to 0 m contains more prevalent greyish purple and medium brown beds and very common mud crack casts; quartz veins cutting this member commonly contain chlorite, minor barite, and hematite, and are bounded by a halo of altered siltite which contains no muscovite and which has been bleached to a pale purplish grey; contact with Lower Member is gradational and based on color change and disappearance of muscovite and chert beds, extent of cleavage development, abundance of Fe- and Mn-oxide coatings, mud chip composition, and the presence of ripple marks.

-2025 to -690 m

Lower Member (Y1b1) : thin to thick bedded, blueish mint green siltite and fine grained quartzite; color caused by abundance of chlorite; siltite beds occur as fining upward couplets from 0.5 to 2.5 cm thick; quartzites fracture conchoidally and occur as massive, well sorted, unlaminated beds from 3.0 to 30.0 cm thick; thick parting surfaces are not generally micaceous; mud chips, mud chip conglomerates, and imbricated mud chip lag deposits are common; mud chips are commonly argillaceous (green), but minor siliceous (grey) mud chips do occur; minor desiccation crack and syneresis crack casts; exhibits well developed joint surfaces; quartz veins cutting this unit contain chlorite, minor barite and

hematite, and are bounded by a halo of altered siltite which has been bleached to a pale purplish grey; marker horizons in the member consist of:

- 1) a thin to medium bedded, red and purple siltite at -765 to -750 m which appears to be identical to overlying and underlying units except for a different iron valence state (Fe^{III} in hematite as opposed to Fe^{II} in chlorite);
- 2) a medium to thick bedded, dark mint green siltite at -900 to -870 m is cemented by calcite;
- 3) four thick bedded to massive, well sorted, fine grained quartzite lenses appear to be semi-continuous across the field area and occur at -1140 to -1110 m, -1410 to -1380 m, -1560 to -1500 m, and -1620 to -1590 m; and
- 4) a thin bedded, dominantly dark grey to black, carbonaceous, fine to medium grained siltite is interbedded with mint green siltite; often individual beds are 0.2 to 1.0 cm thick and are green in the lower half and black in the upper half of the bed; unit occurs at -1995 to -1915;

Below these marker beds, from -2025 to -1995 m, this unit tends to be thinner bedded, contains more dominant purple to greyish purple beds between 2.0 to 20.0 mm

thick, and is gradational with the underlying Bonner Formation; contact with Bonner Formation is marked by the first appearance of muscovite on bedding planes, dominant red color, and increased grain size.

-2565 to -2025 m

Bonner Formation - Helikian (Ybn) : composed of five members;

-2070 to -2025 m

Upper Laminated Quartzite Member (Ybn5) : mainly red, pink, maroon, and purple, fine to very fine grained, hematitic quartzites and medium to coarse grained siltites; bedding is flat laminated and commonly micaceous; minor ripple, lenticular, swaly, and scour and fill bed forms; parting lineation occurs on micaceous parting surfaces; thin to very thin argillaceous and muscovite partings; occasional asymmetric convolute, undulatory, and tuning fork ripple marks and minor symmetric oscillation ripple marks; common heavy mineral lag deposits; beds occur as fining upward couplets commonly 0.5 to 4.0 cm thick; beds also commonly have argillaceous tops with desiccation crack casts; imbricated matrix supported siliceous mud chip conglomerates occur throughout; occasional orange limonite blebs occur disseminated in the coarser grained beds; crosscutting quartz veins contain minor hematite and do not alter the country rock.

-2100 to -2070 m

Upper Crossbedded Quartzite Member (Ybn₄) : mainly white, tan, light pale green, grey to light grey, pale orange, and cream colored, medium to very coarse grained, thick bedded to massive, thick to flaggy parting quartzite; dominant trough and tabular-tangential crossbeds with minor epsilon crossbeds; beds commonly 5.0 cm to 1.0 m thick; coarser grained quartzites are vitreous; finer grained quartzites have micaceous thin partings and are generally flat laminated and fine upward as couplets; finer grained quartzites predominate in upper half of member; finer grained quartzites commonly exhibit symmetrical oscillation and asymmetrical tuning fork and undulatory ripple marks; some parasitic ripple sets; beds of the finer grained quartzites are generally 0.5 to 3.0 cm thick; coarse grained quartzites contain abundant siliceous and phosphatic (apatite) mud chips as lag deposits along foresets; heavy mineral lag deposits occur in both the crossbedded and flat laminated quartzites as foreset lag deposits and lag lamellae, respectively; abundant red limonite (hematite) blebs occur disseminated throughout the unit and hematite often dusts the insides of cavities in the quartzites; minor to rare white claystone also occurs within the unit; crosscutting quartz veins contain minor hematite and do not alter the country rock.

-2250 to -2100 m

Middle Laminated Quartzite Member (Ybn₃) : mainly red, pink, maroon, and purple, very fine grained quartzites and medium to coarse grained siltites; thin to very thin parting; contains abundant muscovite on bedding planes and siliceous mud chips as lag lamellae which are often imbricated; generally lacks heavy mineral beds; parting lineation; very common red argillaceous bed tops which act as parting surfaces; abundant desiccation crack casts; bedding is generally 0.5 to 3.0 cm thick and occurs as fining upward couplets; occasional orange limonite blebs occur disseminated throughout unit; crosscutting quartz veins contain minor hematite and do not alter the country rock.

-2370 to -2250

Lower Crossbedded Quartzite Member (Ybn₂) : generally tan, light orange, white, and cream colored, thick bedded, thick to flaggy parting, medium to very coarse grained, vitreous quartzite; beds from 10.0 to 100.0 cm thick; common trough and tabular-tangential crossbeds and minor epsilon crossbeds; very abundant heavy mineral accumulations as foreset lag deposits; minor red limonite (hematite) blebs occur disseminated throughout; unit from -2370 to -2310 m has three interbedded 5.0 to 10.0 m thick sequences of light pink, red, maroon, and purple, flat laminated to ripple laminated, fine to very

fine grained, micaceous quartzites with parting lineation; ripple marks are dominantly asymmetric tuning fork and undulatory types; beds are 0.5 to 3.0 cm thick and occur as fining upward couplets; most beds contain minor heavy mineral lag deposits along their bases; coarse grained quartzite beds in this section are often pale green immediately (up to 3.0 m) below the finer grained quartzite sequences; crosscutting quartz veins contain minor hematite and do not alter the country rock.

-2565 to -2370

Lower Laminated Quartzite Member (Ybn₁) : gradational with the lower part of the Lower Crossbedded Quartzite Member; consists of light pink, red, maroon, and purple, fine to very fine grained, thin to flaggy parting, thin bedded, flat and ripple laminated, micaceous quartzite; beds consist of fining upward couplets which range from 0.8 to 15.0 cm thick; minor light pink, trough crossbedded, fine grained quartzites occur from -2400 to -2370 m; below -2400 m argillaceous bed tops, partings, and desiccation crack casts are prevalent; minor heavy mineral accumulations occur as foreset lag and bedding plane lag deposits throughout the unit; minor siliceous mud chips occur; from -2490 to -2565 m unit consists of pale green flat laminated micaceous quartzites interbedded with the red micaceous quartzites; green

facies are generally 3.0 to 10.0 m thick while interbedded red micaceous quartzites are generally 10.0 to 25.0 cm thick; abundant tuning fork and undulatory ripple marks present only in red micaceous quartzites; minor disseminated orange limonite blebs occur only in the red beds; crosscutting quartz veins contain minor hematite and do not alter the country rock.

-3705 to -2565 m

Mount Shields Formation - Helikian (Yms) : composed of two members;

-3285 to -2565 m

Siltite Member (Ymss) : composed of 6 units;

-2730 to -2565 m

Upper Red Siltite Unit (Ymss₆) : dominantly red, purple, and maroon, flat laminated, thin parting, medium to coarse grained, micaceous siltite; common climbing ripple cross stratification and asymmetric tuning fork ripple marks; beds occur as 5.0 to 10.0 mm thick sets of fining upward couplets and often have argillaceous tops with desiccation crack casts.

-2880 to -2730 m

Upper Green Siltite Unit (Ymss₅) : pale green, buff, tan, and light grey, medium grained, thin parting, flat

laminated siltite with sedimentary structures similar to the Upper Red Siltite Unit; contains fewer desiccation crack casts and ripple marks; tops of beds are argillaceous but are often oxidized to red and purple colors; beds consist of fining upward couplets from 0.5 to 3.5 cm thick and often have micaceous parting surfaces as well as having muscovite disseminated throughout; occasional light grey to grey chert nodules occur up to 5.0 by 10.0 cm in cross section.

-3000 to -2880 m

Purple and Green Argillite Unit (Ymss₄) : flat laminated, intercalated, purple and green argillite composed of fining upward couplets from 0.5 to 2.0 cm thick; minor pinch and swell, scour and fill, lenticular, and swaly bed forms; green beds commonly have red argillaceous tops and partings; very minor disseminated orange limonite blebs occur along basal bedding planes in the purple couplets.

-3090 to -3000 m

Purple and Green Siltite Unit (Ymss₃) : flat laminated, intercalated, purple and green, medium to fine grained, thin parting siltite and argillite; very micaceous; siltites are dominantly purple while argillites are dominantly green; beds consist of fining upward couplets from 0.2 to 1.2 cm thick; green beds

often have argillaceous red tops and parting surfaces; minor calcite cement occurs in some of the coarser siltites from -3090 to -3050 m (lower 40 m).

-3225 to -3090 m

Mixed Clastic Unit (Ymss₂) : intercalated, purple and green argillite, siltite, and fine to very fine grained quartzite; very micaceous; fining upward couplets range from 5.0 to 80.0 mm thick; common climbing ripple cross stratification; minor symmetric oscillation ripple marks; siltites and quartzite units are dominantly green while argillite is dominantly purple; scour and fill, and pinch and swell bed forms occur occasionally; green units commonly have red argillaceous tops which exhibit abundant desiccation crack casts; coarser units contain minor siliceous mud chip conglomerates; minor carbonate cement occurs in coarser units from -3115 to -3090 m (upper 25 m).

-3285 to -3225 m

Lower Red Siltite Unit (Ymss₁) : dominantly purple and red, flat laminated, medium grained siltite with abundant muscovite along bedding planes; contains climbing ripple cross stratification and siliceous mud chip conglomerates; beds consist of fining upward couplets between 1.0 and 5.0 cm thick; argillaceous tops and desiccation crack casts on bed tops are common; from

-3285 to -3270 m unit grades into pink, very fine grained, flat laminated, and minor symmetric oscillation ripple laminated quartzite.

-3705 to -3285 m

Quartzite Member (Ymsg) : composed of three units;

-3375 to -3285 m

Pyritic Quartzite Unit (Ymsq3) : fine to very fine grained, white, tan, light grey to light green, flat laminated quartzite; minor planar and tabular-tangential crossbeds with low foreset to bedding angles; often appears bleached to milky white; contains bimodal size distribution of authigenic pyrite cubes; one set of pyrite cubes averages 1.0 mm on a side and commonly exhibits red limonite rims; the second set of pyrite cubes averages 5.0 mm on a side and appears to be rimmed with magnetite; beds with large amounts of pyrite which have undergone supergene oxidation are often discolored to red adjacent to the pyrite; unit contains thick, occasionally red argillite partings; siliceous mud chip conglomerates and climbing ripple cross stratification are also present; finer grained beds consist of fining upward sequences from 2.0 to 35.0 cm thick.

-3485 to -3375 m

Lower Green Siltite Unit (Ymsq₂) : green to pale grey, medium to coarse grained, flat laminated, thin parting siltite; highly micaceous; contains some climbing ripple cross stratification and siliceous mud chip conglomerates; highly micaceous units weather to medium brown; very minor purple facies occur but are usually confined to argillaceous tops on green beds; beds consist of fining upward couplets from 0.5 to 2.5 cm thick.

-3705 to -3485 m

Hellgate Quartzite Unit (Ymsq₁) : pink, and dark pink, very fine to medium grained, medium parting quartzite; flat laminated, fining upward couplets range from 5.0 to 20.0 cm thick; minor tabular-tangential and planar crossbeds with low foreset to bedding angles; contains siliceous mud chip conglomerates, minor heavy mineral lag deposits on foresets and bedding planes, and muscovite on finer grained bed tops.

-4205? to -3705 m

Snowslip Formation - Helikian (Ysn) : composed of two members;

-4055 to -3705 m

Purple and Green Argillite Member (Ysn₂) : purple and green, thin parting, intercalated siltite and argillite;

beds commonly 0.5 to 2.0 cm thick; often green argillite has red argillaceous tops; minor desiccation crack casts occur on red argillaceous tops; bedding is flat laminated with minor pinch and swell, and scour and fill bed forms; some red argillaceous mud chip conglomerates present; coarse grained chlorite occurs on major joint and fracture surfaces.

-4205? to -4055 m

Green Argillite Member (Ysn₁) : grey to green, thin parting argillite and minor siltite and argillite; flat laminated with minor pinch and swell bed forms; beds occur as fining upward couplets from 0.5 to 2.0 cm thick; no mud chip conglomerates occur; minor muscovite on parting surfaces; coarse grained chlorite occurs on major joint and fracture surfaces; minor carbonate cement; possibly equivalent to the Upper Member of the Wallace Formation as described by Harrison (1984).

HELIKIAN SECTION - MIDDLE BELT CARBONATE

? to -4205? m

Wallace Formation - Helikian (Ywl) : medium to thick parting, grey, tan, and black, calcareous siltite with minor black cherty argillite beds; interbedded calcite and dolomite cemented quartzites; distinctive wavy bedding from 3.0 to 10.0 cm thick; abundant syneresis molar

tooth type (Eby cycle) structures; minor oscillation ripple marks; probably Middle Member of the Wallace Formation as described by Harrison (1984); base not exposed.

APPENDIX 2

STREAM SEDIMENT SURVEY -
GEOCHEMICAL DATA

	SAMP	EASTING	NORTHING	REPL	REPL	CU(PPM)	PB(PPM)	ZN(PPM)	MO(PPM)	AG(PPM)	CO(PPM)	MN(PPM)	F(PPM)	BA(PPM)	FE(PCT)
	NUM	(FT)	(FT)	NUM	STAT										
1															
2															
3															
4	30675	48493.74	127124.13			92	87	56	1.00	0.60	15	1850	370	620	1.75
5	30701	45764.29	124602.19			29	49	82	2.00	0.40	10	1450	680	710	2.15
6	30702	40043.10	91624.13			33	25	69	1.00	0.40	13	1650	680	970	2.00
7	30703	40651.37	93821.81			7	9	35	1.00	0.20	7	640	350	710	1.20
8	30704	41313.10	92889.06			34	20	55	2.00	0.20	9	620	450	1040	1.80
9	30705	43569.02	93190.88			20	18	64	2.00	0.20	7	450	390	850	1.60
10	30706	44843.59	91746.06			17	17	48	2.00	0.30	8	710	320	710	1.25
11	30707	45703.36	91682.88			15	20	55	2.00	0.20	7	690	480	730	1.40
12	30708	42895.03	94793.44			16	22	40	1.00	0.20	6	750	390	980	1.30
13	30709	42540.07	96134.75			10	22	82	1.00	0.30	5	440	390	800	1.20
14	30710	42124.21	97236.31			26	17	63	1.00	0.20	13	2300	480	740	2.10
15	30711	48086.20	92433.88			25	14	57	1.00	0.20	9	600	450	810	1.95
16	30712	48257.85	95533.25			13	11	49	1.00	0.60	6	300	830	800	1.60
17	30713	47392.82	94196.50			3	9	106	1.00	0.20	3	420	300	800	0.95
18	30714	45873.82	94462.25			23	17	57	2.00	0.20	7	265	430	560	1.95
19	30715	44400.21	101487.81			41	24	55	3.00	0.40	7	1250	300	740	1.60
20	30716	42815.86	100333.75			31	41	68	2.00	0.20	7	1100	390	760	1.80
21	30717	44714.94	110726.69			44	33	88	2.00	0.20	10	1200	430	660	2.05
22	30718	43350.97	109671.81			36	27	83	2.00	0.40	11	1300	500	710	2.15
23	30719	41863.77	107757.38			32	26	79	2.00	0.30	8	690	500	780	2.00
24	30720	42022.94	107536.75			44	22	50	1.00	0.40	6	265	270	720	1.30
25	30721	43784.13	107850.19			28	27	76	2.00	0.20	9	1700	430	710	2.20
26	30722	44474.04	105167.56			28	19	67	1.00	0.20	7	630	430	950	1.75
27	30723	42929.62	103993.31			33	23	61	2.00	0.40	7	705	320	770	1.65
28	30724	40815.35	105521.31			38	25	72	2.00	0.40	8	650	430	770	2.05
29	30725	40728.06	103581.63			36	25	71	2.00	0.20	9	800	700	780	2.30
30	30731	43465.85	124271.50			19	18	47	1.00	0.20	7	1000	400	750	1.40
31	30732	31634.39	132035.50			20	19	68	2.00	0.20	8	920	400	850	1.75
32	30733	31819.64	133274.75			56	29	95	2.00	0.30	9	1500	610	950	2.65
33	30734	33099.04	133129.38			49	23	109	2.00	0.20	6	2050	380	680	2.50
34	30735	34139.11	133144.94			19	17	68	1.00	0.20	7	780	700	820	2.05
35	30738	48158.51	122276.25			39	45	102	2.00	0.40	12	915	470	550	2.25
36	30739	48744.42	124654.75			50	55	65	2.00	0.20	9	2000	330	600	1.60
37	30740	48443.49	124275.56			61	43	63	2.00	0.40	9	1200	380	560	1.85
38	32787	43246.01	124312.31			34	20	46	1.00	0.20	8	1050	450	1050	1.60
39	32801	43692.02	125910.63			26	16	49	1.00	0.20	9	890	300	940	1.50
40	32805	43907.57	125926.75			24	32	62	2.00	0.40	9	1250	860	790	2.25
41	32813	43272.27	126786.25			27	29	59	2.00	0.20	9	870	670	770	2.15
42	32819	43128.82	126428.75			39	13	43	2.00	0.20	9	935	400	1100	1.75
43	32824	42537.61	128049.31			44	29	71	1.00	0.20	10	1650	450	760	2.15
44	32906	42457.27	127969.69			26	21	55	2.00	0.20	10	1100	610	860	1.90
45	32907	40516.46	139468.88			18	20	64	1.00	0.30	6	825	860	730	2.15
46	32910	39826.59	134900.88			13	9	44	1.00	0.20	9	1250	380	710	1.00
47	32917	38910.49	137694.25			7	10	127	1.00	0.20	4	660	210	1230	1.15
48	32921	38949.74	135644.63			5	9	126	1.00	0.20	4	460	240	960	1.05
49	33741	42008.03	130511.63			29	27	60	1.00	0.20	5	730	430	820	1.50
50	33742	42309.04	130750.31			47	45	106	1.00	0.20	8	860	430	710	2.25
51	33743	44603.78	129500.63			5	17	53	1.00	0.20	5	1660	350	870	0.75
52	33744	43017.50	132747.38			44	35	425	1.00	0.20	7	835	430	940	1.70
53	33745	37932.90	131668.88			46	34	104	0.50	0.20	7	1240	410	920	2.10
54	33782	25738.85	114249.00			16	15	49	1.00	0.20	7	460	470	740	1.50
55	33783	20931.70	128130.38	33784	1	9	11	37	2.00	0.20	8	660	230	510	1.70
56	33803	40099.90	133319.75			25	13	85	2.00	0.20	7	1550	350	1020	1.20
57	33805	40719.64	133257.13	33807	1	33	22	66	2.00	0.30	9	1000	350	950	2.10
58	33809	21020.45	125309.25			13	11	57	2.00	0.20	15	660	500	590	2.50

59 60 61	SAMP NUM	EASTING (FT)	NORTHING (FT)	REPL NUM	REPL STAT	CU(PPM)	PB(PPM)	ZN(PPM)	MO(PPM)	AG(PPM)	CO(PPM)	MN(PPM)	F(PPM)	BA(PPM)	FE(PCT)
62	33818	21063.86	107736.13			8	14	39	1.00	0.20	10	680	240	760	1.40
63	33819	22977.58	107288.31			21	39	39	2.00	0.20	11	720	100	1310	1.40
64	33820	24213.08	106970.31			7	12	21	1.00	0.20	5	185	230	570	1.00
65	33822	24285.29	109189.56			13	5	19	2.00	0.20	4	300	210	350	0.75
66	33823	24105.60	109212.19			16	15	44	1.00	0.20	10	500	380	690	1.30
67	33824	20963.34	110457.88			32	22	68	2.00	0.20	11	820	520	1100	2.05
68	33825	23741.04	111657.69			14	14	44	3.00	0.20	4	340	380	390	1.10
69	33827	29523.65	96632.25			10	11	71	2.00	0.20	6	335	310	730	1.65
70	33828	31403.47	97056.63			14	32	82	3.00	0.20	6	2400	570	840	2.05
71	33829	31970.45	98914.50			25	14	59	3.00	0.20	8	1450	280	870	3.00
72	33830	30678.44	101039.38			29	14	68	2.00	0.20	10	1250	430	1130	2.40
73	33831	29413.04	102794.44			15	5	60	2.00	0.20	13	685	550	630	2.30
74	33832	29524.35	103572.94			11	9	35	2.00	0.20	6	450	580	650	1.65
75	33833	27725.46	106419.31			8	18	51	2.00	0.20	6	535	450	690	1.75
76	33834	28296.66	104430.88			11	19	65	3.00	0.20	5	450	500	780	1.95
77	33835	29055.98	102999.69			8	26	176	3.00	0.20	5	430	400	800	1.55
78	33836	29720.71	95049.19			11	10	71	2.00	0.20	6	450	240	800	1.60
79	33837	31151.36	93837.56			11	11	38	2.00	0.20	7	183	500	770	1.65
80	33838	34610.88	93704.56			11	14	122	2.00	0.20	9	770	280	820	1.40
81	33839	37250.56	93614.63			14	29	41	1.00	0.20	16	1200	380	960	1.45
82	33840	35954.18	94579.50			11	11	69	2.00	0.20	7	350	230	1080	1.65
83	33841	33296.69	95249.50			15	71	159	2.00	0.20	13	303	430	330	2.05
84	33842	33198.56	95749.88			16	21	75	1.00	0.20	13	640	400	550	2.50
85	33843	33059.24	95930.38			9	51	163	2.00	0.40	8	440	300	750	1.90
86	33844	35842.53	96799.94	33845	1	20	19	81	2.00	0.20	14	800	520	730	2.50
87	33846	36523.50	97057.38			13	15	46	1.00	0.30	9	530	580	1060	1.40
88	33847	46932.00	104618.31			18	17	58	2.00	0.30	9	430	610	650	1.60
89	33848	47629.21	103875.69			28	25	77	2.00	0.20	10	805	470	530	1.95
90	33849	49484.62	102648.69			17	15	85	3.00	0.20	8	325	280	800	1.60
91	33850	51170.23	104142.38			33	25	52	2.00	0.30	14	1100	450	790	1.60
92	33878	41622.01	133813.25			46	22	63	2.00	0.20	9	1800	520	800	1.90
93	33879	42564.03	134289.25			22	11	50	2.00	0.20	9	735	450	760	1.30
94	33880	41448.25	137025.75			17	16	61	2.00	0.20	6	570	680	720	1.50
95	33884	50951.07	104363.19			21	18	73	1.00	0.20	7	300	800	770	1.70
96	33885	49254.46	105269.63			30	18	61	2.00	0.20	5	305	590	570	1.80
97	33886	49612.95	104868.25			77	40	85	2.00	0.20	9	1250	390	530	2.25
98	33887	50171.90	104586.13			64	37	70	2.00	0.20	9	1200	610	720	2.25
99	33888	52587.74	103477.00			65	30	74	2.00	0.40	9	685	430	760	2.35
100	33889	52458.65	106377.56			20	12	99	2.00	0.20	5	630	680	720	1.60
101	33890	52708.65	103716.56			20	18	96	2.00	0.20	8	1750	620	610	1.90
102	33891	53205.11	102774.75			20	16	128	1.00	0.20	5	1200	880	800	1.85
103	38058	41107.83	103520.19			52	31	82	2.00	0.20	8	1200	590	800	2.60
104	38059	41261.21	101759.63			34	21	90	2.00	0.20	8	830	680	750	2.15
105	38060	36490.27	104177.56			37	25	65	2.00	0.20	8	835	560	700	2.00
106	38061	48398.62	122315.63	30638	1	40	45	99	1.00	0.20	9	835	390	510	2.10
107	38062	49702.18	123752.38			65	46	93	2.00	0.20	10	1100	530	720	2.30
108	38063	49662.83	124012.50			55	47	82	1.00	0.20	7	1350	590	680	1.90
109	38064	50803.68	124349.69			49	36	73	1.00	0.20	8	745	530	750	2.00
110	38065	50864.32	124609.50			43	34	78	1.00	0.30	6	1450	430	690	1.60
111	38066	51457.56	121888.00			27	28	68	1.00	0.20	8	770	970	690	1.70
112	38067	51732.79	120780.38			19	15	63	1.00	0.20	4	615	790	700	1.40
113	38068	52502.85	118137.50			36	27	61	1.00	0.30	4	720	300	580	1.60
114	38070	53819.29	122582.19			21	15	76	1.00	0.20	4	370	1000	730	1.60
115	38071	53858.30	122182.06			22	16	83	0.50	0.30	6	410	1100	690	1.60
116	38072	53903.67	124947.13			25	17	67	0.50	0.30	4	315	530	630	1.55

117	SAMP	EASTING	NORTHING	REPL	REPL	CU(PPM)	PB(PPM)	ZN(PPM)	MO(PPM)	AG(PPM)	CO(PPM)	MN(PPM)	F(PPM)	BA(PPM)	FE(PCT)
118	NUM	(FT)	(FT)	NUM	STAT										
119	-----														
120	46027	21361.95	99330.94			18	24	47	1.00	0.20	8	610	800	860	1.90
121	46049	26642.02	97345.44			13	22	35	1.00	0.20	6	420	400	740	1.40
122	46050	21241.33	96532.38			4	8	21	1.00	0.20	2	305	230	790	0.60
123	46326	31694.59	105335.63			11	11	44	1.00	0.20	5	285	380	720	1.35
124	46327	34117.20	106026.50			12	10	42	1.00	0.20	5	280	410	700	1.35
125	46328	36021.18	107079.31			14	40	74	1.00	0.20	7	425	580	760	1.60
126	46474	49505.59	97588.56			53	30	73	2.00	0.70	8	1300	430	710	2.75
127	46475	50252.35	99385.75			17	24	101	2.00	0.20	7	1110	320	850	1.30
128	46476	51685.65	97600.38			23	19	55	1.00	0.20	5	865	350	750	1.40
129	46485	27087.56	127722.19			12	13	106	2.00	0.20	5	840	120	930	1.20
130	46486	27465.26	129318.06			9	13	70	2.00	0.20	8	260	210	860	1.05
131	46502	48105.86	129985.75			28	20	63	1.00	0.20	7	715	430	750	1.90
132	46503	49182.81	129261.25			11	9	80	1.00	0.40	3	310	230	890	1.15
133	46504	49382.29	129140.38			48	36	70	1.00	0.20	7	1250	410	720	1.80
134	46505	51615.58	127550.94			3	14	102	0.50	0.20	4	1200	300	850	0.90
135	46506	50910.32	131033.94			18	12	83	1.00	0.20	4	500	230	880	1.30
136	46507	53382.76	129243.44			4	11	95	1.00	0.20	2	390	220	740	0.90
137	46508	53484.70	129703.00			51	24	110	1.00	0.20	7	585	390	780	2.60
138	46509	52175.98	132368.63			7	9	87	0.50	0.20	3	350	410	780	1.10
139	46510	49313.18	138479.38			8	9	53	0.50	0.20	3	245	390	710	1.10
140	46511	50012.94	136377.75			5	8	74	0.50	0.20	3	350	800	700	1.00
141	46512	49026.34	134821.94			8	9	56	1.00	0.20	3	360	760	680	1.30
142	46513	50416.84	132576.00			9	8	69	0.50	0.20	2	170	530	720	1.10
143	46514	48780.17	133363.00			12	7	95	1.00	0.20	2	253	370	850	1.10
144	46515	50695.40	132234.88			44	27	85	2.00	0.20	9	780	480	860	2.50
145	46516	49213.21	131721.13			57	30	69	1.00	0.40	9	790	480	760	2.25
146	46517	48674.91	132123.38			47	24	69	1.00	0.20	8	1100	260	710	2.10
147	46518	48617.62	132763.69			47	23	83	1.00	0.20	9	560	370	800	2.35
148	46519	46639.58	133232.00			17	16	63	0.50	0.30	3	590	350	780	1.25
149	46520	45822.48	133915.50			46	29	79	0.50	0.30	7	930	450	730	2.50
150	46521	45925.35	134595.06			44	21	77	0.50	0.20	8	935	450	840	2.30
151	46522	47729.91	135667.44			12	11	57	2.00	0.40	4	255	620	730	1.40
152	46523	47404.58	134408.81			7	12	67	2.00	0.20	4	295	530	760	1.10
153	46524	45591.54	136056.50			15	11	42	1.00	0.30	6	349	350	690	1.25
154	46525	45768.15	135255.75			51	21	79	3.00	0.30	8	1200	550	700	3.05
155	46526	24130.37	99890.81	47400	1	4	11	66	2.00	0.20	5	575	270	740	1.20
156	46527	29641.20	106111.50			29	13	59	3.00	0.20	8	1050	350	990	2.50
157	46528	30297.91	104501.81			23	19	69	2.00	0.40	7	650	480	800	2.40
158	46529	41506.52	142465.63			24	17	56	2.00	0.20	6	550	680	750	1.75
159	46530	43489.47	143338.94			44	22	61	2.00	0.20	5	1400	300	720	1.50
160	46531	45710.43	143012.19			11	11	56	2.00	0.20	3	660	590	730	1.20
161	46532	45731.00	142692.19			12	11	66	1.00	0.20	4	360	500	740	1.25
162	46533	47752.16	142055.88			9	14	53	2.00	0.20	4	1100	550	720	1.30
163	46534	43841.61	140997.75			35	22	38	1.00	0.20	5	415	350	610	1.40
164	46535	43982.01	141117.25			50	33	62	2.00	0.20	8	1600	410	660	2.05
165	46536	45308.23	142383.63			48	22	81	2.00	0.40	6	1000	480	670	2.80
166	46537	45076.70	139533.56			27	20	65	1.00	0.30	8	895	550	700	1.55
167	46538	45061.14	140853.63			23	8	71	4.00	0.20	5	895	350	580	2.15
168	46539	45372.74	138352.56	46540	1	41	18	61	3.00	0.60	8	770	480	700	2.15
169	46541	46954.40	138847.25			14	12	73	4.00	0.20	6	915	590	700	1.60
170	46542	43590.39	137658.56			13	11	86	2.00	0.20	6	660	320	840	1.10
171	46551	49382.29	129140.38			27	26	64	2.00	0.20	9	1100	830	740	1.60
172	46602	22708.57	105292.00			15	9	81	3.00	0.30	5	240	290	960	1.40
173	46603	21052.21	104175.88			6	23	43	2.00	0.20	6	200	390	680	1.30
174	46604	21819.67	104684.88			21	29	91	2.00	0.20	11	1200	490	760	1.80

	SAMP	EASTING	NORTHING	REPL	REPL	CU (PPM)	PB (PPM)	ZN (PPM)	MO (PPM)	AG (PPM)	CO (PPM)	MN (PPM)	F (PPM)	BA (PPM)	FE (PCT)
	NUM	(FT)	(FT)	NUM	STAT										
175															
176															
177															
178	46605	22274.08	102918.06			16	7	30	1.00	0.20	5	294	480	670	1.30
179	46606	25703.26	100768.13			10	7	35	1.00	0.20	5	240	460	680	1.25
180	46607	26775.92	101632.56			14	8	21	5.00	0.20	5	153	840	770	0.80
181	46608	27136.17	101647.38			11	27	97	3.00	0.20	7	684	600	770	1.65
182	46609	49472.28	115328.88			16	34	130	3.00	0.20	10	3450	460	1150	1.60
183	46610	49517.17	116628.69			23	34	100	2.00	0.20	10	1850	400	720	1.70
184	46611	50256.49	116445.88			46	35	77	3.00	0.30	7	825	660	810	2.10
185	46612	51251.14	115022.19			29	33	71	2.00	0.20	6	660	770	800	1.80
186	46613	52894.14	115816.00			10	11	76	2.00	0.20	5	565	930	710	1.35
187	46616	35641.86	123220.88			8	11	44	1.00	0.20	7	905	590	870	1.10
188	46617	34287.28	124666.00			24	21	63	2.00	0.20	5	705	800	1030	1.75
189	46618	32390.91	125633.13			33	20	130	2.00	0.30	6	695	720	1030	2.05
190	46619	35655.98	127678.44			33	23	136	2.00	0.20	8	710	570	1050	2.10
191	46620	35955.98	127677.25	46621	1	37	37	90	2.00	0.20	9	1100	630	860	2.25
192	46622	35941.22	128917.25			39	28	80	2.00	0.40	7	590	590	800	1.80
193	46623	35802.50	129217.88			47	46	119	2.00	0.20	10	1450	570	1060	2.30
194	46624	35085.89	130020.94			46	43	100	2.00	0.20	10	1500	420	730	2.20
195	46625	37090.53	131112.44			34	22	84	2.00	0.20	7	1100	550	890	2.10
196	46633	28098.97	101833.44			11	11	39	2.00	0.40	3	300	600	690	1.25
197	46634	39728.97	119785.50			16	13	72	1.00	0.40	4	840	420	750	1.55
198	46635	41761.96	117917.81			44	39	110	2.00	0.30	10	1400	310	550	2.10
199	46636	39483.78	118406.38			11	10	158	1.00	0.30	7	600	490	890	1.35
200	46637	38735.37	116169.19			32	20	50	3.00	0.60	6	630	210	580	1.60
201	46638	37727.78	114153.00			35	21	53	3.00	0.20	5	615	220	720	1.80
202	46639	36979.36	111915.75			38	19	56	2.00	0.30	7	720	270	970	2.00
203	46640	37550.64	109593.63			43	21	63	3.00	0.30	6	690	360	970	2.40
204	46641	36385.16	108138.00			33	20	69	2.00	0.30	5	440	400	740	1.85
205	46642	38802.74	123449.00			20	36	121	2.00	0.20	11	1200	240	790	1.55
206	46643	38524.09	123810.06			49	29	128	3.00	0.40	9	1350	330	710	2.05
207	46644	39867.24	124645.00			37	32	54	3.00	0.20	6	830	380	720	1.90
208	46645	38112.29	125991.63			52	36	72	3.00	0.20	9	920	310	700	2.10
209	46646	35469.38	125221.56	46647	1	38	22	77	3.00	0.40	7	1200	540	870	2.10
210	46648	36193.06	126198.88			44	42	125	2.00	0.40	9	820	300	590	2.05
211	46649	36193.78	127156.19			52	44	99	2.00	0.20	11	1100	310	720	2.25
212	46651	49566.26	125392.75			3	5	40	1.00	0.30	1	195	360	780	0.95
213	46654	21176.22	101693.81			89	15	60	2.00	0.20	8	930	270	1340	1.90
214	46655	21138.54	101854.38			11	10	42	1.00	0.20	5	560	340	820	1.30
215	46656	26599.39	100495.06			16	8	29	3.00	0.20	6	540	770	520	1.40
216	46657	27177.27	98966.50			5	5	32	3.00	0.20	1	272	250	270	0.90
217	46658	46257.50	116720.94			32	42	59	3.00	0.20	9	725	440	690	1.70
218	46659	46202.22	117981.13			33	25	72	3.00	0.60	7	710	400	650	2.05
219	46660	46482.53	118060.13			35	31	74	2.00	0.40	7	630	520	710	1.95
220	46661	43997.66	116769.44			37	29	83	2.00	0.40	8	585	520	710	2.00
221	46662	42995.63	116233.19			33	26	78	3.00	0.20	8	800	570	720	2.00
222	46663	42887.59	114093.56			37	31	103	3.00	0.30	9	805	660	760	2.40
223	46664	45507.57	114083.69			49	43	108	2.00	0.20	12	1000	460	580	2.05
224	46665	41982.49	112737.00			25	14	51	6.00	0.20	3	395	250	470	0.70
225	46666	40680.39	112181.88			41	25	81	3.00	0.60	9	785	490	830	2.10
226	46667	37850.94	109672.44			35	22	72	2.00	0.20	9	665	460	790	2.05
227	46676	49894.38	121207.25			11	12	64	1.00	0.30	2	470	280	730	1.25
228	46677	30767.84	119499.19			46	29	43	1.00	0.30	7	570	440	750	1.25
229	46678	34430.44	120185.44			93	32	72	3.00	0.30	7	805	160	1300	2.40
230	46679	34130.51	120206.50			62	24	84	2.00	1.00	10	1700	520	950	2.00
231	46680	33838.18	118078.63			69	20	90	2.00	0.20	9	520	230	850	1.75
232	46681	31479.03	117156.50			60	19	72	1.00	0.20	7	445	300	870	1.60

233	SAMP	EASTING	NORTHING	REPL	REPL	CU (PPM)	PB (PPM)	ZN (PPM)	MO (PPM)	AG (PPM)	CO (PPM)	MN (PPM)	F (PPM)	BA (PPM)	FE (PCT)
234	NUM	(FT)	(FT)	NUM	STAT										
235															
236	46689	30692.56	126879.50			39	52	108	2.00	0.40	10	1250	380	970	1.85
237	46690	31236.44	127797.19			33	34	75	2.00	0.20	7	910	350	1060	1.60
238	46691	32042.11	129133.75			25	10	61	2.00	0.20	6	575	400	1020	1.40
239	46692	30353.28	130026.13			21	18	65	2.00	0.20	7	1500	400	710	1.45
240	46693	33770.31	131066.50	46694	1	22	16	64	2.00	0.20	8	865	420	1010	1.55
241	46695	35797.15	132677.94			18	14	57	2.00	0.20	8	565	440	910	1.45
242	46696	35797.91	132857.94			17	15	70	2.00	0.20	8	525	730	850	2.00
243	46697	33747.42	135106.63			28	11	46	2.00	0.20	10	1200	500	630	1.70
244	46698	31328.03	135256.88			14	14	110	2.00	0.20	7	350	700	1130	1.25
245	46699	29714.98	133793.44			65	26	77	2.00	0.40	9	2800	670	1000	2.05
246	46700	31495.58	132316.13			17	14	68	2.00	0.20	9	405	780	980	2.00
247	46701	35710.81	114960.56			9	7	38	1.00	0.20	7	440	330	740	1.05
248	46705	21655.98	124902.19			13	28	324	3.00	0.30	7	800	310	720	1.80
249	46706	22888.60	124228.44			14	10	71	2.00	0.30	9	400	500	610	2.05
250	46726	34272.39	115386.00			163	10	87	3.00	0.20	9	440	210	920	1.40
251	46727	34283.22	112945.88			96	15	51	3.00	0.30	13	680	210	1170	1.80
252	46728	34003.82	113107.00			43	46	66	2.00	0.20	9	805	330	870	1.35
253	46729	31042.57	112778.13			15	20	180	3.00	0.20	10	1000	400	1060	1.55
254	46731	32614.06	110512.19			89	29	64	3.00	0.20	10	930	280	920	1.60
255	46733	31935.87	105674.69			31	12	49	3.00	0.20	14	1350	540	560	2.20
256	46735	37441.11	140859.25			4	7	87	1.00	0.20	6	425	210	940	0.90
257	46736	37156.54	139500.19			136	7	47	2.00	0.30	7	430	360	710	1.05
258	46737	34414.74	138969.38			14	13	67	1.00	0.20	7	750	440	700	1.30
259	46738	33255.42	139173.31			28	33	66	2.00	0.20	8	1550	420	600	1.85
260	46739	32658.72	140155.31			20	21	67	2.00	0.20	6	1050	470	640	1.60
261	46740	33921.32	140931.06			9	8	46	1.00	0.20	7	240	280	780	0.90
262	46741	34520.85	140789.06			18	12	58	2.00	0.20	7	655	300	770	1.35
263	46780	28887.05	133082.56			5	5	75	1.00	0.20	3	430	230	790	0.80
264	46781	28029.10	133272.06			14	7	53	1.00	0.20	4	340	320	670	1.00
265	46782	26716.99	132186.50			39	17	166	2.00	0.20	31	5200	120	1400	6.00
266	46783	24267.17	131313.63			16	12	58	2.00	0.20	4	470	180	990	1.50
267	46784	24224.73	131094.13			13	15	50	2.00	0.20	2	365	400	720	1.70
268	46785	22643.56	129191.50			20	16	66	1.00	0.20	7	735	470	980	1.90
269	46796	25615.38	115797.75			21	65	59	2.00	0.20	5	360	500	710	1.80
270	46797	22013.20	121038.00			19	20	82	2.00	0.20	4	330	290	480	1.20
271	46798	22025.49	125758.13			20	19	73	2.00	0.20	10	440	370	750	2.20
272	46799	27230.50	137853.69			12	14	36	1.00	0.20	6	280	320	760	1.30
273	46800	29384.59	139960.50			48	37	78	2.00	0.30	7	1120	390	760	2.55
274	46808	23133.98	115685.25			10	16	116	2.00	0.30	7	425	520	720	1.60
275	46809	21917.44	114198.63			14	10	62	2.00	0.20	7	305	350	700	1.30
276	46810	20989.46	115288.94			12	11	58	2.00	0.30	5	650	270	750	1.55
277	46811	22461.11	116332.75			10	14	94	2.00	0.20	6	139	320	790	1.60
278	46826	38587.82	142855.38			31	20	61	2.00	0.40	11	1200	260	720	2.30
279	46827	38016.28	145377.38			60	34	81	2.00	0.40	11	1300	380	700	2.95
280	46828	35280.25	140606.50			8	11	52	0.50	0.20	4	235	440	750	1.00
281	46829	35219.44	140366.69			20	11	131	1.00	0.20	5	380	200	870	1.30
282	46830	33746.98	142611.69			20	18	69	2.00	0.20	5	375	440	720	1.45
283	46831	34008.46	143050.75			50	25	132	3.00	0.20	9	850	320	870	3.05
284	46832	35767.63	142804.88			8	9	53	1.00	0.20	4	650	260	770	0.85
285	46833	36909.98	143501.00			75	23	88	3.00	0.60	8	740	440	730	2.80
286	46834	35975.70	145204.19			47	19	74	2.00	0.20	8	1050	280	750	2.50
287	46835	35834.89	144964.69			44	16	73	3.00	0.20	7	925	330	710	2.35
288	46836	33339.48	146333.00	46837	1	15	14	45	2.00	0.20	7	540	270	780	1.25
289	46838	32486.04	142335.88			22	20	138	3.00	0.20	6	1100	300	830	1.75
290	46839	32129.74	143437.13			16	17	51	2.00	0.20	11	840	370	750	1.70

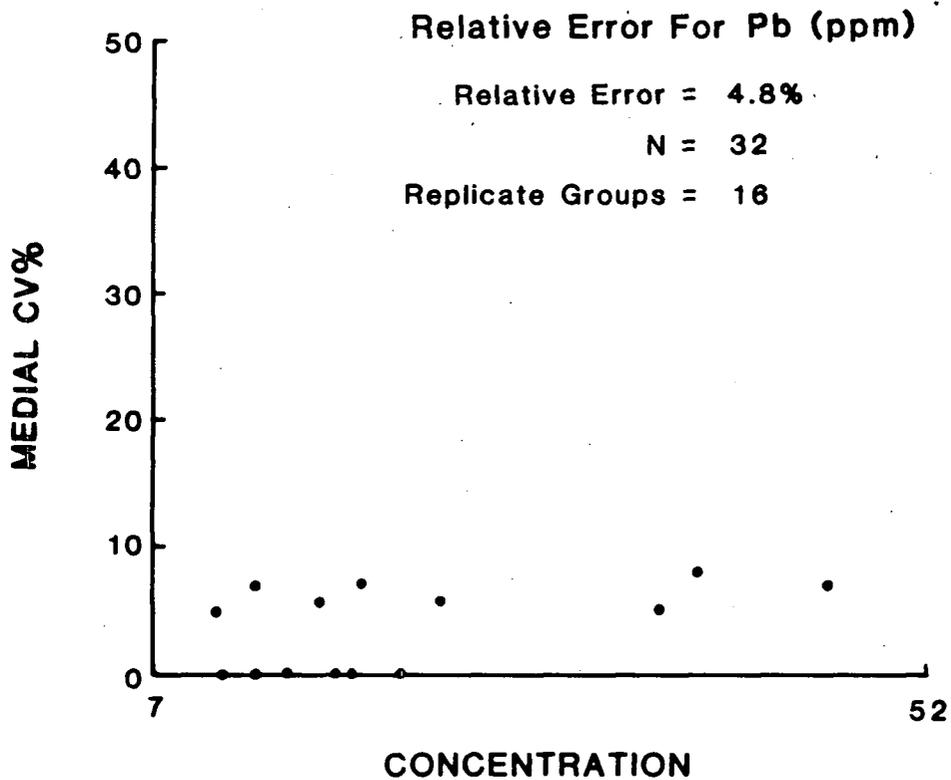
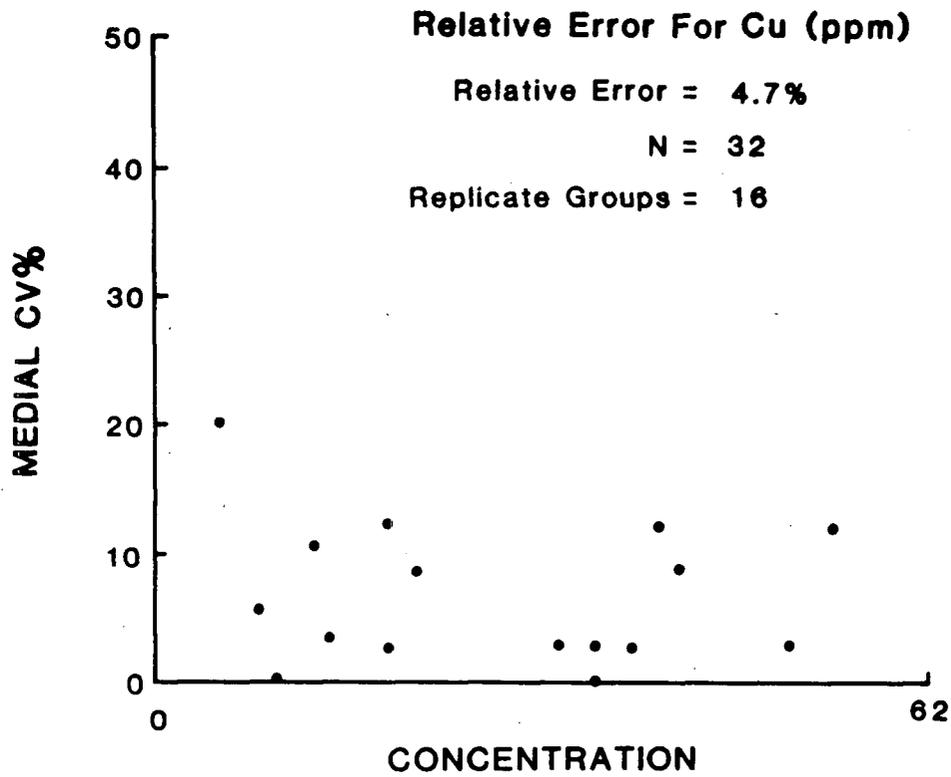
291	SAMP	EASTING	NORTHING	REPL	REPL	CU(PPM)	PB(PPM)	ZN(PPM)	MO(PPM)	AG(PPM)	CO(PPM)	MN(PPM)	F(PPM)	BA(PPM)	FE(PCT)	
292	NUM	(FT)	(FT)	NUM	STAT											
293																
294	46840	34461.82	147029.25			44	24	94	4.00	0.30	7	730	170	760	2.40	
295	46841	34646.38	148388.69			15	13	44	2.00	0.20	6	405	280	680	1.20	
296	46842	25854.06	108330.81	46843	1	10	11	75	3.00	0.20	6	258	400	770	1.65	
297	46844	28764.00	111824.88	46847	1	37	15	62	2.00	0.30	11	1200	460	970	2.40	
298	46845	26322.48	110360.13			5	13	90	1.00	0.20	4	338	270	720	1.00	
299	46846	28630.93	109546.50			7	12	89	2.00	0.20	7	750	270	750	1.55	
300	46848	27714.56	112560.13			35	14	85	2.00	0.20	8	1400	310	910	2.20	
301	46904	38016.62	132548.50			19	21	180	1.00	0.20	6	470	440	810	1.45	
302	46905	35000.12	133381.31			9	22	143	1.00	0.20	3	950	330	790	1.00	
303	46943	28969.58	139555.38			13	14	68	0.50	0.20	6	530	450	780	1.75	
304	46944	28970.02	141135.50			23	15	52	0.50	0.20	4	1020	290	810	1.10	
305	46978	30150.86	115045.06			44	51	83	2.00	0.20	8	550	380	850	1.65	
306	46979	24897.19	115965.69			28	39	75	1.00	0.20	10	1170	420	780	2.30	
307	46980	24213.13	119213.50			23	23	76	2.00	0.20	10	690	450	770	2.05	
308	46981	22239.51	117995.25			31	51	101	2.00	0.20	6	590	340	870	1.90	
309	46982	27372.48	137695.44			14	12	60	0.50	0.20	3	690	310	710	1.20	
310	46983	27925.00	138302.31			29	27	68	1.00	0.20	10	2200	390	720	1.70	
311	46984	28634.55	139151.19			33	12	52	0.50	0.20	5	1050	290	740	1.25	
312	46985	28851.51	141014.00			8	8	98	0.50	0.20	2	400	210	800	1.00	
313	47150	43106.56	134887.00			26	21	58	2.00	0.20	6	500	500	750	1.80	
314	47162	43745.12	134544.25			37	20	65	3.00	0.20	7	805	380	830	2.05	
315	47326	25386.41	118600.44			25	31	72	2.00	0.20	8	1000	570	670	1.95	
316	47327	25402.58	120060.31			34	18	58	2.00	0.20	8	630	460	660	1.80	
317	47328	25510.99	120819.25			28	22	91	2.00	0.20	4	630	480	890	1.95	
318	47330	27586.82	122236.31			36	27	78	2.00	0.30	9	1200	270	1150	2.10	
319	47331	24760.56	121687.56			26	22	67	3.00	0.20	7	1000	600	860	2.00	
320	47332	23506.40	122221.50			13	8	49	2.00	0.20	5	395	380	750	1.40	
321	47333	26001.04	125334.06			15	11	190	3.00	0.20	5	1800	220	1410	1.50	
322	47334	26042.58	125473.63			17	19	86	2.00	0.20	8	950	520	1110	1.95	
323	47335	24123.36	125554.88			20	25	85	3.00	0.20	8	605	670	950	2.10	
324	47351	22551.74	102754.00			13	14	55	2.00	0.20	8	445	330	760	1.70	
325	47352	22268.84	101177.94			10	6	39	2.00	0.20	5	320	280	580	1.15	
326	47353	23198.59	100464.38			13	8	46	2.00	0.20	6	400	300	780	1.35	
327	47354	26506.70	103756.75			6	8	43	3.00	0.20	5	340	170	640	1.35	
328	47355	41356.40	121759.38			12	15	136	3.00	0.20	10	965	140	700	1.60	
329	47356	42829.70	119973.81			28	25	81	3.00	0.20	12	925	470	990	2.20	
330	47357	43069.54	119932.94			41	14	49	2.00	0.20	7	1200	320	1000	1.70	
331	47358	43500.29	117471.25			27	18	71	1.00	0.20	5	645	290	980	1.70	
332	47359	42708.80	114414.25			42	26	97	2.00	0.20	9	805	340	880	2.60	
333	47360	40792.65	115441.50			21	28	90	1.00	0.50	9	1225	320	780	1.75	
334	47361	41724.66	113317.94			49	38	103	3.00	0.60	7	820	320	570	1.95	
335	47362	41762.94	112857.81			49	27	100	2.00	0.30	9	1000	390	920	2.35	
336	47363	39375.44	110866.75			40	24	83	2.00	0.20	9	840	340	820	2.20	
337	47364	49744.98	113387.75			28	22	88	2.00	0.20	7	1800	350	830	1.60	
338	47365	49335.14	110769.31			26	15	107	2.00	0.20	7	1160	320	1000	1.80	
339	47366	49575.66	110908.38			33	14	61	2.00	0.20	8	380	480	800	1.70	
340	47367	46565.83	108299.69			18	19	223	2.00	0.20	6	1900	480	1190	1.65	
341	47368	46619.82	106699.50			12	10	105	1.00	0.20	3	825	460	820	1.40	
342	47369	49037.70	106130.38			56	44	124	2.00	0.20	15	1900	370	750	2.40	
343	47370	49338.37	106309.25			46	25	136	2.00	0.20	7	980	390	740	2.00	
344	47371	49338.37	106309.25			45	23	135	2.00	0.20	9	1600	390	720	2.00	
345	47372	48763.64	107711.44			22	17	73	2.00	0.20	5	1000	390	660	2.00	
346	47373	48925.14	108110.88			36	13	125	2.00	0.20	6	750	440	750	1.70	
347	47374	48908.98	109130.94			13	14	81	3.00	0.20	7	475	270	840	1.30	
348	47375	27476.30	117677.25			25	23	66	3.00	0.20	7	560	620	680	2.00	

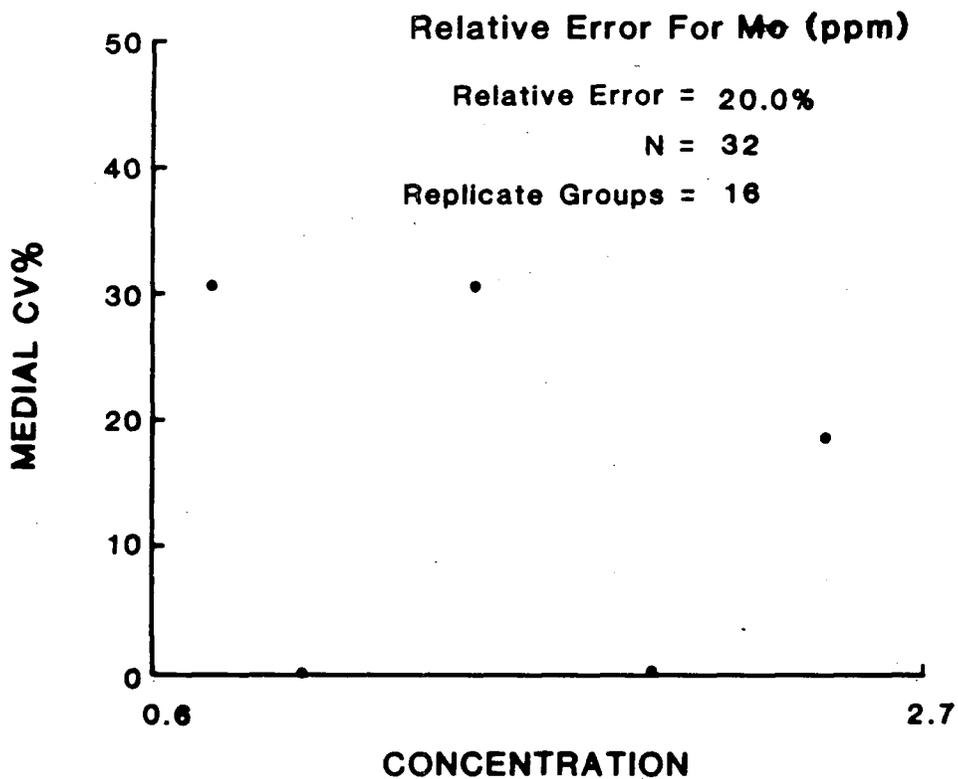
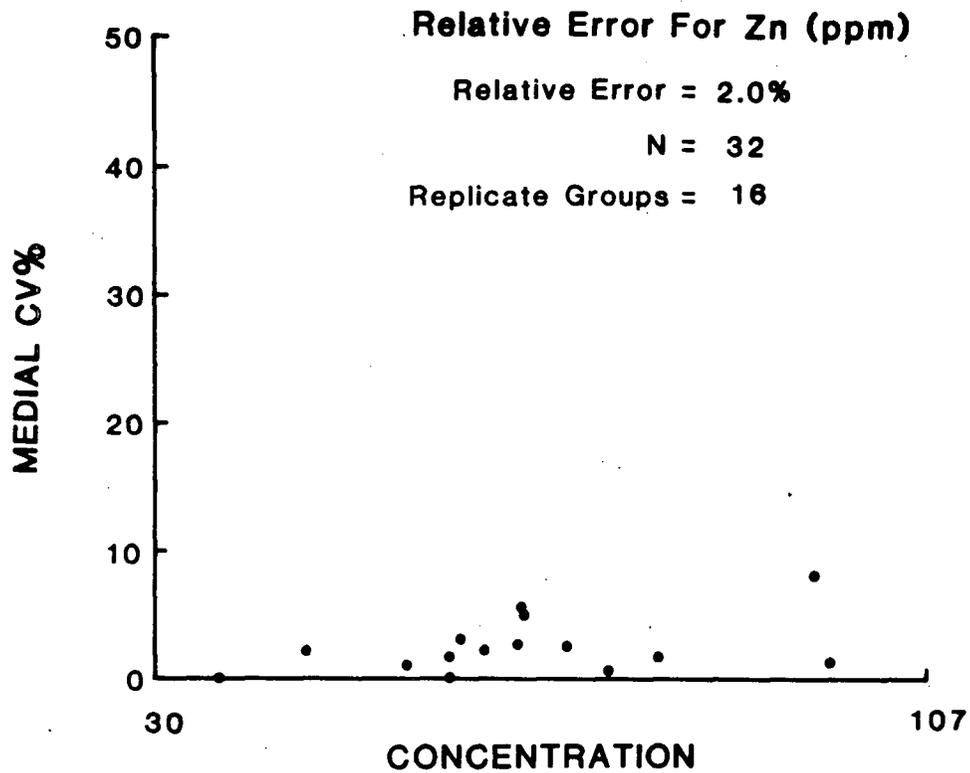
349	SAMP	EASTING	NORTHING	REPL	REPL	CU(PPM)	PB(PPM)	ZN(PPM)	MO(PPM)	AG(PPM)	CO(PPM)	MN(PPM)	F(PPM)	BA(PPM)	FE(PCT)
350	NUM	(FT)	(FT)	NUM	STAT										
351															
352	47377	38235.86	132367.63			51	28	74	4.00	0.20	11	1400	340	1000	2.30
353	47390	24736.36	98921.88			7	6	23	5.00	0.20	3	177	220	270	0.75
354	47391	24373.50	98727.13			9	15	65	2.00	0.20	5	415	270	560	1.40
355	47392	23259.40	97763.25			19	8	44	2.00	0.20	6	395	320	810	1.50
356	47393	23895.56	97494.00			56	26	56	3.00	0.20	11	1400	230	1080	1.85
357	47394	26324.63	98088.56			6	5	80	1.00	0.20	5	298	250	750	1.15
358	47395	25135.41	96095.81			13	18	77	2.00	0.20	8	1400	120	880	1.70
359	47396	23531.80	95859.06			30	19	48	2.00	0.20	6	640	240	900	1.50
360	47397	22976.51	96187.13			17	65	49	2.00	0.20	6	305	250	710	1.35
361	47398	23486.66	99640.13			14	9	32	2.00	0.20	5	410	250	840	1.35
362	47399	24215.29	100229.63			23	19	59	3.00	0.20	9	655	510	840	1.90
363	47401	28019.77	97754.19			15	11	81	3.00	0.20	7	460	300	760	1.80
364	47402	29220.54	99176.88			15	18	74	3.00	0.20	5	930	460	850	1.95
365	47403	29446.31	99573.69			20	24	73	3.00	0.20	7	360	400	780	1.80
366	47404	29765.71	99529.00			12	22	93	3.00	0.20	5	236	300	710	1.55
367	47406	28114.08	104253.50			11	10	43	2.00	0.20	5	319	530	700	1.45
368	47407	26192.85	105561.50			31	12	55	3.00	0.50	10	850	320	720	2.05
369	47408	31735.80	121615.56			22	13	32	1.00	0.20	8	720	260	740	1.10
370	47409	33016.08	121690.75			14	8	63	2.00	0.20	7	425	220	900	1.45
371	47410	33723.38	118308.06			62	55	152	1.00	0.20	8	865	240	1090	2.00
372	47411	31060.31	117498.06			80	32	65	3.00	0.30	11	855	250	850	2.15
373	47412	28419.12	112869.94			62	24	74	3.00	0.20	6	285	130	800	1.30
374	47413	40516.11	127697.88			69	20	75	3.00	0.20	9	1700	340	1250	2.30
375	47414	39787.54	130401.00			61	14	66	3.00	0.20	6	685	300	1160	1.55
376	47416	40078.04	132879.81			73	16	78	3.00	0.20	6	310	320	1160	1.70
377	47417	40018.97	133100.06	47378	1	64	21	65	2.00	0.20	9	1000	460	930	2.05
378	47418	37883.40	134149.13			44	10	53	1.00	0.20	7	860	480	820	1.30
379	47419	37603.57	134190.25			44	24	74	3.00	0.20	9	2200	380	870	2.20
380	47420	35729.68	135638.25	47421	1	55	40	57	2.00	0.20	10	1400	360	750	1.80
381	47422	33574.10	136687.38			13	11	71	2.00	0.20	8	290	310	780	1.10
382	47423	35991.80	136137.13			54	34	115	2.00	0.20	13	1800	460	730	2.55
383	47424	37510.02	137559.00			780	12	58	2.00	0.20	9	2000	260	800	1.50
384	47425	32848.12	119571.31			13	7	51	2.00	0.20	5	349	180	820	1.20
385	47469	45435.13	127457.06			38	71	69	1.00	0.20	14	1700	310	780	1.40
386	61129	20930.32	99234.50	61130	1	19	26	56	1.00	0.20	9	660	570	860	1.80
387	61240	22979.86	99671.94			20	24	55	1.00	0.20	8	650	480	850	1.90
388	68050	51746.32	97780.19			46	37	83	1.00	0.20	12	1325	280	670	1.85
389	68095	36954.33	99935.75			29	20	82	2.00	0.20	8	1560	520	800	1.60
390	68096	34660.50	101584.38			59	41	65	2.00	0.40	9	845	340	740	2.05
391	68097	36540.72	101637.31			34	24	68	3.00	0.30	9	620	540	990	2.00
392	68098	39139.73	101367.63			15	19	54	0.50	0.20	6	640	280	710	1.30
393	68099	39080.33	101527.81			30	29	59	1.00	0.20	10	660	540	750	1.95
394	68131	47539.82	101376.00			24	10	48	1.00	0.20	5	430	390	880	1.10
395	68132	45636.07	100383.13			30	21	59	1.00	0.20	14	1320	590	920	2.00
396	68133	45476.97	100623.75			20	18	64	1.00	0.20	7	505	410	930	1.85
397	68134	44343.37	97008.00			27	32	185	1.00	0.20	8	1700	220	1250	2.10
398	68154	37666.35	97813.06			17	22	51	1.00	0.20	6	555	400	1120	1.50
399	68155	37506.95	97973.69			29	21	69	1.00	0.20	11	555	500	830	2.10
400	68156	40128.21	98303.81			16	18	90	1.00	0.20	8	980	300	930	1.55
401	68157	39968.82	98464.44			119	31	58	1.00	0.20	13	730	270	1030	2.40
402	68158	40149.12	98543.75			27	17	58	1.00	0.20	9	505	440	710	1.50
403	68159	42369.48	98635.44			37	24	71	1.00	0.20	10	690	440	890	2.10
404	68175	49232.81	99509.63			52	28	99	1.00	0.40	8	2650	300	710	2.40
405	68180	50272.43	99405.69			14	15	46	0.50	0.20	4	120	410	760	1.30
406	68200	47548.53	98375.94			32	17	59	1.00	0.20	5	400	300	600	1.85

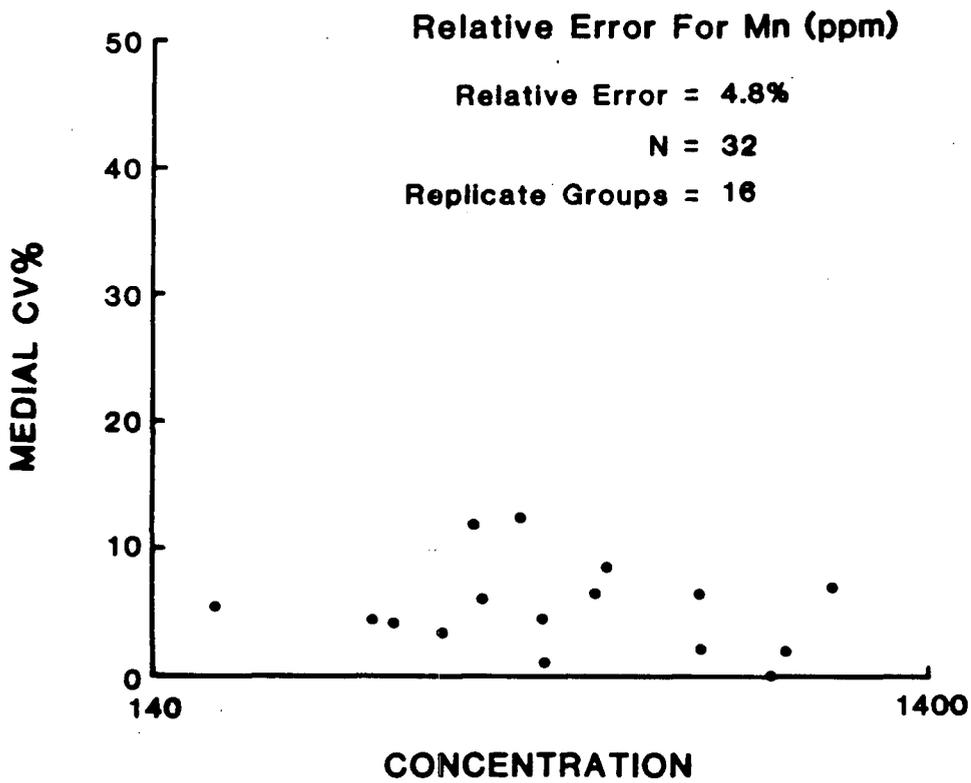
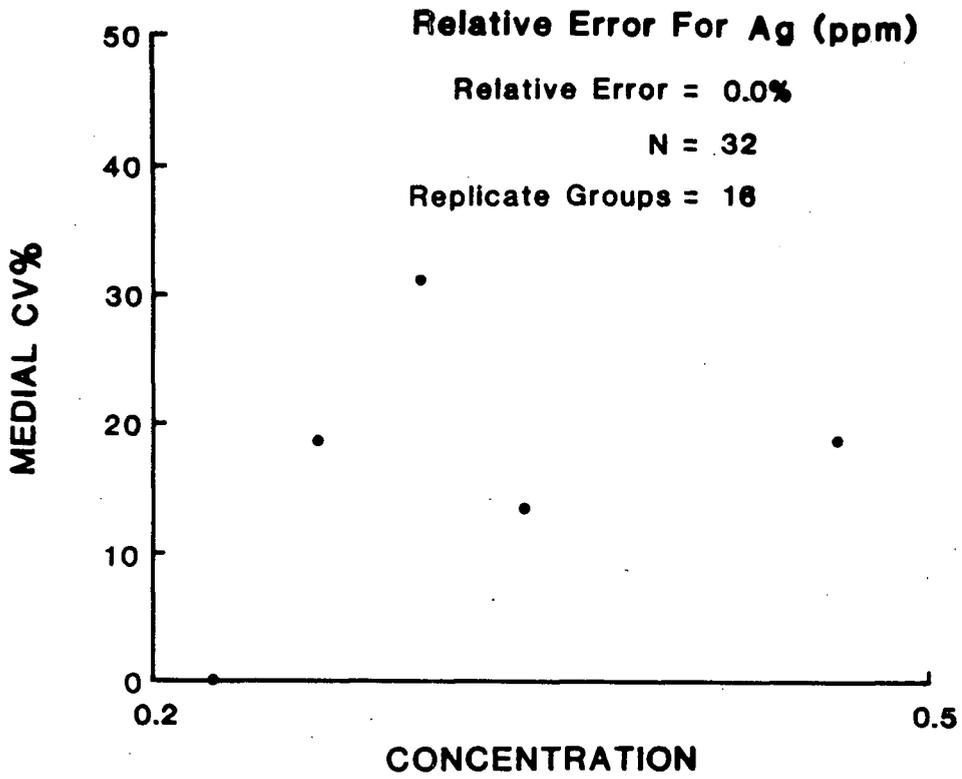
407 408 409	SAMP NUM	EASTING (FT)	NORTHING (FT)	REPL NUM	REPL STAT	CU(PPM)	PB(PPM)	ZN(PPM)	MO(PPM)	AG(PPM)	CO(PPM)	MN(PPM)	F(PPM)	BA(PPM)	FE(PCT)
410	68201	46083.28	96981.44			37	15	60	1.00	0.20	6	670	320	780	1.90
411	68202	44719.23	95906.56			25	11	47	1.00	0.20	5	302	370	710	1.45
412	68663	37988.38	119632.00	68664	1	12	13	72	1.00	0.20	7	770	330	950	1.35
413	68665	36443.28	118277.81			123	25	70	0.50	0.30	7	945	420	740	2.55
414	68666	34348.26	119605.69			12	10	90	0.50	0.20	4	1550	250	1000	1.20
415	68667	31176.48	116477.63			40	17	96	1.00	0.20	6	710	340	1000	1.60
416	68668	47492.62	150225.50			20	19	79	0.50	0.20	4	555	410	740	1.45
417	68680	20922.00	148196.38			21	17	52	1.00	0.30	5	410	380	730	2.40
418	68681	21805.75	144667.06			27	22	52	1.00	0.30	8	775	460	840	2.55
419	68682	24116.02	142235.50			40	50	98	1.00	0.40	6	1875	350	1140	2.00
420	68683	24909.63	139525.06			24	22	63	2.00	0.40	7	480	500	890	2.15
421	68684	26092.56	137679.63			14	16	42	1.00	0.30	7	405	420	710	1.70
422	68685	41013.31	144487.25			34	22	63	1.00	0.40	7	780	320	660	1.85
423	68686	43278.14	145919.69			54	30	62	1.00	0.30	10	1320	350	750	1.90
424	68687	45837.99	145871.06			9	9	72	1.00	0.20	3	410	280	760	1.00
425	68688	44706.40	148374.88			59	18	107	1.00	0.30	4	373	270	720	2.45
426	68689	44666.80	148495.00			13	12	82	1.00	0.20	4	105	270	820	1.30
427	68690	46791.56	149907.88			14	14	129	1.00	0.20	3	1100	670	880	1.20
428	68691	40006.02	148270.69			6	8	20	0.50	0.20	1	165	260	430	0.60
429	68692	39412.81	150292.69			47	22	89	1.00	0.20	3	970	280	760	1.75
430	68694	24830.72	144284.50			30	26	93	1.00	0.50	8	880	550	820	2.40
431	68695	27285.08	144754.88			26	17	96	0.50	0.30	7	710	280	850	1.80
432	68696	27906.82	143002.44			5	6	100	0.50	0.20	2	390	190	730	0.75
433	68697	29948.90	142847.75			12	9	89	1.00	0.20	2	530	210	750	1.00
434	68698	30085.42	143129.38			15	11	48	0.50	0.20	4	435	270	730	1.30
435	68759	26692.90	148967.88			12	17	68	0.50	0.20	5	800	290	730	1.25
436	68760	26477.37	148605.19			17	15	45	0.50	0.20	4	339	420	780	1.40
437	68761	27909.08	149282.94			14	17	51	0.50	0.20	6	780	480	860	1.50
438	68762	29302.35	148220.13			14	16	147	1.00	0.20	4	760	290	890	1.20
439	68979	25655.56	130258.19			20	26	66	1.00	0.20	7	590	300	950	1.40
440	68980	24332.38	128172.75			36	34	76	1.00	0.30	7	610	430	970	2.00

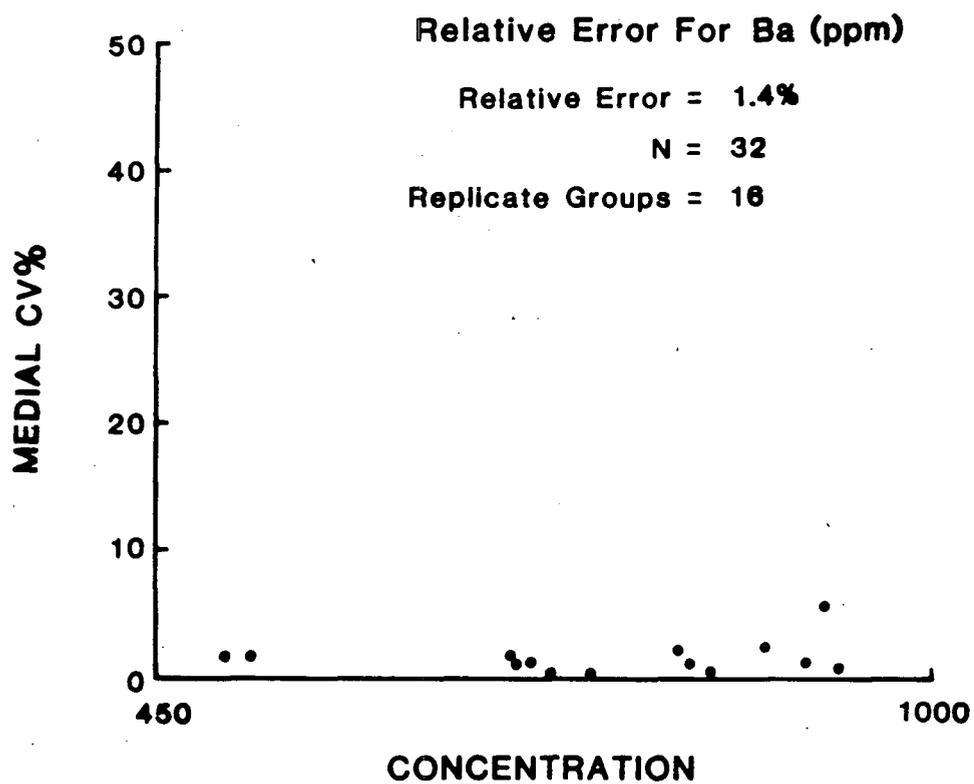
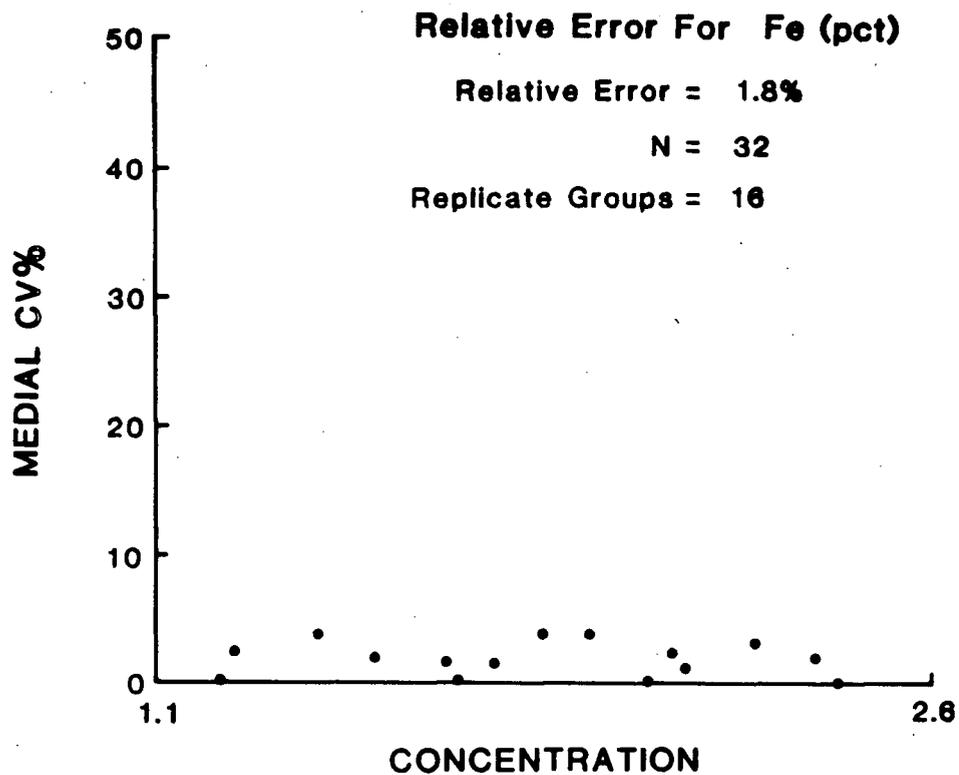
APPENDIX 3

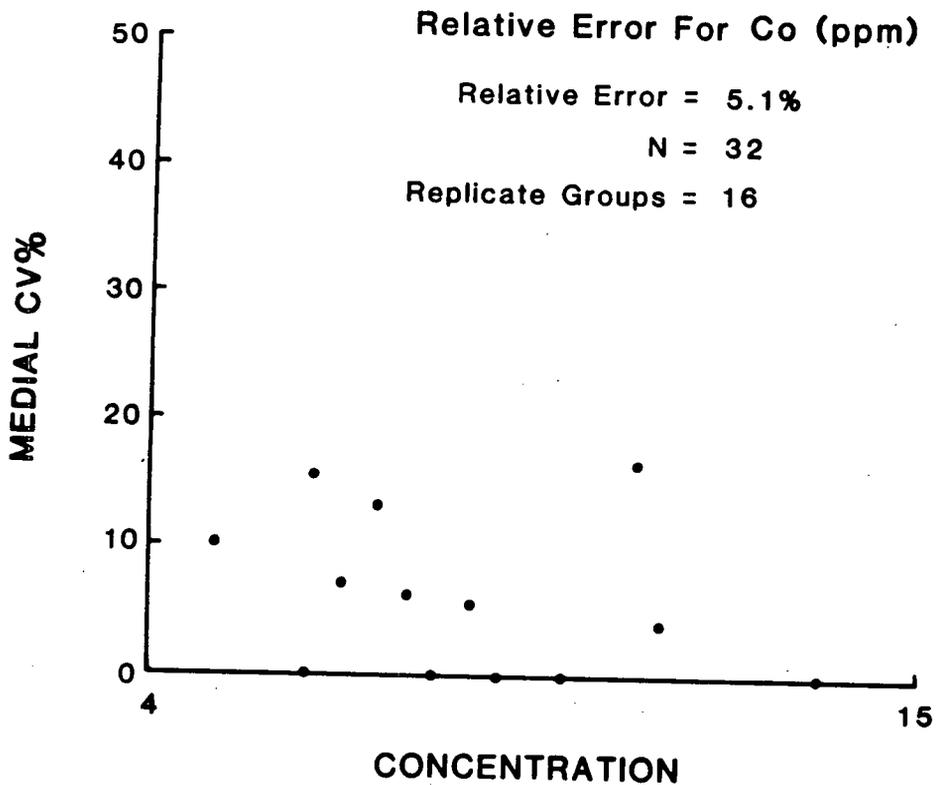
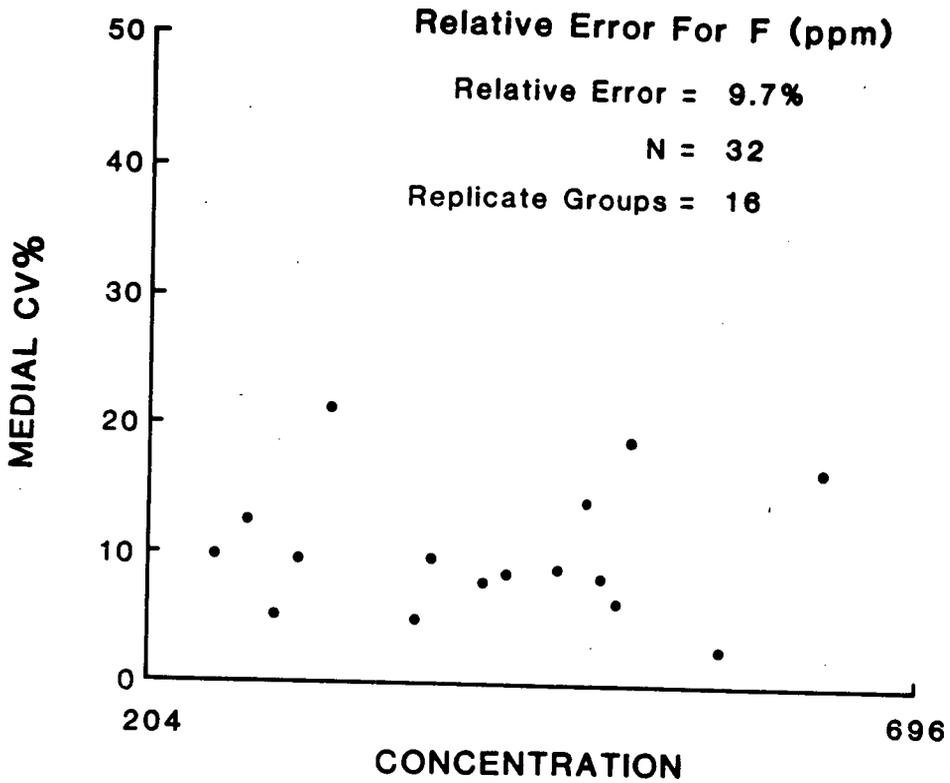
STREAM SEDIMENT SURVEY -
RELATIVE ERROR PLOTS











APPENDIX 4

STREAM SEDIMENT PROBABILITY PLOTS

T I T L E S	PROGRAM: PERCENTAGE HISTOGRAMS	NAME	CU
	DATA: DAISY CREEK RECON STREAM SEDIMENT GEOCHEMISTRY	N	416
	RUN:	X BAR	30.377
	TIME: 25/03/83 11:55:33	STD DEV	42.122

CELL	LOWER LIMIT	NMBR	%	ARITHMETIC VALUES	BAR INTERVAL = 0.25000	STD DEV	ARITH.LIMIT
1		0	0.0				
2	-175.0	0	0.0				
3	-164.4	0	0.0				
4	-153.9	0	0.0				
5	-143.4	0	0.0				
6	-132.8	0	0.0				
7	-122.3	0	0.0				
8	-111.8	0	0.0				
9	-101.3	0	0.0				
10	-90.72	0	0.0				
11	-80.19	0	0.0				
12	-69.66	0	0.0				
13	-59.13	0	0.0				
14	-48.60	0	0.0				
15	-38.07	0	0.0				
16	-27.54	0	0.0				
17	-17.01	0	0.0				
18	-6.480	7	1.7	**			
19	4.051	111	26.7	*****			
20	14.58	103	24.8	*****			
21	25.11	72	17.3	*****			
22	35.64	59	14.2	*****			
23	46.17	33	7.9	*****			
24	56.70	15	3.6	****			
25	67.23	5	1.2	*			
26	77.76	1	0.2				
27	88.30	5	1.2	*			
28	98.83	0	0.0				
29	109.4	1	0.2				
30	119.9	1	0.2				
31	130.4	1	0.2				
32	140.9	0	0.0				
33	151.5	0	0.0				
34	162.0	1	0.2				
35	172.5	0	0.0				
36	183.1	0	0.0				
37	193.6	0	0.0				
38	204.1	0	0.0				
39	214.7	0	0.0				
40	225.2	0	0.0				
41	235.7	1	0.2				

T I T L E S	PROGRAM:	PERCENTAGE CUMULATIVE FREQUENCY PLOTS	NAME	CU
	DATA:	DAISY CREEK RECON STREAM SEDIMENT GEOCHEMISTRY	N	416
	RUN:		X BAR	30.377
	TIME:	25/03/83 11:55:33	STD DEV	42.122

CELL	LOWER LIMIT	NMBR	%	ARITHMETIC VALUES										BAR INTERVAL = 0.12500	STD DEV	ARITH.LIMIT		
1	133.1	3	0.7	*														
2	127.8	0	0.7	*														
3	122.5	1	1.0	*														
4	117.3	1	1.2	*														
5	112.0	0	1.2	*														
6	106.7	0	1.2	*														
7	101.5	0	1.2	*														
8	96.19	0	1.2	*														
9	90.93	3	1.9	*														
10	85.66	2	2.4	*														
11	80.40	0	2.4	*														
12	75.13	2	2.9	*														
13	69.87	2	3.4	*														
14	64.60	5	4.6	*														
15	59.34	9	6.7	*														
16	54.07	8	8.7	*														
17	48.81	19	13.2	*														
18	43.54	32	20.9	*														
19	38.28	17	25.0	*														
20	33.01	31	32.5	*														
21	27.74	45	43.3	*														
22	22.48	37	52.2	*														
23	17.21	46	63.2	*														
24	11.95	86	83.9	*														
25	6.684	49	95.7	*														
26	1.418	18	100.0	*														*
27	-3.847	0	100.0	*														*
28	-9.112	0	100.0	*														*
29	-14.38	0	100.0	*														*
30	-19.64	0	100.0	*														*
31	-24.91	0	100.0	*														*
32	-30.17	0	100.0	*														*
33	-35.44	0	100.0	*														*
34	-40.70	0	100.0	*														*
35	-45.97	0	100.0	*														*
36	-51.23	0	100.0	*														*
37	-56.50	0	100.0	*														*
38	-61.76	0	100.0	*														*
39	-67.03	0	100.0	*														*
40	-72.30	0	100.0	*														*
41		0	100.0	*														*

T I T L E S	PROGRAM:	PERCENTAGE HISTOGRAMS
	DATA:	DAISY CREEK RECON STREAM SEDIMENT GEOCHEMISTRY
	RUN:	
	TIME:	25/03/83 11:55:33

NAME	CU
N	416
X BAR	1.3589
STD DEV	0.31414

CELL	LOWER LIMIT	NMBR	%	LOGARITHMIC VALUES	BAR INTERVAL = 0.25000	STD DEV	ARITH.LIMIT
1		0	0.0				
2	-0.1725	0	0.0				0.6722
3	-0.9397E-01	0	0.0				0.8054
4	-0.1543E-01	0	0.0				0.9651
5	0.6310E-01	0	0.0				1.156
6	0.1416	0	0.0				1.386
7	0.2202	0	0.0				1.660
8	0.2987	0	0.0				1.989
9	0.3772	0	0.0				2.384
10	0.4558	3	0.7	*			2.856
11	0.5343	4	1.0	*			3.422
12	0.6128	0	0.0				4.101
13	0.6914	7	1.7	**			4.913
14	0.7699	11	2.6	***			5.887
15	0.8485	9	2.2	**			7.054
16	0.9270	18	4.3	****			8.453
17	1.006	29	7.0	*****			10.13
18	1.084	37	8.9	*****			12.14
19	1.163	35	8.4	*****			14.54
20	1.241	32	7.7	*****			17.42
21	1.320	36	8.7	*****			20.88
22	1.398	33	7.9	*****			25.01
23	1.477	39	9.4	*****			29.97
24	1.555	36	8.7	*****			35.91
25	1.634	44	10.6	*****			43.03
26	1.712	19	4.6	*****			51.56
27	1.791	11	2.6	***			61.78
28	1.869	3	0.7	*			74.03
29	1.948	5	1.2	*			88.70
30	2.026	2	0.5				106.3
31	2.105	1	0.2				127.4
32	2.184	1	0.2				152.6
33	2.262	0	0.0				182.8
34	2.341	0	0.0				219.1
35	2.419	0	0.0				262.5
36	2.498	0	0.0				314.5
37	2.576	0	0.0				376.9
38	2.655	0	0.0				451.6
39	2.733	0	0.0				541.1
40	2.812	0	0.0				648.3
41	2.890	1	0.2				776.9

T I T L E S	PROGRAM:	PERCENTAGE CUMULATIVE FREQUENCY PLOTS	NAME	CU
	DATA:	DAISY CREEK RECON STREAM SEDIMENT GEOCHEMISTRY	N	416
	RUN:		X BAR	1.9589
	TIME:	25/03/83 11:55:33	STD DEV	0.31414

CELL	LOWER LIMIT	NMBR	%	LOGARITHMIC VALUES										BAR INTERVAL = 0.12500	STD DEV	ARITH. LIMIT		
1	2.125	3	0.7	*														133.2
2	2.085	1	1.0	*														121.7
3	2.046	1	1.2	*														111.2
4	2.007	0	1.2	*														101.6
5	1.968	2	1.7	*														92.80
6	1.928	3	2.4	*														84.78
7	1.889	1	2.6	*														77.45
8	1.850	3	3.4	*														70.75
9	1.810	5	4.6	*														64.64
10	1.771	9	6.7	*														59.05
11	1.732	10	9.1	*														53.95
12	1.693	11	11.8	*														49.28
13	1.653	23	17.3	*														45.02
14	1.614	19	21.9	*														41.13
15	1.575	17	26.0	*														37.57
16	1.536	20	30.8	*														34.33
17	1.496	24	36.5	*														31.36
18	1.457	18	40.9	*														28.65
19	1.418	18	45.2	*														26.17
20	1.379	21	50.2	*														23.91
21	1.339	14	53.6	*														21.84
22	1.300	26	59.9	*														19.95
23	1.261	7	61.5	*														18.23
24	1.221	18	65.9	*														16.65
25	1.182	11	68.5	*														15.21
26	1.143	32	76.2	*														13.90
27	1.104	18	80.5	*														12.70
28	1.064	14	83.9	*														11.60
29	1.025	15	87.5	*														10.60
30	0.9859	8	89.4	*														9.680
31	0.9466	10	91.8	*														8.843
32	0.9074	0	91.8	*														8.079
33	0.8681	9	94.0	*														7.380
34	0.8288	7	95.7	*														6.742
35	0.7896	0	95.7	*														6.160
36	0.7503	4	96.6	*														5.627
37	0.7110	0	96.6	*														5.141
38	0.6717	7	98.3	*														4.696
39	0.6325	0	98.3	*														4.290
40	0.5932	4	99.3	*														3.919
41		3	100.0	*														

T	PROGRAM:	PERCENTAGE HISTOGRAMS
I		
T		
L	DATA:	DAISY CREEK RECON STREAM SEDIMENT GEOCHEMISTRY
E	RUN:	
S	TIME:	25/03/83 11:55:33

NAME	PB
N	416
X BAR	21.130
STD DEV	11.680

CELL	LOWER LIMIT	NMBR	%	ARITHMETIC VALUES	BAR INTERVAL = 0.25000	STD DEV	ARITH.LIMIT
1		0	0.0				
2	-35.81	0	0.0				
3	-32.89	0	0.0				
4	-29.97	0	0.0				
5	-27.05	0	0.0				
6	-24.13	0	0.0				
7	-21.21	0	0.0				
8	-18.29	0	0.0				
9	-15.37	0	0.0				
10	-12.45	0	0.0				
11	-9.531	0	0.0				
12	-6.611	0	0.0				
13	-3.691	0	0.0				
14	-0.7708	0	0.0				
15	2.149	6	1.4	*			
16	5.069	11	2.6	***			
17	7.989	42	10.1	*****			
18	10.91	55	13.2	*****			
19	13.83	53	12.7	*****			
20	16.75	50	12.0	*****			
21	19.67	54	13.0	*****			
22	22.59	36	8.7	*****			
23	25.51	25	6.0	*****			
24	28.43	20	4.8	*****			
25	31.35	18	4.3	****			
26	34.27	10	2.4	**			
27	37.19	7	1.7	**			
28	40.11	7	1.7	**			
29	43.03	5	1.2	*			
30	45.95	4	1.0	*			
31	48.87	5	1.2	*			
32	51.79	1	0.2				
33	54.71	2	0.5				
34	57.63	0	0.0				
35	60.55	0	0.0				
36	63.47	2	0.5				
37	66.39	0	0.0				
38	69.31	2	0.5				
39	72.23	0	0.0				
40	75.15	0	0.0				
41	78.07	1	0.2				

T I T L E S	PROGRAM:	PERCENTAGE CUMULATIVE FREQUENCY PLOTS	NAME	PB
	DATA:	DAISY CREEK RECON STREAM SEDIMENT GEOCHEMISTRY	N	416
	RUN:		X BAR	21.130
	TIME:	25/03/83 11:55:33	STD DEV	11.680

CELL	LOWER LIMIT	NMBR	%	ARITHMETIC VALUES										BAR INTERVAL = 0.12500	STD DEV	ARITH.LIMIT		
1	49.60	12	2.9		*													
2	48.14	1	3.1		*													
3	46.68	1	3.4		*													
4	45.22	3	4.1		*													
5	43.76	5	5.3		*													
6	42.30	3	6.0		*													
7	40.84	4	7.0		*													
8	39.38	3	7.7		*													
9	37.92	4	8.7		*													
10	36.46	4	9.6		*													
11	35.00	4	10.6		*													
12	33.54	10	13.0		*													
13	32.08	4	13.9		*													
14	30.62	11	16.6		*													
15	29.16	4	17.5		*													
16	27.70	17	21.6		*													
17	26.24	9	23.8		*													
18	24.78	23	29.3		*													
19	23.32	14	32.7		*													
20	21.86	32	40.4		*													
21	20.40	14	43.8		*													
22	18.94	33	51.7		*													
23	17.48	16	55.5		*													
24	16.02	18	59.9		*													
25	14.56	30	67.1		*													
26	13.10	23	72.6		*													
27	11.64	27	79.1		*													
28	10.18	28	85.8		*													
29	8.719	28	92.5		*													
30	7.259	14	95.9		*													
31	5.799	11	98.6		*													
32	4.339	6	100.0		*												*	
33	2.879	0	100.0		*												*	
34	1.419	0	100.0		*												*	
35	-0.4087E-01	0	100.0		*												*	
36	-1.501	0	100.0		*												*	
37	-2.961	0	100.0		*												*	
38	-4.421	0	100.0		*												*	
39	-5.881	0	100.0		*												*	
40	-7.341	0	100.0		*												*	
41		0	100.0		*												*	

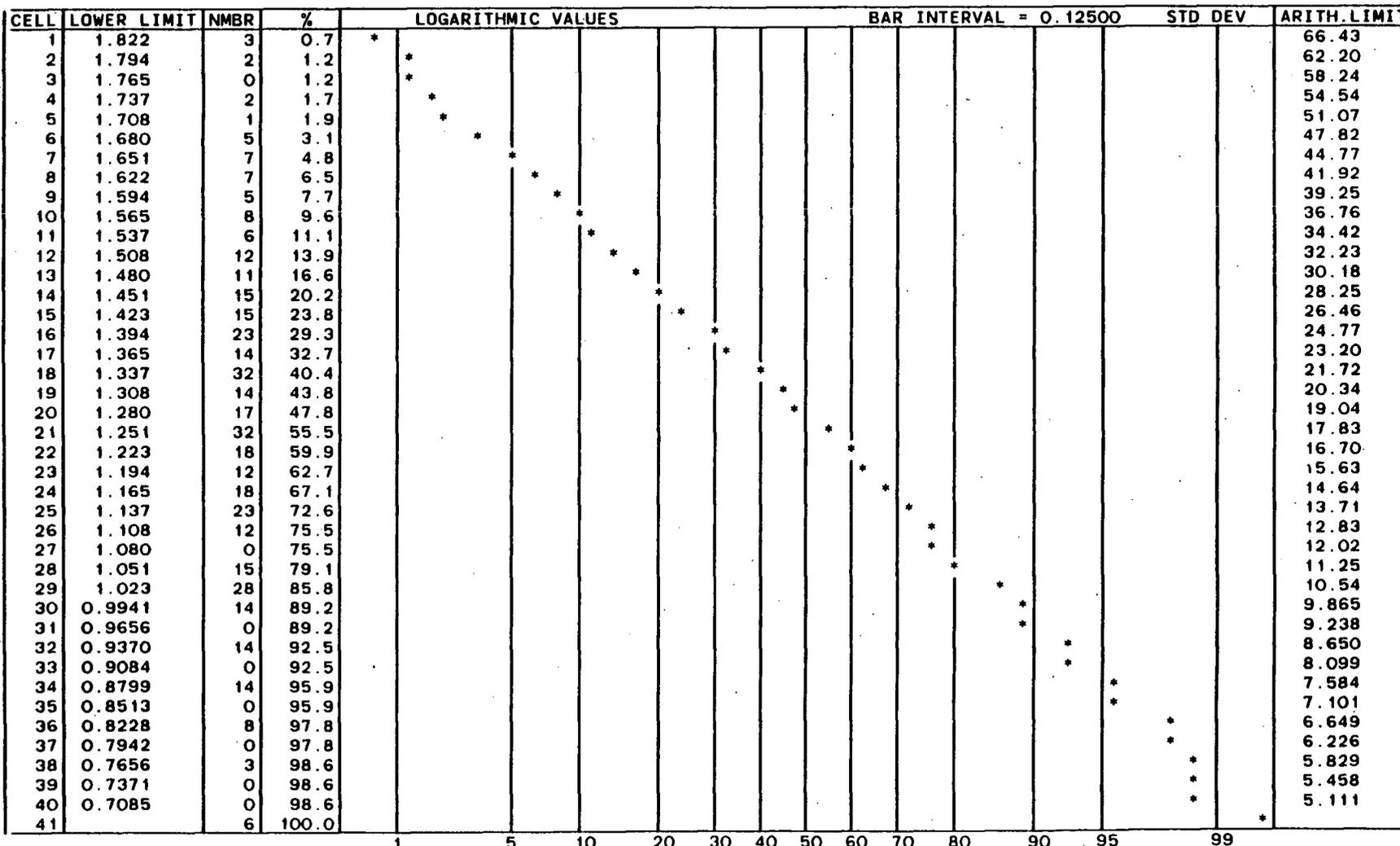
T I T L E S	PROGRAM:	PERCENTAGE HISTOGRAMS
	DATA:	DAISY CREEK RECON STREAM SEDIMENT GEOCHEMISTRY
	RUN:	
	TIME:	25/03/83 11:55:33

NAME	PB
N	416
X BAR	1.2655
STD DEV	0.22849

CELL	LOWER LIMIT	NMBR	%	LOGARITHMIC VALUES	BAR INTERVAL = 0.25000	STD DEV	ARITH. LIMIT
1		0	0.0				
2	0.1516	0	0.0				1.418
3	0.2087	0	0.0				1.617
4	0.2658	0	0.0				1.844
5	0.3229	0	0.0				2.103
6	0.3800	0	0.0				2.399
7	0.4372	0	0.0				2.736
8	0.4943	0	0.0				3.121
9	0.5514	0	0.0				3.560
10	0.6085	0	0.0				4.060
11	0.6657	6	1.4	*			4.631
12	0.7228	3	0.7	*			5.282
13	0.7799	0	0.0				6.024
14	0.8370	8	1.9	**			6.871
15	0.8942	14	3.4	***			7.837
16	0.9513	28	6.7	*****			8.939
17	1.008	28	6.7	*****			10.20
18	1.066	27	6.5	*****			11.63
19	1.123	41	9.9	*****			13.26
20	1.180	30	7.2	*****			15.13
21	1.237	32	7.7	*****			17.25
22	1.294	54	13.0	*****			19.68
23	1.351	36	8.7	*****			22.45
24	1.408	36	8.7	*****			25.60
25	1.465	19	4.6	*****			29.20
26	1.523	18	4.3	****			33.30
27	1.580	14	3.4	***			37.99
28	1.637	10	2.4	**			43.33
29	1.694	7	1.7	**			49.42
30	1.751	0	0.0				56.36
31	1.808	4	1.0	*			64.29
32	1.865	0	0.0				73.32
33	1.922	1	0.2				83.63
34	1.979	0	0.0				95.38
35	2.037	0	0.0				108.8
36	2.094	0	0.0				124.1
37	2.151	0	0.0				141.5
38	2.208	0	0.0				161.4
39	2.265	0	0.0				184.1
40	2.322	0	0.0				210.0
41	2.379	0	0.0				239.5

T I T L E S	PROGRAM:	PERCENTAGE CUMULATIVE FREQUENCY PLOTS
	DATA:	DAISY CREEK RECON STREAM SEDIMENT GEOCHEMISTRY
	RUN:	
	TIME:	25/03/83 11:55:33

NAME	PB
N	416
X BAR	1.2655
STD DEV	0.22849



T I T L E S	PROGRAM:	PERCENTAGE HISTOGRAMS
	DATA:	DAISY CREEK RECON STREAM SEDIMENT GEOCHEMISTRY
	RUN:	
	TIME:	25/03/83 11:55:33

NAME	ZN
N	416
X BAR	75.305
STD DEV	35.611

CELL	LOWER LIMIT	NMBR	%	ARITHMETIC VALUES	BAR INTERVAL = 0.25000	STD DEV	ARITH.LIMIT
1		0	0.0				
2	-98.30	0	0.0				
3	-89.39	0	0.0				
4	-80.49	0	0.0				
5	-71.59	0	0.0				
6	-62.69	0	0.0				
7	-53.78	0	0.0				
8	-44.88	0	0.0				
9	-35.98	0	0.0				
10	-27.08	0	0.0				
11	-18.17	0	0.0				
12	-9.270	0	0.0				
13	-0.3673	0	0.0				
14	8.535	0	0.0				
15	17.44	6	1.4	*			
16	26.34	9	2.2	**			
17	35.24	28	6.7	*****			
18	44.15	47	11.3	*****			
19	53.05	55	13.2	*****			
20	61.95	74	17.8	*****			
21	70.85	63	15.1	*****			
22	79.76	43	10.3	*****			
23	88.66	25	6.0	*****			
24	97.56	21	5.0	*****			
25	106.5	9	2.2	**			
26	115.4	5	1.2	*			
27	124.3	11	2.6	***			
28	133.2	5	1.2	*			
29	142.1	2	0.5				
30	151.0	3	0.7	*			
31	159.9	2	0.5				
32	168.8	1	0.2				
33	177.7	3	0.7	*			
34	186.6	1	0.2				
35	195.5	0	0.0				
36	204.4	0	0.0				
37	213.3	0	0.0				
38	222.2	1	0.2				
39	231.1	0	0.0				
40	240.0	0	0.0				
41	248.9	2	0.5				

T
I PROGRAM: PERCENTAGE CUMULATIVE FREQUENCY PLOTS
L
E DATA: DAISY CREEK RECON STREAM SEDIMENT GEOCHEMISTRY
S RUN:
 TIME: 25/03/83 11:55:33

NAME	ZN
N	416
X BAR	75.305
STD DEV	35.611

CELL	LOWER LIMIT	NMBR	%	ARITHMETIC VALUES										BAR INTERVAL = 0.12500	STD DEV	ARITH.LIMIT			
1	162.1	10	2.4		*														
2	157.7	2	2.9		*														
3	153.2	0	2.9		*														
4	148.8	1	3.1		*														
5	144.3	1	3.4		*														
6	139.8	1	3.6		*														
7	135.4	4	4.6		*														
8	130.9	3	5.3		*														
9	126.5	6	6.7		*	*													
10	122.0	4	7.7		*	*													
11	117.6	3	8.4		*	*	*												
12	113.1	2	8.9		*	*	*												
13	108.7	4	9.9		*	*	*	*											
14	104.2	8	11.8		*	*	*	*	*										
15	99.79	11	14.4		*	*	*	*	*	*									
16	95.34	11	17.1		*	*	*	*	*	*	*								
17	90.88	10	19.5		*	*	*	*	*	*	*	*							
18	86.43	16	23.3		*	*	*	*	*	*	*	*	*						
19	81.98	26	29.6		*	*	*	*	*	*	*	*	*	*					
20	77.53	20	34.4		*	*	*	*	*	*	*	*	*	*					
21	73.08	26	40.6		*	*	*	*	*	*	*	*	*	*	*				
22	68.63	43	51.0		*	*	*	*	*	*	*	*	*	*	*	*			
23	64.18	36	59.6		*	*	*	*	*	*	*	*	*	*	*	*			
24	59.73	37	68.5		*	*	*	*	*	*	*	*	*	*	*	*	*		
25	55.27	31	76.0		*	*	*	*	*	*	*	*	*	*	*	*	*	*	
26	50.82	30	83.2		*	*	*	*	*	*	*	*	*	*	*	*	*	*	*
27	46.37	19	87.7		*	*	*	*	*	*	*	*	*	*	*	*	*	*	*
28	41.92	24	93.5		*	*	*	*	*	*	*	*	*	*	*	*	*	*	*
29	37.47	10	95.9		*	*	*	*	*	*	*	*	*	*	*	*	*	*	*
30	33.02	6	97.4		*	*	*	*	*	*	*	*	*	*	*	*	*	*	*
31	28.57	5	98.6		*	*	*	*	*	*	*	*	*	*	*	*	*	*	*
32	24.11	0	98.6		*	*	*	*	*	*	*	*	*	*	*	*	*	*	*
33	19.66	5	99.8		*	*	*	*	*	*	*	*	*	*	*	*	*	*	*
34	15.21	1	100.0		*	*	*	*	*	*	*	*	*	*	*	*	*	*	*
35	10.76	0	100.0		*	*	*	*	*	*	*	*	*	*	*	*	*	*	*
36	6.310	0	100.0		*	*	*	*	*	*	*	*	*	*	*	*	*	*	*
37	1.858	0	100.0		*	*	*	*	*	*	*	*	*	*	*	*	*	*	*
38	-2.593	0	100.0		*	*	*	*	*	*	*	*	*	*	*	*	*	*	*
39	-7.044	0	100.0		*	*	*	*	*	*	*	*	*	*	*	*	*	*	*
40	-11.50	0	100.0		*	*	*	*	*	*	*	*	*	*	*	*	*	*	*
41		0	100.0		*	*	*	*	*	*	*	*	*	*	*	*	*	*	*

196

T I T L E S	PROGRAM: PERCENTAGE HISTOGRAMS
	DATA: DAISY CREEK RECON STREAM SEDIMENT GEOCHEMISTRY
	RUN:
	TIME: 25/03/83 11:55:33

NAME	ZN
N	416
X BAR	1.8422
STD DEV	0.16888

CELL	LOWER LIMIT	NMBR	%	LOGARITHMIC VALUES	BAR INTERVAL = 0.25000	STD DEV	ARITH. LIMIT
1		0	0.0				
2	1.019	0	0.0				10.45
3	1.061	0	0.0				11.51
4	1.103	0	0.0				12.69
5	1.146	0	0.0				13.98
6	1.188	0	0.0				15.41
7	1.230	0	0.0				16.98
8	1.272	2	0.5				18.72
9	1.314	3	0.7	*			20.63
10	1.357	1	0.2				22.73
11	1.399	0	0.0				25.06
12	1.441	2	0.5				27.61
13	1.483	3	0.7	*			30.43
14	1.526	5	1.2	*			33.54
15	1.568	10	2.4	**			36.96
16	1.610	17	4.1	****			40.74
17	1.652	23	5.5	*****			44.90
18	1.694	26	6.3	*****			49.48
19	1.737	44	10.6	*****			54.53
20	1.779	51	12.3	*****			60.10
21	1.821	52	12.5	*****			66.23
22	1.863	46	11.1	*****			73.00
23	1.906	40	9.6	*****			80.45
24	1.948	25	6.0	*****			88.66
25	1.990	23	5.5	*****			97.71
26	2.032	8	1.9	**			107.7
27	2.074	13	3.1	***			118.7
28	2.117	8	1.9	**			130.8
29	2.159	3	0.7	*			144.2
30	2.201	3	0.7	*			158.9
31	2.243	5	1.2	*			175.1
32	2.285	0	0.0				193.0
33	2.328	1	0.2				212.7
34	2.370	0	0.0				234.4
35	2.412	0	0.0				258.3
36	2.454	0	0.0				284.7
37	2.497	1	0.2				313.7
38	2.539	0	0.0				345.8
39	2.581	0	0.0				381.1
40	2.623	1	0.2				420.0
41	2.665	0	0.0				462.9

T I T L	PROGRAM: PERCENTAGE CUMULATIVE FREQUENCY PLOTS
E	DATA: DAISY CREEK RECON STREAM SEDIMENT GEOCHEMISTRY
S	RUN:
	TIME: 25/03/83 11:55:33

NAME	ZN
N	416
X BAR	1.8422
STD DEV	0.16888

CELL	LOWER LIMIT	NMBR	%	LOGARITHMIC VALUES	BAR INTERVAL = 0.12500	STD DEV	ARITH. LIMIT
1	2.254	7	1.7	*			179.4
2	2.233	1	1.9	*			170.9
3	2.212	2	2.4	*			162.8
4	2.190	2	2.9	*			155.1
5	2.169	1	3.1	*			147.7
6	2.148	2	3.6	*			140.7
7	2.127	5	4.8	*			134.0
8	2.106	7	6.5	*			127.7
9	2.085	6	7.9	*			121.6
10	2.064	3	8.7	*			115.8
11	2.043	1	8.9	*			110.3
12	2.022	11	11.5	*			105.1
13	2.001	8	13.5	*			100.1
14	1.979	15	17.1	*			95.37
15	1.958	10	19.5	*			90.84
16	1.937	16	23.3	*			86.53
17	1.916	18	27.6	*			82.43
18	1.895	23	33.2	*			78.52
19	1.874	22	38.5	*			74.79
20	1.853	27	45.0	*			71.24
21	1.832	36	53.6	*			67.86
22	1.811	25	59.6	*			64.64
23	1.789	23	65.1	*			61.58
24	1.768	23	70.7	*			58.66
25	1.747	22	76.0	*			55.87
26	1.726	10	78.4	*			53.22
27	1.705	20	83.2	*			50.70
28	1.684	11	85.8	*			48.29
29	1.663	8	87.7	*			46.00
30	1.642	15	91.3	*			43.82
31	1.621	9	93.5	*			41.74
32	1.599	3	94.2	*			39.76
33	1.578	7	95.9	*			37.87
34	1.557	1	96.2	*			36.08
35	1.536	5	97.4	*			34.36
36	1.515	0	97.4	*			32.73
37	1.494	3	98.1	*			31.18
38	1.473	1	98.3	*			29.70
39	1.452	1	98.6	*			28.29
40	1.431	0	98.6	*			26.95
41		6	100.0	*			

T I T L E S	PROGRAM: PERCENTAGE HISTOGRAMS
	DATA: DAISY CREEK RECON STREAM SEDIMENT GEOCHEMISTRY
	RUN:
	TIME: 25/03/83 11:55:33

NAME	MO
N	416
X BAR	1.7752
STD DEV	0.83628

CELL	LOWER LIMIT	NMBR	%	ARITHMETIC VALUES	BAR INTERVAL = 0.25000	STD DEV	ARITH.LIMIT
1		0	0.0				
2	-2.302	0	0.0				
3	-2.093	0	0.0				
4	-1.883	0	0.0				
5	-1.674	0	0.0				
6	-1.465	0	0.0				
7	-1.256	0	0.0				
8	-1.047	0	0.0				
9	-0.8381	0	0.0				
10	-0.6291	0	0.0				
11	-0.4200	0	0.0				
12	-0.2109	0	0.0				
13	-0.1847E-02	0	0.0				
14	0.2072	0	0.0				
15	0.4163	29	7.0	*****			
16	0.6254	0	0.0				
17	0.8344	132	31.7	*****			
18	1.043	0	0.0				
19	1.253	0	0.0				
20	1.462	0	0.0				
21	1.671	0	0.0				
22	1.880	184	44.2	*****			
23	2.089	0	0.0				
24	2.298	0	0.0				
25	2.507	0	0.0				
26	2.716	0	0.0				
27	2.925	64	15.4	*****			
28	3.134	0	0.0				
29	3.343	0	0.0				
30	3.552	0	0.0				
31	3.761	0	0.0				
32	3.970	4	1.0	*			
33	4.180	0	0.0				
34	4.389	0	0.0				
35	4.598	0	0.0				
36	4.807	2	0.5				
37	5.016	0	0.0				
38	5.225	0	0.0				
39	5.434	0	0.0				
40	5.643	0	0.0				
41	5.852	1	0.2				

T I T L E S	PROGRAM:	PERCENTAGE CUMULATIVE FREQUENCY PLOTS
	DATA:	DAISY CREEK RECON STREAM SEDIMENT GEOCHEMISTRY
	RUN:	
	TIME:	25/03/83 11:55:33

NAME	MO
N	416
X BAR	1.7752
STD DEV	0.83628

CELL	LOWER LIMIT	NMBR	%	ARITHMETIC VALUES										BAR INTERVAL = 0.12500	STD DEV	ARITH.LIMIT		
1	3.814	7	1.7	*														
2	3.709	0	1.7	*														
3	3.605	0	1.7	*														
4	3.500	0	1.7	*														
5	3.396	0	1.7	*														
6	3.291	0	1.7	*														
7	3.186	0	1.7	*														
8	3.082	0	1.7	*														
9	2.977	64	17.1	*														
10	2.873	0	17.1	*														
11	2.768	0	17.1	*														
12	2.664	0	17.1	*														
13	2.559	0	17.1	*														
14	2.455	0	17.1	*														
15	2.350	0	17.1	*														
16	2.246	0	17.1	*														
17	2.141	0	17.1	*														
18	2.037	0	17.1	*														
19	1.932	184	61.3	*														
20	1.828	0	61.3	*														
21	1.723	0	61.3	*														
22	1.618	0	61.3	*														
23	1.514	0	61.3	*														
24	1.409	0	61.3	*														
25	1.305	0	61.3	*														
26	1.200	0	61.3	*														
27	1.096	0	61.3	*														
28	0.9912	132	93.0	*														
29	0.8867	0	93.0	*														
30	0.7822	0	93.0	*														
31	0.6776	0	93.0	*														
32	0.5731	0	93.0	*														
33	0.4686	29	100.0	*														
34	0.3640	0	100.0	*														
35	0.2595	0	100.0	*														
36	0.1550	0	100.0	*														
37	0.5042E-01	0	100.0	*														
38	-0.5412E-01	0	100.0	*														
39	-0.1587	0	100.0	*														
40	-0.2632	0	100.0	*														
41		0	100.0	*														

T I T L E S	PROGRAM: PERCENTAGE HISTOGRAMS
	DATA: DAISY CREEK RECON STREAM SEDIMENT GEOCHEMISTRY
	RUN:
	TIME: 25/03/83 11:55:33

NAME	MD
N	416
X BAR	0.19658
STD DEV	0.22402

CELL	LOWER LIMIT	NMBR	%	LOGARITHMIC VALUES	BAR INTERVAL = 0.25000	STD DEV	ARITH. LIMIT
1		0	0.0				
2	-0.8955	0	0.0				0.1272
3	-0.8395	0	0.0				0.1447
4	-0.7835	0	0.0				0.1646
5	-0.7275	0	0.0				0.1873
6	-0.6715	0	0.0				0.2131
7	-0.6155	0	0.0				0.2424
8	-0.5595	0	0.0				0.2758
9	-0.5035	0	0.0				0.3137
10	-0.4475	0	0.0				0.3569
11	-0.3915	0	0.0				0.4060
12	-0.3355	29	7.0	*****			0.4619
13	-0.2794	0	0.0				0.5255
14	-0.2234	0	0.0				0.5978
15	-0.1674	0	0.0				0.6801
16	-0.1114	0	0.0				0.7737
17	-0.5543E-01	132	31.7	*****			0.8802
18	0.5705E-03	0	0.0				1.001
19	0.5657E-01	0	0.0				1.139
20	0.1126	0	0.0				1.296
21	0.1686	0	0.0				1.474
22	0.2246	0	0.0				1.677
23	0.2806	184	44.2	*****			1.908
24	0.3366	0	0.0				2.171
25	0.3926	0	0.0				2.469
26	0.4486	64	15.4	*****			2.809
27	0.5046	0	0.0				3.196
28	0.5606	4	1.0	*			3.636
29	0.6166	0	0.0				4.136
30	0.6726	2	0.5				4.706
31	0.7286	1	0.2				5.353
32	0.7846	0	0.0				6.090
33	0.8406	0	0.0				6.928
34	0.8966	0	0.0				7.882
35	0.9526	0	0.0				8.967
36	1.009	0	0.0				10.20
37	1.065	0	0.0				11.60
38	1.121	0	0.0				13.20
39	1.177	0	0.0				15.02
40	1.233	0	0.0				17.09
41	1.289	0	0.0				19.44

- 161 -

T I T L E S	PROGRAM:	PERCENTAGE CUMULATIVE FREQUENCY PLOTS	NAME	MO
	DATA:	DAISY CREEK RECON STREAM SEDIMENT GEOCHEMISTRY	N	416
	RUN:		X BAR	0.19658
	TIME:	25/03/83 11:55:33	STD DEV	0.22402

CELL	LOWER LIMIT	NMBR	%	LOGARITHMIC VALUES										BAR INTERVAL = 0.12500	STD DEV	ARITH. LIMIT	
1	0.7426	1	0.2	*													5.529
2	0.7146	0	0.2	*													5.183
3	0.6866	2	0.7	*													4.860
4	0.6586	0	0.7	*													4.556
5	0.6306	0	0.7	*													4.272
6	0.6026	0	0.7	*													4.005
7	0.5746	4	1.7	*													3.755
8	0.5466	0	1.7	*													3.521
9	0.5186	0	1.7	*													3.301
10	0.4906	0	1.7	*													3.095
11	0.4626	64	17.1				*										2.901
12	0.4346	0	17.1				*										2.720
13	0.4066	0	17.1				*										2.550
14	0.3786	0	17.1				*										2.391
15	0.3506	0	17.1				*										2.242
16	0.3226	0	17.1				*										2.102
17	0.2946	184	61.3						*								1.971
18	0.2666	0	61.3						*								1.848
19	0.2386	0	61.3						*								1.732
20	0.2106	0	61.3						*								1.624
21	0.1826	0	61.3						*								1.523
22	0.1546	0	61.3						*								1.428
23	0.1266	0	61.3						*								1.338
24	0.9858E-01	0	61.3						*								1.255
25	0.7057E-01	0	61.3						*								1.176
26	0.4257E-01	0	61.3						*								1.103
27	0.1457E-01	0	61.3						*								1.034
28	-0.1343E-01	132	93.0									*					0.9695
29	-0.4143E-01	0	93.0									*					0.9090
30	-0.6943E-01	0	93.0									*					0.8522
31	-0.9744E-01	0	93.0									*					0.7990
32	-0.1254	0	93.0									*					0.7491
33	-0.1534	0	93.0									*					0.7024
34	-0.1814	0	93.0									*					0.6585
35	-0.2094	0	93.0									*					0.6174
36	-0.2374	0	93.0									*					0.5788
37	-0.2654	0	93.0									*					0.5427
38	-0.2934	0	93.0									*					0.5088
39	-0.3215	29	100.0											*			0.4770
40	-0.3495	0	100.0											*			0.4472
41		0	100.0											*			

T I T L E S	PROGRAM:	PERCENTAGE HISTOGRAMS	NAME	AG
	DATA:	DAISY CREEK RECON STREAM SEDIMENT GEOCHEMISTRY	N	416
	RUN:		X BAR	0.24182
	TIME:	25/03/83 11:55:33	STD DEV	0.93040E-01

CELL	LOWER LIMIT	NMBR	%	ARITHMETIC VALUES	BAR INTERVAL = 0.25000	STD DEV	ARITH.LIMIT
1		0	0.0				
2	-0.2117	0	0.0				
3	-0.1885	0	0.0				
4	-0.1652	0	0.0				
5	-0.1420	0	0.0				
6	-0.1187	0	0.0				
7	-0.9545E-01	0	0.0				
8	-0.7219E-01	0	0.0				
9	-0.4893E-01	0	0.0				
10	-0.2567E-01	0	0.0				
11	-0.2406E-02	0	0.0				
12	0.2085E-01	0	0.0				
13	0.4411E-01	0	0.0				
14	0.6737E-01	0	0.0				
15	0.9063E-01	0	0.0				
16	0.1139	0	0.0				
17	0.1372	0	0.0				
18	0.1604	0	0.0				
19	0.1837	317	76.2	*****			
20	0.2069	0	0.0				
21	0.2302	0	0.0				
22	0.2535	0	0.0				
23	0.2767	0	0.0				
24	0.3000	52	12.5	*****			
25	0.3232	0	0.0				
26	0.3465	0	0.0				
27	0.3698	0	0.0				
28	0.3930	34	8.2	*****			
29	0.4163	0	0.0				
30	0.4395	0	0.0				
31	0.4628	0	0.0				
32	0.4861	3	0.7	*			
33	0.5093	0	0.0				
34	0.5326	0	0.0				
35	0.5558	0	0.0				
36	0.5791	8	1.9	**			
37	0.6024	0	0.0				
38	0.6256	0	0.0				
39	0.6489	0	0.0				
40	0.6721	0	0.0				
41	0.6954	2	0.5				

T I T L E S	PROGRAM:	PERCENTAGE CUMULATIVE FREQUENCY PLOTS
	DATA:	DAISY CREEK RECON STREAM SEDIMENT GEOCHEMISTRY
	RUN:	
	TIME:	25/03/83 11:55:33

NAME	AG
N	416
X BAR	0.24182
STD DEV	0.93040E-01

CELL	LOWER LIMIT	NMBR	%	ARITHMETIC VALUES										BAR INTERVAL = 0.12500	STD DEV	ARITH.LIMIT		
1	0.4686	13	3.1	*														
2	0.4570	0	3.1	*														
3	0.4453	0	3.1	*														
4	0.4337	0	3.1	*														
5	0.4221	0	3.1	*														
6	0.4105	0	3.1	*														
7	0.3988	34	11.3	*	*													
8	0.3872	0	11.3	*	*													
9	0.3756	0	11.3	*	*													
10	0.3639	0	11.3	*	*													
11	0.3523	0	11.3	*	*													
12	0.3407	0	11.3	*	*													
13	0.3290	0	11.3	*	*													
14	0.3174	0	11.3	*	*													
15	0.3058	0	11.3	*	*													
16	0.2942	52	23.8	*	*	*												
17	0.2825	0	23.8	*	*	*												
18	0.2709	0	23.8	*	*	*												
19	0.2593	0	23.8	*	*	*												
20	0.2476	0	23.8	*	*	*												
21	0.2360	0	23.8	*	*	*												
22	0.2244	0	23.8	*	*	*												
23	0.2127	0	23.8	*	*	*												
24	0.2011	0	23.8	*	*	*												
25	0.1895	317	100.0	*	*	*	*	*	*	*	*	*	*	*	*	*	*	*
26	0.1779	0	100.0	*	*	*	*	*	*	*	*	*	*	*	*	*	*	*
27	0.1662	0	100.0	*	*	*	*	*	*	*	*	*	*	*	*	*	*	*
28	0.1546	0	100.0	*	*	*	*	*	*	*	*	*	*	*	*	*	*	*
29	0.1430	0	100.0	*	*	*	*	*	*	*	*	*	*	*	*	*	*	*
30	0.1313	0	100.0	*	*	*	*	*	*	*	*	*	*	*	*	*	*	*
31	0.1197	0	100.0	*	*	*	*	*	*	*	*	*	*	*	*	*	*	*
32	0.1081	0	100.0	*	*	*	*	*	*	*	*	*	*	*	*	*	*	*
33	0.9645E-01	0	100.0	*	*	*	*	*	*	*	*	*	*	*	*	*	*	*
34	0.8482E-01	0	100.0	*	*	*	*	*	*	*	*	*	*	*	*	*	*	*
35	0.7319E-01	0	100.0	*	*	*	*	*	*	*	*	*	*	*	*	*	*	*
36	0.6156E-01	0	100.0	*	*	*	*	*	*	*	*	*	*	*	*	*	*	*
37	0.4993E-01	0	100.0	*	*	*	*	*	*	*	*	*	*	*	*	*	*	*
38	0.3830E-01	0	100.0	*	*	*	*	*	*	*	*	*	*	*	*	*	*	*
39	0.2667E-01	0	100.0	*	*	*	*	*	*	*	*	*	*	*	*	*	*	*
40	0.1504E-01	0	100.0	*	*	*	*	*	*	*	*	*	*	*	*	*	*	*
41		0	100.0	*	*	*	*	*	*	*	*	*	*	*	*	*	*	*

T I T L E S	PROGRAM:	PERCENTAGE HISTOGRAMS
	DATA:	DAISY CREEK RECON STREAM SEDIMENT GEOCHEMISTRY
	RUN:	
	TIME:	25/03/83 11:55:33

NAME	AG
N	416
X BAR	-0.63730
STD DEV	0.12226

CELL	LOWER LIMIT	NMBR	%	LOGARITHMIC VALUES	BAR INTERVAL = 0.25000	STD DEV	ARITH. LIMIT
1		0	0.0				
2	-1.233	0	0.0				0.5844E-01
3	-1.203	0	0.0				0.6270E-01
4	-1.172	0	0.0				0.6727E-01
5	-1.142	0	0.0				0.7218E-01
6	-1.111	0	0.0				0.7744E-01
7	-1.080	0	0.0				0.8308E-01
8	-1.050	0	0.0				0.8914E-01
9	-1.019	0	0.0				0.9564E-01
10	-0.9888	0	0.0				0.1026
11	-0.9582	0	0.0				0.1101
12	-0.9277	0	0.0				0.1181
13	-0.8971	0	0.0				0.1267
14	-0.8665	0	0.0				0.1360
15	-0.8360	0	0.0				0.1459
16	-0.8054	0	0.0				0.1565
17	-0.7748	0	0.0				0.1679
18	-0.7443	0	0.0				0.1802
19	-0.7137	317	76.2	*****			0.1933
20	-0.6831	0	0.0				0.2074
21	-0.6526	0	0.0				0.2225
22	-0.6220	0	0.0				0.2388
23	-0.5914	0	0.0				0.2562
24	-0.5609	0	0.0				0.2749
25	-0.5303	52	12.5	*****			0.2949
26	-0.4998	0	0.0				0.3164
27	-0.4692	0	0.0				0.3395
28	-0.4386	0	0.0				0.3642
29	-0.4081	34	8.2	*****			0.3908
30	-0.3775	0	0.0				0.4193
31	-0.3469	0	0.0				0.4499
32	-0.3164	3	0.7	*			0.4827
33	-0.2858	0	0.0				0.5178
34	-0.2552	0	0.0				0.5556
35	-0.2247	8	1.9	**			0.5961
36	-0.1941	0	0.0				0.6396
37	-0.1635	1	0.2				0.6862
38	-0.1330	0	0.0				0.7362
39	-0.1024	0	0.0				0.7899
40	-0.7185E-01	0	0.0				0.8475
41	-0.4129E-01	1	0.2				C.9093

T I T L E S	PROGRAM: PERCENTAGE CUMULATIVE FREQUENCY PLOTS
	DATA: DAISY CREEK RECON STREAM SEDIMENT GEOCHEMISTRY
	RUN:
	TIME: 25/03/83 11:55:33

NAME	AG
N	416
X BAR	-0.63730
STD DEV	0.12226

CELL	LOWER LIMIT	NMBR	%	LOGARITHMIC VALUES										BAR INTERVAL = 0.12500	STD DEV	ARITH. LIMIT	
1	-0.3393	13	3.1	*													0.4578
2	-0.3546	0	3.1	*													0.4420
3	-0.3699	0	3.1	*													0.4267
4	-0.3851	0	3.1	*													0.4120
5	-0.4004	34	11.3	*													0.3977
6	-0.4157	0	11.3	*													0.3840
7	-0.4310	0	11.3	*													0.3707
8	-0.4463	0	11.3	*													0.3579
9	-0.4616	0	11.3	*													0.3455
10	-0.4768	0	11.3	*													0.3336
11	-0.4921	0	11.3	*													0.3220
12	-0.5074	0	11.3	*													0.3109
13	-0.5227	0	11.3	*													0.3001
14	-0.5380	52	23.8	*				*									0.2898
15	-0.5532	0	23.8	*				*									0.2797
16	-0.5685	0	23.8	*				*									0.2701
17	-0.5838	0	23.8	*				*									0.2607
18	-0.5991	0	23.8	*				*									0.2517
19	-0.6144	0	23.8	*				*									0.2430
20	-0.6297	0	23.8	*				*									0.2346
21	-0.6449	0	23.8	*				*									0.2265
22	-0.6602	0	23.8	*				*									0.2187
23	-0.6755	0	23.8	*				*									0.2111
24	-0.6908	0	23.8	*				*									0.2038
25	-0.7061	317	100.0	*				*								*	0.1968
26	-0.7214	0	100.0	*				*								*	0.1900
27	-0.7366	0	100.0	*				*								*	0.1834
28	-0.7519	0	100.0	*				*								*	0.1770
29	-0.7672	0	100.0	*				*								*	0.1709
30	-0.7825	0	100.0	*				*								*	0.1650
31	-0.7978	0	100.0	*				*								*	0.1593
32	-0.8130	0	100.0	*				*								*	0.1538
33	-0.8283	0	100.0	*				*								*	0.1485
34	-0.8436	0	100.0	*				*								*	0.1433
35	-0.8589	0	100.0	*				*								*	0.1384
36	-0.8742	0	100.0	*				*								*	0.1336
37	-0.8895	0	100.0	*				*								*	0.1290
38	-0.9047	0	100.0	*				*								*	0.1245
39	-0.9200	0	100.0	*				*								*	0.1202
40	-0.9353	0	100.0	*				*								*	0.1161
41		0	100.0	*				*								*	

- 196 -

T I T L E S	PROGRAM:	PERCENTAGE HISTOGRAMS
	DATA:	DAISY CREEK RECON STREAM SEDIMENT GEOCHEMISTRY
	RUN:	
	TIME:	25/03/83 11:55:33

NAME	CO
N	416
X BAR	7.2188
STD DEV	2.8518

CELL	LOWER LIMIT	NMBR	%	ARITHMETIC VALUES	BAR INTERVAL = 0.25000	STD DEV	ARITH.LIMIT
1		0	0.0				
2	-6.684	0	0.0				
3	-5.971	0	0.0				
4	-5.258	0	0.0				
5	-4.545	0	0.0				
6	-3.832	0	0.0				
7	-3.119	0	0.0				
8	-2.406	0	0.0				
9	-1.693	0	0.0				
10	-0.9802	0	0.0				
11	-0.2672	0	0.0				
12	0.4457	3	0.7	*			
13	1.159	0	0.0				
14	1.872	9	2.2	**			
15	2.585	18	4.3	****			
16	3.298	31	7.5	*****			
17	4.010	0	0.0				
18	4.723	49	11.8	*****			
19	5.436	53	12.7	*****			
20	6.149	0	0.0				
21	6.862	77	18.5	*****			
22	7.575	55	13.2	*****			
23	8.288	56	13.5	*****			
24	9.001	0	0.0				
25	9.714	29	7.0	*****			
26	10.43	14	3.4	***			
27	11.14	0	0.0				
28	11.85	4	1.0	*			
29	12.57	8	1.9	**			
30	13.28	0	0.0				
31	13.99	5	1.2	*			
32	14.70	3	0.7	*			
33	15.42	1	0.2				
34	16.13	0	0.0				
35	16.84	0	0.0				
36	17.56	0	0.0				
37	18.27	0	0.0				
38	18.98	0	0.0				
39	19.70	0	0.0				
40	20.41	0	0.0				
41	21.12	1	0.2				

T I T L E S	PROGRAM:	PERCENTAGE CUMULATIVE FREQUENCY PLOTS	NAME	CO
	DATA:	DAISY CREEK RECON STREAM SEDIMENT GEOCHEMISTRY	N	416
	RUN:		X BAR	7.2188
	TIME:	25/03/83 11:55:33	STD DEV	2.8518

CELL	LOWER LIMIT	NMBR	%	ARITHMETIC VALUES											BAR INTERVAL = 0.12500	STD DEV	ARITH.LIMIT	
1	14.17	5	1.2	*														
2	13.81	5	2.4	*														
3	13.46	0	2.4	*														
4	13.10	0	2.4	*														
5	12.74	8	4.3		*													
6	12.39	0	4.3		*													
7	12.03	0	4.3		*													
8	11.67	4	5.3		*													
9	11.32	0	5.3		*													
10	10.96	14	8.7			*												
11	10.61	0	8.7			*												
12	10.25	0	8.7			*												
13	9.892	29	15.6				*											
14	9.536	0	15.6				*											
15	9.179	0	15.6				*											
16	8.823	56	29.1					*										
17	8.466	0	29.1					*										
18	8.110	0	29.1					*										
19	7.753	55	42.3						*									
20	7.397	0	42.3						*									
21	7.041	0	42.3						*									
22	6.684	77	60.8							*								
23	6.328	0	60.8							*								
24	5.971	53	73.6								*							
25	5.615	0	73.6								*							
26	5.258	0	73.6								*							
27	4.902	49	85.3									*						
28	4.545	0	85.3									*						
29	4.189	0	85.3									*						
30	3.832	31	92.8										*					
31	3.476	0	92.8										*					
32	3.119	0	92.8										*					
33	2.763	18	97.1											*				
34	2.406	0	97.1											*				
35	2.050	0	97.1											*				
36	1.693	9	99.3												*			
37	1.337	0	99.3												*			
38	0.9804	3	100.0													*		
39	0.6240	0	100.0													*		
40	0.2675	0	100.0													*		
41		0	100.0													*		

T I T L E S	PROGRAM:	PERCENTAGE HISTOGRAMS	NAME	CO
	DATA:	DAISY CREEK RECON STREAM SEDIMENT GEOCHEMISTRY	N	416
	RUN:		X BAR	0.82302
	TIME:	25/03/83 11:55:33	STD DEV	0.18623

CELL	LOWER LIMIT	NMBR	%	LOGARITHMIC VALUES	BAR INTERVAL = 0.25000	STD DEV	ARITH.LIMIT
1		0	0.0				
2	-0.8485E-01	0	0.0				0.8225
3	-0.3830E-01	3	0.7 *				0.9156
4	0.8262E-02	0	0.0				1.019
5	0.5482E-01	0	0.0				1.135
6	0.1014	0	0.0				1.263
7	0.1479	0	0.0				1.406
8	0.1945	0	0.0				1.565
9	0.2411	0	0.0				1.742
10	0.2876	9	2.2 **				1.939
11	0.3342	0	0.0				2.159
12	0.3807	0	0.0				2.403
13	0.4273	0	0.0				2.675
14	0.4738	18	4.3 ****				2.977
15	0.5204	0	0.0				3.314
16	0.5670	31	7.5 *****				3.689
17	0.6135	0	0.0				4.107
18	0.6601	49	11.8 *****				4.572
19	0.7066	0	0.0				5.089
20	0.7532	53	12.7 *****				5.665
21	0.7997	77	18.5 *****				6.306
22	0.8463	0	0.0				7.019
23	0.8929	55	13.2 *****				7.814
24	0.9394	56	13.5 *****				8.698
25	0.9860	29	7.0 *****				9.682
26	1.033	14	3.4 ***				10.78
27	1.079	12	2.9 ***				12.00
28	1.126	5	1.2 *				13.36
29	1.172	4	1.0 *				14.87
30	1.219	0	0.0				16.55
31	1.265	0	0.0				18.42
32	1.312	0	0.0				20.51
33	1.358	0	0.0				22.83
34	1.405	0	0.0				25.41
35	1.452	1	0.2				28.28
36	1.498	0	0.0				31.49
37	1.545	0	0.0				35.05
38	1.591	0	0.0				39.01
39	1.638	0	0.0				43.43
40	1.684	0	0.0				48.34
41	1.731	0	0.0				53.81

T
I
T
L
E
S

PROGRAM: PERCENTAGE CUMULATIVE FREQUENCY PLOTS
DATA: DAISY CREEK RECON STREAM SEDIMENT GEOCHEMISTRY
RUN:
TIME: 25/03/83 11:55:33

NAME	CO
N	416
X BAR	0.82302
STD DEV	0.18623

CELL	LOWER LIMIT	NMBR	%	LOGARITHMIC VALUES	BAR INTERVAL = 0.12500	STD DEV	ARITH. LIMIT
1	1.277	1	0.2	*			18.92
2	1.254	0	0.2	*			17.93
3	1.230	0	0.2	*			17.00
4	1.207	0	0.2	*			16.11
5	1.184	1	0.5	*			15.27
6	1.161	3	1.2	*			14.47
7	1.137	5	2.4	*			13.72
8	1.114	0	2.4	*			13.00
9	1.091	8	4.3	*			12.32
10	1.067	4	5.3	*			11.68
11	1.044	0	5.3	*			11.07
12	1.021	14	8.7	*			10.49
13	0.9976	29	15.6	*			9.945
14	0.9743	0	15.6	*			9.426
15	0.9511	56	29.1	*			9.934
16	0.9278	0	29.1	*			6.468
17	0.9045	0	29.1	*			8.026
18	0.8812	55	42.3	*			7.607
19	0.8579	0	42.3	*			7.210
20	0.8347	77	60.8	*			6.834
21	0.8114	0	60.8	*			6.477
22	0.7881	0	60.8	*			6.139
23	0.7648	53	73.6	*			5.819
24	0.7415	0	73.6	*			5.515
25	0.7183	0	73.6	*			5.227
26	0.6950	49	85.3	*			4.954
27	0.6717	0	85.3	*			4.696
28	0.6484	0	85.3	*			4.451
29	0.6252	0	85.3	*			4.218
30	0.6019	31	92.8	*			3.998
31	0.5786	0	92.8	*			3.790
32	0.5553	0	92.8	*			3.592
33	0.5320	0	92.8	*			3.404
34	0.5088	0	92.8	*			3.227
35	0.4855	0	92.8	*			3.058
36	0.4622	18	97.1	*			2.899
37	0.4389	0	97.1	*			2.747
38	0.4156	0	97.1	*			2.604
39	0.3924	0	97.1	*			2.468
40	0.3691	0	97.1	*			2.339
41		12	100.0	*			

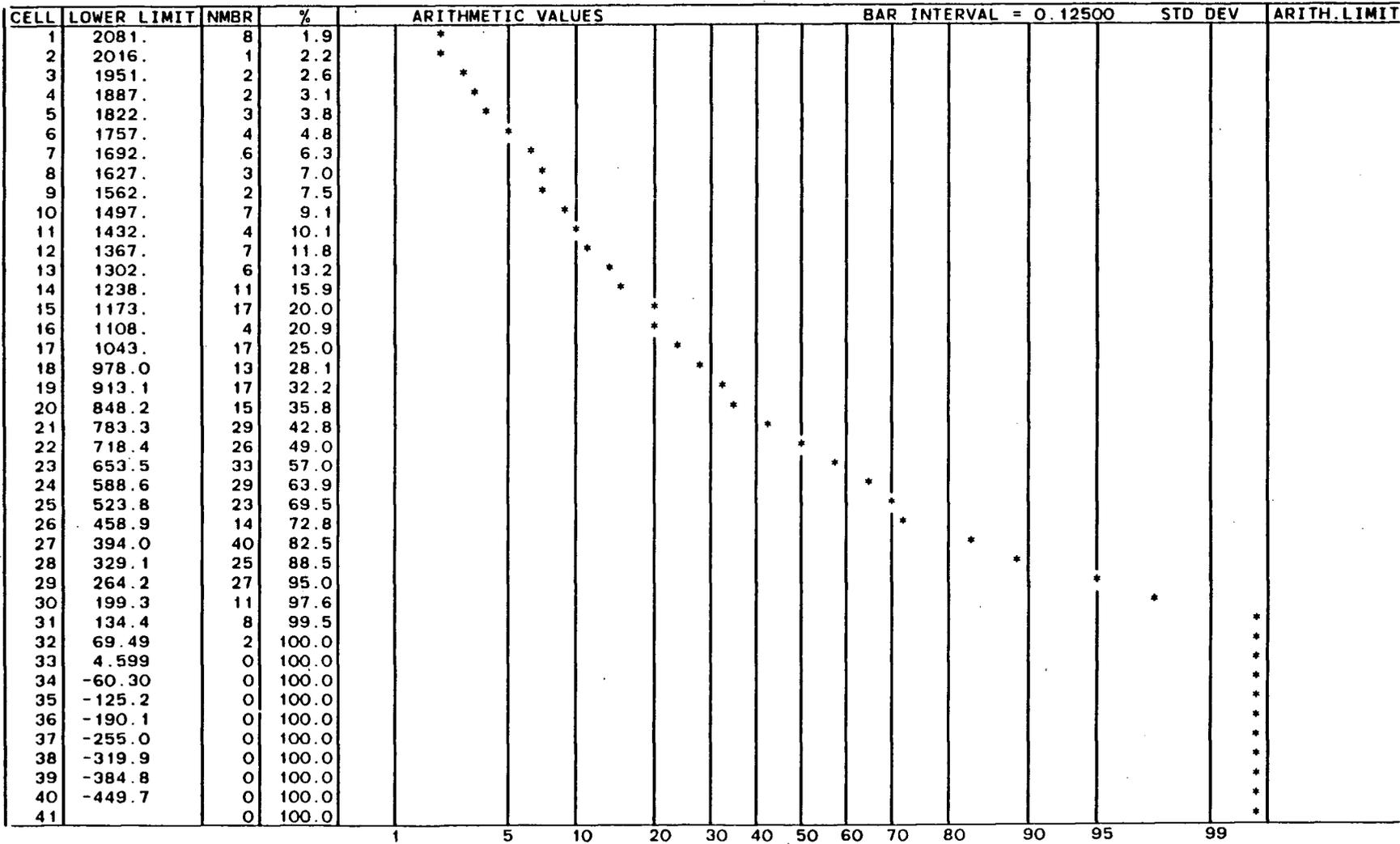
T I T L E S	PROGRAM: PERCENTAGE HISTOGRAMS
	DATA: DAISY CREEK RECON STREAM SEDIMENT GEOCHEMISTRY
	RUN:
	TIME: 25/03/83 11:55:33

NAME	MN
N	416
X BAR	815.78
STD DEV	519.16

CELL	LOWER LIMIT	NMBR	%	ARITHMETIC VALUES	BAR INTERVAL = 0.25000	STD DEV	ARITH.LIMIT
1		0	0.0				
2	-1715.	0	0.0				
3	-1585.	0	0.0				
4	-1456.	0	0.0				
5	-1326.	0	0.0				
6	-1196.	0	0.0				
7	-1066.	0	0.0				
8	-936.4	0	0.0				
9	-806.6	0	0.0				
10	-676.8	0	0.0				
11	-547.0	0	0.0				
12	-417.2	0	0.0				
13	-287.4	0	0.0				
14	-157.6	0	0.0				
15	-27.85	0	0.0				
16	101.9	11	2.6	***			
17	231.7	54	13.0	*****			
18	361.5	56	13.5	*****			
19	491.3	42	10.1	*****			
20	621.1	65	15.6	*****			
21	750.9	49	11.8	*****			
22	880.7	34	8.2	*****			
23	1010.	20	4.8	*****			
24	1140.	27	6.5	*****			
25	1270.	9	2.2	**			
26	1400.	14	3.4	***			
27	1530.	8	1.9	**			
28	1659.	7	1.7	**			
29	1789.	9	2.2	**			
30	1919.	2	0.5				
31	2049.	1	0.2				
32	2179.	3	0.7	*			
33	2308.	1	0.2				
34	2438.	0	0.0				
35	2568.	1	0.2				
36	2698.	1	0.2				
37	2828.	0	0.0				
38	2957.	0	0.0				
39	3087.	0	0.0				
40	3217.	0	0.0				
41	3347.	2	0.5				

T I T L E S	PROGRAM: PERCENTAGE CUMULATIVE FREQUENCY PLOTS
	DATA: DAISY CREEK RECON STREAM SEDIMENT GEOCHEMISTRY
	RUN:
	TIME: 25/03/83 11:55:33

NAME	MN
N	416
X BAR	815.78
STD DEV	519.16



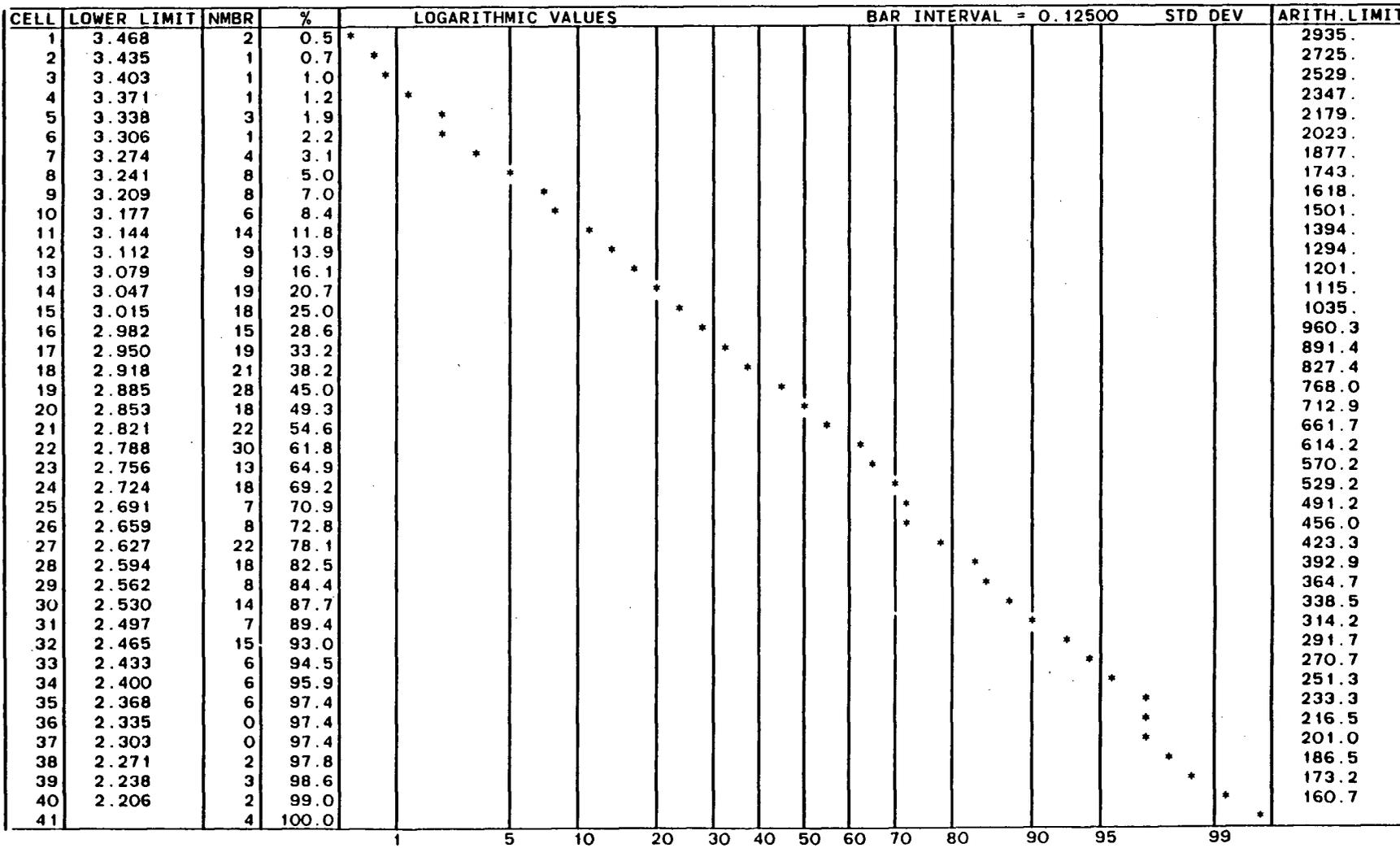
T I T L E S	PROGRAM:	PERCENTAGE HISTOGRAMS
	DATA:	DAISY CREEK RECON STREAM SEDIMENT GEOCHEMISTRY
	RUN:	
	TIME:	25/03/83 11:55:33

NAME	MN
N	416
X BAR	2.8369
STD DEV	0.25878

CELL	LOWER LIMIT	NMBR	%	LOGARITHMIC VALUES	BAR INTERVAL = 0.25000	STD DEV	ARITH. LIMIT
1		0	0.0				
2	1.575	0	0.0				37.61
3	1.640	0	0.0				43.65
4	1.705	0	0.0				50.66
5	1.769	0	0.0				58.80
6	1.834	0	0.0				68.25
7	1.899	0	0.0				79.21
8	1.963	1	0.2				91.93
9	2.028	1	0.2				106.7
10	2.093	1	0.2				123.8
11	2.158	2	0.5				143.7
12	2.222	4	1.0	*			166.8
13	2.287	2	0.5				193.6
14	2.352	10	2.4	***			224.7
15	2.416	16	3.8	*****			260.8
16	2.481	24	5.8	*****			302.7
17	2.546	22	5.3	*****			351.3
18	2.610	36	8.7	*****			407.8
19	2.675	15	3.6	****			473.3
20	2.740	35	8.4	*****			549.3
21	2.805	53	12.7	*****			637.5
22	2.869	48	11.5	*****			740.0
23	2.934	31	7.5	*****			858.8
24	2.999	30	7.2	*****			996.8
25	3.063	33	7.9	*****			1157.
26	3.128	20	4.8	****			1343.
27	3.193	16	3.8	****			1558.
28	3.257	8	1.9	**			1809.
29	3.322	4	1.0	*			2099.
30	3.387	2	0.5				2437.
31	3.451	0	0.0				2828.
32	3.516	1	0.2				3282.
33	3.581	0	0.0				3809.
34	3.646	0	0.0				4421.
35	3.710	1	0.2				5132.
36	3.775	0	0.0				5956.
37	3.840	0	0.0				6913.
38	3.904	0	0.0				8023.
39	3.969	0	0.0				9312.
40	4.034	0	0.0				0.1081E+05
41	4.098	0	0.0				0.1254E+05

T I T L E S	PROGRAM:	PERCENTAGE CUMULATIVE FREQUENCY PLOTS
	DATA:	DAISY CREEK RECON STREAM SEDIMENT GEOCHEMISTRY
	RUN:	
	TIME:	25/03/83 11:55:33

NAME	MN
N	416
X BAR	2.0369
STD DEV	0.25878

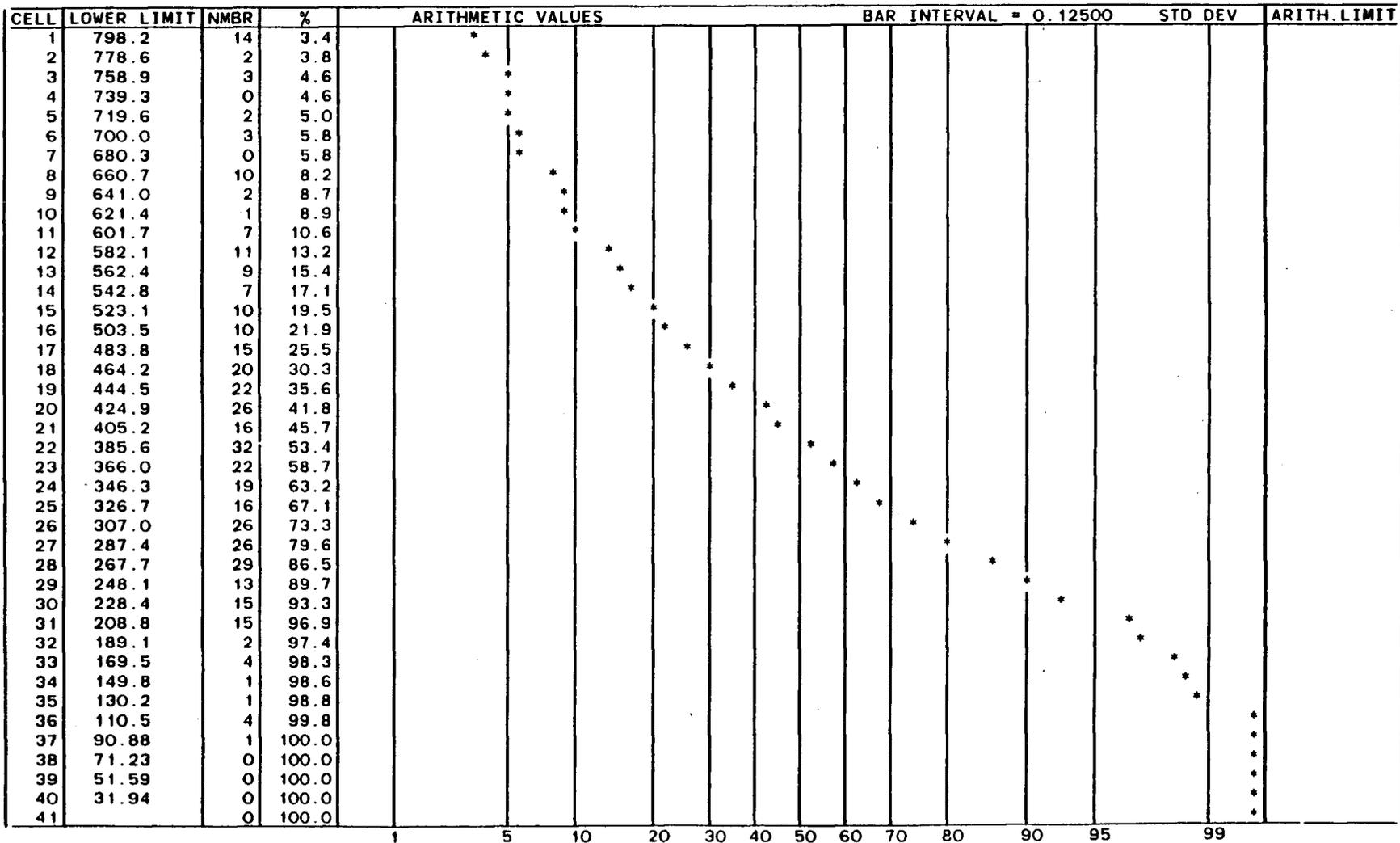


T I T L E S	PROGRAM: PERCENTAGE HISTOGRAMS
	DATA: DAISY CREEK RECON STREAM SEDIMENT GEOCHEMISTRY
	RUN:
	TIME: 25/03/83 11:55:33

NAME	F
N	416
X BAR	415.07
STD DEV	157.18

CELL	LOWER LIMIT	NMBR	%	ARITHMETIC VALUES	BAR INTERVAL = 0.25000	STD DEV	ARITH.LIMIT
1		0	0.0				
2	-351.2	0	0.0				
3	-311.9	0	0.0				
4	-272.6	0	0.0				
5	-233.3	0	0.0				
6	-194.0	0	0.0				
7	-154.7	0	0.0				
8	-115.4	0	0.0				
9	-76.12	0	0.0				
10	-36.83	0	0.0				
11	2.467	0	0.0				
12	41.76	0	0.0				
13	81.06	4	1.0 *				
14	120.4	2	0.5				
15	159.7	6	1.4 *				
16	198.9	25	6.0 *****				
17	238.2	34	8.2 *****				
18	277.5	49	11.8 *****				
19	316.8	48	11.5 *****				
20	356.1	42	10.1 *****				
21	395.4	47	11.3 *****				
22	434.7	39	9.4 *****				
23	474.0	30	7.2 *****				
24	513.3	25	6.0 *****				
25	552.6	18	4.3 ****				
26	591.9	11	2.6 ***				
27	631.2	6	1.4 *				
28	670.5	9	2.2 **				
29	709.8	2	0.5				
30	749.1	4	1.0 *				
31	788.4	5	1.2 *				
32	827.7	5	1.2 *				
33	867.0	1	0.2				
34	906.3	1	0.2				
35	945.6	1	0.2				
36	984.9	1	0.2				
37	1024.	0	0.0				
38	1063.	1	0.2				
39	1103.	0	0.0				
40	1142.	0	0.0				
41	1181.	0	0.0				

T I T L E S	PROGRAM:	PERCENTAGE CUMULATIVE FREQUENCY PLOTS	NAME	F
	DATA:	DAISY CREEK RECON STREAM SEDIMENT GEOCHEMISTRY	N	416
	RUN:		X BAR	415.07
	TIME:	25/03/83 11:55:33	STD DEV	157.18

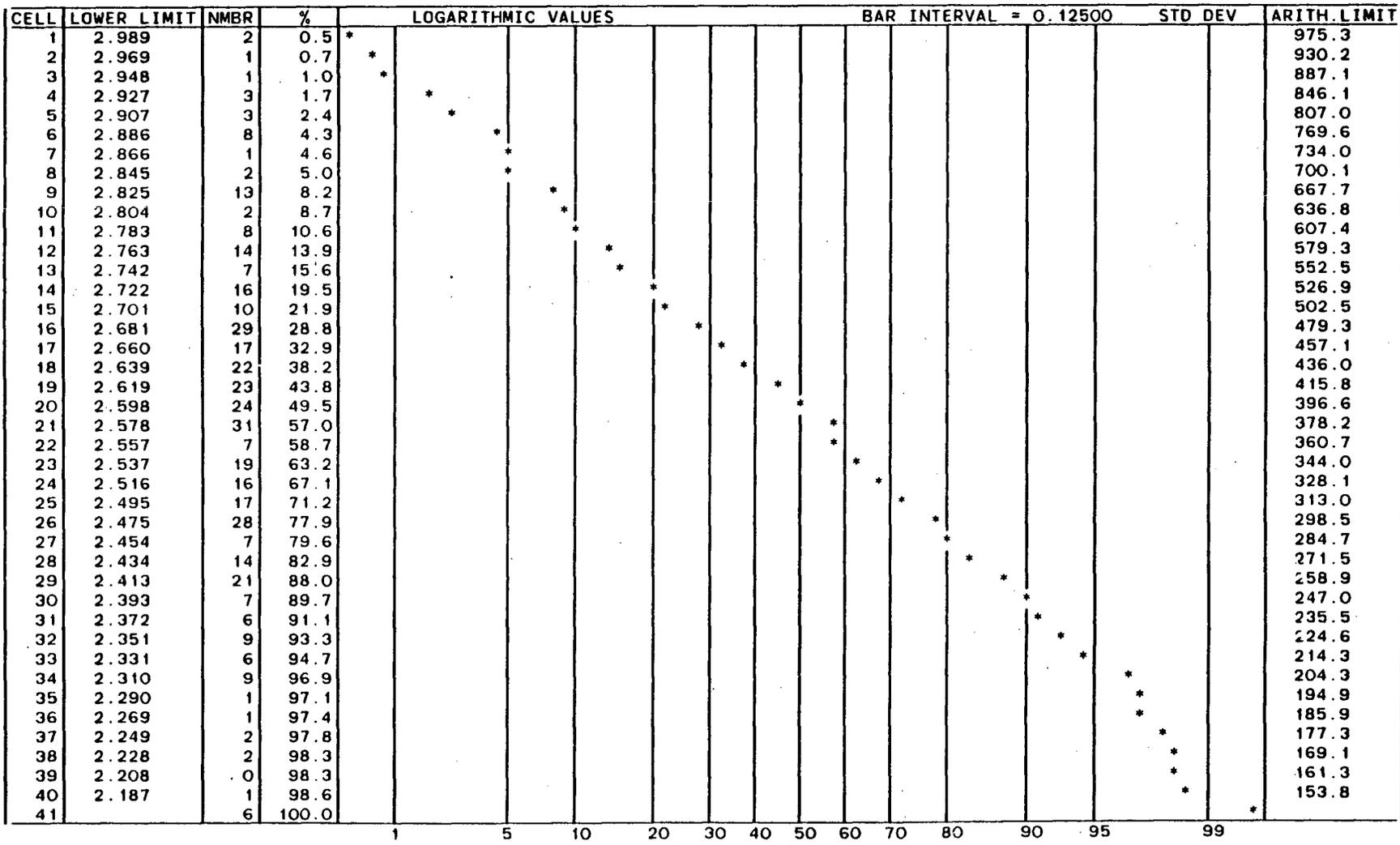


T I T L E S	PROGRAM:	PERCENTAGE HISTOGRAMS
	DATA:	DAISY CREEK RECON STREAM SEDIMENT GEOCHEMISTRY
	RUN:	
	TIME:	25/03/83 11:55:33

NAME	F
N	416
X BAR	2.5880
STD DEV	0.16456

CELL	LOWER LIMIT	NMBR	%	LOGARITHMIC VALUES	BAR INTERVAL = 0.25000	STD DEV	ARITH.LIMIT
1		0	0.0				
2	1.786	0	0.0				61.07
3	1.827	0	0.0				67.14
4	1.868	0	0.0				73.81
5	1.909	0	0.0				81.14
6	1.950	0	0.0				89.21
7	1.992	1	0.2				98.07
8	2.033	0	0.0				107.8
9	2.074	4	1.0	*			118.5
10	2.115	1	0.2				130.3
11	2.156	0	0.0				143.2
12	2.197	3	0.7	*			157.5
13	2.238	3	0.7	*			173.1
14	2.280	1	0.2				190.3
15	2.321	24	5.8	*****			209.2
16	2.362	13	3.1	***			230.0
17	2.403	21	5.0	*****			252.9
18	2.444	40	9.6	*****			278.0
19	2.485	34	8.2	*****			305.6
20	2.526	27	6.5	*****			336.0
21	2.567	54	13.0	*****			369.4
22	2.609	42	10.1	*****			406.1
23	2.650	45	10.8	*****			446.4
24	2.691	28	6.7	*****			490.8
25	2.732	28	6.7	*****			539.5
26	2.773	11	2.6	***			593.2
27	2.814	15	3.6	****			652.1
28	2.855	6	1.4	*			716.9
29	2.897	10	2.4	**			788.1
30	2.938	2	0.5				866.4
31	2.979	2	0.5				952.5
32	3.020	1	0.2				1047.
33	3.061	0	0.0				1151.
34	3.102	0	0.0				1266.
35	3.143	0	0.0				1391.
36	3.185	0	0.0				1530.
37	3.226	0	0.0				1681.
38	3.267	0	0.0				1849.
39	3.308	0	0.0				2032.
40	3.349	0	0.0				2234.
41	3.390	0	0.0				2456.

T I T L E S	PROGRAM:	PERCENTAGE CUMULATIVE FREQUENCY PLOTS	NAME	F
	DATA:	DAISY CREEK RECON STREAM SEDIMENT GEOCHEMISTRY	N	416
	RUN:		X BAR	2.5880
	TIME:	25/03/83 11:55:33	STD DEV	0.16456



T I T L E S	PROGRAM:	PERCENTAGE HISTOGRAMS
	DATA:	DAISY CREEK RECON STREAM SEDIMENT GEOCHEMISTRY
	RUN:	
	TIME:	25/03/83 11:55:33

NAME	BA
N	416
X BAR	795.70
STD DEV	158.27

CELL	LOWER LIMIT	NMBR	%	ARITHMETIC VALUES	BAR INTERVAL = 0.25000	STD DEV	AR:TH.LIMIT
1		0	0.0				
2	24.15	0	0.0				
3	63.72	0	0.0				
4	103.3	0	0.0				
5	142.8	0	0.0				
6	182.4	0	0.0				
7	222.0	0	0.0				
8	261.5	2	0.5				
9	301.1	1	0.2				
10	340.7	1	0.2				
11	380.2	1	0.2				
12	419.8	1	0.2				
13	459.4	2	0.5				
14	498.9	5	1.2	*			
15	538.5	10	2.4	**			
16	578.1	13	3.1	***			
17	617.6	9	2.2	**			
18	657.2	25	6.0	*****			
19	696.8	81	19.5	*****			
20	736.3	69	16.6	*****			
21	775.9	49	11.8	*****			
22	815.5	34	8.2	*****			
23	855.0	30	7.2	*****			
24	894.6	13	3.1	***			
25	934.2	19	4.6	****			
26	973.7	15	3.6	****			
27	1013.	8	1.9	**			
28	1053.	7	1.7	**			
29	1092.	6	1.4	*			
30	1132.	6	1.4	*			
31	1172.	1	0.2				
32	1211.	3	0.7	*			
33	1251.	0	0.0				
34	1290.	2	0.5				
35	1330.	1	0.2				
36	1369.	1	0.2				
37	1409.	1	0.2				
38	1449.	0	0.0				
39	1488.	0	0.0				
40	1528.	0	0.0				
41	1567.	0	0.0				

T I T L E S	PROGRAM:	PERCENTAGE HISTOGRAMS
	DATA:	DAISY CREEK RECON STREAM SEDIMENT GEOCHEMISTRY
	RUN:	
	TIME:	25/03/83 11:55:33

NAME	BA
N	416
X BAR	2.8919
STD DEV	0.91444E-01

CELL	LOWER LIMIT	NMBR	%	LOGARITHMIC VALUES	BAR INTERVAL = 0.25000	STD DEV	ARITH. LIMIT
1		2	0.5				
2	2.446	0	0.0				279.3
3	2.469	0	0.0				294.4
4	2.492	0	0.0				310.3
5	2.515	1	0.2				327.1
6	2.538	1	0.2				344.8
7	2.560	0	0.0				363.4
8	2.583	1	0.2				383.1
9	2.606	0	0.0				403.8
10	2.629	1	0.2				425.6
11	2.652	1	0.2				448.6
12	2.675	1	0.2				472.9
13	2.698	3	0.7	*			498.4
14	2.720	5	1.2	**			525.4
15	2.743	12	2.9	***			553.7
16	2.766	8	1.9	**			583.7
17	2.789	6	1.4	*			615.2
18	2.812	20	4.8	*****			648.5
19	2.835	73	17.5	*****			683.5
20	2.858	56	13.5	*****			720.5
21	2.880	73	17.5	*****			759.4
22	2.903	27	6.5	*****			800.4
23	2.926	35	8.4	*****			843.7
24	2.949	20	4.8	*****			889.3
25	2.972	23	5.5	*****			937.4
26	2.995	17	4.1	****			988.0
27	3.018	9	2.2	**			1041.
28	3.040	9	2.2	**			1098.
29	3.063	4	1.0	*			1157.
30	3.086	3	0.7	*			1220.
31	3.109	3	0.7	*			1286.
32	3.132	2	0.5				1355.
33	3.155	0	0.0				1428.
34	3.178	0	0.0				1505.
35	3.201	0	0.0				1587.
36	3.223	0	0.0				1673.
37	3.246	0	0.0				1763.
38	3.269	0	0.0				1858.
39	3.292	0	0.0				1959.
40	3.315	0	0.0				2065.
41	3.338	0	0.0				2176.

T I T L E S	PROGRAM:	PERCENTAGE CUMULATIVE FREQUENCY PLOTS
	DATA:	DAISY CREEK RECON STREAM SEDIMENT GEOCHEMISTRY
	RUN:	
	TIME:	25/03/83 11:55:33

NAME	BA
N	416
X BAR	2.8919
STD DEV	0.91444E-01

CELL	LOWER LIMIT	NMBR	%	LOGARITHMIC VALUES										BAR INTERVAL = 0.12500	STD DEV	ARITH. LIMIT		
1	3.115	4	1.0	*														1303.
2	3.103	1	1.2	*														1269.
3	3.092	2	1.7	*	*													1236.
4	3.080	1	1.9	*	*													1204.
5	3.069	1	2.2	*	*													1172.
6	3.058	5	3.4	*	*	*												1142.
7	3.046	4	4.3	*	*	*	*											1112.
8	3.035	4	5.3	*	*	*	*	*										1083.
9	3.023	6	6.7	*	*	*	*	*	*									1055.
10	3.012	6	8.2	*	*	*	*	*	*	*								1028.
11	3.000	3	8.9	*	*	*	*	*	*	*	*							1001.
12	2.989	14	12.3	*	*	*	*	*	*	*	*	*						975.1
13	2.978	16	16.1	*	*	*	*	*	*	*	*	*	*					949.8
14	2.966	7	17.8	*	*	*	*	*	*	*	*	*	*	*				925.1
15	2.955	7	19.5	*	*	*	*	*	*	*	*	*	*	*	*			901.1
16	2.943	14	22.8	*	*	*	*	*	*	*	*	*	*	*	*	*		877.7
17	2.932	18	27.2	*	*	*	*	*	*	*	*	*	*	*	*	*	*	854.9
18	2.920	19	31.7	*	*	*	*	*	*	*	*	*	*	*	*	*	*	832.7
19	2.909	15	35.3	*	*	*	*	*	*	*	*	*	*	*	*	*	*	811.0
20	2.898	31	42.8	*	*	*	*	*	*	*	*	*	*	*	*	*	*	790.0
21	2.886	29	49.8	*	*	*	*	*	*	*	*	*	*	*	*	*	*	769.5
22	2.875	42	59.9	*	*	*	*	*	*	*	*	*	*	*	*	*	*	749.5
23	2.863	16	63.7	*	*	*	*	*	*	*	*	*	*	*	*	*	*	730.0
24	2.852	41	73.6	*	*	*	*	*	*	*	*	*	*	*	*	*	*	711.0
25	2.840	40	83.2	*	*	*	*	*	*	*	*	*	*	*	*	*	*	692.6
26	2.829	15	86.8	*	*	*	*	*	*	*	*	*	*	*	*	*	*	674.6
27	2.818	10	89.2	*	*	*	*	*	*	*	*	*	*	*	*	*	*	657.1
28	2.806	5	90.4	*	*	*	*	*	*	*	*	*	*	*	*	*	*	640.0
29	2.795	3	91.1	*	*	*	*	*	*	*	*	*	*	*	*	*	*	623.4
30	2.783	4	92.1	*	*	*	*	*	*	*	*	*	*	*	*	*	*	607.2
31	2.772	3	92.8	*	*	*	*	*	*	*	*	*	*	*	*	*	*	591.4
32	2.760	7	94.5	*	*	*	*	*	*	*	*	*	*	*	*	*	*	576.0
33	2.749	3	95.2	*	*	*	*	*	*	*	*	*	*	*	*	*	*	561.1
34	2.738	7	96.9	*	*	*	*	*	*	*	*	*	*	*	*	*	*	546.5
35	2.726	0	96.9	*	*	*	*	*	*	*	*	*	*	*	*	*	*	532.3
36	2.715	3	97.6	*	*	*	*	*	*	*	*	*	*	*	*	*	*	518.5
37	2.703	2	98.1	*	*	*	*	*	*	*	*	*	*	*	*	*	*	505.0
38	2.692	0	98.1	*	*	*	*	*	*	*	*	*	*	*	*	*	*	491.9
39	2.680	1	98.3	*	*	*	*	*	*	*	*	*	*	*	*	*	*	479.1
40	2.669	1	98.6	*	*	*	*	*	*	*	*	*	*	*	*	*	*	466.7
41		6	100.0	*	*	*	*	*	*	*	*	*	*	*	*	*	*	

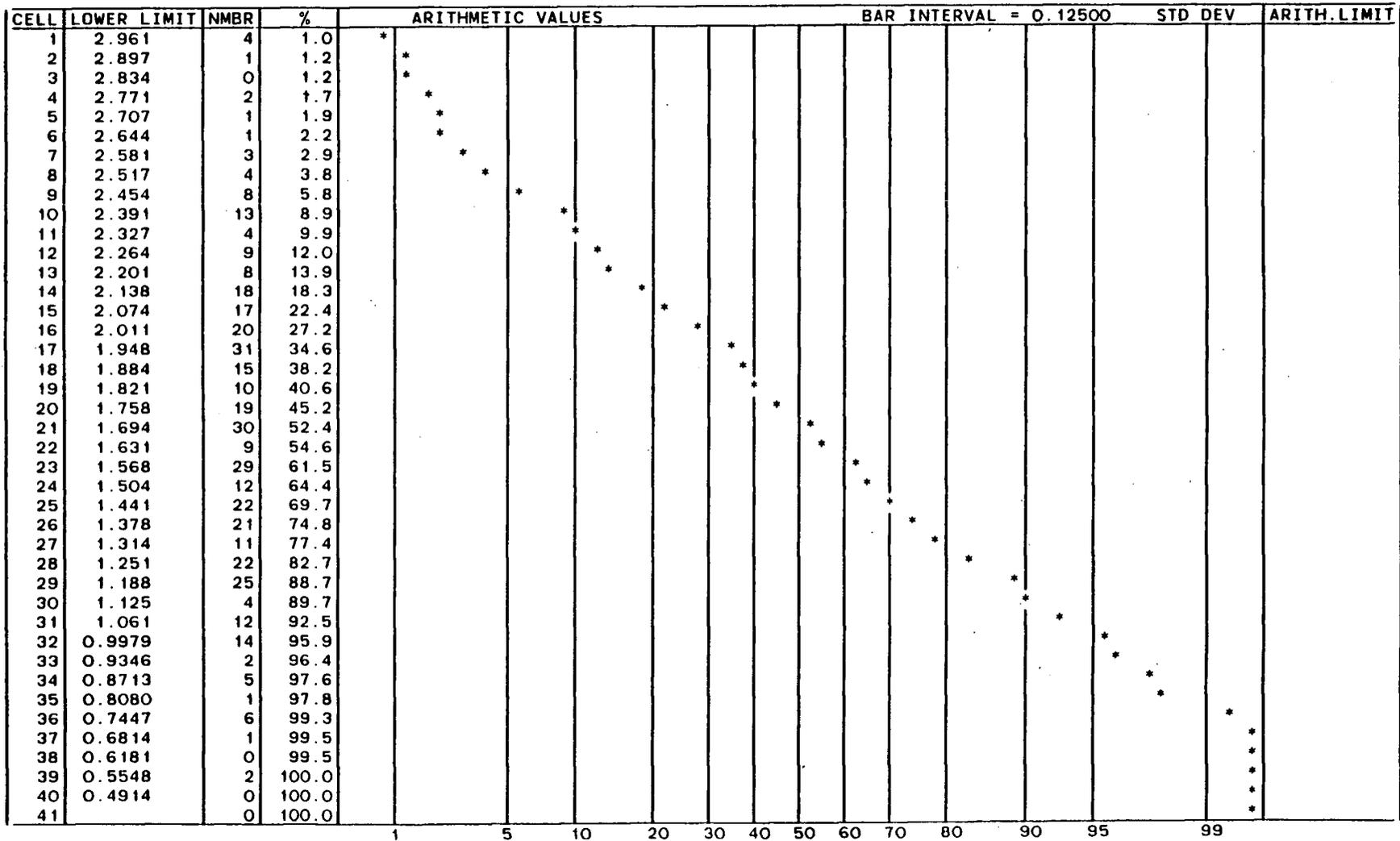
T I T L E S	PROGRAM: PERCENTAGE HISTOGRAMS
	DATA: DAISY CREEK RECON STREAM SEDIMENT GEOCHEMISTRY
	RUN:
	TIME: 25/03/83 11:55:33

NAME	FE
N	416
X BAR	1.7260
STD DEV	0.50650

CELL	LOWER LIMIT	NMBR	%	ARITHMETIC VALUES	BAR INTERVAL = 0.25000	STD DEV	ARITH.LIMIT
1		0	0.0				
2	-0.7431	0	0.0				
3	-0.6165	0	0.0				
4	-0.4899	0	0.0				
5	-0.3633	0	0.0				
6	-0.2366	0	0.0				
7	-0.1100	0	0.0				
8	0.1660E-01	0	0.0				
9	0.1432	0	0.0				
10	0.2698	0	0.0				
11	0.3965	0	0.0				
12	0.5231	2	0.5				
13	0.6497	5	1.2	*			
14	0.7763	8	1.9	**			
15	0.9030	12	2.9	***			
16	1.030	20	4.8	*****			
17	1.156	25	6.0	*****			
18	1.283	54	13.0	*****			
19	1.409	22	5.3	*****			
20	1.536	50	12.0	*****			
21	1.663	30	7.2	*****			
22	1.789	44	10.6	*****			
23	1.916	31	7.5	*****			
24	2.043	47	11.3	*****			
25	2.169	16	3.8	****			
26	2.296	25	6.0	*****			
27	2.422	9	2.2	**			
28	2.549	8	1.9	**			
29	2.676	3	0.7	*			
30	2.802	0	0.0				
31	2.929	4	1.0	*			
32	3.056	0	0.0				
33	3.182	0	0.0				
34	3.309	0	0.0				
35	3.435	0	0.0				
36	3.562	0	0.0				
37	3.689	0	0.0				
38	3.815	0	0.0				
39	3.942	0	0.0				
40	4.069	0	0.0				
41	4.195	1	0.2				

T I T L	PROGRAM: PERCENTAGE CUMULATIVE FREQUENCY PLOTS
L	DATA: DAISY CREEK RECON STREAM SEDIMENT GEOCHEMISTRY
E	RUN:
S	TIME: 25/03/83 11:55:33

NAME	FE
N	416
X BAR	1.7260
STD DEV	0.50650



T I T L E S	PROGRAM:	PERCENTAGE HISTOGRAMS
	DATA:	DAISY CREEK RECON STREAM SEDIMENT GEOCHEMISTRY
	RUN:	
	TIME:	25/03/83 11:55:33

NAME	FE
N	416
X BAR	0.21889
STD DEV	0.12766

CELL	LOWER LIMIT	NMBR	%	LOGARITHMIC VALUES	BAR INTERVAL = 0.25000	STD DEV	ARITH. LIMIT
1		0	0.0				
2	-0.4035	0	0.0				0.3949
3	-0.3716	0	0.0				0.4251
4	-0.3396	0	0.0				0.4575
5	-0.3077	0	0.0				0.4924
6	-0.2758	0	0.0				0.5299
7	-0.2439	2	0.5				0.5703
8	-0.2120	0	0.0				0.6138
9	-0.1801	1	0.2				0.6606
10	-0.1481	4	1.0 *				0.7110
11	-0.1162	2	0.5				0.7652
12	-0.8431E-01	1	0.2				0.8235
13	-0.5239E-01	7	1.7 **				0.8864
14	-0.2048E-01	10	2.4 **				0.9539
15	0.1144E-01	16	3.8 ****				1.027
16	0.4335E-01	4	1.0 *				1.105
17	0.7527E-01	25	6.0 *****				1.189
18	0.1072	33	7.9 *****				1.280
19	0.1391	30	7.2 *****				1.378
20	0.1710	25	6.0 *****				1.483
21	0.2029	56	13.5 *****				1.596
22	0.2349	31	7.5 *****				1.717
23	0.2668	36	8.7 *****				1.848
24	0.2987	57	13.7 *****				1.989
25	0.3306	35	8.4 *****				2.141
26	0.3625	17	4.1 ****				2.304
27	0.3944	16	3.8 ****				2.480
28	0.4263	3	0.7 *				2.669
29	0.4583	4	1.0 *				2.873
30	0.4902	0	0.0				3.092
31	0.5221	0	0.0				3.327
32	0.5540	0	0.0				3.581
33	0.5859	0	0.0				3.854
34	0.6178	0	0.0				4.148
35	0.6498	0	0.0				4.464
36	0.6817	0	0.0				4.805
37	0.7136	0	0.0				5.171
38	0.7455	0	0.0				5.566
39	0.7774	1	0.2				5.990
40	0.8093	0	0.0				6.447
41	0.8413	0	0.0				6.938

TITLE	PROGRAM:	PERCENTAGE CUMULATIVE FREQUENCY PLOTS	NAME	FE
	DATA:	DAISY CREEK RECON STREAM SEDIMENT GEOCHEMISTRY	N	416
	RUN:		X BAR	0.21889
	TIME:	25/03/83 11:55:33	STD DEV	0.12766

CELL	LOWER LIMIT	NMBR	%	LOGARITHMIC VALUES										BAR INTERVAL = 0.12500	STD DEV	ARITH. LIMIT		
1	0.5301	1	0.2	*														3.389
2	0.5141	0	0.2	*														3.267
3	0.4982	0	0.2	*														3.149
4	0.4822	2	0.7	*	*													3.035
5	0.4662	2	1.2	*	*													2.926
6	0.4503	0	1.2	*	*													2.820
7	0.4343	3	1.9	*	*	*												2.718
8	0.4184	1	2.2	*	*													2.620
9	0.4024	7	3.8	*	*	*												2.526
10	0.3865	9	6.0	*	*	*	*											2.435
11	0.3705	16	9.9	*	*	*	*	*										2.347
12	0.3545	9	12.0	*	*	*	*	*	*									2.262
13	0.3386	16	15.9	*	*	*	*	*	*	*								2.181
14	0.3226	10	18.3	*	*	*	*	*	*	*	*							2.102
15	0.3067	37	27.2	*	*	*	*	*	*	*	*	*						2.026
16	0.2907	20	32.0	*	*	*	*	*	*	*	*	*	*					1.953
17	0.2747	26	38.2	*	*	*	*	*	*	*	*	*	*	*				1.883
18	0.2588	10	40.6	*	*	*	*	*	*	*	*	*	*	*	*			1.815
19	0.2428	31	48.1	*	*	*	*	*	*	*	*	*	*	*	*	*		1.749
20	0.2269	18	52.4	*	*	*	*	*	*	*	*	*	*	*	*	*		1.686
21	0.2109	9	54.6	*	*	*	*	*	*	*	*	*	*	*	*	*		1.625
22	0.1950	29	61.5	*	*	*	*	*	*	*	*	*	*	*	*	*	*	1.567
23	0.1790	12	64.4	*	*	*	*	*	*	*	*	*	*	*	*	*	*	1.510
24	0.1630	13	67.5	*	*	*	*	*	*	*	*	*	*	*	*	*	*	1.456
25	0.1471	9	69.7	*	*	*	*	*	*	*	*	*	*	*	*	*	*	1.403
26	0.1311	21	74.8	*	*	*	*	*	*	*	*	*	*	*	*	*	*	1.352
27	0.1152	11	77.4	*	*	*	*	*	*	*	*	*	*	*	*	*	*	1.304
28	0.9921E-01	22	82.7	*	*	*	*	*	*	*	*	*	*	*	*	*	*	1.257
29	0.8325E-01	12	85.6	*	*	*	*	*	*	*	*	*	*	*	*	*	*	1.211
30	0.6729E-01	13	88.7	*	*	*	*	*	*	*	*	*	*	*	*	*	*	1.168
31	0.5133E-01	4	89.7	*	*	*	*	*	*	*	*	*	*	*	*	*	*	1.125
32	0.3538E-01	12	92.5	*	*	*	*	*	*	*	*	*	*	*	*	*	*	1.085
33	0.1942E-01	4	93.5	*	*	*	*	*	*	*	*	*	*	*	*	*	*	1.046
34	0.3459E-02	0	93.5	*	*	*	*	*	*	*	*	*	*	*	*	*	*	1.008
35	-0.1250E-01	10	95.9	*	*	*	*	*	*	*	*	*	*	*	*	*	*	0.9716
36	-0.2846E-01	2	96.4	*	*	*	*	*	*	*	*	*	*	*	*	*	*	0.9366
37	-0.4442E-01	0	96.4	*	*	*	*	*	*	*	*	*	*	*	*	*	*	0.9028
38	-0.6037E-01	5	97.6	*	*	*	*	*	*	*	*	*	*	*	*	*	*	0.8702
39	-0.7633E-01	1	97.8	*	*	*	*	*	*	*	*	*	*	*	*	*	*	0.8388
40	-0.9229E-01	0	97.8	*	*	*	*	*	*	*	*	*	*	*	*	*	*	0.8086
41		9	100.0	*	*	*	*	*	*	*	*	*	*	*	*	*	*	0.3086

APPENDIX 5

STREAM SEDIMENT REGRESSION OUTPUT

FUNCTIONAL REGRESSION
REDUCED MAJOR AXIS

DAISY CREEK STREAM SEDIMENT DATA - ALL DATA INCLUDED

BACKWARDS STEPWISE MULTIPLE REGRESSION ON CU

ITERATION NUMBER 1

NUMBER OF SAMPLES REGRESSED (N) = 388

CORRELATION COEFFICIENT (R) = 0.7778

STANDARD ERROR (E) = 0.1878

VARIABLE NAME	BETA VALUE	SIGNIFICANCE LEVEL
PB	0.5523	0.0
ZN	-0.1379	0.0186
MO	0.2120	0.0000
AG	0.3405	0.0001
CO	0.0302	0.3463
MN	0.3760	0.0000
F	0.2017	0.0007
BA	0.1068	0.1905
FE	0.5808	0.0000
CONSTANT	-0.9877	0.0

FUNCTIONAL REGRESSION
REDUCED MAJOR AXIS

DAISY CREEK STREAM SEDIMENT DATA - ALL DATA INCLUDED

BACKWARDS STEPWISE MULTIPLE REGRESSION ON CU

ITERATION NUMBER 2

NUMBER OF SAMPLES REGRESSED (N) = 388

CORRELATION COEFFICIENT (R) = 0.7781

STANDARD ERROR (E) = 0.1874

VARIABLE NAME	BETA VALUE	SIGNIFICANCE LEVEL
-----	-----	-----
PB	0.5577	0.0
ZN	-0.1440	0.0126
MO	0.2150	0.0000
AG	0.3364	0.0001
MN	0.3806	0.0000
F	0.2011	0.0007
BA	0.1138	0.1722
FE	0.6011	0.0000
CONSTANT	-0.9980	0.0

FUNCTIONAL REGRESSION
REDUCED MAJOR AXIS

DAISY CREEK STREAM SEDIMENT DATA - ALL DATA INCLUDED

BACKWARDS STEPWISE MULTIPLE REGRESSION ON CU

ITERATION NUMBER 3

NUMBER OF SAMPLES REGRESSED (N) = 388

CORRELATION COEFFICIENT (R) = 0.7788

STANDARD ERROR (E) = 0.1868

VARIABLE NAME	BETA VALUE	SIGNIFICANCE LEVEL
PB	0.5512	0.0
ZN	-0.1300	0.0185
MO	0.2090	0.0000
AG	0.3299	0.0001
MN	0.3891	0.0000
F	0.1985	0.0008
FE	0.6116	0.0000
CONSTANT	-0.7097	0.0

DAISY CREEK STREAM SEDIMENT DATA - ALL DATA INCLUDED

TYPE I A ANOMALIES

SAMPLES WHICH ARE ABOVE THE UPPER LIMIT OF THE
ZONE OF MIXING OF THE TWO POPULATIONS
AND WHICH ARE SIGNIFICANTLY GREATER THAN PREDICTED

SAMPLE NUMBER	X-COORD- INATE	Y-COORD- INATE	DEPENDENT VARIABLE
46726	34272.39	115385.97	163.000
47424	37510.02	137559.00	780.000

DAISY CREEK STREAM SEDIMENT DATA - ALL DATA INCLUDED

TYPE I B ANOMALIES

SAMPLES WHICH ARE ABOVE THE UPPER LIMIT OF THE
ZONE OF MIXING OF THE TWO POPULATIONS
AND WHICH ARE SIGNIFICANTLY LESS THAN PREDICTED

SAMPLE NUMBER	X-COORD- INATE	Y-COORD- INATE	DEPENDENT VARIABLE
30675	48493.74	127124.13	92.000

DAISY CREEK STREAM SEDIMENT DATA - ALL DATA INCLUDED

TYPE I C ANOMALIES

SAMPLES WHICH ARE ABOVE THE UPPER LIMIT OF THE
ZONE OF MIXING OF THE TWO POPULATIONS
AND WHICH ARE NOT SIGNIFICANTLY DIFFERANT THAN BACKGROUND

SAMPLE NUMBER	X-COORD- INATE	Y-COORD- INATE	DEPENDENT VARIABLE
46678	34430.44	120185.41	93.000
46727	34283.22	112945.90	96.000
68157	39968.82	98464.43	119.000
68665	36443.28	118277.84	123.000
46736	37156.54	139500.20	136.000

DAISY CREEK STREAM SEDIMENT DATA - ALL DATA INCLUDED

TYPE II A ANOMALIES

SAMPLES WHICH ARE PART OF THE MIXED POPULATION
BUT WHICH ARE SIGNIFICANTLY GREATER THAN PREDICTED

SAMPLE NUMBER	X-COORD- INATE	Y-COORD- INATE	DEPENDENT VARIABLE
------------------	-------------------	-------------------	-----------------------

NO SAMPLES RECOGNIZED

DAISY CREEK STREAM SEDIMENT DATA - ALL DATA INCLUDED

TYPE II B ANOMALIES

SAMPLES WHICH ARE PART OF THE MIXED POPULATION
BUT WHICH ARE SIGNIFICANTLY LESS THAN PREDICTED

SAMPLE NUMBER	X-COORD- INATE	Y-COORD- INATE	DEPENDENT VARIABLE
46827	38016.28	145377.36	60.000
30740	48443.49	124275.54	61.000
46679	34130.51	120206.53	62.000
33887	50171.90	104586.13	64.000
47417	40018.97	133100.08	64.000
46699	29714.98	133793.46	65.000
38062	49702.18	123752.40	65.000
33888	52587.74	103477.02	65.000
46833	36909.98	143501.02	75.000
33886	49612.95	104868.22	77.000

DAISY CREEK STREAM SEDIMENT DATA - ALL DATA INCLUDED

TYPE II C ANOMALIES

SAMPLES WHICH ARE PART OF THE MIXED POPULATION
BUT WHICH ARE NOT SIGNIFICANTLY DIFFERENT FROM BACKGROUND

SAMPLE NUMBER	X-COORD- INATE	Y-COORD- INATE	DEPENDENT VARIABLE
46681	31479.03	117156.50	60.000
47414	39787.54	130401.02	61.000
47410	33723.38	118308.06	62.000
47412	28419.12	112869.94	62.000
46680	33838.18	118078.60	69.000
47413	40516.11	127697.90	69.000
47416	40078.04	132879.82	73.000
47411	31060.31	117498.06	80.000
46731	32614.06	110512.19	89.000
46654	21176.22	101693.84	89.000

DAISY CREEK STREAM SEDIMENT DATA - ALL DATA INCLUDED

TYPE III A ANOMALIES

SAMPLES WHICH ARE PART OF THE BACKGROUND POPULATION
BUT WHICH ARE SIGNIFICANTLY GREATER THAN PREDICTED

SAMPLE NUMBER -----	X-COORD- INATE -----	Y-COORD- INATE -----	DEPENDENT VARIABLE -----
33822	24285.29	109189.54	13.000
46829	35219.44	140366.70	20.000
68131	47539.82	101376.00	24.000
68202	44719.23	95906.56	25.000
46665	41982.49	112736.98	25.000
46984	28634.55	139151.16	33.000
68667	31176.48	116477.63	40.000
30720	42022.94	107536.78	44.000
47418	37883.40	134149.12	44.000
68688	44706.40	148374.88	59.000

DAISY CREEK STREAM SEDIMENT DATA - ALL DATA INCLUDED

TYPE III B ANOMALIES

SAMPLES WHICH ARE PART OF THE BACKGROUND POPULATION
BUT WHICH ARE SIGNIFICANTLY LESS THAN PREDICTED

SAMPLE NUMBER	X-COORD- INATE	Y-COORD- INATE	DEPENDENT VARIABLE
46505	51615.58	127550.91	3.000
46526	24130.37	99890.82	4.000
33743	44603.78	129500.62	5.000
33833	27725.46	106419.34	8.000
33835	29055.98	102999.66	8.000
33843	33059.24	95930.40	9.000
46608	27136.17	101647.38	11.000
33828	31403.47	97056.62	14.000
46609	49472.28	115328.85	16.000
32805	43907.57	125926.78	24.000
30701	45764.29	124602.19	29.000

DAISY CREEK STREAM SEDIMENT DATA - ALL DATA INCLUDED

BACKWARDS STEPWISE MULTIPLE REGRESSION ON CU

ANOMALY RECOGNITION SUMMARY

TOTAL # OF SAMPLES = 416

TOTAL # OF COMPLETE SAMPLES = 416

TOTAL # OF ANOMALOUS SAMPLES (> T2) = 8

TOTAL # OF MIXED SAMPLES (< T2 AND > T1) = 20

TOTAL # OF BACKGROUND SAMPLES (< T1) = 388

LOWER THRESHOLD (T1) = 60.0000

UPPER THRESHOLD (T2) = 90.0000

TOTAL # OF CLASS I A ANOMALIES = 2

TOTAL # OF CLASS I B ANOMALIES = 1

TOTAL # OF CLASS I C ANOMALIES = 5

TOTAL # OF CLASS II A ANOMALIES = 0

TOTAL # OF CLASS II B ANOMALIES = 10

TOTAL # OF CLASS II C ANOMALIES = 10

TOTAL # OF CLASS III A ANOMALIES = 10

TOTAL # OF CLASS III B ANOMALIES = 11

1	MIXED POPULATION				
2	46827	38016.28	145377.36	-0.641	60.000
3	30740	48443.49	124275.54	-0.664	61.000
4	46679	34130.51	120206.53	-0.763	62.000
5	33887	50171.90	104586.13	-0.627	64.000
6	47417	40018.97	133100.08	-0.405	64.000
7	46699	29714.98	133793.46	-0.728	65.000
8	38062	49702.18	123752.40	-0.630	65.000
9	33888	52587.74	103477.02	-0.525	65.000
10	46833	36909.98	143501.02	-0.477	75.000
11	33886	49612.95	104868.22	-0.439	77.000
12	46681	31479.03	117156.50	-0.115	60.000
13	47414	39787.54	130401.02	-0.150	61.000
14	47410	33723.38	118308.06	-0.333	62.000
15	47412	28419.12	112869.94	0.032	62.000
16	46680	33838.18	118078.60	-0.102	69.000
17	47413	40516.11	127697.90	-0.355	69.000
18	47416	40078.04	132879.82	-0.043	73.000
19	47411	31060.31	117498.06	-0.362	80.000
20	46731	32614.06	110512.19	-0.257	89.000
21	46654	21176.22	101693.84	-0.018	89.000
22	ANOMALOUS POPULATION				
23	46726	34272.39	115385.97	0.437	163.000
24	47424	37510.02	137559.00	0.900	780.000
25	30675	48493.74	127124.13	-0.772	92.000
26	46678	34430.44	120185.41	-0.091	93.000
27	46727	34283.22	112945.90	-0.078	96.000
28	68157	39968.82	98464.43	-0.066	119.000
29	68665	36443.28	118277.84	-0.020	123.000
30	46736	37156.54	139500.20	0.289	136.000

APPENDIX 6

**SOIL SURVEY -
GEOCHEMICAL DATA**

	SAMP	EAST	NRTH	AG	CU	PB	ZN	CO	MO	AS	HG	BA
	NUM	(M)	(M)	(PPM)	(PPM)	(PPM)	(PPM)	(PPM)	(PPM)	(PPM)	(PPB)	(PPM)
1												
2												
3												
4	88000	495	420	0.1	35.0	2.5	45.0	15.0	1.0	10.0	30	670
5	88002	390	420	0.1	35.0	20.0	35.0	5.0	1.0	5.0	10	440
6	88004	420	420	0.1	50.0	15.0	60.0	2.5	1.0	5.0	30	610
7	88006	435	420	0.1	65.0	145.0	70.0	2.5	1.0	5.0	30	620
8	88008	450	420	0.1	70.0	60.0	75.0	10.0	1.0	5.0	30	630
9	88010	465	420	0.4	150.0	115.0	70.0	2.5	1.0	5.0	50	690
10	88012	435	405	0.1	65.0	45.0	70.0	10.0	1.0	5.0	40	590
11	88014	495	405	0.1	30.0	5.0	45.0	10.0	1.0	5.0	40	650
12	88016	510	405	0.1	20.0	2.5	50.0	10.0	1.0	5.0	20	600
13	88018	525	405	0.1	15.0	2.5	40.0	10.0	1.0	5.0	50	570
14	88019	540	405	0.1	30.0	2.5	35.0	10.0	1.0	5.0	40	540
15	88022	465	405	0.1	35.0	40.0	60.0	10.0	1.0	5.0	30	730
16	88024	450	405	0.1	55.0	60.0	95.0	15.0	1.0	5.0	20	940
17	88026	420	510	0.1	10.0	2.5	25.0	10.0	1.0	5.0	10	420
18	88028	435	510	0.1	50.0	10.0	35.0	2.5	1.0	5.0	10	590
19	88030	390	510	0.1	65.0	10.0	45.0	2.5	1.0	20.0	20	590
20	88032	420	525	0.1	15.0	2.5	35.0	2.5	1.0	5.0	20	480
21	88034	435	525	0.1	45.0	2.5	40.0	2.5	1.0	5.0	10	490
22	88036	450	525	0.8	520.0	5.0	70.0	5.0	1.0	5.0	40	650
23	88038	465	525	0.1	275.0	2.5	45.0	2.5	1.0	5.0	50	550
24	88040	480	525	0.1	135.0	10.0	90.0	10.0	1.0	5.0	40	700
25	88042	495	525	0.1	35.0	5.0	55.0	10.0	1.0	5.0	30	710
26	88802	480	480	0.4	245.0	5.0	55.0	10.0	1.0	5.0	30	920
27	88804	480	465	0.4	280.0	2.5	70.0	10.0	1.0	5.0	40	740
28	88806	480	450	0.7	365.0	10.0	95.0	2.5	1.0	5.0	40	830
29	88808	480	435	0.7	180.0	2.5	75.0	10.0	1.0	5.0	40	700
30	88810	480	420	0.4	125.0	2.5	70.0	10.0	1.0	5.0	50	850
31	88812	480	405	0.8	50.0	5.0	55.0	10.0	1.0	5.0	30	690
32	88814	480	390	0.1	30.0	10.0	55.0	2.5	1.0	10.0	30	650
33	88816	480	375	0.1	20.0	5.0	45.0	15.0	1.0	5.0	30	630
34	88818	480	360	0.1	40.0	5.0	55.0	10.0	1.0	10.0	30	640
35	88820	480	345	0.1	25.0	2.5	50.0	15.0	1.0	5.0	40	610
36	88822	495	480	0.1	35.0	15.0	75.0	10.0	1.0	5.0	40	750
37	88824	510	480	0.1	40.0	2.5	65.0	10.0	1.0	5.0	50	650
38	88826	525	480	0.1	15.0	2.5	45.0	10.0	1.0	5.0	10	610
39	88828	540	480	0.1	20.0	2.5	45.0	10.0	1.0	5.0	20	650
40	88830	570	480	0.1	35.0	5.0	55.0	5.0	1.0	5.0	30	690
41	88832	465	480	1.7	950.0	20.0	65.0	2.5	1.0	10.0	50	710
42	88834	450	480	0.1	370.0	210.0	40.0	10.0	1.0	5.0	20	560
43	88836	435	480	0.1	110.0	1150.0	60.0	10.0	1.0	10.0	20	630
44	88838	420	480	0.1	10.0	10.0	50.0	5.0	1.0	5.0	20	670
45	88840	390	480	0.1	20.0	10.0	40.0	2.5	1.0	5.0	20	700
46	88842	495	465	0.1	30.0	5.0	50.0	5.0	1.0	5.0	40	240
47	88844	510	465	0.1	25.0	2.5	50.0	10.0	1.0	10.0	50	710
48	88846	525	465	0.1	15.0	5.0	50.0	5.0	1.0	5.0	40	630
49	88848	540	465	0.1	15.0	5.0	55.0	10.0	1.0	5.0	20	570
50	88850	420	465	0.1	15.0	10.0	45.0	2.5	1.0	5.0	10	570
51	88852	495	360	0.1	20.0	2.5	40.0	10.0	1.0	5.0	30	630
52	88854	510	360	0.1	20.0	2.5	35.0	5.0	1.0	5.0	10	630
53	88856	525	360	0.1	10.0	2.5	30.0	10.0	1.0	5.0	20	580
54	88858	540	360	0.1	25.0	2.5	50.0	2.5	1.0	5.0	30	670
55	88860	540	363	0.1	25.0	2.5	45.0	10.0	1.0	5.0	20	850
56	88862	537	357	0.1	20.0	2.5	45.0	10.0	1.0	5.0	30	650
57	88864	543	357	0.1	25.0	2.5	45.0	10.0	1.0	5.0	20	680
58	88866	570	360	0.1	25.0	2.5	50.0	10.0	1.0	5.0	20	670

	SAMP NUM	EAST (M)	NRTH (M)	AG (PPM)	CU (PPM)	PB (PPM)	ZN (PPM)	CO (PPM)	MO (PPM)	AS (PPM)	HG (PPB)	BA (PPM)
59												
60												
61												
62	88868	540	390	0.1	20.0	2.5	30.0	10.0	1.0	5.0	20	540
63	88870	570	390	0.1	15.0	2.5	35.0	10.0	1.0	5.0	20	570
64	88872	525	390	0.1	20.0	2.5	35.0	10.0	1.0	5.0	20	480
65	88874	510	390	0.1	10.0	5.0	35.0	10.0	1.0	5.0	20	600
66	88876	495	390	0.1	40.0	10.0	65.0	15.0	1.0	5.0	40	690
67	88878	495	375	0.8	30.0	2.5	50.0	2.5	1.0	5.0	30	620
68	88880	510	375	0.4	25.0	5.0	40.0	10.0	1.0	5.0	30	710
69	88882	525	375	0.1	20.0	2.5	45.0	15.0	1.0	5.0	20	850
70	88884	540	375	0.1	15.0	2.5	30.0	2.5	1.0	5.0	40	550
71	88886	483	402	0.1	30.0	2.5	45.0	10.0	1.0	5.0	50	780
72	88888	480	408	0.1	75.0	2.5	80.0	2.5	1.0	5.0	60	750
73	88890	477	402	0.1	315.0	10.0	70.0	10.0	1.0	5.0	40	870
74	88892	480	510	0.1	145.0	10.0	65.0	10.0	1.0	5.0	50	760
75	88894	510	510	0.1	25.0	2.5	60.0	10.0	1.0	5.0	30	730
76	88896	540	510	0.1	30.0	2.5	50.0	2.5	1.0	5.0	40	540
77	88898	570	510	0.1	25.0	2.5	50.0	10.0	1.0	30.0	40	620
78	88899	450	510	0.4	560.0	10.0	60.0	10.0	1.0	5.0	40	570
79	88900	465	510	1.1	620.0	2.5	45.0	2.5	1.0	5.0	30	480
80	88902	435	465	0.1	50.0	670.0	40.0	2.5	1.0	5.0	10	500
81	88904	450	465	0.1	180.0	360.0	45.0	10.0	1.0	5.0	20	560
82	88906	465	465	0.1	370.0	25.0	50.0	2.5	1.0	5.0	50	640
83	88908	495	450	0.1	30.0	2.5	50.0	2.5	1.0	5.0	60	530
84	88910	510	450	0.1	15.0	2.5	30.0	2.5	1.0	5.0	10	520
85	88912	525	450	0.1	65.0	2.5	55.0	10.0	1.0	5.0	20	670
86	88914	540	450	0.1	20.0	2.5	30.0	2.5	1.0	5.0	10	500
87	88916	570	450	0.1	20.0	2.5	45.0	2.5	1.0	5.0	20	570
88	88918	390	450	0.1	45.0	20.0	45.0	2.5	1.0	5.0	20	530
89	88920	420	450	0.1	20.0	10.0	50.0	10.0	1.0	5.0	30	550
90	88922	435	450	0.1	35.0	50.0	70.0	10.0	1.0	5.0	50	750
91	88924	450	450	0.1	315.0	940.0	85.0	15.0	1.0	5.0	30	760
92	88926	465	450	0.4	455.0	40.0	80.0	10.0	1.0	5.0	40	760
93	88928	540	435	0.1	2.5	2.5	20.0	2.5	1.0	5.0	30	440
94	88930	525	435	0.1	20.0	2.5	40.0	2.5	1.0	5.0	50	710
95	88932	510	435	0.1	30.0	2.5	45.0	2.5	1.0	5.0	50	640
96	88934	495	435	0.1	20.0	5.0	50.0	10.0	1.0	5.0	40	710
97	88936	420	435	0.1	35.0	10.0	50.0	2.5	1.0	5.0	20	540
98	88938	435	435	0.1	105.0	55.0	60.0	10.0	1.0	5.0	20	610
99	88940	450	435	0.1	80.0	240.0	55.0	15.0	1.0	5.0	30	750
100	88942	465	435	0.1	285.0	70.0	65.0	10.0	1.0	5.0	30	740
101	88944	570	420	0.1	20.0	5.0	35.0	10.0	1.0	5.0	20	570
102	88946	540	420	0.1	10.0	5.0	35.0	10.0	1.0	5.0	10	570
103	88948	525	420	0.1	20.0	5.0	30.0	10.0	1.0	5.0	10	510
104	88950	510	420	0.1	10.0	5.0	30.0	5.0	1.0	5.0	20	420
105	88952	420	405	0.1	25.0	35.0	60.0	10.0	1.0	10.0	10	590
106	88954	390	390	0.1	30.0	20.0	60.0	2.5	1.0	10.0	10	680
107	88956	420	390	0.1	35.0	10.0	45.0	10.0	1.0	5.0	20	470
108	88958	435	390	0.1	85.0	30.0	55.0	10.0	1.0	5.0	20	670
109	88960	450	390	0.1	140.0	40.0	85.0	10.0	1.0	5.0	30	820
110	88962	465	390	0.1	55.0	25.0	75.0	2.5	1.0	5.0	40	750
111	88964	465	375	0.1	65.0	20.0	105.0	10.0	1.0	5.0	30	950
112	88966	450	375	0.1	140.0	30.0	80.0	10.0	1.0	5.0	30	830
113	88968	435	375	0.1	95.0	25.0	65.0	10.0	1.0	5.0	20	640
114	88970	420	375	0.1	30.0	20.0	65.0	10.0	1.0	5.0	30	640
115	88972	390	360	0.1	30.0	15.0	45.0	5.0	1.0	5.0	10	570
116	88974	387	357	0.1	30.0	10.0	70.0	10.0	1.0	5.0	20	660

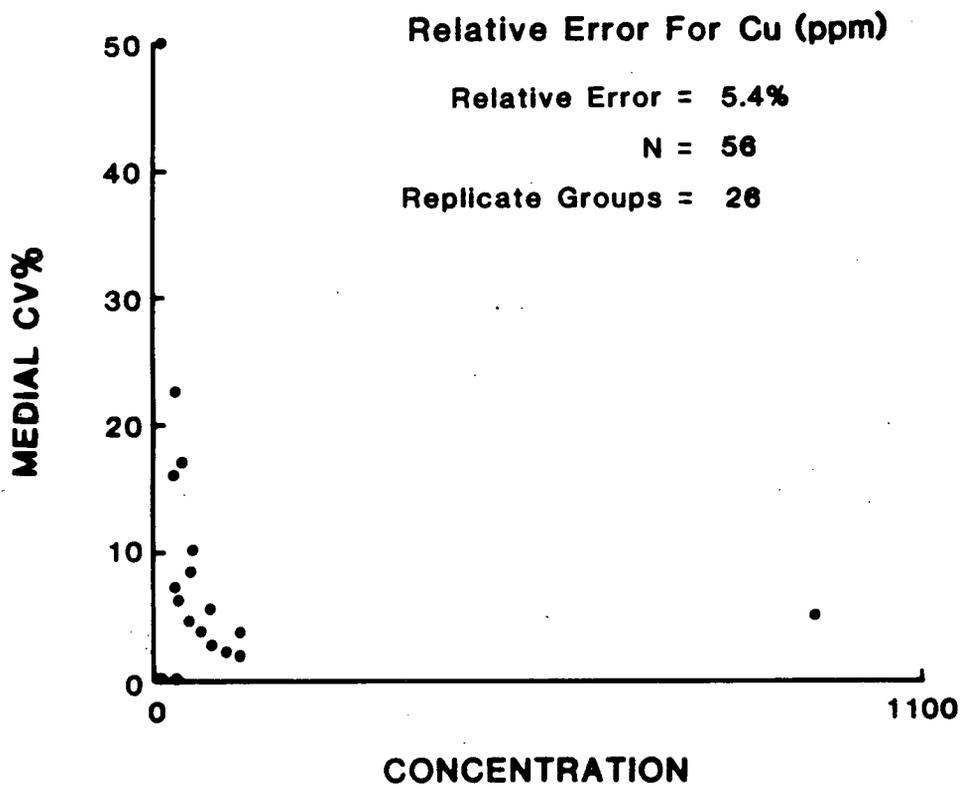
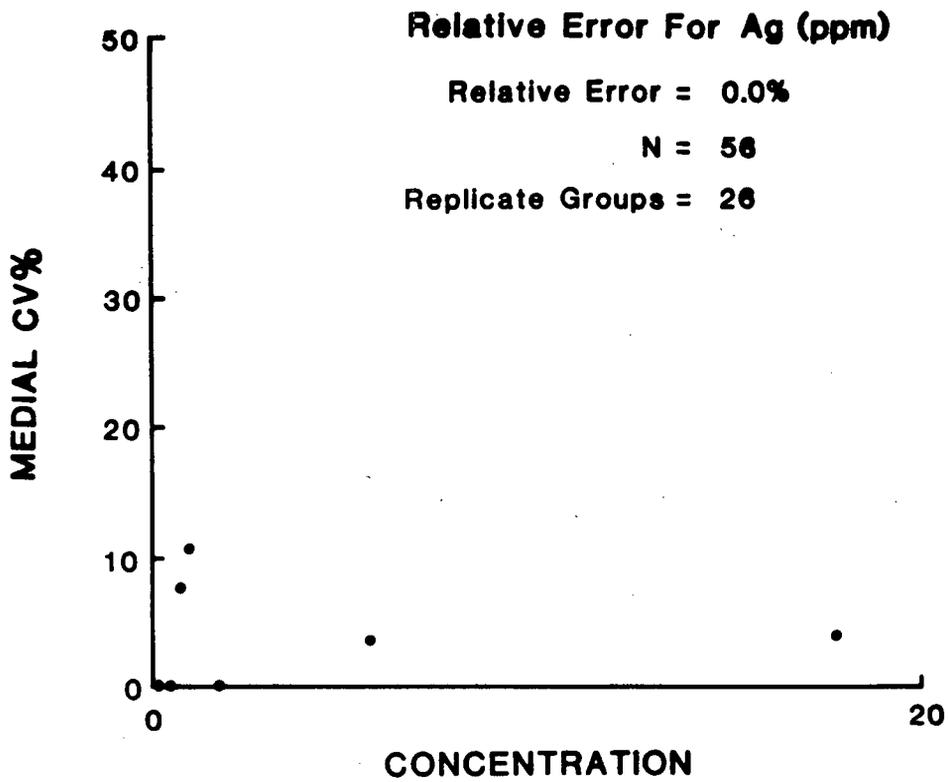
- 233 -

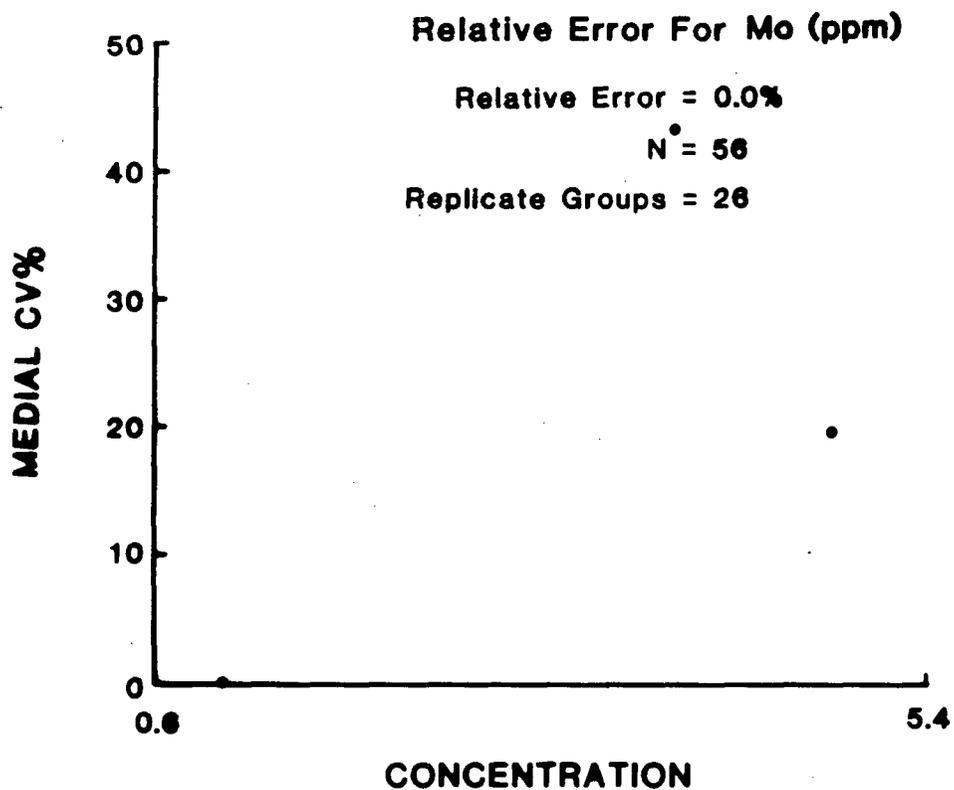
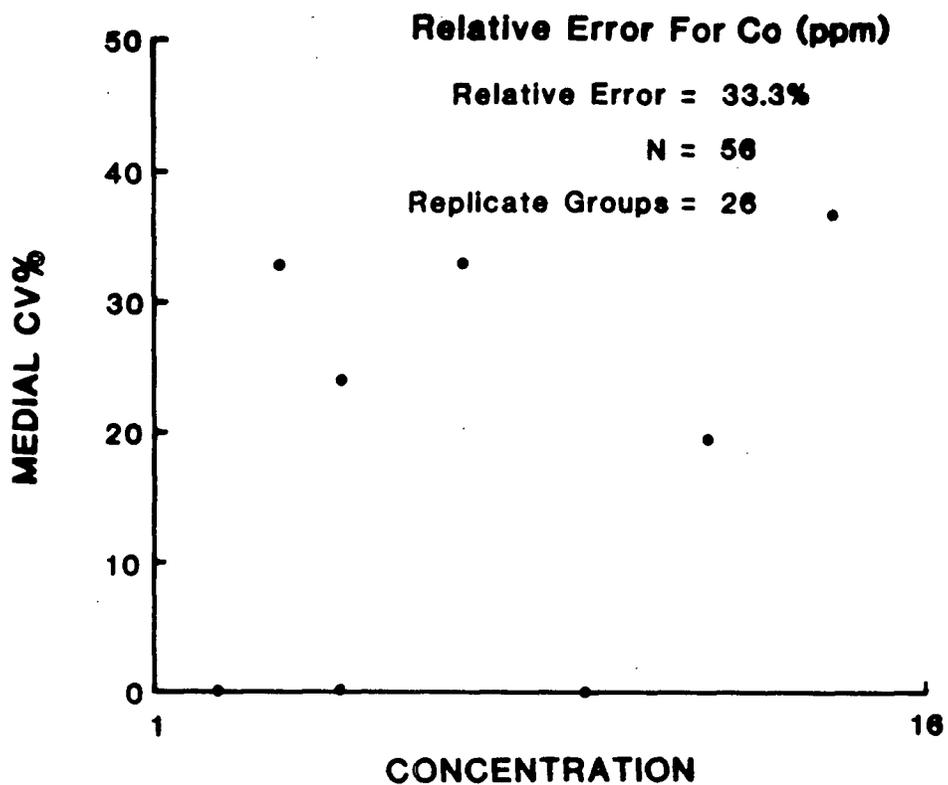
Listing of SOILGD at 17:33:18 on SEP 24, 1984 for CCID=MZAB Page 3

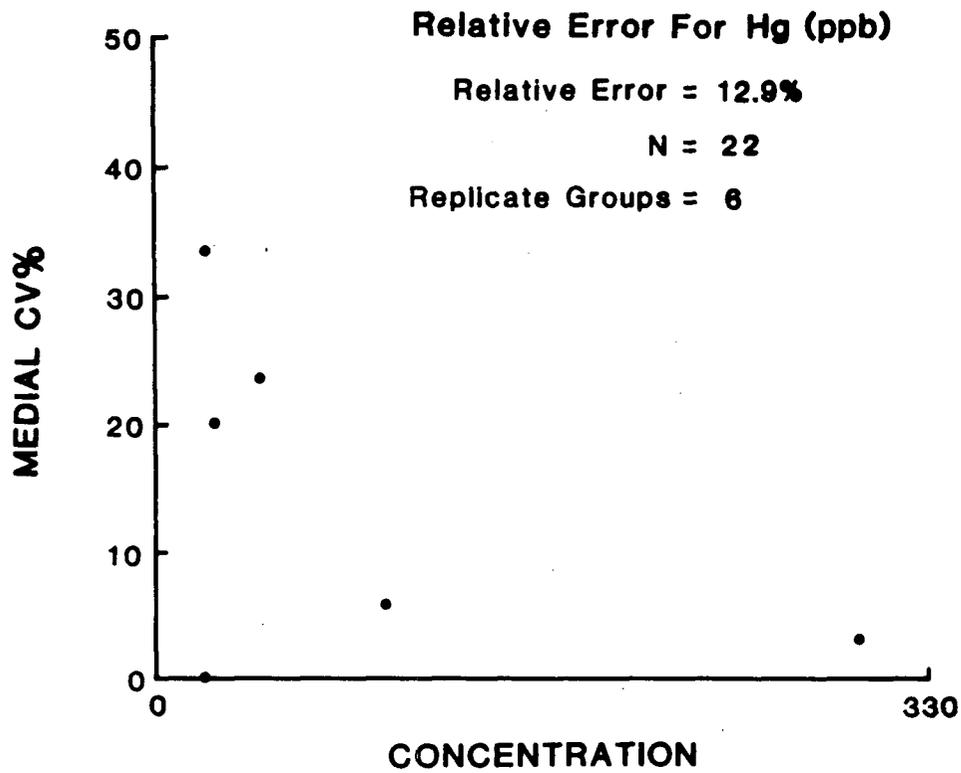
SAMP NUM	EAST (M)	NRTH (M)	AG (PPM)	CU (PPM)	PB (PPM)	ZN (PPM)	CO (PPM)	MO (PPM)	AS (PPM)	HG (PPB)	BA (PPM)	
117	88976	390	363	0.1	25.0	20.0	45.0	5.0	1.0	10.0	30	470
118	88978	393	357	0.1	50.0	15.0	90.0	5.0	1.0	5.0	20	650
119	88980	420	360	0.1	55.0	25.0	50.0	10.0	1.0	5.0	20	630
120	88982	435	360	0.1	195.0	20.0	50.0	10.0	1.0	5.0	30	600
121	88984	450	360	0.1	130.0	15.0	70.0	5.0	1.0	5.0	50	780
122	88986	465	360	0.1	55.0	20.0	115.0	5.0	1.0	5.0	30	930

APPENDIX Z

**SOIL SURVEY -
RELATIVE ERROR PLOTS**







APPENDIX B

SOIL SURVEY PROBABILITY PLOTS

T	PROGRAM:	PERCENTAGE HISTOGRAMS
I		
T	DATA:	DAISY CREEK SC HORIZON SOIL GEOCHEMISTRY
L	RUN:	
E	TIME:	25/03/83 14:08:04
S		

NAME	AG-SC
N	116
X BAR	0.16896
STD DEV	0.22512

CELL	LOWER LIMIT	NMBR	%	ARITHMETIC VALUES	BAR INTERVAL = 0.25000	STD DEV	ARITH.LIMIT
1		0	0.0				
2	-0.9285	0	0.0				
3	-0.8722	0	0.0				
4	-0.8160	0	0.0				
5	-0.7597	0	0.0				
6	-0.7034	0	0.0				
7	-0.6471	0	0.0				
8	-0.5908	0	0.0				
9	-0.5345	0	0.0				
10	-0.4783	0	0.0				
11	-0.4220	0	0.0				
12	-0.3657	0	0.0				
13	-0.3094	0	0.0				
14	-0.2531	0	0.0				
15	-0.1969	0	0.0				
16	-0.1406	0	0.0				
17	-0.8430E-01	0	0.0				
18	-0.2802E-01	0	0.0				
19	0.2826E-01	0	0.0				
20	0.8454E-01	102	87.9	*****			
21	0.1408	0	0.0				
22	0.1971	0	0.0				
23	0.2534	0	0.0				
24	0.3097	0	0.0				
25	0.3659	7	6.0	*****			
26	0.4222	0	0.0				
27	0.4785	0	0.0				
28	0.5348	0	0.0				
29	0.5911	0	0.0				
30	0.6474	2	1.7	**			
31	0.7036	0	0.0				
32	0.7599	3	2.6	***			
33	0.8162	0	0.0				
34	0.8725	0	0.0				
35	0.9288	0	0.0				
36	0.9850	0	0.0				
37	1.041	0	0.0				
38	1.098	1	0.9	*			
39	1.154	0	0.0				
40	1.210	0	0.0				
41	1.266	1	0.9	*			

T I T L E S	PROGRAM:	PERCENTAGE CUMULATIVE FREQUENCY PLOTS
	DATA:	DAISY CREEK SC HORIZON SOIL GEOCHEMISTRY
	RUN:	
	TIME:	25/03/83 14:08:04

NAME	AG-SC
N	116
X BAR	0.16896
STD DEV	0.22512

CELL	LOWER LIMIT	NMBR	%	ARITHMETIC VALUES											BAR INTERVAL = 0.12500	STD DEV	ARITH.LIMIT		
1	0.7177	5	4.3		*														
2	0.6896	2	6.0		*														
3	0.6614	0	6.0		*														
4	0.6333	0	6.0		*														
5	0.6051	0	6.0		*														
6	0.5770	0	6.0		*														
7	0.5489	0	6.0		*														
8	0.5207	0	6.0		*														
9	0.4926	0	6.0		*														
10	0.4644	0	6.0		*														
11	0.4363	0	6.0		*														
12	0.4082	0	6.0		*														
13	0.3800	7	12.1		*														
14	0.3519	0	12.1		*														
15	0.3237	0	12.1		*														
16	0.2956	0	12.1		*														
17	0.2675	0	12.1		*														
18	0.2393	0	12.1		*														
19	0.2112	0	12.1		*														
20	0.1830	0	12.1		*														
21	0.1549	0	12.1		*														
22	0.1268	0	12.1		*														
23	0.9861E-01	102	100.0															*	
24	0.7047E-01	0	100.0															*	
25	0.4233E-01	0	100.0															*	
26	0.1419E-01	0	100.0															*	
27	-0.1395E-01	0	100.0															*	
28	-0.4209E-01	0	100.0															*	
29	-0.7023E-01	0	100.0															*	
30	-0.9837E-01	0	100.0															*	
31	-0.1265	0	100.0															*	
32	-0.1547	0	100.0															*	
33	-0.1828	0	100.0															*	
34	-0.2109	0	100.0															*	
35	-0.2391	0	100.0															*	
36	-0.2672	0	100.0															*	
37	-0.2954	0	100.0															*	
38	-0.3235	0	100.0															*	
39	-0.3516	0	100.0															*	
40	-0.3798	0	100.0															*	
41		0	100.0															*	

- 243 -

T I T L E S	PROGRAM:	PERCENTAGE HISTOGRAMS
	DATA:	DAISY CREEK SC HORIZON SOIL GEOCHEMISTRY
	RUN:	
	TIME:	25/03/83 14:08:04

NAME	AG-SC
N	116
X BAR	-0.90616
STD DEV	0.26358

CELL	LOWER LIMIT	NMBR	%	LOGARITHMIC VALUES	BAR INTERVAL = 0.25000	STD DEV	ARITH. LIMIT
1		0	0.0				
2	-2.191	0	0.0				0.6440E-02
3	-2.125	0	0.0				0.7495E-02
4	-2.059	0	0.0				0.8723E-02
5	-1.993	0	0.0				0.1015E-01
6	-1.928	0	0.0				0.1182E-01
7	-1.862	0	0.0				0.1375E-01
8	-1.796	0	0.0				0.1601E-01
9	-1.730	0	0.0				0.1863E-01
10	-1.664	0	0.0				0.2168E-01
11	-1.598	0	0.0				0.2523E-01
12	-1.532	0	0.0				0.2937E-01
13	-1.466	0	0.0				0.3418E-01
14	-1.400	0	0.0				0.3978E-01
15	-1.334	0	0.0				0.4629E-01
16	-1.269	0	0.0				0.5388E-01
17	-1.203	0	0.0				0.6271E-01
18	-1.137	0	0.0				0.7298E-01
19	-1.071	0	0.0				0.8494E-01
20	-1.005	102	87.9	*****			0.9886E-01
21	-0.9391	0	0.0				0.1151
22	-0.8732	0	0.0				0.1339
23	-0.8073	0	0.0				0.1558
24	-0.7414	0	0.0				0.1814
25	-0.6755	0	0.0				0.2111
26	-0.6096	0	0.0				0.2457
27	-0.5437	0	0.0				0.2859
28	-0.4778	0	0.0				0.3328
29	-0.4119	7	6.0	*****			0.3873
30	-0.3460	0	0.0				0.4508
31	-0.2801	0	0.0				0.5246
32	-0.2143	2	1.7	**			0.6106
33	-0.1484	3	2.6	***			0.7106
34	-0.8247E-01	0	0.0				0.8271
35	-0.1657E-01	1	0.9	*			0.9626
36	0.4932E-01	0	0.0				1.120
37	0.1152	0	0.0				1.304
38	0.1811	1	0.9	*			1.517
39	0.2470	0	0.0				1.766
40	0.3129	0	0.0				2.055
41	0.3788	0	0.0				2.392

T I T L E S	PROGRAM:	PERCENTAGE CUMULATIVE FREQUENCY PLOTS
	DATA:	DAISY CREEK SC HORIZON SOIL GEOCHEMISTRY
	RUN:	
	TIME:	25/03/83 14:08:04

NAME	AG-SC
N	116
X BAR	-0.90616
STD DEV	0.26358

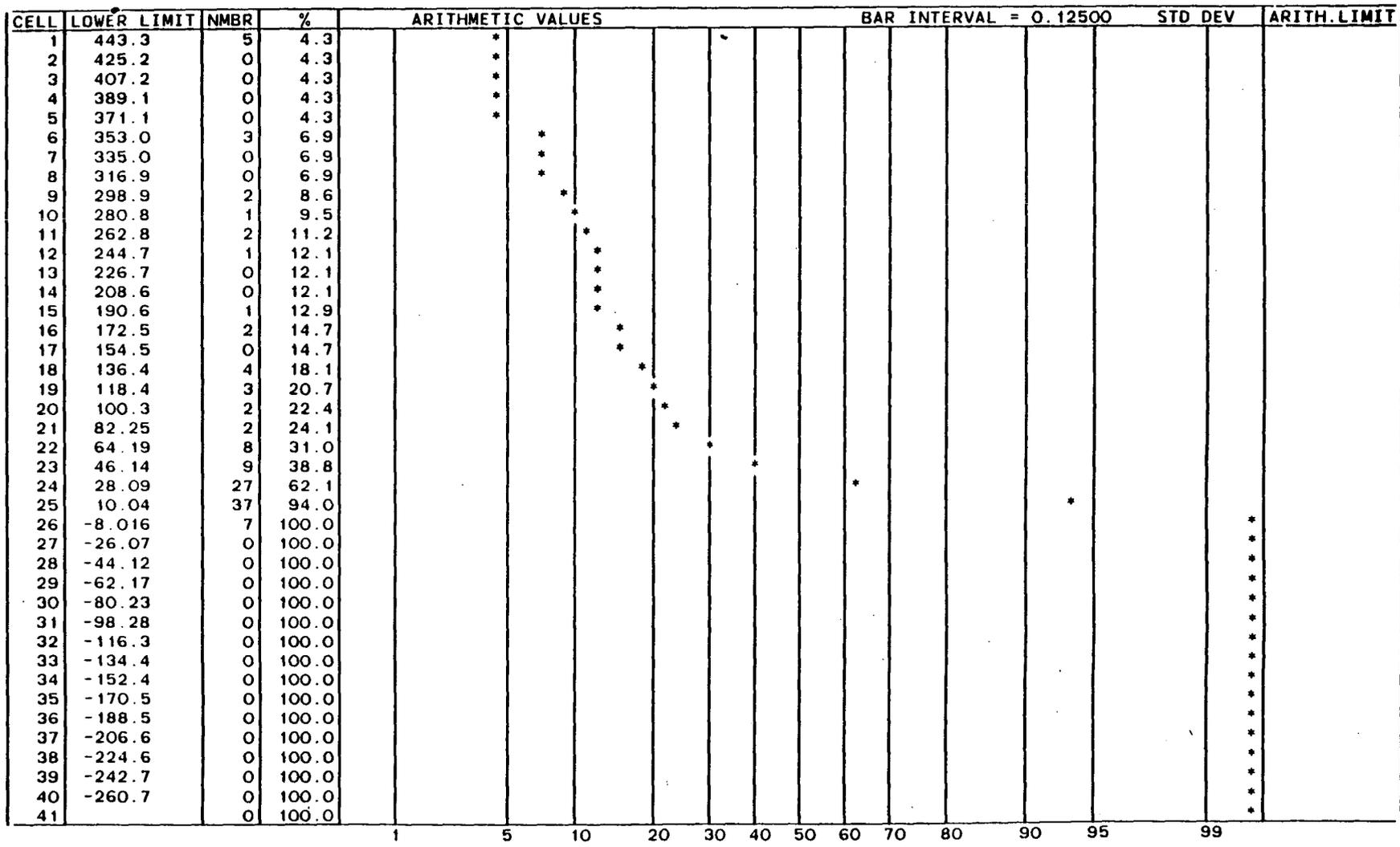
CELL	LOWER LIMIT	NMBR	%	LOGARITHMIC VALUES										BAR INTERVAL = 0.12500	STD DEV	ARITH. LIMIT	
1	-0.2637	7	6.0			*											0.5449
2	-0.2966	0	6.0			*											0.5051
3	-0.3296	0	6.0			*											0.4682
4	-0.3625	0	6.0			*											0.4340
5	-0.3955	0	6.0			*											0.4023
6	-0.4284	7	12.1			*											0.3729
7	-0.4614	0	12.1			*											0.3456
8	-0.4943	0	12.1			*											0.3204
9	-0.5273	0	12.1			*											0.2970
10	-0.5602	0	12.1			*											0.2753
11	-0.5932	0	12.1			*											0.2552
12	-0.6261	0	12.1			*											0.2365
13	-0.6590	0	12.1			*											0.2193
14	-0.6920	0	12.1			*											0.2032
15	-0.7249	0	12.1			*											0.1884
16	-0.7579	0	12.1			*											0.1746
17	-0.7908	0	12.1			*											0.1619
18	-0.8238	0	12.1			*											0.1500
19	-0.8567	0	12.1			*											0.1391
20	-0.8897	0	12.1			*											0.1289
21	-0.9226	0	12.1			*											0.1195
22	-0.9556	0	12.1			*											0.1108
23	-0.9885	0	12.1			*											0.1027
24	-1.021	102	100.0													*	0.9518E-01
25	-1.054	0	100.0													*	0.8822E-01
26	-1.087	0	100.0													*	0.8178E-01
27	-1.120	0	100.0													*	0.7580E-01
28	-1.153	0	100.0													*	0.7027E-01
29	-1.186	0	100.0													*	0.6513E-01
30	-1.219	0	100.0													*	0.6037E-01
31	-1.252	0	100.0													*	0.5596E-01
32	-1.285	0	100.0													*	0.5187E-01
33	-1.318	0	100.0													*	0.4808E-01
34	-1.351	0	100.0													*	0.4457E-01
35	-1.384	0	100.0													*	0.4131E-01
36	-1.417	0	100.0													*	0.3830E-01
37	-1.450	0	100.0													*	0.3550E-01
38	-1.483	0	100.0													*	0.3291E-01
39	-1.516	0	100.0													*	0.3050E-01
40	-1.549	0	100.0													*	0.2827E-01
41		0	100.0													*	

TITLE	PROGRAM:	PERCENTAGE HISTOGRAMS	NAME	CU-SC
	DATA:	DAISY CREEK SC HORIZON SOIL GEOCHEMISTRY	N	116
	RUN:		X BAR	91.272
	TIME:	25/03/83 14:08:04	STD DEV	144.42

CELL	LOWER LIMIT	NMBR	%	ARITHMETIC VALUES	BAR INTERVAL = 0.25000	STD DEV	ARITH.LIMIT
1		0	0.0				
2	-612.8	0	0.0				
3	-576.7	0	0.0				
4	-540.6	0	0.0				
5	-504.5	0	0.0				
6	-468.4	0	0.0				
7	-432.2	0	0.0				
8	-396.1	0	0.0				
9	-360.0	0	0.0				
10	-323.9	0	0.0				
11	-287.8	0	0.0				
12	-251.7	0	0.0				
13	-215.6	0	0.0				
14	-179.5	0	0.0				
15	-143.4	0	0.0				
16	-107.3	0	0.0				
17	-71.20	0	0.0				
18	-35.09	0	0.0				
19	1.012	66	56.9	*****			
20	37.12	20	17.2	*****			
21	73.22	5	4.3	****			
22	109.3	7	6.0	*****			
23	145.4	3	2.6	***			
24	181.5	1	0.9	*			
25	217.6	1	0.9	*			
26	253.7	3	2.6	***			
27	289.8	2	1.7	**			
28	326.0	0	0.0				
29	362.1	3	2.6	***			
30	398.2	0	0.0				
31	434.3	1	0.9	*			
32	470.4	0	0.0				
33	506.5	1	0.9	*			
34	542.6	1	0.9	*			
35	578.7	0	0.0				
36	614.8	1	0.9	*			
37	650.9	0	0.0				
38	687.0	0	0.0				
39	723.1	0	0.0				
40	759.2	0	0.0				
41	795.3	1	0.9	*			

T I T L E S	PROGRAM:	PERCENTAGE CUMULATIVE FREQUENCY PLOTS
	DATA:	DAISY CREEK SC HORIZON SOIL GEOCHEMISTRY
	RUN:	
	TIME:	25/03/83 14:08:04

NAME	CU-SC
N	116
X BAR	91.272
STD DEV	144.42



- 247 -

T I T L E S	PROGRAM: PERCENTAGE HISTOGRAMS
	DATA: DAISY CREEK SC HORIZON SOIL GEOCHEMISTRY
	RUN:
	TIME: 25/03/83 14:08:04

NAME	CU-SC
N	116
X BAR	1.6542
STD DEV	0.47411

CELL	LOWER LIMIT	NMBR	%	LOGARITHMIC VALUES	BAR INTERVAL = 0.25000	STD DEV	ARITH. LIMIT
1		0	0.0				
2	-0.6570	0	0.0				0.2203
3	-0.5385	0	0.0				0.2894
4	-0.4200	0	0.0				0.3802
5	-0.3014	0	0.0				0.4995
6	-0.1829	0	0.0				0.6563
7	-0.6439E-01	0	0.0				0.8622
8	0.5413E-01	0	0.0				1.133
9	0.1727	0	0.0				1.488
10	0.2912	1	0.9 *				1.955
11	0.4097	0	0.0				2.569
12	0.5282	0	0.0				3.375
13	0.6468	0	0.0				4.434
14	0.7653	0	0.0				5.825
15	0.8838	6	5.2 *****				7.653
16	1.002	0	0.0				10.05
17	1.121	9	7.8 *****				13.21
18	1.239	17	14.7 *****				17.35
19	1.358	11	9.5 *****				22.80
20	1.476	22	19.0 *****				29.95
21	1.595	10	8.6 *****				39.35
22	1.714	9	7.8 *****				51.70
23	1.832	4	3.4 ***				67.93
24	1.951	3	2.6 ***				89.24
25	2.069	7	6.0 *****				117.2
26	2.188	3	2.6 ***				154.0
27	2.306	1	0.9 *				202.4
28	2.425	5	4.3 *****				265.9
29	2.543	4	3.4 ***				349.3
30	2.662	2	1.7 **				458.9
31	2.780	1	0.9 *				602.9
32	2.899	1	0.9 *				792.1
33	3.017	0	0.0				1041.
34	3.136	0	0.0				1367.
35	3.254	0	0.0				1796.
36	3.373	0	0.0				2360.
37	3.491	0	0.0				3100.
38	3.610	0	0.0				4073.
39	3.728	0	0.0				5351.
40	3.847	0	0.0				7030.
41	3.966	0	0.0				9236.

T I T L E S	PROGRAM:	PERCENTAGE CUMULATIVE FREQUENCY PLOTS
	DATA:	DAISY CREEK SC HORIZON SOIL GEOCHEMISTRY
	RUN:	
	TIME:	25/03/83 14:08:04

NAME	CU-SC
N	116
X BAR	1.6542
STD DEV	0.4411

CELL	LOWER LIMIT	NMBR	%	LOGARITHMIC VALUES										BAR INTERVAL = 0.12500	STD DEV	ARITH. LIMIT		
1	2.810	1	0.9	*														645.5
2	2.751	1	1.7	*	*													563.1
3	2.691	2	3.4		*	*												491.3
4	2.632	1	4.3		*	*												428.6
5	2.573	0	4.3		*	*												374.0
6	2.514	3	6.9		*	*	*											326.3
7	2.454	3	9.5		*	*	*	*										284.6
8	2.395	2	11.2		*	*	*	*	*									248.3
9	2.336	1	12.1		*	*	*	*	*	*								216.7
10	2.277	1	12.9		*	*	*	*	*	*	*							189.0
11	2.217	2	14.7		*	*	*	*	*	*	*	*						164.9
12	2.158	2	16.4		*	*	*	*	*	*	*	*	*					143.9
13	2.099	4	19.8		*	*	*	*	*	*	*	*	*	*				125.5
14	2.039	2	21.6		*	*	*	*	*	*	*	*	*	*	*			109.5
15	1.980	1	22.4		*	*	*	*	*	*	*	*	*	*	*	*		95.54
16	1.921	2	24.1		*	*	*	*	*	*	*	*	*	*	*	*	*	83.35
17	1.862	2	25.9		*	*	*	*	*	*	*	*	*	*	*	*	*	72.72
18	1.802	6	31.0		*	*	*	*	*	*	*	*	*	*	*	*	*	63.45
19	1.743	0	31.0		*	*	*	*	*	*	*	*	*	*	*	*	*	55.35
20	1.684	9	38.8		*	*	*	*	*	*	*	*	*	*	*	*	*	48.29
21	1.625	2	40.5		*	*	*	*	*	*	*	*	*	*	*	*	*	42.13
22	1.565	3	43.1		*	*	*	*	*	*	*	*	*	*	*	*	*	36.76
23	1.506	9	50.9		*	*	*	*	*	*	*	*	*	*	*	*	*	32.07
24	1.447	13	62.1		*	*	*	*	*	*	*	*	*	*	*	*	*	27.98
25	1.388	11	71.6		*	*	*	*	*	*	*	*	*	*	*	*	*	24.41
26	1.328	0	71.6		*	*	*	*	*	*	*	*	*	*	*	*	*	21.30
27	1.269	17	86.2		*	*	*	*	*	*	*	*	*	*	*	*	*	18.58
28	1.210	0	86.2		*	*	*	*	*	*	*	*	*	*	*	*	*	16.21
29	1.151	9	94.0		*	*	*	*	*	*	*	*	*	*	*	*	*	14.14
30	1.091	0	94.0		*	*	*	*	*	*	*	*	*	*	*	*	*	12.34
31	1.032	0	94.0		*	*	*	*	*	*	*	*	*	*	*	*	*	10.76
32	0.9727	6	99.1		*	*	*	*	*	*	*	*	*	*	*	*	*	9.391
33	0.9135	0	99.1		*	*	*	*	*	*	*	*	*	*	*	*	*	8.193
34	0.8542	0	99.1		*	*	*	*	*	*	*	*	*	*	*	*	*	7.148
35	0.7949	0	99.1		*	*	*	*	*	*	*	*	*	*	*	*	*	6.236
36	0.7357	0	99.1		*	*	*	*	*	*	*	*	*	*	*	*	*	5.441
37	0.6764	0	99.1		*	*	*	*	*	*	*	*	*	*	*	*	*	4.747
38	0.6171	0	99.1		*	*	*	*	*	*	*	*	*	*	*	*	*	4.141
39	0.5579	0	99.1		*	*	*	*	*	*	*	*	*	*	*	*	*	3.613
40	0.4986	0	99.1		*	*	*	*	*	*	*	*	*	*	*	*	*	3.152
41		1	100.0		*	*	*	*	*	*	*	*	*	*	*	*	*	

- 249 -

T	PROGRAM:	PERCENTAGE HISTOGRAMS
I		
T		
L	DATA:	DAISY CREEK SC HORIZON SOIL GEOCHEMISTRY
E	RUN:	
S	TIME:	25/03/83 14:08:04

NAME	PB-SC
N	116
X BAR	43.966
STD DEV	154.17

CELL	LOWER LIMIT	NMBR	%	ARITHMETIC VALUES	BAR INTERVAL = 0.25000	STD DEV	ARITH.LIMIT
1		0	0.0				
2	-707.6	0	0.0				
3	-669.1	0	0.0				
4	-630.5	0	0.0				
5	-592.0	0	0.0				
6	-553.4	0	0.0				
7	-514.9	0	0.0				
8	-476.4	0	0.0				
9	-437.8	0	0.0				
10	-399.3	0	0.0				
11	-360.7	0	0.0				
12	-322.2	0	0.0				
13	-283.6	0	0.0				
14	-245.1	0	0.0				
15	-206.6	0	0.0				
16	-168.0	0	0.0				
17	-129.5	0	0.0				
18	-90.93	0	0.0				
19	-52.39	0	0.0				
20	-13.85	92	79.3	*****			
21	24.70	15	12.9	*****			
22	63.24	1	0.9	*			
23	101.8	1	0.9	*			
24	140.3	1	0.9	*			
25	178.9	1	0.9	*			
26	217.4	1	0.9	*			
27	255.9	0	0.0				
28	294.5	0	0.0				
29	333.0	1	0.9	*			
30	371.6	0	0.0				
31	410.1	0	0.0				
32	448.7	0	0.0				
33	487.2	0	0.0				
34	525.7	0	0.0				
35	564.3	0	0.0				
36	602.8	0	0.0				
37	641.4	1	0.9	*			
38	679.9	0	0.0				
39	718.5	0	0.0				
40	757.0	0	0.0				
41	795.5	2	1.7	**			

T I T L E S	PROGRAM: PERCENTAGE CUMULATIVE FREQUENCY PLOTS
	DATA: DAISY CREEK SC HORIZON SOIL GEDCHEMISTRY
	RUN:
	TIME: 25/03/83 14:08:04

NAME	PB-SC
N	116
X BAR	43.966
STD DEV	154.17

CELL	LOWER LIMIT	NMBR	%	ARITHMETIC VALUES										BAR INTERVAL = 0.12500	STD DEV	ARITH.LIMIT		
1	419.8	3	2.6		*													
2	400.5	0	2.6		*													
3	381.2	0	2.6		*													
4	361.9	0	2.6		*													
5	342.7	1	3.4		*													
6	323.4	0	3.4		*													
7	304.1	0	3.4		*													
8	284.9	0	3.4		*													
9	265.6	0	3.4		*													
10	246.3	0	3.4		*													
11	227.0	1	4.3		*													
12	207.8	1	5.2		*													
13	188.5	0	5.2		*													
14	169.2	0	5.2		*													
15	150.0	0	5.2		*													
16	130.7	1	6.0		*													
17	111.4	1	6.9		*													
18	92.14	0	6.9		*													
19	72.87	0	6.9		*													
20	53.60	4	10.3		*													
21	34.33	6	15.5		*													
22	15.06	15	28.4		*													
23	-4.212	83	100.0		*												*	
24	-23.48	0	100.0		*												*	
25	-42.75	0	100.0		*												*	
26	-62.03	0	100.0		*												*	
27	-81.30	0	100.0		*												*	
28	-100.6	0	100.0		*												*	
29	-119.8	0	100.0		*												*	
30	-139.1	0	100.0		*												*	
31	-158.4	0	100.0		*												*	
32	-177.7	0	100.0		*												*	
33	-196.9	0	100.0		*												*	
34	-216.2	0	100.0		*												*	
35	-235.5	0	100.0		*												*	
36	-254.7	0	100.0		*												*	
37	-274.0	0	100.0		*												*	
38	-293.3	0	100.0		*												*	
39	-312.6	0	100.0		*												*	
40	-331.8	0	100.0		*												*	
41		0	100.0		*												*	

- 251 -

T I T L E S	PROGRAM:	PERCENTAGE HISTOGRAMS
	DATA:	DAISY CREEK SC HORIZON SOIL GEOCHEMISTRY
	RUN:	
	TIME:	25/03/83 14:08:04

NAME	PB-SC
N	116
X BAR	0.94320
STD DEV	0.61899

CELL	LOWER LIMIT	NMBR	%	LOGARITHMIC VALUES	BAR INTERVAL = 0.25000	STD DEV	ARITH. LIMIT
1		0	0.0				
2	-2.074	0	0.0				0.8426E-02
3	-1.920	0	0.0				0.1203E-01
4	-1.765	0	0.0				0.1718E-01
5	-1.610	0	0.0				0.2454E-01
6	-1.455	0	0.0				0.3505E-01
7	-1.301	0	0.0				0.5005E-01
8	-1.146	0	0.0				0.7147E-01
9	-0.9911	0	0.0				0.1021
10	-0.8364	0	0.0				0.1458
11	-0.6816	0	0.0				0.2081
12	-0.5269	0	0.0				0.2972
13	-0.3721	0	0.0				0.4245
14	-0.2174	0	0.0				0.6062
15	-0.6265E-01	0	0.0				0.8657
16	0.9209E-01	0	0.0				1.236
17	0.2468	44	37.9	*****			1.765
18	0.4016	0	0.0				2.521
19	0.5563	18	15.5	*****			3.600
20	0.7111	0	0.0				5.141
21	0.8658	16	13.8	*****			7.342
22	1.021	0	0.0				10.49
23	1.175	14	12.1	*****			14.97
24	1.330	6	5.2	*****			21.38
25	1.485	4	3.4	***			30.54
26	1.640	5	4.3	****			43.61
27	1.794	1	0.9	*			62.27
28	1.949	1	0.9	*			88.93
29	2.104	1	0.9	*			127.0
30	2.259	2	1.7	**			181.4
31	2.413	1	0.9	*			259.0
32	2.568	0	0.0				369.9
33	2.723	1	0.9	*			528.2
34	2.878	1	0.9	*			754.3
35	3.032	1	0.9	*			1077.
36	3.187	0	0.0				1538.
37	3.342	0	0.0				2197.
38	3.497	0	0.0				3137.
39	3.651	0	0.0				4480.
40	3.806	0	0.0				6397.
41	3.961	0	0.0				9136.

T I T L E S	PROGRAM:	PERCENTAGE CUMULATIVE FREQUENCY PLOTS
	DATA:	DAISY CREEK SC HORIZON SOIL GEOCHEMISTRY
	RUN:	
	TIME:	25/03/83 14:08:04

NAME	PB-SC
N	116
X BAR	0.94320
STD DEV	0.61899

CELL	LOWER LIMIT	NMBR	%	LOGARITHMIC VALUES										BAR INTERVAL = 0.12500	STD DEV	ARITH. LIMIT	
1	2.452	4	3.4		*												283.1
2	2.375	1	4.3		*												236.9
3	2.297	1	5.2		*												198.3
4	2.220	0	5.2		*												165.9
5	2.142	1	6.0		*												138.8
6	2.065	0	6.0		*												116.2
7	1.988	1	6.9		*												97.21
8	1.910	0	6.9		*												81.35
9	1.833	1	7.8		*												68.07
10	1.756	2	9.5		*												56.96
11	1.678	2	11.2		*												47.67
12	1.601	4	14.7		*												39.89
13	1.523	1	15.5		*												33.38
14	1.446	2	17.2		*												27.93
15	1.369	4	20.7		*												23.37
16	1.291	9	28.4		*												19.56
17	1.214	0	28.4		*												16.37
18	1.137	5	32.8		*												13.70
19	1.059	0	32.8		*												11.46
20	0.9819	16	46.6		*												9.591
21	0.9045	0	46.6		*												8.026
22	0.8271	0	46.6		*												6.716
23	0.7498	0	46.6		*												5.620
24	0.6724	18	62.1		*												4.703
25	0.5950	0	62.1		*												3.936
26	0.5176	0	62.1		*												3.293
27	0.4403	0	62.1		*												2.756
28	0.3629	44	100.0		*												2.306
29	0.2855	0	100.0		*												1.930
30	0.2082	0	100.0		*												1.615
31	0.1308	0	100.0		*												1.351
32	0.5340E-01	0	100.0		*												1.131
33	-0.2397E-01	0	100.0		*												0.9463
34	-0.1013	0	100.0		*												0.7919
35	-0.1787	0	100.0		*												0.6627
36	-0.2561	0	100.0		*												0.5545
37	-0.3335	0	100.0		*												0.4640
38	-0.4108	0	100.0		*												0.3883
39	-0.4882	0	100.0		*												0.3249
40	-0.5656	0	100.0		*												0.2719
41		0	100.0		*												

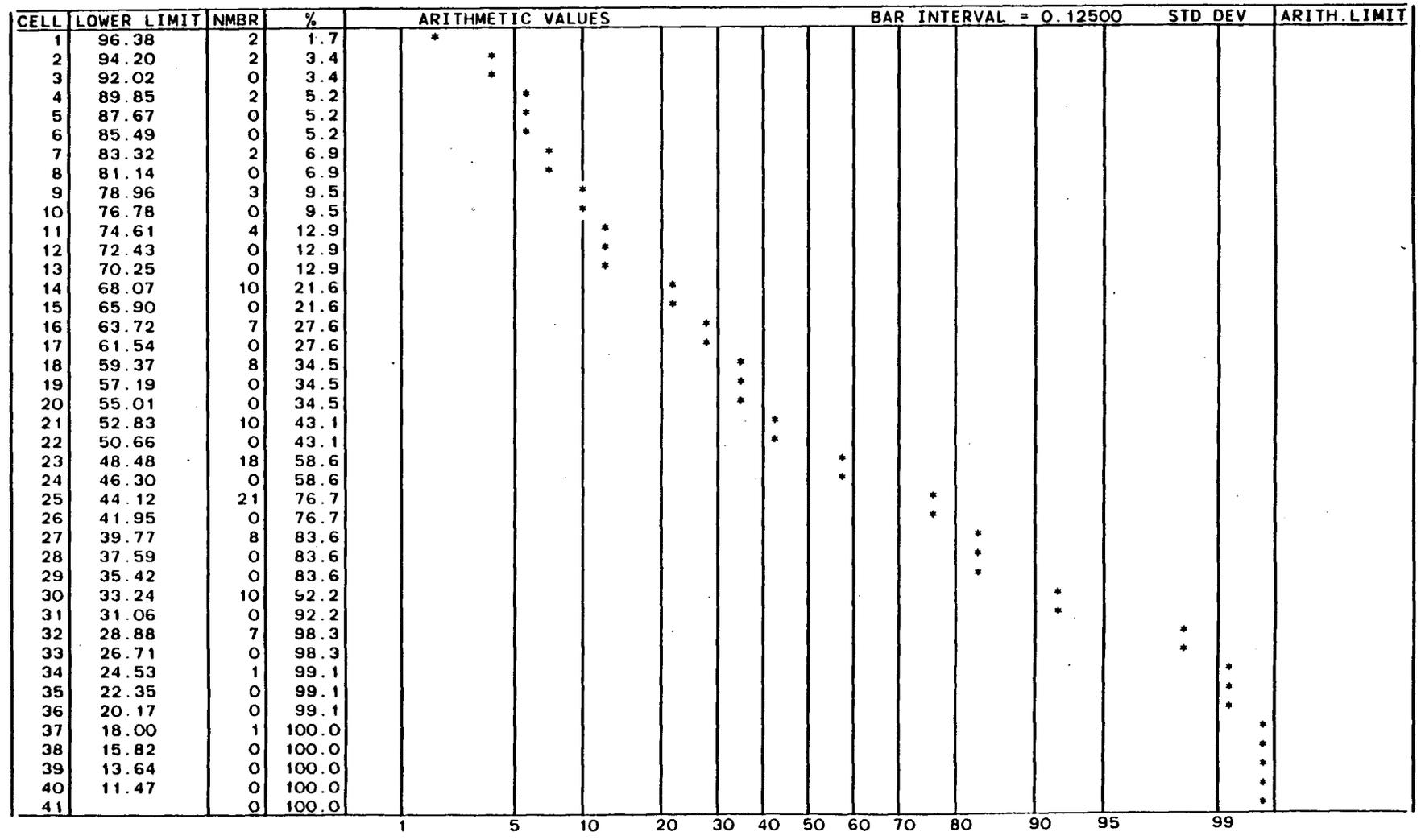
T I T L E S	PROGRAM:	PERCENTAGE HISTOGRAMS
	DATA:	DAISY CREEK SC HORIZON SOIL GEOCHEMISTRY
	RUN:	
	TIME:	25/03/83 14:08:04

NAME	ZN-SC
N	116
X BAR	53.922
STD DEV	17.418

CELL	LOWER LIMIT	NMBR	%	ARITHMETIC VALUES	BAR INTERVAL = 0.25000	STD DEV	ARITH.LIMIT
1		0	0.0				
2	-30.99	0	0.0				
3	-26.64	0	0.0				
4	-22.28	0	0.0				
5	-17.93	0	0.0				
6	-13.57	0	0.0				
7	-9.219	0	0.0				
8	-4.864	0	0.0				
9	-0.5094	0	0.0				
10	3.845	0	0.0				
11	8.200	0	0.0				
12	12.55	0	0.0				
13	16.91	1	0.9 *				
14	21.26	1	0.9 *				
15	25.62	0	0.0				
16	29.97	7	6.0 *****				
17	34.33	10	8.6 *****				
18	38.68	8	6.9 *****				
19	43.04	21	18.1 *****				
20	47.39	18	15.5 *****				
21	51.75	10	8.6 *****				
22	56.10	8	6.9 *****				
23	60.45	0	0.0				
24	64.81	7	6.0 *****				
25	69.16	10	8.6 *****				
26	73.52	4	3.4 ***				
27	77.87	3	2.6 ***				
28	82.23	2	1.7 **				
29	86.58	2	1.7 **				
30	90.94	2	1.7 **				
31	95.29	0	0.0				
32	99.64	0	0.0				
33	104.0	1	0.9 *				
34	108.4	0	0.0				
35	112.7	1	0.9 *				
36	117.1	0	0.0				
37	121.4	0	0.0				
38	125.8	0	0.0				
39	130.1	0	0.0				
40	134.5	0	0.0				
41	138.8	0	0.0				

T I T L E S	PROGRAM:	PERCENTAGE CUMULATIVE FREQUENCY PLOTS
	DATA:	DAISY CREEK SC HORIZON SOIL GEOCHEMISTRY
	RUN:	
	TIME:	25/03/83 14:08:04

NAME	ZN-SC
N	116
X BAR	53.922
STD DEV	17.418



T I T L E S	PROGRAM: PERCENTAGE HISTOGRAMS	NAME	ZN-SC
	DATA: DAISY CREEK SC HORIZON SOIL GEOCHEMISTRY	N	116
	RUN:	X BAR	1.7102
	TIME: 25/03/83 14:08:04	STD DEV	0.13782

CELL	LOWER LIMIT	NMBR	%	LOGARITHMIC VALUES	BAR INTERVAL = 0.25000	STD DEV	ARITH. LIMIT
1		0	0.0				
2	1.038	0	0.0				10.92
3	1.073	0	0.0				11.82
4	1.107	0	0.0				12.80
5	1.142	0	0.0				13.86
6	1.176	0	0.0				15.00
7	1.211	0	0.0				16.24
8	1.245	0	0.0				17.58
9	1.279	1	0.9 *				19.03
10	1.314	0	0.0				20.60
11	1.348	0	0.0				22.30
12	1.383	1	0.9 *				24.15
13	1.417	0	0.0				26.14
14	1.452	7	6.0 *****				28.30
15	1.486	0	0.0				30.63
16	1.521	10	8.6 *****				33.16
17	1.555	0	0.0				35.90
18	1.590	8	6.9 *****				38.86
19	1.624	21	18.1 *****				42.07
20	1.658	0	0.0				45.55
21	1.693	18	15.5 *****				49.31
22	1.727	10	8.6 *****				53.38
23	1.762	8	6.9 *****				57.79
24	1.796	7	6.0 *****				62.56
25	1.831	10	8.6 *****				67.72
26	1.865	4	3.4 ***				73.32
27	1.900	5	4.3 ****				79.37
28	1.934	2	1.7 **				85.92
29	1.969	2	1.7 **				93.02
30	2.003	1	0.9 *				100.7
31	2.037	1	0.9 *				109.0
32	2.072	0	0.0				118.0
33	2.106	0	0.0				127.8
34	2.141	0	0.0				138.3
35	2.175	0	0.0				149.7
36	2.210	0	0.0				162.1
37	2.244	0	0.0				175.5
38	2.279	0	0.0				190.0
39	2.313	0	0.0				205.6
40	2.348	0	0.0				222.6
41	2.382	0	0.0				241.0

T I T L E S	PROGRAM:	PERCENTAGE CUMULATIVE FREQUENCY PLOTS
	DATA:	DAISY CREEK SC HORIZON SOIL GEOCHEMISTRY
	RUN:	
	TIME:	25/03/83 14:08:04

NAME	ZN-SC
N	116
X BAR	1.7102
STD DEV	0.13782

CELL	LOWER LIMIT	NMBR	%	LOGARITHMIC VALUES										BAR INTERVAL = 0.12500	STD DEV	ARITH. LIMIT		
1	2.046	1	0.9	*														111.2
2	2.029	0	0.9	*														106.9
3	2.012	1	1.7		*													102.7
4	1.994	0	1.7		*													98.72
5	1.977	2	3.4			*												94.88
6	1.960	0	3.4			*												91.19
7	1.943	2	5.2			*												87.64
8	1.925	2	6.9			*	*											84.24
9	1.908	0	6.9			*	*											80.96
10	1.891	3	9.5			*												77.81
11	1.874	4	12.9			*	*											74.78
12	1.857	0	12.9			*	*											71.88
13	1.839	10	21.6			*	*											69.08
14	1.822	0	21.6			*	*											66.39
15	1.805	7	27.6			*	*	*										63.81
16	1.788	0	27.6			*	*	*										61.33
17	1.770	8	34.5			*	*	*	*									58.95
18	1.753	0	34.5			*	*	*	*									56.65
19	1.736	10	43.1			*	*	*	*	*								54.45
20	1.719	0	43.1			*	*	*	*	*								52.33
21	1.702	0	43.1			*	*	*	*	*								50.30
22	1.684	18	58.6			*	*	*	*	*	*							48.34
23	1.667	0	58.6			*	*	*	*	*	*							46.46
24	1.650	21	76.7			*	*	*	*	*	*	*						44.65
25	1.633	0	76.7			*	*	*	*	*	*	*						42.92
26	1.615	0	76.7			*	*	*	*	*	*	*						41.25
27	1.598	8	83.6			*	*	*	*	*	*	*	*					39.64
28	1.581	0	83.6			*	*	*	*	*	*	*	*					38.10
29	1.564	0	83.6			*	*	*	*	*	*	*	*					36.62
30	1.546	0	83.6			*	*	*	*	*	*	*	*					35.20
31	1.529	10	92.2			*	*	*	*	*	*	*	*	*				33.83
32	1.512	0	92.2			*	*	*	*	*	*	*	*	*				32.51
33	1.495	0	92.2			*	*	*	*	*	*	*	*	*				31.25
34	1.478	0	92.2			*	*	*	*	*	*	*	*	*				30.03
35	1.460	7	98.3			*	*	*	*	*	*	*	*	*	*			28.86
36	1.443	0	98.3			*	*	*	*	*	*	*	*	*	*			27.74
37	1.426	0	98.3			*	*	*	*	*	*	*	*	*	*			26.66
38	1.409	0	98.3			*	*	*	*	*	*	*	*	*	*			25.63
39	1.391	1	99.1			*	*	*	*	*	*	*	*	*	*	*		24.63
40	1.374	0	99.1			*	*	*	*	*	*	*	*	*	*	*		23.67
41		1	100.0			*	*	*	*	*	*	*	*	*	*	*	*	

057

T I T L E S	PROGRAM: PERCENTAGE HISTOGRAMS
	DATA: DAISY CREEK SC HORIZON SOIL GEOCHEMISTRY
	RUN:
	TIME: 25/03/83 14:08:04

NAME	CO-SC
N	116
X BAR	7.7155
STD DEV	3.8669

CELL	LOWER LIMIT	NMBR	%	ARITHMETIC VALUES	BAR INTERVAL = 0.25000	STD DEV	ARITH.LIMIT
1		0	0.0				
2	-11.14	0	0.0				
3	-10.17	0	0.0				
4	-9.202	0	0.0				
5	-8.236	0	0.0				
6	-7.269	0	0.0				
7	-6.302	0	0.0				
8	-5.335	0	0.0				
9	-4.369	0	0.0				
10	-3.402	0	0.0				
11	-2.435	0	0.0				
12	-1.468	0	0.0				
13	-0.5017	0	0.0				
14	0.4650	0	0.0				
15	1.432	0	0.0				
16	2.398	32	27.6	*****			
17	3.365	0	0.0				
18	4.332	13	11.2	*****			
19	5.299	0	0.0				
20	6.265	0	0.0				
21	7.232	0	0.0				
22	8.199	0	0.0				
23	9.166	63	54.3	*****			
24	10.13	0	0.0				
25	11.10	0	0.0				
26	12.07	0	0.0				
27	13.03	0	0.0				
28	14.00	0	0.0				
29	14.97	8	6.9	*****			
30	15.93	0	0.0				
31	16.90	0	0.0				
32	17.87	0	0.0				
33	18.83	0	0.0				
34	19.80	0	0.0				
35	20.77	0	0.0				
36	21.73	0	0.0				
37	22.70	0	0.0				
38	23.67	0	0.0				
39	24.63	0	0.0				
40	25.60	0	0.0				
41	26.57	0	0.0				

T I T L E S	PROGRAM:	PERCENTAGE HISTOGRAMS	NAME	CO-SC
	DATA:	DAISY CREEK SC HORIZON SOIL GEOCHEMISTRY	N	116
	RUN:		X BAR	0.81231
	TIME:	25/03/83 14:08:04	STD DEV	0.27882

CELL	LOWER LIMIT	NMBR	%	LOGARITHMIC VALUES	BAR INTERVAL = 0.25000	STD DEV	ARITH. LIMIT
1		0	0.0				
2	-0.5469	0	0.0				0.2838
3	-0.4772	0	0.0				0.3332
4	-0.4075	0	0.0				0.3913
5	-0.3378	0	0.0				0.4594
6	-0.2681	0	0.0				0.5394
7	-0.1984	0	0.0				0.6333
8	-0.1287	0	0.0				0.7435
9	-0.5901E-01	0	0.0				0.8730
10	0.1070E-01	0	0.0				1.025
11	0.8041E-01	0	0.0				1.203
12	0.1501	0	0.0				1.413
13	0.2198	0	0.0				1.659
14	0.2895	0	0.0				1.948
15	0.3592	32	27.6	*****			2.287
16	0.4289	0	0.0				2.685
17	0.4986	0	0.0				3.152
18	0.5683	0	0.0				3.701
19	0.6380	13	11.2	*****			4.346
20	0.7078	0	0.0				5.102
21	0.7775	0	0.0				5.990
22	0.8472	0	0.0				7.033
23	0.9169	0	0.0				8.258
24	0.9866	63	54.3	*****			9.696
25	1.056	0	0.0				11.38
26	1.126	8	6.9	*****			13.37
27	1.196	0	0.0				15.69
28	1.265	0	0.0				18.42
29	1.335	0	0.0				21.63
30	1.405	0	0.0				25.40
31	1.475	0	0.0				29.82
32	1.544	0	0.0				35.01
33	1.614	0	0.0				41.11
34	1.684	0	0.0				48.26
35	1.753	0	0.0				56.67
36	1.823	0	0.0				66.53
37	1.893	0	0.0				78.12
38	1.962	0	0.0				91.72
39	2.032	0	0.0				107.7
40	2.102	0	0.0				126.4
41	2.172	0	0.0				148.4

T I T L E S	PROGRAM:	PERCENTAGE CUMULATIVE FREQUENCY PLOTS
	DATA:	DAISY CREEK SC HORIZON SOIL GEOCHEMISTRY
	RUN:	
	TIME:	25/03/83 14:08:04

NAME	CO-SC
N	116
X BAR	0.81231
STD DEV	0.27882

CELL	LOWER LIMIT	NMBR	%	LOGARITHMIC VALUES										BAR INTERVAL = 0.12500	STD DEV	ARITH. LIMIT	
1	1.492	0	0.0	*													31.04
2	1.457	0	0.0	*													28.65
3	1.422	0	0.0	*													26.44
4	1.387	0	0.0	*													24.40
5	1.353	0	0.0	*													22.52
6	1.318	0	0.0	*													20.78
7	1.283	0	0.0	*													19.18
8	1.248	0	0.0	*													17.70
9	1.213	0	0.0	*													16.33
10	1.178	0	0.0	*													15.08
11	1.143	8	6.9	*	*												13.91
12	1.109	0	6.9	*	*												12.84
13	1.074	0	6.9	*	*												11.85
14	1.039	0	6.9	*	*												10.94
15	1.004	0	6.9	*	*												10.09
16	0.9691	63	61.2	*	*					*							9.314
17	0.9343	0	61.2	*	*					*							8.596
18	0.8994	0	61.2	*	*					*							7.933
19	0.8646	0	61.2	*	*					*							7.321
20	0.8297	0	61.2	*	*					*							6.757
21	0.7949	0	61.2	*	*					*							6.236
22	0.7600	0	61.2	*	*					*							5.755
23	0.7252	0	61.2	*	*					*							5.311
24	0.6903	13	72.4	*	*					*							4.901
25	0.6555	0	72.4	*	*					*							4.523
26	0.6206	0	72.4	*	*					*							4.175
27	0.5858	0	72.4	*	*					*							3.853
28	0.5509	0	72.4	*	*					*							3.556
29	0.5161	0	72.4	*	*					*							3.281
30	0.4812	0	72.4	*	*					*							3.028
31	0.4464	0	72.4	*	*					*							2.795
32	0.4115	0	72.4	*	*					*							2.579
33	0.3767	32	100.0	*	*					*							2.380
34	0.3418	0	100.0	*	*					*							2.197
35	0.3069	0	100.0	*	*					*							2.027
36	0.2721	0	100.0	*	*					*							1.871
37	0.2372	0	100.0	*	*					*							1.727
38	0.2024	0	100.0	*	*					*							1.594
39	0.1675	0	100.0	*	*					*							1.471
40	0.1327	0	100.0	*	*					*							1.357
41		0	100.0	*	*					*							*

176

T I T L E S	PROGRAM: PERCENTAGE HISTOGRAMS
	DATA: DAISY CREEK SC HORIZON SOIL GEOCHEMISTRY
	RUN:
	TIME: 25/03/83 14:08:04

NAME	MO-SC
N	116
X BAR	1.0000
STD DEV	0.0

CELL	LOWER LIMIT	NMBR	%	ARITHMETIC VALUES	BAR INTERVAL = 0.25000	STD DEV	ARITH.LIMIT
1		116	100.0	*****			
2	1.000	0	0.0				
3	1.000	0	0.0				
4	1.000	0	0.0				
5	1.000	0	0.0				
6	1.000	0	0.0				
7	1.000	0	0.0				
8	1.000	0	0.0				
9	1.000	0	0.0				
10	1.000	0	0.0				
11	1.000	0	0.0				
12	1.000	0	0.0				
13	1.000	0	0.0				
14	1.000	0	0.0				
15	1.000	0	0.0				
16	1.000	0	0.0				
17	1.000	0	0.0				
18	1.000	0	0.0				
19	1.000	0	0.0				
20	1.000	0	0.0				
21	1.000	0	0.0				
22	1.000	0	0.0				
23	1.000	0	0.0				
24	1.000	0	0.0				
25	1.000	0	0.0				
26	1.000	0	0.0				
27	1.000	0	0.0				
28	1.000	0	0.0				
29	1.000	0	0.0				
30	1.000	0	0.0				
31	1.000	0	0.0				
32	1.000	0	0.0				
33	1.000	0	0.0				
34	1.000	0	0.0				
35	1.000	0	0.0				
36	1.000	0	0.0				
37	1.000	0	0.0				
38	1.000	0	0.0				
39	1.000	0	0.0				
40	1.000	0	0.0				
41	1.000	0	0.0				

TITLE S	PROGRAM:	PERCENTAGE CUMULATIVE FREQUENCY PLOTS	NAME	MO-SC
	DATA:	DAISY CREEK SC HORIZON SOIL GEOCHEMISTRY	N	116
	RUN:		X BAR	1.0000
	TIME:	25/03/83 14:08:04	STD DEV	0.0

CELL	LOWER LIMIT	NMBR	%	ARITHMETIC VALUES										BAR INTERVAL = 0.12500	STD DEV	ARITH.LIMIT	
1	1.000	0	0.0	*													
2	1.000	0	0.0	*													
3	1.000	0	0.0	*													
4	1.000	0	0.0	*													
5	1.000	0	0.0	*													
6	1.000	0	0.0	*													
7	1.000	0	0.0	*													
8	1.000	0	0.0	*													
9	1.000	0	0.0	*													
10	1.000	0	0.0	*													
11	1.000	0	0.0	*													
12	1.000	0	0.0	*													
13	1.000	0	0.0	*													
14	1.000	0	0.0	*													
15	1.000	0	0.0	*													
16	1.000	0	0.0	*													
17	1.000	0	0.0	*													
18	1.000	0	0.0	*													
19	1.000	0	0.0	*													
20	1.000	0	0.0	*													
21	1.000	0	0.0	*													
22	1.000	0	0.0	*													
23	1.000	0	0.0	*													
24	1.000	0	0.0	*													
25	1.000	0	0.0	*													
26	1.000	0	0.0	*													
27	1.000	0	0.0	*													
28	1.000	0	0.0	*													
29	1.000	0	0.0	*													
30	1.000	0	0.0	*													
31	1.000	0	0.0	*													
32	1.000	0	0.0	*													
33	1.000	0	0.0	*													
34	1.000	0	0.0	*													
35	1.000	0	0.0	*													
36	1.000	0	0.0	*													
37	1.000	0	0.0	*													
38	1.000	0	0.0	*													
39	1.000	0	0.0	*													
40	1.000	0	0.0	*													
41	1.000	116	100.0	*													*
					1	5	10	20	30	40	50	60	70	80	90	95	99

T I T L E S	PROGRAM:	PERCENTAGE HISTOGRAMS
	DATA:	DAISY CREEK SC HORIZON SOIL GEOCHEMISTRY
	RUN:	
	TIME:	25/03/83 14:08:04

NAME	MO-SC
N	116
X BAR	0.0
STD DEV	0.0

CELL	LOWER LIMIT	NMBR	%	LOGARITHMIC VALUES	BAR INTERVAL = 0.25000	STD DEV	ARITH.LIMIT
1		116	100.0	*****			1.000
2		0	0.0				1.000
3	0.0	0	0.0				1.000
4	0.0	0	0.0				1.000
5	0.0	0	0.0				1.000
6	0.0	0	0.0				1.000
7	0.0	0	0.0				1.000
8	0.0	0	0.0				1.000
9	0.0	0	0.0				1.000
10	0.0	0	0.0				1.000
11	0.0	0	0.0				1.000
12	0.0	0	0.0				1.000
13	0.0	0	0.0				1.000
14	0.0	0	0.0				1.000
15	0.0	0	0.0				1.000
16	0.0	0	0.0				1.000
17	0.0	0	0.0				1.000
18	0.0	0	0.0				1.000
19	0.0	0	0.0				1.000
20	0.0	0	0.0				1.000
21	0.0	0	0.0				1.000
22	0.0	0	0.0				1.000
23	0.0	0	0.0				1.000
24	0.0	0	0.0				1.000
25	0.0	0	0.0				1.000
26	0.0	0	0.0				1.000
27	0.0	0	0.0				1.000
28	0.0	0	0.0				1.000
29	0.0	0	0.0				1.000
30	0.0	0	0.0				1.000
31	0.0	0	0.0				1.000
32	0.0	0	0.0				1.000
33	0.0	0	0.0				1.000
34	0.0	0	0.0				1.000
35	0.0	0	0.0				1.000
36	0.0	0	0.0				1.000
37	0.0	0	0.0				1.000
38	0.0	0	0.0				1.000
39	0.0	0	0.0				1.000
40	0.0	0	0.0				1.000
41	0.0	0	0.0				1.000

T I T L E S	PROGRAM: PERCENTAGE CUMULATIVE FREQUENCY PLOTS
	DATA: DAISY CREEK SC HORIZON SOIL GEOCHEMISTRY
	RUN:
	TIME: 25/03/83 14:08:04

NAME	MO-SC
N	116
X BAR	0.0
STD DEV	0.0

CELL	LOWER LIMIT	NMBR	%	LOGARITHMIC VALUES										BAR INTERVAL = 0.12500	STD DEV	ARITH. LIMIT	
1	0.0	0	0.0	*													1.000
2	0.0	0	0.0	*													1.000
3	0.0	0	0.0	*													1.000
4	0.0	0	0.0	*													1.000
5	0.0	0	0.0	*													1.000
6	0.0	0	0.0	*													1.000
7	0.0	0	0.0	*													1.000
8	0.0	0	0.0	*													1.000
9	0.0	0	0.0	*													1.000
10	0.0	0	0.0	*													1.000
11	0.0	0	0.0	*													1.000
12	0.0	0	0.0	*													1.000
13	0.0	0	0.0	*													1.000
14	0.0	0	0.0	*													1.000
15	0.0	0	0.0	*													1.000
16	0.0	0	0.0	*													1.000
17	0.0	0	0.0	*													1.000
18	0.0	0	0.0	*													1.000
19	0.0	0	0.0	*													1.000
20	0.0	0	0.0	*													1.000
21	0.0	0	0.0	*													1.000
22	0.0	0	0.0	*													1.000
23	0.0	0	0.0	*													1.000
24	0.0	0	0.0	*													1.000
25	0.0	0	0.0	*													1.000
26	0.0	0	0.0	*													1.000
27	0.0	0	0.0	*													1.000
28	0.0	0	0.0	*													1.000
29	0.0	0	0.0	*													1.000
30	0.0	0	0.0	*													1.000
31	0.0	0	0.0	*													1.000
32	0.0	0	0.0	*													1.000
33	0.0	0	0.0	*													1.000
34	0.0	0	0.0	*													1.000
35	0.0	0	0.0	*													1.000
36	0.0	0	0.0	*													1.000
37	0.0	0	0.0	*													1.000
38	0.0	0	0.0	*													1.000
39	0.0	0	0.0	*													1.000
40	0.0	0	0.0	*													1.000
41		116	100.0	*													1.000
					1	5	10	20	30	40	50	60	70	80	90	95	99

T	PROGRAM:	PERCENTAGE HISTOGRAMS
I		
T		
L	DATA:	DAISY CREEK SC HORIZON SOIL GEOCHEMISTRY
E	RUN:	
S	TIME:	25/03/83 14:08:04

NAME	AS-SC
N	116
X BAR	5.7328
STD DEV	2.9675

CELL	LOWER LIMIT	NMBR	%	ARITHMETIC VALUES	BAR INTERVAL = 0.25000	STD DEV	ARITH.LIMIT
1		0	0.0				
2	-8.734	0	0.0				
3	-7.992	0	0.0				
4	-7.250	0	0.0				
5	-6.508	0	0.0				
6	-5.766	0	0.0				
7	-5.025	0	0.0				
8	-4.283	0	0.0				
9	-3.541	0	0.0				
10	-2.799	0	0.0				
11	-2.057	0	0.0				
12	-1.315	0	0.0				
13	-0.5732	0	0.0				
14	0.1686	0	0.0				
15	0.9105	0	0.0				
16	1.652	0	0.0				
17	2.394	0	0.0				
18	3.136	0	0.0				
19	3.878	0	0.0				
20	4.620	105	90.5	*****			
21	5.362	0	0.0				
22	6.104	0	0.0				
23	6.846	0	0.0				
24	7.587	0	0.0				
25	8.329	0	0.0				
26	9.071	0	0.0				
27	9.813	9	7.8	*****			
28	10.55	0	0.0				
29	11.30	0	0.0				
30	12.04	0	0.0				
31	12.78	0	0.0				
32	13.52	0	0.0				
33	14.26	0	0.0				
34	15.01	0	0.0				
35	15.75	0	0.0				
36	16.49	0	0.0				
37	17.23	0	0.0				
38	17.97	0	0.0				
39	18.72	0	0.0				
40	19.46	1	0.9	*			
41	20.20	1	0.9	*			

T I T L E S	PROGRAM:	PERCENTAGE CUMULATIVE FREQUENCY PLOTS
	DATA:	DAISY CREEK SC HORIZON SOIL GEOCHEMISTRY
	RUN:	
	TIME:	25/03/83 14:08:04

NAME	AS-SC
N	116
X BAR	5.7328
STD DEV	2.9675

CELL	LOWER LIMIT	NMBR	%	ARITHMETIC VALUES										BAR INTERVAL = 0.12500	STD DEV	ARITH.LIMIT	
1	12.97	2	1.7	*													
2	12.60	0	1.7	*													
3	12.22	0	1.7	*													
4	11.85	0	1.7	*													
5	11.48	0	1.7	*													
6	11.11	0	1.7	*													
7	10.74	0	1.7	*													
8	10.37	0	1.7	*													
9	9.999	9	9.5	*													
10	9.628	0	9.5	*													
11	9.257	0	9.5	*													
12	8.886	0	9.5	*													
13	8.515	0	9.5	*													
14	8.144	0	9.5	*													
15	7.773	0	9.5	*													
16	7.402	0	9.5	*													
17	7.031	0	9.5	*													
18	6.660	0	9.5	*													
19	6.289	0	9.5	*													
20	5.918	0	9.5	*													
21	5.547	0	9.5	*													
22	5.176	0	9.5	*													
23	4.805	105	100.0	*													*
24	4.434	0	100.0	*													*
25	4.064	0	100.0	*													*
26	3.693	0	100.0	*													*
27	3.322	0	100.0	*													*
28	2.951	0	100.0	*													*
29	2.580	0	100.0	*													*
30	2.209	0	100.0	*													*
31	1.838	0	100.0	*													*
32	1.467	0	100.0	*													*
33	1.096	0	100.0	*													*
34	0.7251	0	100.0	*													*
35	0.3541	0	100.0	*													*
36	-0.1683E-01	0	100.0	*													*
37	-0.3878	0	100.0	*													*
38	-0.7587	0	100.0	*													*
39	-1.130	0	100.0	*													*
40	-1.501	0	100.0	*													*
41		0	100.0	*													*

1 5 10 20 30 40 50 60 70 80 90 95 99

T I T L E S	PROGRAM:	PERCENTAGE HISTOGRAMS
	DATA:	DAISY CREEK SC HORIZON SOIL GEOCHEMISTRY
	RUN:	
	TIME:	25/03/83 14:08:04

NAME	AS-SC
N	116
X BAR	0.73422
STD DEV	0.11944

CELL	LOWER LIMIT	NMBR	%	LOGARITHMIC VALUES	BAR INTERVAL = 0.25000	STD DEV	ARITH. LIMIT
1		0	0.0				1.419
2	0.1519	0	0.0				1.520
3	0.1818	0	0.0				1.628
4	0.2117	0	0.0				1.744
5	0.2415	0	0.0				1.868
6	0.2714	0	0.0				2.001
7	0.3012	0	0.0				2.143
8	0.3311	0	0.0				2.296
9	0.3610	0	0.0				2.459
10	0.3908	0	0.0				2.634
11	0.4207	0	0.0				2.822
12	0.4505	0	0.0				3.023
13	0.4804	0	0.0				3.238
14	0.5103	0	0.0				3.468
15	0.5401	0	0.0				3.715
16	0.5700	0	0.0				3.980
17	0.5998	0	0.0				4.263
18	0.6297	0	0.0				4.566
19	0.6596	0	0.0				4.891
20	0.6894	105	90.5	*****			5.239
21	0.7193	0	0.0				5.612
22	0.7491	0	0.0				6.012
23	0.7790	0	0.0				6.440
24	0.8089	0	0.0				6.898
25	0.8387	0	0.0				7.389
26	0.8686	0	0.0				7.915
27	0.8984	0	0.0				8.478
28	0.9283	0	0.0				9.082
29	0.9582	0	0.0				9.728
30	0.9880	9	7.8	*****			10.42
31	1.018	0	0.0				11.16
32	1.048	0	0.0				11.96
33	1.078	0	0.0				12.81
34	1.107	0	0.0				13.72
35	1.137	0	0.0				14.70
36	1.167	0	0.0				15.74
37	1.197	0	0.0				16.86
38	1.227	0	0.0				18.06
39	1.257	0	0.0				19.35
40	1.287	1	0.9	*			20.72
41	1.316	1	0.9	*			

T I T L E S	PROGRAM:	PERCENTAGE CUMULATIVE FREQUENCY PLOTS
	DATA:	DAISY CREEK SC HORIZON SOIL GEOCHEMISTRY
	RUN:	
	TIME:	25/03/83 14:08:04

NAME	AS-SC
N	116
X BAR	0.73422
STD DEV	0.11944

CELL	LOWER LIMIT	NMBR	%	LOGARITHMIC VALUES										BAR INTERVAL = 0.12500	STD DEV	ARITH. LIMIT	
1	1.025	2	1.7	*													10.60
2	1.010	0	1.7	*													10.24
3	0.9955	9	9.5	*													9.897
4	0.9806	0	9.5	*													9.562
5	0.9656	0	9.5	*													9.239
6	0.9507	0	9.5	*													8.927
7	0.9358	0	9.5	*													8.625
8	0.9208	0	9.5	*													8.334
9	0.9059	0	9.5	*													8.052
10	0.8910	0	9.5	*													7.780
11	0.8760	0	9.5	*													7.517
12	0.8611	0	9.5	*													7.263
13	0.8462	0	9.5	*													7.018
14	0.8313	0	9.5	*													6.780
15	0.8163	0	9.5	*													6.551
16	0.8014	0	9.5	*													6.330
17	0.7865	0	9.5	*													6.116
18	0.7715	0	9.5	*													5.909
19	0.7566	0	9.5	*													5.710
20	0.7417	0	9.5	*													5.517
21	0.7267	0	9.5	*													5.330
22	0.7118	0	9.5	*													5.150
23	0.6969	105	100.0												*		4.976
24	0.6820	0	100.0												*		4.808
25	0.6670	0	100.0												*		4.645
26	0.6521	0	100.0												*		4.488
27	0.6372	0	100.0												*		4.337
28	0.6222	0	100.0												*		4.190
29	0.6073	0	100.0												*		4.049
30	0.5924	0	100.0												*		3.912
31	0.5775	0	100.0												*		3.780
32	0.5625	0	100.0												*		3.652
33	0.5476	0	100.0												*		3.529
34	0.5327	0	100.0												*		3.409
35	0.5177	0	100.0												*		3.294
36	0.5028	0	100.0												*		3.183
37	0.4879	0	100.0												*		3.075
38	0.4729	0	100.0												*		2.971
39	0.4580	0	100.0												*		2.871
40	0.4431	0	100.0												*		2.774
41		0	100.0												*		

T I T L E S	PROGRAM:	PERCENTAGE HISTOGRAMS	NAME	HG-SC
	DATA:	DAISY CREEK SC HORIZON SOIL GEOCHEMISTRY	N	116
	RUN:		X BAR	29.310
	TIME:	25/03/83 14:08:04	STD DEV	12.767

CELL	LOWER LIMIT	NMBR	%	ARITHMETIC VALUES	BAR INTERVAL = 0.25000	STD DEV	ARITH.LIMIT
1		0	0.0				
2	-32.93	0	0.0				
3	-29.74	0	0.0				
4	-26.55	0	0.0				
5	-23.35	0	0.0				
6	-20.16	0	0.0				
7	-16.97	0	0.0				
8	-13.78	0	0.0				
9	-10.59	0	0.0				
10	-7.395	0	0.0				
11	-4.203	0	0.0				
12	-1.011	0	0.0				
13	2.180	0	0.0				
14	5.372	0	0.0				
15	8.564	15	12.9	*****			
16	11.76	0	0.0				
17	14.95	0	0.0				
18	18.14	33	28.4	*****			
19	21.33	0	0.0				
20	24.52	0	0.0				
21	27.71	31	26.7	*****			
22	30.91	0	0.0				
23	34.10	0	0.0				
24	37.29	21	18.1	*****			
25	40.48	0	0.0				
26	43.67	0	0.0				
27	46.87	14	12.1	*****			
28	50.06	0	0.0				
29	53.25	0	0.0				
30	56.44	0	0.0				
31	59.63	2	1.7	**			
32	62.82	0	0.0				
33	66.02	0	0.0				
34	69.21	0	0.0				
35	72.40	0	0.0				
36	75.59	0	0.0				
37	78.78	0	0.0				
38	81.97	0	0.0				
39	85.17	0	0.0				
40	88.36	0	0.0				
41	91.55	0	0.0				

T I T L E S	PROGRAM:	PERCENTAGE CUMULATIVE FREQUENCY PLOTS
	DATA:	DAISY CREEK SC HORIZON SOIL GEOCHEMISTRY
	RUN:	
	TIME:	25/03/83 14:08:04

NAME	HG-SC
N	116
X BAR	29.310
STD DEV	12.767

CELL	LOWER LIMIT	NMBR	%	ARITHMETIC VALUES										BAR INTERVAL = 0.12500	STD DEV	ARITH.LIMIT		
1	60.43	0	0.0	*														
2	58.83	2	1.7	*														
3	57.24	0	1.7	*														
4	55.64	0	1.7	*														
5	54.05	0	1.7	*														
6	52.45	0	1.7	*														
7	50.85	0	1.7	*														
8	49.26	14	13.8	*														
9	47.66	0	13.8	*														
10	46.07	0	13.8	*														
11	44.47	0	13.8	*														
12	42.88	0	13.8	*														
13	41.28	0	13.8	*														
14	39.68	21	31.9	*														
15	38.09	0	31.9	*														
16	36.49	0	31.9	*														
17	34.90	0	31.9	*														
18	33.30	0	31.9	*														
19	31.70	0	31.9	*														
20	30.11	0	31.9	*														
21	28.51	31	58.6	*														
22	26.92	0	58.6	*														
23	25.32	0	58.6	*														
24	23.72	0	58.6	*														
25	22.13	0	58.6	*														
26	20.53	0	58.6	*														
27	18.94	33	87.1	*														
28	17.34	0	87.1	*														
29	15.75	0	87.1	*														
30	14.15	0	87.1	*														
31	12.55	0	87.1	*														
32	10.96	0	87.1	*														
33	9.362	15	100.0	*														
34	7.766	0	100.0	*														
35	6.170	0	100.0	*														
36	4.574	0	100.0	*														
37	2.978	0	100.0	*														
38	1.382	0	100.0	*														
39	-0.2136	0	100.0	*														
40	-1.809	0	100.0	*														
41		0	100.0	*														

1 5 10 20 30 40 50 60 70 80 90 95 99

T I T L E S	PROGRAM: PERCENTAGE HISTOGRAMS
	DATA: DAISY CREEK SC HORIZON SOIL GEOCHEMISTRY
	RUN:
	TIME: 25/03/83 14:08:04

NAME	HG-SC
N	116
X BAR	1.4199
STD DEV	0.21383

CELL	LOWER LIMIT	NMBR	%	LOGARITHMIC VALUES	BAR INTERVAL = 0.25000	STD DEV	ARITH. LIMIT
1		0	0.0				
2	0.3775	0	0.0				2.385
3	0.4309	0	0.0				2.697
4	0.4844	0	0.0				3.051
5	0.5379	0	0.0				3.450
6	0.5913	0	0.0				3.902
7	0.6448	0	0.0				4.413
8	0.6982	0	0.0				4.992
9	0.7517	0	0.0				5.645
10	0.8051	0	0.0				6.385
11	0.8586	0	0.0				7.221
12	0.9121	0	0.0				8.167
13	0.9655	15	12.9	*****			9.237
14	1.019	0	0.0				10.45
15	1.072	0	0.0				11.81
16	1.126	0	0.0				13.36
17	1.179	0	0.0				15.11
18	1.233	0	0.0				17.09
19	1.286	33	28.4	*****			19.33
20	1.340	0	0.0				21.86
21	1.393	0	0.0				24.73
22	1.447	31	26.7	*****			27.97
23	1.500	0	0.0				31.63
24	1.554	21	18.1	*****			35.77
25	1.607	0	0.0				40.46
26	1.660	14	12.1	*****			45.76
27	1.714	0	0.0				51.75
28	1.767	2	1.7	**			58.53
29	1.821	0	0.0				66.20
30	1.874	0	0.0				74.87
31	1.928	0	0.0				84.67
32	1.981	0	0.0				95.76
33	2.035	0	0.0				108.3
34	2.088	0	0.0				122.5
35	2.142	0	0.0				138.5
36	2.195	0	0.0				156.7
37	2.248	0	0.0				177.2
38	2.302	0	0.0				200.4
39	2.355	0	0.0				226.7
40	2.409	0	0.0				256.4
41	2.462	0	0.0				289.9

T I T L E S	PROGRAM:	PERCENTAGE CUMULATIVE FREQUENCY PLOTS
	DATA:	DAISY CREEK SC HORIZON SOIL GEOCHEMISTRY
	RUN:	
	TIME:	25/03/83 14:08:04

NAME	HG-SC
N	116
X BAR	1.4199
STD DEV	0.21383

CELL	LOWER LIMIT	NMBR	%	LOGARITHMIC VALUES										BAR INTERVAL = 0.12500	STD DEV	ARITH. LIMIT	
1	1.941	0	0.0	*													87.32
2	1.914	0	0.0	*													82.11
3	1.888	0	0.0	*													77.21
4	1.861	0	0.0	*													72.60
5	1.834	0	0.0	*													68.27
6	1.807	0	0.0	*													64.19
7	1.781	0	0.0	*													60.36
8	1.754	2	1.7	*	*												56.76
9	1.727	0	1.7	*	*												53.37
10	1.701	0	1.7	*	*												50.18
11	1.674	14	13.8			*											47.19
12	1.647	0	13.8			*											44.37
13	1.620	0	13.8			*											41.72
14	1.594	21	31.9				*										39.23
15	1.567	0	31.9				*										36.89
16	1.540	0	31.9				*										34.69
17	1.513	0	31.9				*										32.62
18	1.487	0	31.9				*										30.67
19	1.460	31	58.6					*									28.84
20	1.433	0	58.6					*									27.12
21	1.407	0	58.6					*									25.50
22	1.380	0	58.6					*									23.98
23	1.353	0	58.6					*									22.55
24	1.326	0	58.6					*									21.20
25	1.300	33	87.1						*								19.94
26	1.273	0	87.1						*								18.75
27	1.246	0	87.1						*								17.63
28	1.219	0	87.1						*								16.57
29	1.193	0	87.1						*								15.59
30	1.166	0	87.1						*								14.65
31	1.139	0	87.1						*								13.78
32	1.113	0	87.1						*								12.96
33	1.086	0	87.1						*								12.18
34	1.059	0	87.1						*								11.46
35	1.032	0	87.1						*								10.77
36	1.006	0	87.1						*								10.13
37	0.9789	15	100.0						*								9.525
38	0.9522	0	100.0						*								8.957
39	0.9254	0	100.0						*								8.422
40	0.8987	0	100.0						*								7.919
41		0	100.0						*								

- 273 -

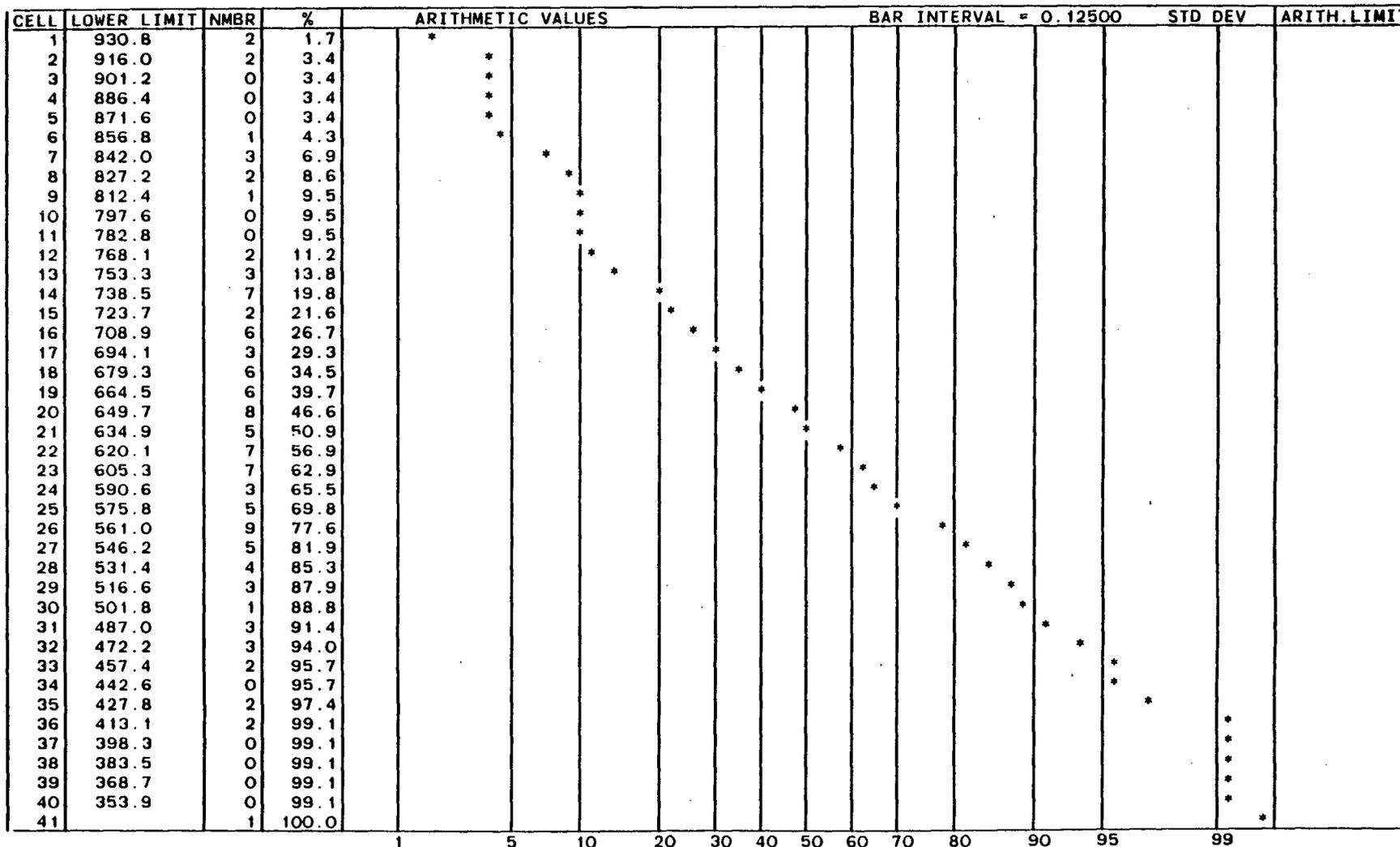
T I T L E S	PROGRAM:	PERCENTAGE HISTOGRAMS
	DATA:	DAISY CREEK SC HORIZON SOIL GEOCHEMISTRY
	RUN:	
	TIME:	25/03/83 14:08:04

NAME	BA-SC
N	116
X BAR	642.33
STD DEV	118.33

CELL	LOWER LIMIT	NMBR	%	ARITHMETIC VALUES	BAR INTERVAL = 0.25000	STD DEV	ARITH.LIMIT
1		0	0.0				
2	65.45	0	0.0				
3	95.04	0	0.0				
4	124.6	0	0.0				
5	154.2	0	0.0				
6	183.8	0	0.0				
7	213.4	1	0.9 *				
8	243.0	0	0.0				
9	272.5	0	0.0				
10	302.1	0	0.0				
11	331.7	0	0.0				
12	361.3	0	0.0				
13	390.9	2	1.7 **				
14	420.5	2	1.7 **				
15	450.0	2	1.7 **				
16	479.6	6	5.2 *****				
17	509.2	4	3.4 ***				
18	538.8	9	7.8 *****				
19	568.4	14	12.1 *****				
20	598.0	10	8.6 *****				
21	627.5	19	16.4 *****				
22	657.1	9	7.8 *****				
23	686.7	13	11.2 *****				
24	716.3	4	3.4 ***				
25	745.9	8	6.9 *****				
26	775.5	2	1.7 **				
27	805.0	3	2.6 ***				
28	834.6	3	2.6 ***				
29	864.2	1	0.9 *				
30	893.8	1	0.9 *				
31	923.4	3	2.6 ***				
32	953.0	0	0.0				
33	982.5	0	0.0				
34	1012.	0	0.0				
35	1042.	0	0.0				
36	1071.	0	0.0				
37	1101.	0	0.0				
38	1130.	0	0.0				
39	1160.	0	0.0				
40	1190.	0	0.0				
41	1219.	0	0.0				

T I T L E S	PROGRAM:	PERCENTAGE CUMULATIVE FREQUENCY PLOTS
	DATA:	DAISY CREEK SC HORIZON SOIL GEOCHEMISTRY
	RUN:	
	TIME:	25/03/83 14:08:04

NAME	BA-SC
N	116
X BAR	642.33
STD DEV	118.33

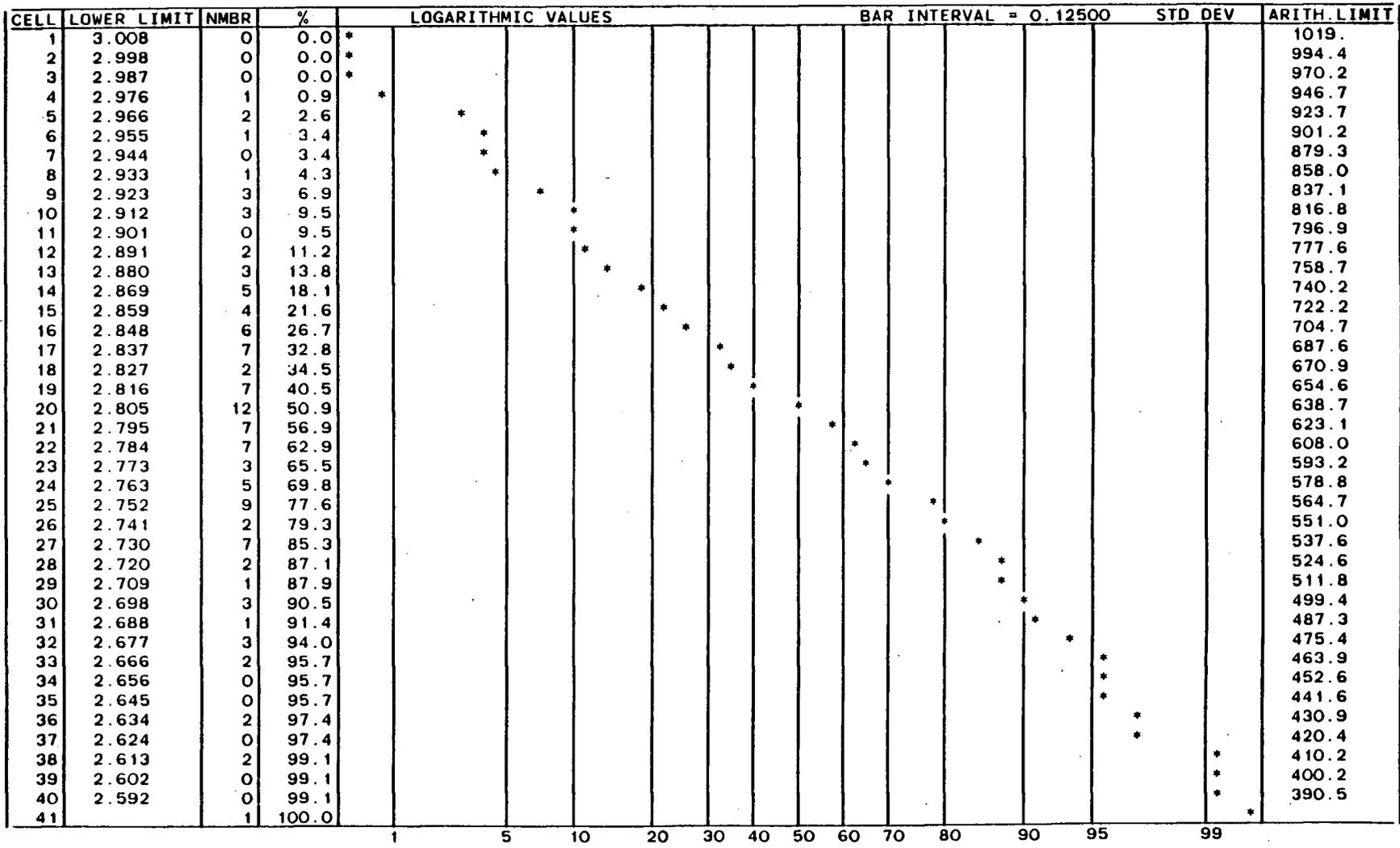


T I T L E S	PROGRAM: PERCENTAGE HISTOGRAMS	NAME	BA-SC
	DATA: DAISY CREEK SC HORIZON SOIL GEOCHEMISTRY	N	116
	RUN:	X BAR	2.7999
	TIME: 25/03/83 14:08:04	STD DEV	0.85470E-01

CELL	LOWER LIMIT	NMBR	%	LOGARITHMIC VALUES	BAR INTERVAL = 0.25000	STD DEV	PRITH. LIMIT
1		1	0.9	*			
2	2.383	0	0.0				241.7
3	2.405	0	0.0				253.9
4	2.426	0	0.0				266.7
5	2.447	0	0.0				280.1
6	2.469	0	0.0				294.3
7	2.490	0	0.0				309.1
8	2.511	0	0.0				324.7
9	2.533	0	0.0				341.1
10	2.554	0	0.0				358.3
11	2.576	0	0.0				376.3
12	2.597	0	0.0				395.3
13	2.618	2	1.7	**			415.3
14	2.640	2	1.7	**			436.2
15	2.661	5	4.3	****			458.2
16	2.682	3	2.6	***			481.3
17	2.704	4	3.4	***			505.6
18	2.725	7	6.0	*****			531.1
19	2.747	12	10.3	*****			557.8
20	2.768	11	9.5	*****			586.0
21	2.789	15	12.9	*****			615.5
22	2.811	14	12.1	*****			646.6
23	2.832	15	12.9	*****			679.2
24	2.853	4	3.4	***			713.4
25	2.875	10	8.6	*****			749.4
26	2.896	1	0.9	*			787.2
27	2.917	5	4.3	****			826.9
28	2.939	1	0.9	*			868.6
29	2.960	4	3.4	***			912.4
30	2.982	0	0.0				958.4
31	3.003	0	0.0				1007.
32	3.024	0	0.0				1058.
33	3.046	0	0.0				1111.
34	3.067	0	0.0				1167.
35	3.088	0	0.0				1226.
36	3.110	0	0.0				1288.
37	3.131	0	0.0				1352.
38	3.152	0	0.0				1421.
39	3.174	0	0.0				1492.
40	3.195	0	0.0				1568.
41	3.217	0	0.0				1647.

T I T L E S	PROGRAM:	PERCENTAGE CUMULATIVE FREQUENCY PLOTS
	DATA:	DAISY CREEK SC HORIZON SOIL GEOCHEMISTRY
	RUN:	
	TIME:	25/03/83 14:08:04

NAME	BA-SC
N	116
X BAR	2.7999
STD DEV	0.85470E-01



977